

1 **Molybdenum-based diazotrophy in a *Sphagnum* peatland in northern Minnesota**

2 Melissa J. Warren^{1*}, Xueju Lin^{2**}, John C. Gaby^{2***}, Cecilia B. Kretz^{1****}, Max Kolton², Peter L.
3 Morton³, Jennifer Pett-Ridge⁴, David J. Weston⁵, Christopher W. Schadt⁵, Joel E. Kostka^{1,2},
4 Jennifer B. Glass^{1,2#}

5
6 ¹School of Earth and Atmospheric Sciences, Georgia Institute of Technology, Atlanta, GA, USA

7 ²School of Biology, Georgia Institute of Technology, Atlanta, GA, USA

8 ³Department of Earth, Ocean and Atmospheric Science, Florida State University, Tallahassee,
9 Florida, USA

10 ⁴Physical and Life Sciences Directorate, Lawrence Livermore National Laboratory, Livermore,
11 CA, USA

12 ⁵Biosciences Division, Oak Ridge National Laboratory, Oak Ridge, TN, USA

13

14 Now at: *CH2M, Atlanta, GA, USA; **InstantLabs, Baltimore, MD, USA; ***Faculty of

15 Chemistry, Biotechnology and Food Science, Norwegian University of Life Sciences, Oslo,

16 Norway; ****Division of Bacterial Diseases, National Center for Immunization and Respiratory

17 Diseases, Centers for Disease Control and Prevention, Atlanta, GA, USA

18 **#Corresponding author:** Jennifer B. Glass, 404-894-3942; jennifer.glass@eas.gatech.edu

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20 **Running Title:** Nitrogen fixation in a *Sphagnum* peat bog

21 **Abstract**

22 Microbial N₂ fixation (diazotrophy) represents an important nitrogen source to oligotrophic
23 peatland ecosystems, which are important sinks for atmospheric CO₂ and susceptible to changing
24 climate. The objectives of this study were: (i) to determine the active microbial group and type of
25 nitrogenase mediating diazotrophy in a ombrotrophic *Sphagnum*-dominated peat bog (the S1
26 peat bog, Marcell Experimental Forest, Minnesota, USA); and (ii) to determine the effect of
27 environmental parameters (light, O₂, CO₂, CH₄) on potential rates of diazotrophy measured by
28 acetylene (C₂H₂) reduction and ¹⁵N₂ incorporation. Molecular analysis of metabolically active
29 microbial communities suggested that diazotrophy in surface peat was primarily mediated by
30 *Alphaproteobacteria* (*Bradyrhizobiaceae* and *Beijerinckiaceae*). Despite higher dissolved
31 vanadium (V; 11 nM) than molybdenum (Mo; 3 nM) in surface peat, a combination of
32 metagenomic, amplicon sequencing and activity measurements indicated that Mo-containing
33 nitrogenases dominate over the V-containing form. Acetylene reduction was only detected in
34 surface peat exposed to light, with the highest rates observed in peat collected from hollows with
35 the highest water content. Incorporation of ¹⁵N₂ was suppressed 90% by O₂ and 55% by C₂H₂,
36 and was unaffected by CH₄ and CO₂ amendments. These results suggest that peatland
37 diazotrophy is mediated by a combination of C₂H₂-sensitive and C₂H₂-insensitive microbes that
38 are more active at low O₂ and show similar activity at high and low CH₄.

39 **Importance**

40 Previous studies indicate that diazotrophy provides an important nitrogen source and is linked to
41 methanotrophy in *Sphagnum*-dominated peatlands. However, the environmental controls and
42 enzymatic pathways of peatland diazotrophy, as well as the metabolically active microbial
43 populations that catalyze this process remain in question. Our findings indicate that oxygen
44 levels and photosynthetic activity override low nutrient availability in limiting diazotrophy, and
45 that members of the *Alphaproteobacteria* (*Rhizobiales*) catalyze this process at the bog surface
46 using the molybdenum-based form of the nitrogenase enzyme.

47

48 **Introduction**

49 High-latitude peatlands store approximately one-third of global soil carbon and may pose a
50 climatic threat if rising global temperatures accelerate the release of this stored carbon in gaseous
51 forms, as either carbon dioxide or methane (35, 64, 106). Mineral-poor (ombrotrophic) peatlands
52 receive most of their nutrient inputs from atmospheric deposition and contain *Sphagnum* moss as
53 their primary plant cover (12, 64). The peatmoss *Sphagnum* is a keystone genus in these
54 ecosystems and is responsible for much of the primary production and recalcitrant dead organic
55 matter (13, 103). *Sphagnum* mosses also host complex microbiomes (9, 58, 78, 79), including
56 N₂-fixers (diazotrophs) that are significant nitrogen sources to peatland ecosystems (5).

57 Despite decades of research, there is still much debate on the identity of the dominant
58 diazotrophs in ombrotrophic peatlands. Early work implicated *Cyanobacteria* (2, 36, 37) or
59 heterotrophic bacteria (88) based primarily on microscopic studies, while more recent molecular
60 analyses argue for the importance of methanotrophic *Beijerinckiaceae* (19) as major diazotrophs
61 in *Sphagnum* peat bogs (8, 22, 44, 99). Possible contributions from other potential diazotrophs,
62 such as strictly anaerobic methanogenic *Euryarchaeota*, remain unknown. However, it is quite
63 possible that diverse diazotrophs exist within defined niches of peatland environments (60).

64 Diazotrophy is catalyzed by the nitrogenase metalloenzyme, a complex of three subunits
65 (H, D and K) that contains abundant iron as Fe-S clusters. This enzyme is extremely O₂ sensitive
66 (105), and must be protected from exposure to O₂ for diazotrophy to occur (25). The most
67 common form of nitrogenase, encoded by *nif* genes, contains molybdenum (Mo) as its cofactor.
68 When Mo is scarce, some species of Bacteria and Archaea express nitrogenases containing
69 vanadium (V; *vnf* genes) or iron (Fe; *anf* genes) in place of Mo, but these “alternative”
70 nitrogenases are less efficient than the Mo form (74, 101). The most conserved nitrogenase gene,

71 *nifH* (77), has become the marker gene of choice for environmental diazotrophy (28, 29, 107).
72 Phylogenetic studies show five *nifH* clusters: aerobic bacteria (cluster I); alternative nitrogenases
73 (cluster II); anaerobic bacteria and archaea (cluster III); uncharacterized sequences (cluster IV),
74 and paralogs related to chlorophyll biosynthesis (cluster V; 85). Because *vnfH* and *anfH* genes in
75 cluster II cannot be differentiated by sequence alone, the D-subunit (*nifD/vnfD/anfD*) has
76 become the preferred marker gene for studies of alternative nitrogenases (4). Consistent with
77 higher concentrations of V than Mo in most rocks (102), microbes from diverse soils have been
78 shown to contain *vnfD* genes (4, 7, 18, 45, 70, 73). Given that oligotrophic conditions dominate
79 in peatlands, trace metals may limit diazotrophy. However, little is known about trace metal
80 availability and the role of alternative nitrogenase pathways in ombrotrophic peatlands.

81 Similarly, methane monooxygenase (MMO, the enzyme that catalyzes the first step of
82 methane oxidation) occurs in particulate (copper (Cu)-containing pMMO) and soluble (Fe-
83 containing sMMO) forms. While pMMO has more specific substrate requirements, pathways
84 that employ sMMO can use a wider range of compounds (15). Both forms of MMO are inhibited
85 by acetylene (C₂H₂) (14, 82). In organisms with both sets of genes, pMMO is expressed when Cu
86 is abundant, whereas Cu limitation induces sMMO expression (90). The dominant peatland
87 methanotrophs in *Alphaproteobacteria* and *Gammaproteobacteria* tend to possess both MMOs
88 (11, 24, 39, 52, 68), although *Methylocella* and *Methyloferula* species containing solely sMMO
89 have been isolated from peat bogs (20, 21, 100). While most studies have primarily targeted the
90 *pmoA* gene (24, 52), *mmoX* genes and transcripts have also been reported in peatlands (63, 68,
91 84), raising questions about the relative importance of each form for peatland methane oxidation.

92 The acetylene reduction assay (ARA) is commonly used as a proxy for diazotroph
93 activity (41, 42). This assay is effective for capturing the potential activity of diazotrophs that are

94 not inhibited by C₂H₂, such as *Cyanobacteria* and non-methanotrophic *Proteobacteria* (e.g.
95 *Bradyrhizobiaceae*) (50). However, a number of functional guilds of microorganisms, including
96 methanotrophs, methanogens, sulfate reducers, and nitrifiers, are inhibited by C₂H₂ (16, 49, 80,
97 81, 92, 94). If these or other C₂H₂-sensitive microbes perform diazotrophy and/or provide
98 substrates to other diazotrophs (see Fulweiler, et al. (27)), ARA may underestimate diazotrophy
99 in that system. Thus, recent studies have shifted to tracking diazotrophy by incorporation of the
100 stable isotope tracer, ¹⁵N₂ (55, 60, 61, 99).

101 In this study of the S1 peat bog at the Marcell Experimental Forest in northern
102 Minnesota, USA, dissolved macro- (NH₄⁺, NO₃⁻, PO₄³⁻) and micro- (Fe, Cu, V, Mo) nutrients
103 were profiled along with the community composition and abundance of diazotrophic
104 microorganisms. We also performed separate laboratory incubation experiments to measure
105 potential rates of ARA and ¹⁵N₂ incorporation in order to: (a) assess environmental controls
106 (light, O₂, CH₄) on diazotrophy; (b) quantify the effect of C₂H₂ on rates of diazotrophy and
107 methanotrophy; and (c) search for diagnostic markers for alternative nitrogenase activity such as
108 a low conversion factor of ARA to ¹⁵N₂-incorporation (4) and C₂H₂ reduction to ethane (23).
109 Finally, we make recommendations on universal *nifH* primers for amplicon sequencing and
110 quantitative PCR based on our findings.

111

112 **Materials and Methods**

113 ***Site description and sample collection.*** Samples were collected from the S1 (black spruce-
114 *Sphagnum* spp.) peat bog at Marcell Experimental Forest (MEF; 47°30.476' N; 93°27.162' W),
115 the site of the DOE SPRUCE (Spruce and Peatland Responses Under Climatic and
116 Environmental Change) experiment in northern Minnesota, USA (40). The S1 bog is

117 ombrotrophic and acidic (average pH 3.5-4; 57, 89). Over the summer months, the water table is
118 ± 5 cm from the hollow surface (38, 89). Dissolved O₂ levels decrease to below detection (~20
119 ppb) within the top 5 cm of the bog. Three locations were sampled along S1 bog transect 3 (T3)
120 at near, middle and far sites (see Lin, et al. (69) for further details). Surface (0-10 cm depth) peat
121 was collected from hollows dominated by a mixture of *Sphagnum fallax* and *S. angustifolium*,
122 and hummocks dominated by *S. magellanicum*. Peat depth cores (0-200 cm) were sampled from
123 hollows where the water level reached the surface of the *Sphagnum* layer.

124 **Macronutrients.** Peat porewater was collected using piezometers from 0, 10, 25, 50, 75,
125 150 and 200 cm depth. Piezometers were recharged the same day as collection, and porewater
126 was pumped to the surface, filtered through sterile 0.2 μm polyethersulfone membrane filters,
127 and stored frozen until analysis. Nitrate (NO₃⁻) and nitrite (NO₂⁻) were analyzed using the
128 spectrophotometric assay as described by García-Robledo, et al. (31). Ammonium (NH₄⁺)
129 concentrations were determined with the indophenol blue assay (95). Phosphate concentrations
130 were measured with the molybdate-antimony ascorbic acid colorimetric assay (75).

131 **Micronutrients.** Peat porewater was collected from two locations in the S1 bog from
132 cores at 0-30 cm, 30-50 cm, and 100-150 cm depths by filtration through 0.15 μm Rhizon soil
133 samplers (Rhizosphere Research Products). All plastics were washed with HCl prior to sampling;
134 Rhizon soil samplers were cleaned by pumping 10 mL of 1 N HCl through them, followed by a
135 rinsing with ultrapure water until the pH returned to neutral (~100 mL/filter). After collection,
136 samples were acidified to 0.32 M HNO₃ (Fisher Optima) and analyzed using a Thermo
137 ELEMENT 2 HR-ICP-MS (National High Magnetic Field Laboratory, Florida State University).
138 Initial analyses resulted in frequent clogging of the nebulizer, likely due to abundant dissolved
139 organic carbon. Therefore, samples were diluted 1:10 to minimize interruptions from nebulizer

140 clogs. Concentrations were quantified with a 7-point external calibration using standards
141 prepared in 0.32 N HNO₃ from a multi-element standard mix (High Purity Standards).

142 In order to generate an organic-free sample matrix suitable for ICP-MS analysis without
143 contaminating or diluting the sample, subsequent samples were digested as follows: 1 mL
144 aliquots of the porewater samples were heated in 15-mL Teflon beakers (Savillex) with 1 mL of
145 16 N HNO₃ (Ultrex II, JT Baker) and 100 µL of 30% H₂O₂ (Ultrex II, JT Baker) for 36 h at
146 230°C in a trace metal clean, polypropylene exhaust hood. The HNO₃/H₂O₂ mixture oxidizes any
147 DOM to CO₂, but the resulting matrix is too acidic for direct ICP-MS introduction. Therefore,
148 samples were evaporated to near dryness and resuspended in a 0.32 N HNO₃ matrix suitable for
149 ICP-MS analysis, and analyzed by ELEMENT2 ICP-MS along with parallel blank solutions.

150 ***Quantification and sequencing of gene and transcript amplicons.*** Peat was frozen on
151 dry ice at the field site in July 2013, or in liquid N₂ after 7 day incubations at 25°C in the light
152 under degassed (80% N₂ + 20% CO₂) headspace with 1% C₂H₂, with or without 1% CH₄ for June
153 2014 incubations (see ARA section). DNA and RNA were extracted with MoBio PowerSoil
154 DNA and total RNA extraction kits, respectively, as described in Lin, et al. (68). RNA was
155 cleaned with a TURBO DNA-free kit (Ambion). Nucleic acid purity was analyzed for 260/280
156 absorbance ratio (1.8-2.0) on a NanoDrop spectrophotometer. cDNA was synthesized using the
157 GoScript reverse transcription system (Promega) according to the manufacturer's protocol.

158 Plasmid standards for qPCR were constructed according to Lin, et al. (66). Primer pairs
159 are given in Table 1. The gene fragments of *nifH*, *pmoA* and *mcrA* for constructing plasmid
160 standards for qPCR were amplified from genomic DNA of *Rhodobacter sphaeroides*,
161 *Methylococcus capsulatus* Bath, and S1 peat bog peat soil, respectively. To prepare cDNA
162 standards, plasmid DNA with a positive gene insert was linearized with NcoI restriction enzyme

163 following the manufacturer's protocol (Promega), and purified by MinElute PCR purification kit
164 (Qiagen). RNA was synthesized from the linearized plasmid DNA using the Riboprobe *in vitro*
165 transcription system according to the manufacturer's protocol (Promega), followed by cDNA
166 synthesis using the GoScript reverse transcription system (Promega) according to the
167 manufacturer's protocol.

168 The abundance of functional gene transcripts were quantified in samples run in duplicate
169 on a StepOnePlus Real-Time PCR System (ABI) using Power SYBR Green PCR master mix.
170 Reaction mixtures were 20- μ L reaction mixtures with 2- μ L of template cDNA (10-100 ng/ μ L)
171 added to 10 μ L of SYBR green master mix, 0.5-1.6 μ L of each forward and reverse primer (0.3-
172 0.8 μ M final concentration; Table 1), and 4.8-6.5 μ L of PCR-grade water. Samples were run
173 against a cDNA standard curve (10^1 to 10^7 copies of plasmid gene fragment) on a StepOnePlus
174 qPCR instrument with 96 wells with an initial denaturation step of 2-5 min at 95°C and 40 cycles
175 of denaturation at 95°C for 15-30 s, annealing at 55-64°C for 30-45 s, extension at 72°C for 30-
176 45 s, and data acquisition at 83-86.5°C for 16-30 s. To minimize the effects of inhibitors in
177 assays, peat DNA was diluted to 1/40 of original concentrations, and duplicate 20- μ L reaction
178 mixtures, each containing 2- μ L of diluted DNA, were run for each sample. Functional gene and
179 transcript copy numbers were normalized to dry weight of peat, or 16S rRNA transcript copies
180 for incubation samples. Amplicons were sent to the University of Illinois at Chicago for DNA
181 sequencing using a 454 platform. Raw sequences were demultiplexed, trimmed, and quality
182 filtered in CLCbio. The phylogeny of *vnfD/anfD* sequences was inferred using the Maximum
183 Likelihood method based on the Kimura 2-parameter model in MEGA5 (96).

184 ***Acetylene reduction and methane consumption rates.*** Samples of bulk peat (*Sphagnum*
185 spp. and surrounding soil) were collected from 0-10 and 10-30 cm depths in September 2013,

186 April 2014, June 2014, September 2014 and August 2015, and stored at 4°C until the start of
187 laboratory incubations. Samples from 0-10 cm depth were gently homogenized so as not to
188 rupture *Sphagnum* spp. tissues, while peat samples from 10-30 cm depth were fully
189 homogenized. For each sample, 5 g of bulk peat were placed in 70 mL glass serum bottles,
190 stoppered with black bromobutyl stoppers (Geo-Microbial Technologies, pre-treated by boiling
191 3x in 0.1 M NaOH) and sealed with an aluminum crimp seal. Headspaces were oxic (room air,
192 80% N₂ + 20% O₂) or degassed (100% N₂ or 80% N₂ + 20% CO₂) with or without 1% C₂H₂ or
193 1% CH₄. Treatments were incubated for one week at 25°C in the light or dark. A gas
194 chromatograph with a flame ionization detector (SRI Instruments) equipped with a HayeSep N
195 column was used to quantify CH₄, C₂H₂ and C₂H₄. Samples were measured for C₂H₄ production
196 daily until C₂H₄ production was linear (~7 days). Controls not amended with C₂H₂ did not
197 produce ethylene (C₂H₄). Incubations of hollow peat from June 2014 incubated in oxic
198 headspace with and without 1% C₂H₂ were also monitored for consumption of 1% CH₄.
199 Statistical analysis was performed with JMP Pro (v. 12.1.0) using the Tukey-Kramer HSD
200 comparison of all means.

201 ¹⁵N₂ *incorporation rates*. In September 2014, samples were quantified for N₂ fixation
202 rates by ¹⁵N₂ incorporation in parallel with ARA measurements. Incubations were set up as
203 described above and supplemented with 7 mL of 98% ¹⁵N₂ (Cambridge Isotope Laboratories,
204 Tewksbury, MA, USA). After 7 days, samples were dried at 80°C, homogenized into a fine
205 powder, and analyzed for N content and δ¹⁵N by isotope-ratio mass spectrometry (IRMS) with a
206 MICRO cube elemental analyzer and IsoPrime100 IRMS (Elementar) at the University of
207 California, Berkeley, corrected relative to National Institute of Standards and Technology (NIST,
208 Gaithersburg, MD, USA) standards.

209 **Metagenomic analyses.** Metagenomes were generated in a previous study (67).
210 Diazotrophic and methanotrophic pathways were investigated using the following bioinformatics
211 approaches. Briefly, Illumina reads were filtered by quality (Phred33 score threshold of Q25)
212 using Trim Galore (Babraham Bioinformatics) and a minimum sequence length cut off of 100
213 bp. The sequences were then queried using RAPSearch2 (109) against the NCBI-nr database of
214 non-redundant protein sequences as of November 2013. Sequences with a bit-score of 50 and
215 higher were retained to determine the total number of functional genes for normalization across
216 the different samples. The taxonomic composition of protein-coding sequences was determined
217 based on the taxonomic annotation of each gene according to the NCBI-nr taxonomy in
218 MEGAN5 (47) (min score: 50; max expected: 0.01; top percent: 10; min complexity: 0.3).

219 To classify sequences by nitrogenase cluster type, genes were analyzed using BLASTX
220 (e-value 0.1; bit-score 50) versus a custom *nifH* database that includes a phylogenetic tree to
221 distinguish the principal clusters (I, V, III) in the *nifH* phylogeny, as well as paralogous cluster
222 IV, *nifH*-like sequences (28). *nifH* genes from the four clusters were normalized to total protein-
223 coding genes from RAPSearch2 output sequences. The relative abundance of particulate (*pmoA*)
224 vs. soluble (*mmoX*) methane monooxygenase was based on previous analyses in Lin, et al. (68).

225 **Nucleotide sequence accession numbers.** Metagenomes were reported in a previous
226 study (67) and deposited in BioProject PRJNA382698 (SAMN06712535-06712540). *pmoA*
227 cDNA amplicons were reported in a previous study (24) and deposited in BioProject
228 PRJNA311735. *nifH*, *mcrA*, *nifD*, and *vnfD/anfD* cDNA amplicons were deposited in
229 BioProjects PRJNA382268, PRJNA382282, PRJNA382288 and PRJNA382295, respectively.

230 **Results**

231 **Macro- and micro-nutrient profiles.** In S1 bog hollows, NH_4^+ was $\sim 2 \mu\text{M}$ from the surface to 25
232 cm depth, and increased at greater depths (Fig. 1; Fig. S1a). Nitrate was $<1 \mu\text{M}$ in surface peat
233 and decreased with depth (Fig. S1b). Phosphate was $<0.1 \mu\text{M}$ from the surface to 25 cm depth,
234 and then increased with depth (Fig. S1c). For the metal pairs of greatest interest to this study, V
235 ($5\text{-}21 \text{ nM}$) was consistently more abundant than Mo ($1\text{-}7 \text{ nM}$; Fig. 3), and Fe ($7\text{-}35 \mu\text{M}$) was
236 three orders of magnitude higher than Cu ($7\text{-}38 \text{ nM}$; Fig. S3) at all three depth intervals (0-30,
237 30-50 and 100-150 cm), and at three sampling dates (September 2014, June 2015, and September
238 2015; data shown for June 2015 in Figs. 3 and S2). Other trace nutrients were in the nano- to
239 micro-molar range: Co ($5\text{-}20 \text{ nM}$), Ni ($10\text{-}80 \text{ nM}$), Zn ($50\text{-}250 \text{ nM}$) and Mn ($60\text{-}2220 \text{ nM}$; Table
240 S1). Macro- and micro-nutrient profiles were essentially identical in profiles from Zim bog,
241 another ombrotrophic bog $\sim 80 \text{ km}$ southeast of the S1 bog, sampled in September 2014, with the
242 exception of lower Fe at the surface (data not shown).

243 **Nitrogenase expression and phylogeny.** With the polF/polR primer pair, we measured
244 1.2×10^7 copies g^{-1} *nifH* genes at 1 and 20 cm, and 0.2×10^7 copies g^{-1} at 30 and 75 cm; *nifH*
245 transcripts (12.2×10^7 copies g^{-1}) were only detected at 1 cm (10:1 transcript: gene ratio), and not
246 at deeper depths (Fig. 1; Fig. S3a). Sequencing of cDNA from surface peat amplified with
247 polF/polR (*nifH*) and nifD820F/nifD1331R (*nifD*) primers showed that the majority of
248 nitrogenase transcripts belonged to cluster I (*Alphaproteobacteria*), with *Beijerinckiaceae*
249 dominating *nifH* sequences and *Bradyrhizobiaceae* dominating *nifD* sequences (Fig. 2).
250 Additional alphaproteobacterial *nif* transcripts matched to *Rhodospirillaceae*, *Rhizobiaceae*,
251 *Rhodobacteraceae*, *Methylocystaceae* and *Xanthobacteraceae* (Fig. 2). *Gammaproteobacteria*,
252 *Cyanobacteria* (*Oscillatoriothycideae*) and *Nitrospira* were also observed at lower abundance in
253 cDNA amplicon libraries (data not shown).

254 In metagenomes, *nifH* genes were roughly equally distributed between Mo-dependent
255 clusters I and III, and cluster IV/V paralogs (Fig. 2; Table S2). Sequences from Cluster II
256 (alternative nitrogenases, *vnfH/anfH*) were scarce at all depths (<5% overall); two *vnfD* genes
257 from metagenomes showed phylogenetic similarity to those from soil *Proteobacteria* (Fig. S4).
258 Attempts to amplify *vnfD/anfD* from cDNA yielded few reads; those recovered were most
259 similar to *Alphaproteobacteria anfD* from *Rhodopseudomonas* species (Fig. S4).

260 ***Methane-related gene expression and phylogeny.*** Like *nifH*, particulate methane
261 monooxygenase (*pmoA*) and methyl coenzyme M reductase (*mcrA*) transcripts showed the
262 highest abundance in surface peat (Fig. S3b,c). Surface *pmoA* transcripts mapped to
263 *Methylocystaceae* (75%) and *Methylococcaceae* (25%). Surface *mcrA* transcripts mapped to
264 *Methanosarcina* (58%), *Methanocella* (28%), and *Methanoregula* (11%). Attempts to amplify
265 *mmoX* from cDNA were unsuccessful (data not shown). In metagenomes, genes for *pmoA* were
266 dominant in surface peat, whereas the relative abundance of *mmoX* sequences increased with
267 depth (Fig. S2).

268 ***Rates of diazotrophy and methanotrophy.*** Potential rates of acetylene reduction were
269 measured for peat collected from S1 bog hollows and hummocks in April, June, August and
270 September 2013-2015 and incubated for 1 week at 25°C. Acetylene reduction to ethylene was
271 only detected in surface (0-10 cm) peat samples incubated in the light, and not in deep (10-30
272 cm) peat, nor in surface peat incubated in the dark. Acetylene reduction to ethane was not
273 detected in any incubation (data not shown). *Sphagnum* in peat incubations exposed to light
274 became visibly greener over the course of the incubation (Fig. 4).

275 In hollows, where surface peat was dominantly covered by a mix of *S. fallax* and *S.*
276 *angustifolium*, ARA rates were higher and more variable in degassed vs. oxic incubations (0-163

277 vs. 2-23 $\mu\text{mol C}_2\text{H}_4 \text{ g}^{-1} \text{ h}^{-1}$), and were unaffected by the presence or absence of 20% CO_2 . ARA
278 in hollows incubated with degassed headspace was positively correlated ($P < 0.0001$) with peat
279 water content (93 to 96%). In both hollows and hummocks, ARA rates were not affected by
280 addition of 1% CH_4 . In hummocks with lower water content (90-91%), surface peat was
281 dominantly covered by *S. magellanicum*, and oxic and degassed treatments had similarly low
282 ARA rates (0-8 $\mu\text{mol C}_2\text{H}_4 \text{ g}^{-1} \text{ h}^{-1}$). *nifH* transcripts in surface peat from hollows incubated under
283 degassed headspace with 1% C_2H_2 , with or without 1% CH_4 , ranged from $10^4 - 10^7$ copies g^{-1}
284 (July 2014; data not shown), which was 1-4 orders of magnitude lower than field samples from
285 the previous summer (10^8 copies g^{-1} ; Fig. 1), and higher and more variable in hollows than *nifH*
286 transcripts from hummock incubations (Fig. S5).

287 $^{15}\text{N}_2$ incorporation showed similar overall trends as ARA (e.g. 90% higher rates in
288 degassed vs. oxic conditions, and no significant CH_4 effect; Fig. 5). In degassed treatments, 1%
289 C_2H_2 inhibited $^{15}\text{N}_2$ incorporation by 55% but had no effect on oxic treatments. In oxic
290 treatments, CH_4 consumption rates were 100 times higher than $^{15}\text{N}_2$ incorporation rates, and 1%
291 C_2H_2 addition suppressed CH_4 oxidation rates by 95% (Fig. 5). Using the four sites measured
292 with both methods, a conversion factor of 3.9 for $^{15}\text{N}_2$ -to-ARA was calculated (Fig. S6). In sum,
293 laboratory incubations of native peats revealed that diazotrophy was stimulated by light,
294 suppressed by O_2 , and minimally affected by CH_4 and CO_2 .

295

296 Discussion

297 ***Diazotrophs are active in surface peat.*** By definition, the only source of nutrients to
298 ombrotrophic peat bogs is the atmosphere (65). Scarce (low μM) dissolved nitrogen and
299 phosphorus in S1 peat pore water suggests oligotrophic conditions (65, 89, 97), consistent with

300 low rates of atmospheric deposition in Minnesota peat bogs (43). Low nutrient concentrations
301 add further evidence to previous suggestions that these nutrients may limit *Sphagnum*
302 productivity (60) and/or complex mechanisms may exist for nutrient scavenging at ultra-low
303 concentrations. Diazotroph activity (ARA and *nifH* transcription) was solely detected in surface
304 peat samples incubated in the light. The apparent absence of diazotrophy at greater depths is
305 consistent with previous reports (86), and may be due to light limitation and/or remineralization
306 of organic nitrogen to ammonium, which is preferentially used as a nitrogen source by microbes.

307 ***Alphaproteobacteria (Rhizobiales) dominate peatland nitrogenase sequences.***

308 Diazotrophic methanotrophs have the potential to serve as a methane biofilter and a nitrogen
309 source for peatland ecosystems. Previous work showed that *Alphaproteobacteria (Rhizobiales)*,
310 including Type II methanotrophs (63, 99), were the dominant diazotrophs in *Sphagnum*-
311 dominated peatlands and may provide the unaccounted nitrogen input resulting from an
312 imbalance in atmospheric nitrogen deposition and accumulation in *Sphagnum* mosses (44, 54,
313 60, 83, 99). However, the carbon metabolism of peatland diazotrophs remains unclear because
314 the two dominant *Rhizobiales* families grow on both complex organics (*Bradyrhizobiaceae* (32))
315 as well as simple alkanes and C1 compounds (*Beijerinckiaceae* (19)). Consistent with previous
316 studies, *Rhizobiales* showed the highest relative abundance in transcript libraries in this study.
317 *Beijerinckiaceae* and *Bradyrhizobiaceae* dominated *nifH* and *nifD* amplicons, respectively.
318 However, the complex taxonomic classification of *Rhizobiales nifH* and *nifD* genes (22)
319 prevented distinguishing methanotrophic *Beijerinckiaceae* from heterotrophic *Beijerinckia*
320 *indica* and *Bradyrhizobiaceae* solely on the basis of *nifH* and *nifD* phylogenies.

321 ***Diazotrophy is catalyzed by the molybdenum form of nitrogenase.*** Peatland conditions
322 such as low pH, nitrogen, and temperature would be expected to favor diazotrophy by alternative

323 nitrogenases. Molybdenum sorption to peat is enhanced at low pH (6), biological requirements
324 for Mo are higher when bacteria are fixing N₂ than growing on other nitrogen compounds (33),
325 and alternative nitrogenases have higher activity and expression at lower temperatures (74, 101).
326 Below 10 nM Mo, diazotrophy is limited in laboratory cultures (1, 26, 34, 87, 108) and
327 alternative nitrogenases, if present, are expressed (3). Transcription of alternative nitrogenase
328 genes has been reported for *Peltigera cyanolichens* (45), but to our knowledge has not previously
329 been investigated for *Sphagnum* peatlands.

330 Molybdenum concentrations in S1 bog porewaters were 1-7 nM, which is within the
331 same range as Mo in other oligotrophic freshwaters (33), and lower than V (5-21 nM) and Fe (7-
332 35 µM). Similar metal concentrations between: (a) sampling dates in September 2014, June
333 2015, and September 2015; (b) in another ombrotrophic Minnesota peatland, Zim bog, ~80 km
334 from the S1 bog; and (c) in acidic peatlands in Northern Europe (53) suggest that the values we
335 measured are spatially and temporally representative for diverse northern peatlands.

336 Intriguingly, despite the presence of conditions that would be expected to favor
337 alternative nitrogenase expression (low pH, long winters, and Fe>V>Mo), our collective
338 evidence suggests that diazotrophy at the S1 bog was catalyzed by the Mo-containing form of the
339 nitrogenase enzyme. The majority of *nifH* genes retrieved from metagenomes belonged to Mo-
340 containing clusters I and III. A significant number of sequences came from uncharacterized
341 Cluster IV, recently shown to contain a functional nitrogenase (110) that likely binds a Mo-Fe
342 cofactor (72). The metagenomes contained very few cluster II *vnfD/anfD* genes, and minimal
343 *vnfD/anfD* transcripts were amplified from peat cDNA. Ethane, a biomarker of alternative
344 nitrogenase, was undetectable in ARA incubations. Finally, the ¹⁵N₂-to-ARA conversion factor
345 (3.9) was within the same range (3-4) as other peat bogs (41, 62, 99), and matched Mo-

346 nitrogenase in pure culture experiments as opposed to the lower values measured for alternative
347 nitrogenases (4). The question of how peatland diazotrophs access scarce Mo remains uncertain;
348 it is possible that Mo bound to peat organic matter can be scavenged by diazotrophs as is the case
349 in forest soils (71, 104).

350 ***Methanotrophs are active in incubations and inhibited by acetylene.*** We performed
351 bottle experiments to test whether methanotrophs were active. Complete CH₄ consumption in air
352 amended with 1% CH₄ showed that methanotrophs were active in our incubations, and that the
353 black bromobutyl stoppers we used were not toxic to peatland methanotrophs, whereas non-
354 halogenated stoppers are toxic to aquatic aerobic methanotrophs (76). Acetylene fully inhibited
355 CH₄ consumption, demonstrating that the methanotrophs in our incubations were C₂H₂ sensitive,
356 similar to laboratory strains tested in previous studies (16, 94). Like diazotrophy, methanotrophy
357 was apparently also mediated by the enzyme requiring the scarcer metal; dissolved Fe was
358 consistently orders of magnitude higher than Cu in peat porewater, yet *pmoA* sequences were
359 more abundant than *mmoX* sequences in surface peat where the highest CH₄ consumption was
360 observed (24, 68). This finding is also consistent with higher transcription of *pmoA* vs. *mmoX* in
361 other acidic peatlands (10, 53, 63). In laboratory studies, the “copper switch” from sMMO to
362 pMMO supported growth occurs when Cu > 1 μM (90), which is several orders of magnitude
363 higher than Cu concentrations measured at the S1 bog (7-38 nM). It is possible that the “copper
364 switch” occurs at a lower threshold in peat bogs, or that there are other factors controlling the
365 type of MMO expressed, such as CH₄ or O₂ availability. Additionally, the inherent nature of the
366 peat matrix, with characteristically high levels of particulate and dissolved organic matter, likely
367 also affects metal bioavailability in complex ways (98) not addressed in our study.

368 *Surface peatland diazotrophy is sensitive to oxygen and acetylene.* If methanotrophs
369 were the dominant diazotrophs in peatlands, as previously proposed (60, 99), CH₄ addition
370 should have stimulated ¹⁵N₂ incorporation in our bottle experiments, and C₂H₂ should have
371 inhibited it. Instead, this and previous work (61) found that ¹⁵N₂ incorporation was not affected
372 by CH₄ addition. Acetylene partially inhibited ¹⁵N₂ incorporation under degassed conditions (as
373 in Kox, et al. (59)) and had minimal effect on oxic ¹⁵N₂ incorporation, suggesting that C₂H₂- and
374 O₂-sensitive microbial clades contributed approximately half of the diazotrophic activity in our
375 incubations. This finding highlights the importance of quantifying peatland diazotrophy by ¹⁵N₂
376 incorporation instead of, or in addition to, ARA.

377 Based on *nif*, *pmoA* and *mcrA* phylogeny from native peat, the O₂- and C₂H₂-sensitive
378 diazotrophs were likely methanotrophic *Beijerinckiaceae* that can only fix N₂ under micro-oxic
379 conditions, and/or strict anaerobes in cluster III, such as methanogenic *Euryarchaeota*. These
380 two families are the most active in surface peat based on transcript and amplicon sequences.
381 Since *nifH* transcripts were 1-4 orders of magnitude lower in our incubations than native peat, it
382 is likely that diazotrophs were stressed, possibly due to O₂ exposure during the sampling process.
383 Indeed, we observed highest ARA rates in hollows where the water table was typically at the bog
384 surface, limiting O₂ penetration into the peat. The other half of O₂-sensitive diazotrophic activity
385 was likely performed by C₂H₂-insensitive, heterotrophic *Bradyrhizobiaceae* and/or
386 *Beijerinckiaceae* (51, 91), or C₂H₂-insensitive, methanotrophic *Methylocystaceae*, which can
387 adapt to a wide range of CH₄ concentrations (17). Less O₂-sensitive microbes, including
388 heterotrophic *Beijerinckiaceae* (91) and/or photosynthetic *Oscillatoriothrix* (25), likely
389 contributed to the minor amount of ARA activity in the presence of O₂.

390 ***Molecular markers for diazotrophy.*** We end with a word of caution with regard to the
391 molecular detection of diazotrophs. The majority of studies in peatlands have employed PCR
392 amplification and sequencing of the *nifH* marker gene for studying the dynamics of diazotrophs
393 in peatlands. A wide range of *nifH* primer sets exist, with varying universality (29). Peat bog
394 sequencing efforts have used polF/polR (this study; 63), F1/R6 (99), FGPH19/polR+polF/AQER
395 (61), and 19F/nifH3+nifH1(1)/nifH2(2) with nested PCR (8, 59). *In silico* evaluation predicts
396 that the polF/polR primer set will not amplify the majority of *Proteobacteria* and/or
397 *Cyanobacteria* and Group III *nifH* sequences (29), however this primer set yields the highest
398 efficiency for qPCR (30). Of the *nifH* primer sets used previously, F1/R6, 19F/nifH3 and
399 nifH1(1)/nifH2 are predicted to have the highest coverage for soils (>80% predicted primer
400 binding for sequences from soil ecosystems). However, it is important to be aware that the F1/R6
401 primer set contains a number of mismatches with cluster III in peatlands, including
402 methanogenic *Euryarchaeota* represented (Fig. S7). In order to maximize sequence coverage, we
403 suggest primer sets that can amplify *nifH* from cluster III, such as IGK(3)/DVV (29), for future
404 studies.

405 ***Conclusions.*** This study revealed that peatland diazotrophs preferentially transcribed the
406 Mo-based, rather than the V-based, form of the nitrogenase enzyme, despite the dominance of V
407 over Mo in the environment. It also highlighted the sensitivity of diazotrophic peatland
408 communities to O₂ exposure during sample collection, and quantified the inhibitory effect of
409 C₂H₂ addition on peatland diazotrophy. Under our experimental conditions in lab incubations, we
410 did not observe CH₄-stimulated diazotrophy. However, quantification of the relative
411 contributions of methanotrophic and heterotrophic diazotrophy *in situ* awaits further
412 investigation.

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419 **Table 1. qPCR and sequencing primers used in this study.** PCR conditions were based on
420 refs. (24), (63), (93), (29), and (4) for *pmoA*, *mmoX*, *mcrA*, *nifH*, and *nifD/vnfD/anfD*,
421 respectively.

Gene	Primer set	Reference
<i>mmoX</i>	mmoX-206f/mmoX-886r	(48)
<i>pmoA</i>	A189f/Mb661r	(46)
<i>mcrA</i>	Mlas/mcrA-rev	(93)
<i>nifH</i>	polF/polR	(56)
<i>nifD</i>	1) nifD820F/nifD1389R 2) nifD820F/nifD1331R	(4)
<i>vnfD/anfD</i>	1) vnfD_anfD548F/vnfD_anfD1337R 2) vnfD_anfD548F/vnfD_anfD1291R	(4)

422

423 **Figure Captions**

424 **Figure 1. Depth profiles of NH_4^+ concentrations (black circles) and *nifH* gene copies (white**
425 **squares) and transcripts (white triangle) in units of copies per gram of dry weight for the**
426 **S1 bog T3 mid site. Ammonium concentrations are means of measurements from May, June and**
427 **September 2014. *nifH* copy numbers are from July 2013. Error bars are standard errors. *nifH***
428 **transcripts were not detected at 20, 30 and 75 cm depths. Surface (0-10 cm) and deep (10-30 cm)**
429 **peat depth intervals used for rate measurements are designated by dashed lines. The maximum**
430 **water table depth at the S1 site in July 2013 was 2 cm below the hollow surface (dotted line)**
431 **(38).**

432
433 **Figure 2. Numbers of cDNA amplicon sequence reads for *nifH* and *nifD***
434 **alphaproteobacterial transcripts. Primer sets were polF/polR and nifD820F/nifD1331R for**
435 ***nifH* and *nifD*, respectively.**

436
437 **Figure 3. Dissolved molybdenum (white) and vanadium (gray) concentrations in pore water**
438 **from three depths in S1 peat hollows (mid and far sites along T3 transect) from June 2015.**
439 **Pie charts show the relative abundance of genes encoding the five nitrogenase H-subunit clusters**
440 **from metagenomes for each depth; clusters I and III encode Mo-Fe nitrogenases (*nifH*); cluster II**
441 **encodes alternative (*vnfH*, *anfH*) nitrogenases; cluster IV encode nitrogenase paralogs. Deepest**
442 **metagenomes were from 75 cm; insufficient numbers of nitrogenase H-subunit sequences were**
443 **recovered from the far site for cluster analysis.**

444

445 **Figure 4. Acetylene reduction rates for hummocks (90-91% water content) and hollows**
446 **(93-96% water content) at the S1 bog, T3 transect (0-10 cm depth incubated in the light at**
447 **25°C for 7 days).** ARA units are μmol ethylene produced per gram of dry weight per hour.
448 Photo insets show dominant *Sphagnum* species in (a) hummocks (*S. magellanicum*) and (b)
449 hollows (*S. fallax* and *S. angustifolium*). Photo inset (c) shows *Sphagnum* greening after
450 incubation of hollow samples in the light for 7 days at 25°C; bottle B2* (April 2014) received air
451 headspace with 1% CH_4 and treatment D2 (Sept 2013) received air headspace without CH_4 . The
452 vertical dotted line divides the hummock samples (dominated by *S. magellanicum*) from the
453 hollow samples (dominated by *S. fallax/angustifolium*) in terms of water content (92%).

454

455 **Figure 5. Effect of 1% C_2H_2 on $^{15}\text{N}_2$ incorporation and CH_4 consumption for S1 bog surface**
456 **peat.** Rates were measured for samples collected from the NW S1 bog transect in September
457 2014. $^{15}\text{N}_2$ incorporation treatment conditions were 80% N_2 +20% CO_2 ±1% CH_4 (shaded bars) or
458 80% N_2 +20% O_2 ±1% CH_4 (white bars), with (a) and without (b) 1% C_2H_2 ; units are $\text{nmol } ^{15}\text{N}_2$
459 incorporated per gram of dry weight per hour. CH_4 consumption treatments were 80% N_2 +20%
460 CO_2 +1% CH_4 (black circles) with (a) and without (b) 1% C_2H_2 ; units are $\text{nmol } \text{CH}_4$ consumed
461 per gram of dry weight per hour. Error bars are standard errors. Different letters indicate
462 statistically different elemental contents ($p < 0.05$ based on Tukey-Kramer HSD test).

463

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