

1 Running title: Biological sedimentary iron oxides

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3 **A biological source of marine sedimentary iron oxides**

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## Abstract

25           The biogeochemical cycle of iron is intricately linked to numerous element cycles.  
26    Although reductive biological processes that bridge the iron cycle to other element cycles are  
27    established, little is known about microbial oxidative processes on iron cycling in sedimentary  
28    environments—resulting in the formation of iron oxides. Here, we show that a major source of  
29    sedimentary iron oxides originates from the metabolic activity of iron-oxidizing bacteria from  
30    the class Zetaproteobacteria, stimulated by burrowing animals in coastal sediments.  
31    Zetaproteobacteria were estimated to be a global total of  $10^{26}$  cells in coastal, bioturbated  
32    sediments and would equate to an annual production of approximately  $7.9 \times 10^{15}$  grams of  
33    sedimentary iron oxides—twenty-five times larger than the annual flux of iron oxides by rivers.  
34    These data suggest that iron-oxidizing Zetaproteobacteria are keystone organisms in marine  
35    sedimentary environments given their low numerical abundance; yet exert a profound impact via  
36    the production of iron oxides.

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42    Keywords: iron-oxidizing bacteria/Zetaproteobacteria/bioturbation/iron  
43    biogeochemistry/keystone organism

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## Main

45           Iron oxides are important components of coastal and continental shelf sediments, and are  
46 thought to originate primarily by river deposition (Poulton and Raiswell, 2002). Authigenic iron  
47 oxides can also be formed by the oxidation of ferrous iron [Fe(II)], which has been largely  
48 attributed to chemical oxidation of Fe(II) in pore waters in areas with significant sediment  
49 mixing and irrigation by animals—bioturbation and bioirrigation—and subsequent reaction with  
50 oxygen in anoxic sediments (Aller, 1982; Canfield, 1989). Although sedimentary chemical  
51 oxidation of iron is important under saturated oxygen conditions at neutral pH (Canfield, 1989),  
52 bioirrigated sediments contain microenvironments—formed by burrowing animals—that are well  
53 below saturation (Kristensen and Kostka, 2005) where the biological contribution to iron  
54 oxidation is quantitatively more significant (Emerson *et al.*, 2010). Under low-oxygen conditions  
55 ( $< 100 \mu\text{M O}_2$ ) and without rapid replenishment of highly reactive, poorly-crystalline iron oxides  
56 (i.e., ferrihydrite and lepidocrocite), they would be quickly exhausted by hydrogen sulfide and by  
57 bacterial iron reduction, and form iron sulfides or be released to the water column. Collectively,  
58 these findings suggest that biology is involved and important in marine sedimentary iron  
59 oxidation under ferruginous conditions commonly observed in bioturbated sediments.

60           The Zetaproteobacteria represent a class of iron-oxidizing bacteria (FeOB) that are  
61 exclusively found in marine or saline-influenced environments that contain high ferrous iron  
62 [Fe(II)] concentrations (McAllister *et al.*, 2011; McBeth *et al.*, 2013; Scott *et al.*, 2015, 2017).  
63 Coastal marine sediments can have Fe(II) pore water concentrations ranging from  $\sim 1\text{-}2,000 \mu\text{mol}$   
64  $\text{L}^{-1}$  that are capable of supporting lithoautotrophic populations of Zetaproteobacteria (Emerson,  
65 2016). Recent studies have identified Zetaproteobacteria in surface openings of benthic  
66 macrofauna in the Mediterranean Sea (Rubin-Blum *et al.*, 2014), worm burrows in sub-marine

67 groundwater discharge into sands in Delaware (McAllister *et al.*, 2015), coastal sediments in  
68 Denmark (Laufer *et al.*, 2016), and Baltic and North Sea sediments (Reyes *et al.*, 2016). These  
69 recent studies suggest that Zetaproteobacteria may play a significant role in iron oxidation in  
70 marine sediments—a quantitative estimate of their abundance is necessary to determine their  
71 biogeochemical role on a global scale.

72 We analyzed the coastal sediment microbial communities to determine the extent of  
73 Zetaproteobacteria from geographically diverse sites (n=90; Supplementary Table 1) utilizing  
74 16S rRNA gene sequencing, which highlights their importance on a global scale. Iron oxidation  
75 appears ubiquitous in the Zetaproteobacteria (Field *et al.*, 2015; Scott *et al.*, 2015; Barco *et al.*,  
76 2015), thus the 16S rRNA gene can be used to infer this specific metabolism. We also enriched  
77 for environmentally relevant iron-oxidizing bacteria from coastal sediments, which provided  
78 further metabolic evidence of the importance of iron oxidation in marine sediments. A meta-  
79 analysis of 16S rRNA gene studies revealed the extent and importance of Zetaproteobacteria on  
80 the global sedimentary iron biogeochemical cycle.

81 We identified Zetaproteobacteria in 60 % of our samples (Supplementary Table S1), and  
82 their median relative abundance was 1.1 percent of the total microbial community (range = 0.04-  
83 15 %) in worm (e.g., polychaetes) burrows in coastal marine sediments (Fig. 1a).  
84 Zetaproteobacteria were ten times less abundant in bulk sediments (Fig. 1a) with a median near  
85 zero percent (range = 0-1 %), and were statistically different from worm burrows (p-value=9.2 x  
86 10<sup>-7</sup>, Wilcoxon test). The large ranges (0.04-15 %) and non-normal distribution of  
87 Zetaproteobacteria relative abundance in worm burrows (see Supplementary Materials and  
88 Methods) was most likely a combination of differences in burrow ventilation rates and  
89 efficiencies (Kristensen and Kostka, 2005), differences in sediment physicochemical conditions

90 (Supplementary Table S1), and sampling bias (for example, residual sediment on worm  
91 burrows). Bioirrigation by benthic animals increases the extent of oxidative processes in these  
92 sediments, thus biological iron oxidation can occur at greater depths (10s of centimeters) than  
93 typical oxygen penetration of a few millimeters into coastal surface sediments. The abundance of  
94 Zetaproteobacteria at the burrow walls correlated with the concentration of pore water ferrous  
95 iron [Fe(II)] (Fig. 1b), which is their main energy source—resulting in the production of solid  
96 phase iron oxides around worm burrows (Fig. 1c). Quantitatively, highly reactive, iron oxides—  
97 operationally extractable by sodium dithionite (Poulton and Canfield, 2005)—were 3 times  
98 higher at burrow walls, which accounted for 20-40 % of the iron oxides with depth  
99 (Supplementary Figure S1 and S2). These freshly-produced iron oxides are important substrates  
100 for iron-reducing microorganisms that release Fe(II) into pore waters, and are essential to the  
101 supply of dissolved iron (dFe) to phytoplankton in the water column in coastal communities and  
102 continental shelves (Severmann *et al.*, 2010). Iron oxides are also important to the mineralization  
103 of organic matter in marine sediments by iron reducers (Canfield, 1989) substrates for early  
104 pyrite diagenesis (Berner, 1984), enhance organic matter burial (Lalonde *et al.*, 2012), and  
105 inhibit accumulation of pore water hydrogen sulfide, preventing conditions detrimental to  
106 benthic animals (Kristensen and Kostka, 2005), perhaps functioning as a local firewall against  
107 euxinic conditions(Seitaj *et al.*, 2015).

108         The relative abundance of Zetaproteobacteria in worm burrows resembles the Fe(II)  
109 concentration profile (Supplementary Figure S3), and both were at their maximum values around  
110 2-3 cm. The high Fe(II) (~40-140  $\mu$ M) and low oxygen (~20-60  $\mu$ M) conditions present in  
111 bioturbated sediment pore waters (Supplementary Figure S3) were ideal habitats for  
112 microaerophilic Zetaproteobacteria to thrive (Emerson *et al.*, 2010). Zetaproteobacteria relative

113 abundance decreased with depth in both burrows and sediments (Supplementary Fig. S3), likely  
114 due to the decrease in Fe(II) with depth and increase in hydrogen sulfide production by sulfate-  
115 reducing bacteria. Although there is oxygen in these sediments at depth, hydrogen sulfide may  
116 inhibit oxygen respiratory machinery under these conditions. The formation of iron sulfide  
117 minerals with increasing depth by biogenic sulfide may also compete with Zetaproteobacteria for  
118 access to Fe(II). Under these ferruginous settings, biotic rates of Fe(II) oxidation exceed abiotic  
119 chemical oxidation (Emerson *et al.*, 2010).

120 Two Zetaproteobacteria Operational Taxonomic Units (herein, referred to as ZetaOTUs)  
121 dominated the Zetaproteobacterial diversity in worm burrows (Supplementary Table S2). The  
122 dominant ZetaOTU across all samples was ZetaOTU14, which comprised 32 % of all ZetaOTUs  
123 (Supplementary Table S2), and is represented by four single cell amplified genomes (SAGs)  
124 from diffuse flow vent systems (Field *et al.*, 2015; Scott *et al.*, 2015, 2017). We isolated the first  
125 member of ZetaOTU14, strain CSS-1 from iron oxide surface flocculent in a laboratory  
126 bioturbation microcosm (Supplementary Figure S4). This strain grew best under low oxygen  
127 (~60  $\mu\text{M O}_2$ ) and high Fe(II) concentrations similar to those measured from sediment pore  
128 waters (Supplementary Figure S3). Strain CSS-1 produced stalks encrusted with poorly-  
129 crystalline iron oxides under laboratory conditions (Supplementary Figure S4). These iron oxides  
130 are consistent with those produced by other Zetaproteobacteria (Chan *et al.*, 2010), as well as in  
131 naturally occurring iron mats associated with hydrothermal vents, which are highly reactive, and  
132 resistant to undergoing diagenesis to more crystalline oxides (e.g., goethite) (Picard *et al.*, 2015).  
133 Single cell genomes from ZetaOTU14 representatives contained genes essential for growth on  
134 iron and low oxygen conditions (Supplementary Table S2). The second most abundant OTU was  
135 ZetaOTU9 (22 %) and is represented by two cultured isolates (*Ghiorsea bivora* strains TAG-1

136 and SV108) (Mori *et al.*, 2017), as well as 5 SAGs from deep-sea vents (Supplementary Table  
137 S2). ZetaOTU9 isolates also had genes necessary for growth on iron and low oxygen  
138 (Supplementary Table S2), and have also been shown to oxidize hydrogen, which may explain  
139 the ubiquity of this OTU in sediments and other environments (see below). There was no clear  
140 distribution of ZetaOTUs 14 and 9 with respect to depth (Supplementary Figure S3) in worm  
141 burrows and sediments—although it is likely that they are adapted for specific Fe(II) and O<sub>2</sub>  
142 concentrations, which were hypothesized for other Zetaproteobacteria (Field *et al.*, 2015).

143 We searched for Zetaproteobacterial 16S rRNA gene sequences in marine sediment  
144 datasets (Supplementary Table S3), and identified them in numerous sediments on a global scale  
145 (Figure 2). We found a pattern consistent with our samples—ZetaOTUs 14 and 9 were present  
146 and generally the most abundant ZetaOTUs in coastal and shelf sediments (Figure 2).  
147 Zetaproteobacteria relative abundance was not found to exceed one percent in other studies, as  
148 microenvironments were not considered, which are abundant in bioturbated sediments  
149 (Kristensen and Kostka, 2005). Accordingly, we hypothesize that when the abundance of  
150 Zetaproteobacteria exceeds ~0.1 % in sediments, there is active growth and iron oxidation  
151 associated with bioturbating and bioirrigating animals. We estimated a median global population  
152 size of Zetaproteobacteria to be  $1.05 \times 10^{26}$  cells (range =  $3.83 \times 10^{24}$ - $1.44 \times 10^{27}$  cells) from our  
153 measurements (Fig. 1a) and from other studies (Fig. 2) in continental shelf sediments using  
154 cellular abundance in the upper 10 centimeters—the worldwide average depth of bioturbation  
155 (Boudreau, 1998)—of continental shelf environments (total cells =  $10^{29}$ ; <150 m water depth  
156 (Kallmeyer *et al.*, 2012)). The global Zetaproteobacteria abundance estimate was then used with  
157 recent iron oxide production rate measurements from diffuse flow vents ( $\sim 1.3 \times 10^{-16}$  mol Fe cell<sup>-1</sup>  
158 hr<sup>-1</sup>) (Emerson *et al.*, 2016), which could result in the production of ~7 petagrams of iron

159 oxides per year (range = 0.1-70 petagrams per year). Recent two-dimensional, sub-millimeter  
160 Fe(II) measurements in bioturbated sediments revealed extensive Fe(II) oxidation occurring  
161 within the immediate vicinity of worm burrows and a rapid re-oxidation rate of  $3.78 \pm 1.4$  mmol  
162 Fe m<sup>-2</sup> day<sup>-1</sup> (de Chanvalon *et al.*, 2017). These chemical rate measurements combined with an  
163 estimate of the global volume of bioturbated coastal sediments 10 cm deep ( $\sim 2.1 \times 10^{13}$  m<sup>3</sup>)  
164 (Teal *et al.*, 2008) would equate to an annual production of  $1.6 \pm 1.1$  petagrams of iron oxides.  
165 These two independent estimations of iron oxide production rates are well within the range of  
166 one another. Based on these estimates, the annual biological oxidation of iron in sediments—  
167 forming iron oxides—could exceed the annual flux of iron oxides from rivers to coastal  
168 sediments (Poulton and Raiswell, 2002) up to a factor of twenty-five.

169         Zetaproteobacteria exert a profound impact on global sedimentary biogeochemistry via  
170 the production of biogenic, highly-reactive iron oxides despite their low global abundance ( $\sim 0.11$   
171 %)—effectively functioning as keystone organisms in coastal sediments stimulated by burrowing  
172 animals. Zetaproteobacteria contribute significantly to the rapid rates of Fe(II) re-oxidation  
173 measured and observed in coastal sediments (Chanvalon *et al.*, 2017). Climate change outcomes  
174 such as coastal hypoxia may have positive or negative effects on the sedimentary iron  
175 biogeochemical cycle—either stimulating microaerobic bacterial iron oxidation resulting in an  
176 increase in iron oxide production, thus enhancing dFe release or inhibiting oxidation by the  
177 increase in hydrogen sulfide production, precipitating Fe as iron sulfides. The result of an  
178 increase or decrease in dFe flux would be enhanced or reduced primary productivity by  
179 phytoplankton, respectively. Thus, sedimentary iron oxide formation by Zetaproteobacteria may  
180 have a direct impact on important water column processes such as carbon and nitrogen fixation.  
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## References

- 194 Aller RC. (1982). The Effects of Macrobenthos on Chemical Properties of Marine Sediment and  
195 Overlying Water. In: McCall PL, Tevesz MJS (eds). *Animal-Sediment Relations: The Biogenic  
196 Alteration of Sediments*. Springer US: Boston, MA, pp 53–102.
- 197 Barco RA, Emerson D, Sylvan JB, Orcutt BN, Jacobson Meyers ME, Ramírez GA, *et al.* (2015).  
198 New insight into microbial iron oxidation as revealed by the proteomic profile of an obligate  
199 iron-oxidizing chemolithoautotroph. *Appl Environ Microbiol* **81**: 5927–5937.
- 200 Berner A. (1984). Sedimentary pyrite formation: An update. *Geochim Cosmochim Acta*
- 201 Boudreau BP. (1998). Mean mixed depth of sediments: The wherefore and the why. *Limnol  
202 Oceanogr* **43**: 524–526.
- 203 Canfield DE. (1989). Reactive iron in maine sediments. *Geochim Cosmochim Acta* **53**: 619–632.
- 204 Chan CS, Fakra SC, Emerson D, Fleming EJ, Edwards KJ. (2010). Lithotrophic iron-oxidizing

205 bacteria produce organic stalks to control mineral growth: implications for biosignature  
206 formation. *ISME J* **5**: 717–727.

207 de Chanvalon AT, Metzger E, Mouret A, Knoery J, Geslin E, Meysman FJR. (2017). Two  
208 dimensional mapping of iron release in marine sediments at submillimetre scale. *Mar Chem.*  
209 **191**: 34-49.

210 Emerson D. (2016). The irony of iron - Biogenic iron oxides as an iron source to the ocean.  
211 *Front Microbiol* **6**: 1–6.

212 Emerson D, Fleming EJ, McBeth JM. (2010). Iron-oxidizing bacteria: an environmental and  
213 genomic perspective. *Annu Rev Microbiol* **64**: 561–583.

214 Emerson D, Scott JJ, Leavitt AH, Fleming E, Moyer CL. (2016). In situ estimates of iron-  
215 oxidation and accretion rates for iron-oxidizing bacterial mats at Loihi Seamount. *bioRxiv*.

216 Field EK, Sczyrba A, Lyman AE, Harris CC, Woyke T, Stepanauskas R, *et al.* (2015). Genomic  
217 insights into the uncultivated marine Zetaproteobacteria at Loihi Seamount. *ISME J* **9**: 857–870.

218 Kallmeyer J, Pockalny R, Adhikari RR, Smith DC, D’Hondt S. (2012). Global distribution of  
219 microbial abundance and biomass in subseafloor sediment. *Proc Natl Acad Sci* **109**: 16213–  
220 16216.

221 Kristensen E, Kostka JE. (2005). Macrofaunal Burrows and Irrigation in Marine Sediment:  
222 Microbiological and Biogeochemical Interactions. In: Kristensen E (ed) *Interact Between*  
223 *Macro- Microorg Mar Sediments*. American Geophysical Union Washington, DC, pp. 125–157.

224 Lalonde K, Mucci A, Ouellet A, Gelin Y. (2012). Preservation of organic matter in sediments  
225 promoted by iron. *Nature* **483**: 198–200.

226 Laufer K, Nordhoff M, Schmidt C, Behrens S, Jørgensen BB, Kappler A. (2016). Co-existence  
227 of microaerophilic, nitrate-reducing, and phototrophic Fe(II)-oxidizers and Fe(III)-reducers in

- 228 coastal marine sediment. *Appl Environ Microbiol* **82**: 1433–1447.
- 229 McAllister SM, Barnett JM, Heiss JW, Findlay AJ, MacDonald DJ, Dow CL, *et al.* (2015).  
230 Dynamic hydrologic and biogeochemical processes drive microbially enhanced iron and sulfur  
231 cycling within the intertidal mixing zone of a beach aquifer. *Limnol Oceanogr* **60**: 329–345.
- 232 McAllister SM, Davis RE, McBeth JM, Tebo BM, Emerson D, Moyer CL. (2011). Biodiversity  
233 and emerging biogeography of the neutrophilic iron-oxidizing Zetaproteobacteria. *Appl Environ*  
234 *Microbiol* **77**: 5445–5457.
- 235 McBeth JM, Fleming EJ, Emerson D. (2013). The transition from freshwater to marine iron-  
236 oxidizing bacterial lineages along a salinity gradient on the Sheepscot River, Maine, USA.  
237 *Environ Microbiol Rep* **5**: 453–463.
- 238 Mori JF, Scott JJ, Hager KW, Moyer CL, Küsel K, Emerson D (2017). Physiological and  
239 ecological implications of an iron- and hydrogen-oxidizing member of the Zetaproteobacteria,  
240 *Ghiorsea bivora*, gen. nov. sp. nov. Accepted.
- 241 Picard A, Kappler A, Schmid G, Quaroni L, Obst M. (2015). Experimental diagenesis of organo-  
242 mineral structures formed by microaerophilic Fe(II)-oxidizing bacteria. *Nat Commun* **6**: 6277.
- 243 Poulton SW, Canfield DE. (2005). Development of a sequential extraction procedure for iron:  
244 Implications for iron partitioning in continentally derived particulates. *Chem Geol* **214**: 209–221.
- 245 Poulton SW, Raiswell R. (2002). The low-temperature geochemical cycle of iron: From  
246 continental fluxes to marine sediment deposition. *Am J Sci* **302**: 774–805.
- 247 Reyes C, Dellwig O, Dahnke K, Gehre M, Noriega-Ortega BE, Bottcher ME, *et al.* (2016).  
248 Bacterial communities potentially involved in iron-cycling in Baltic Sea and North Sea  
249 sediments revealed by pyrosequencing. *FEMS Microbiol Ecol* **92**.
- 250 Rubin-Blum M, Antler G, Tsadok R, Shemesh E, Austin JA, Coleman DF, *et al.* (2014). First

251 evidence for the presence of iron oxidizing zetaproteobacteria at the levantine continental  
252 margins. *PLoS One* **9**: 1–10.

253 Scott JJ, Breier JA, Luther GW, Emerson D. (2015). Microbial iron mats at the mid-atlantic ridge  
254 and evidence that zetaproteobacteria may be restricted to iron-oxidizing marine systems. *PLoS*  
255 *One* **10**: 1–19.

256 Scott JJ, Glazer BT, Emerson D. (2017). Bringing microbial diversity into focus: high-resolution  
257 analysis of iron mats from the Lō‘ihi Seamount. *Environ Microbiol* **19**: 301-316.

258 Seitaj D, Schauer R, Sulu-Gambari F, Hidalgo-Martinez S, Malkin SY, Burdorf LDW, *et al.*  
259 (2015). Cable bacteria generate a firewall against euxinia in seasonally hypoxic basins. *Proc Natl*  
260 *Acad Sci U S A* **112**: 13278-13283.

261 Severmann S, McManus J, Berelson WM, Hammond DE. (2010). The continental shelf benthic  
262 iron flux and its isotope composition. *Geochim Cosmochim Acta* **74**: 3984–4004.

263 Teal LR, Bulling MT, Parker ER, Solan M. (2008). Global patterns of bioturbation intensity and  
264 mixed depth of marine soft sediments. *Aquat Biol* **2**: 207–218.

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266

### Figure Legends

267 **Figure 1.** Boxplots of the relative abundance of Zetaproteobacteria (**a**) in iron-oxide lined worm  
268 burrow walls (n=29) and surrounding sediments (n=61). Notches are representative of 95 %  
269 confidence interval and the medians (solid black lines) between worm burrows and sediments  
270 (1.1 % and 0 %, respectively) are statistically different (p-value= $9.2 \times 10^{-7}$ , Wilcoxon test).  
271 Filled circles represent individual data points and open circles indicate outliers.

272 Zetaproteobacteria relative abundance (%) as a function of pore water ferrous iron [Fe(II)]  
273 concentration ( $\mu\text{mol L}^{-1}$ ) (**b**) from worm burrows (blue circles, fitted blue line, orange fill = 95 %

274 confidence interval) and sediments (black circles, fitted black line, grey fill = 95 % confidence  
275 interval) (see Methods for details on line fits). Characteristic iron oxide lined worm burrow walls  
276 (c) from “The Eddy”, Sheepscoot River, Maine, USA (image from 27 August 2015). Burrow  
277 walls are likely created by the polychaete, *Nereis diversicolor* or hemichordate, *Saccoglossus*  
278 *kowalevskii*, which are both common to these intertidal sediments in Maine.

279 **Figure 2.** Global distribution of Zetaproteobacteria in marine sediments (circles) and non-  
280 sediment sites (triangles) such as hydrothermal vents. The relative abundance of  
281 Zetaproteobacteria in sediments from other 16S rRNA gene studies was never above 1 % and  
282 was typically within the range measured from bulk marine sediments from Maine. Sequences are  
283 from numerous studies (Supplementary Table S3) that include Sanger, 454, and Illumina  
284 sequencing technologies.

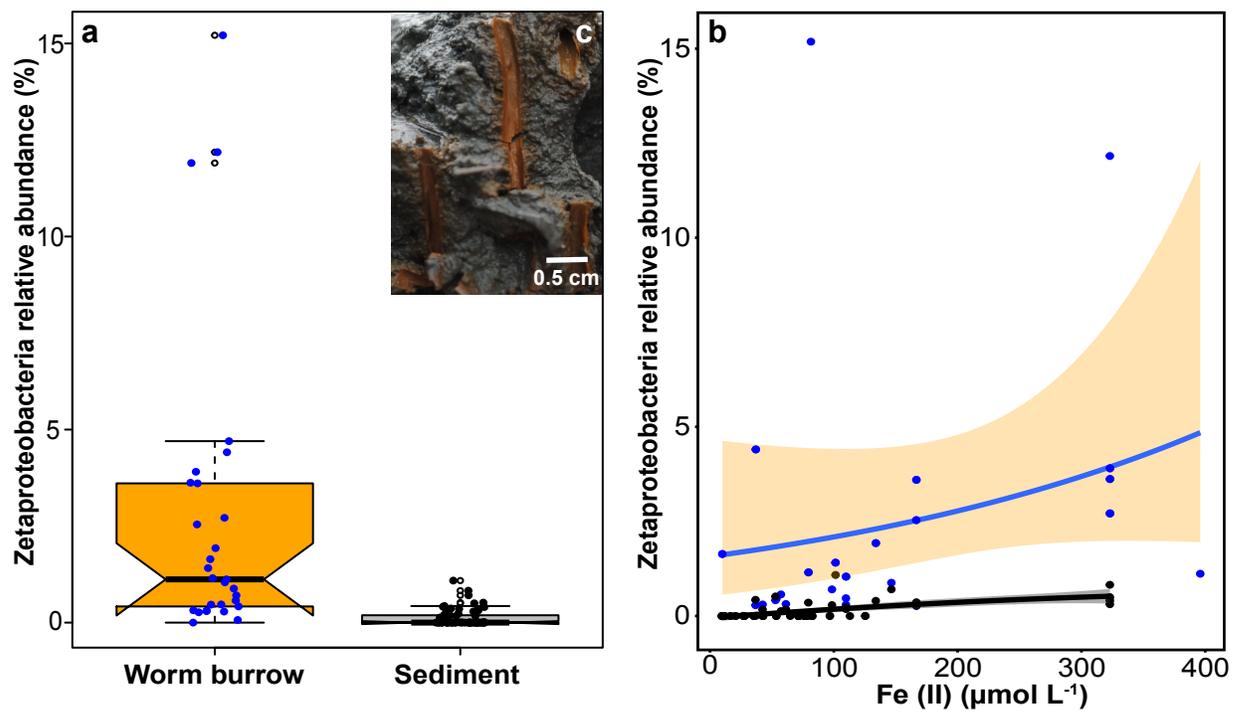
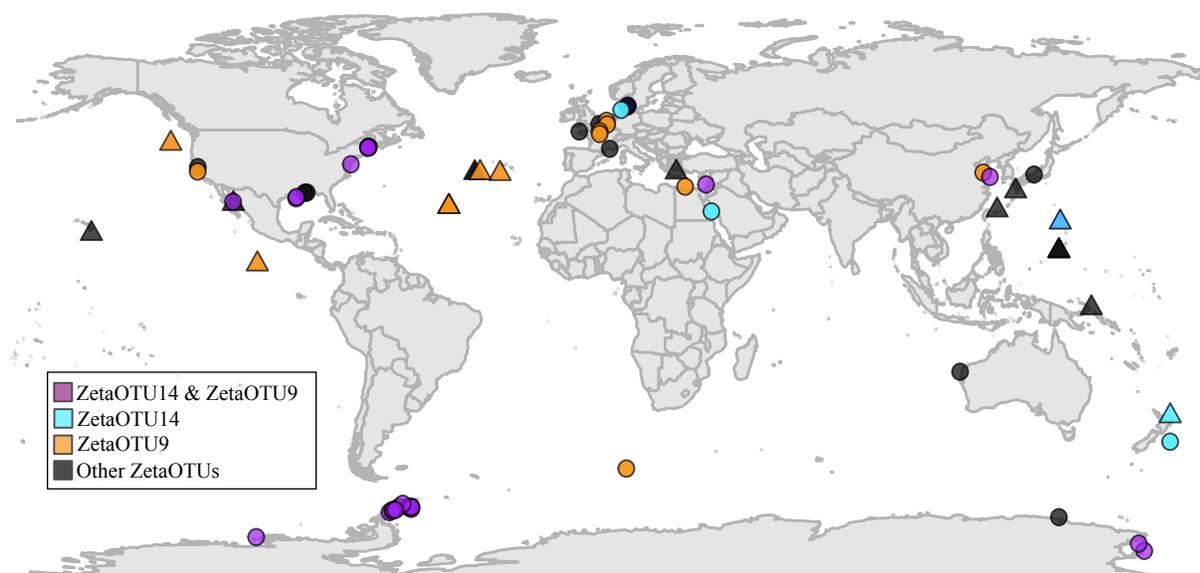


Figure 1



**Figure 2**