

RUNNING TITLE: OC oxidation processes across vegetation

Carbon inputs from riparian vegetation limit oxidation of physically-bound organic carbon via biochemical and thermodynamic processes

Emily B. Graham¹, Malak M. Tfaily², Alex R. Crump¹, Amy E. Goldman¹, Lisa Bramer¹, Evan Arntzen¹, Elvira Romero¹, C. Tom Resch¹, David W. Kennedy¹, and James C. Stegen¹

¹Pacific Northwest National Laboratory, Richland, WA USA

²Environmental Molecular Science Laboratory, Richland, WA USA

Correspondence: Emily B. Graham, Pacific Northwest National Laboratory, PO Box 999, Richland, WA 99352, 509-372-6049, emily.graham@pnnl.gov

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Key points.

- Riparian vegetation protects bound-OC stocks
- Biochemical processes associated with OC oxidation vary with vegetation conditions
- Common thermodynamic principles underlie OC oxidation regardless of vegetation conditions

1 **Abstract.**

2 In light of increasing terrestrial carbon (C) transport across aquatic boundaries, the
3 mechanisms governing organic carbon (OC) oxidation along terrestrial-aquatic interfaces are
4 crucial to future climate predictions. Here, we investigate the biochemistry, metabolic pathways,
5 and thermodynamics corresponding to OC oxidation in the Columbia River corridor using ultra-
6 high resolution C characterization. We leverage natural vegetative differences to encompass
7 variation in terrestrial C inputs. Our results suggest that decreases in terrestrial C deposition
8 associated with diminished riparian vegetation induce oxidation of physically -bound OC. We
9 also find that contrasting metabolic pathways oxidize OC in the presence and absence of
10 vegetation and—in direct conflict with the ‘priming’ concept—that inputs of water-soluble and
11 thermodynamically favorable terrestrial OC protects bound-OC from oxidation. In both
12 environments, the most thermodynamically favorable compounds appear to be preferentially
13 oxidized regardless of which OC pool microbiomes metabolize. In turn, we suggest that the
14 extent of riparian vegetation causes sediment microbiomes to locally adapt to oxidize a particular
15 pool of OC, but that common thermodynamic principles govern the oxidation of each pool (*i.e.*,
16 water-soluble or physically-bound). Finally, we propose a mechanistic conceptualization of OC
17 oxidation along terrestrial-aquatic interfaces that can be used to model heterogeneous patterns of
18 OC loss under changing land cover distributions.

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22 1. Introduction

23 Soils and nearshore sediments comprise a carbon (C) reservoir that is 3.2 times larger
24 than the atmospheric C pool [*Burd et al.*, 2016], yet Earth System Models (ESMs) struggle to
25 integrate mechanisms of OC oxidation in these environments into predictions of atmospheric
26 carbon dioxide concentrations [*Todd-Brown et al.*, 2013; *Wieder et al.*, 2013; *Wieder et al.*,
27 2015]. In particular, OC oxidation in nearshore habitats constitutes a significant uncertainty in
28 atmospheric C flux [*Aalto et al.*, 2003; *Battin et al.*, 2009] and knowledge about C cycling along
29 these transitional ecosystems is necessary to accurately predict global C cycling [*Burd et al.*,
30 2016]. Terrestrial C inputs into aquatic systems have nearly doubled since pre-industrial times;
31 an estimated 2.9 Pg C now crosses terrestrial-aquatic interfaces annually (vs. 0.9 Pg C yr⁻¹ stored
32 within forested ecosystems) [*Battin et al.*, 2008; *Regnier et al.*, 2013]. The magnitude of this flux
33 has garnered significant recent attention [*Battin et al.*, 2008; *Battin et al.*, 2009; *Regnier et al.*,
34 2013], yet the biochemical, metabolic, and thermodynamic mechanisms governing OC oxidation
35 along terrestrial-aquatic interfaces remain a crucial uncertainty in climate predictions. New
36 molecular techniques are providing insight into OC dynamics [*Mason et al.*, 2016; *Malak M*
37 *Tfaily et al.*, 2015; *M.M. Tfaily et al.*, 2017], but we still lack an understanding of why some OC
38 remains stabilized for millennia whereas other OC is rapidly oxidized [*Schmidt et al.*, 2011].

39 The ability of microorganisms to oxidize complex OC is an important constraint on C
40 cycling because OC is a mixture of compounds with different propensities for biotic oxidation
41 [*Hedges and Oades*, 1997; *Hedges et al.*, 2000]. Within terrestrial research, OC oxidation is
42 often framed within the concept of ‘priming,’ whereby microbial oxidation of chemically-
43 complex, less bioavailable OC is fueled by the addition of more bioavailable and
44 thermodynamically favorable OC compounds [*Kuzyakov*, 2010]. However, the applicability of

45 priming in aquatic environments is unclear [*Bengtsson et al.*, 2014; *Bianchi*, 2011; *Guenet et al.*,
46 2010]. Aquatic systems, and in particular nearshore environments, frequently experience mixing
47 of terrestrial and aquatic C sources with distinct chemical character, providing a theoretical basis
48 for priming expectations [*Bengtsson et al.*, 2014; *Guenet et al.*, 2010]. Consistent with priming,
49 *Guenet et al.* [2010] proposed that this mixing generates “hotspots” or “hot moments” of
50 biological activity facilitated by complementary C resources. Alternatively, OC stabilization in
51 sediments is tightly linked to organomineral interactions, which provide physical protection from
52 extracellular enzyme activity [*Hedges and Keil*, 1995; *Hunter et al.*, 2016; *Rothman and Forney*,
53 2007], and the strength of these interactions may override any influence of priming. Early
54 investigations of priming effects in aquatic systems have been inconclusive, with evidence both
55 for [*Dorado-García et al.*, 2016] and against [*Bengtsson et al.*, 2014; *Catalán et al.*, 2015]
56 priming mechanisms.

57 Several new perspectives have attempted to move beyond frameworks, such as priming,
58 that depend on strict chemical definitions to predict OC oxidation [*Burd et al.*, 2016; *Cotrufo et*
59 *al.*, 2013; *Lehmann and Kleber*, 2015]. Recent work proposes that the probability of OC
60 oxidation is related to a spectrum of chemical properties and that even very complex OC can be
61 oxidized when more thermodynamically favorable OC is depleted or isolated from
62 microorganisms. For example, *Lehmann and Kleber* [2015] have proposed a ‘soil continuum
63 hypothesis’ whereby OC is a gradient of continually decomposing compounds that are variably
64 accessible for biotic oxidation, with no notion of chemically labile versus recalcitrant
65 compounds. Similarly, *Burd et al.* [2016] have suggested that OC oxidation is a ‘logistical
66 problem’ involving the ability of microorganisms to access and metabolize compounds. Both

67 concepts capture the emerging belief that chemically-complex, less thermodynamically favorable
68 OC can be oxidized when more favorable compounds are inaccessible.

69 Here, we address a critical knowledge gap in predicting the global C balance [Aalto *et al.*,
70 2003; Battin *et al.*, 2009; Burd *et al.*, 2016; Regnier *et al.*, 2013]—mechanisms governing OC
71 oxidation along terrestrial-aquatic interfaces. Specifically, we investigate the biochemistry,
72 microbial metabolism, and thermodynamics of OC oxidation in nearshore water-soluble and
73 physically-bound (*i.e.*, mineral and microbial) OC pools along a freshwater terrestrial-aquatic
74 interface. We leverage natural variation in riparian vegetation along the Columbia River in
75 Eastern Washington State, the largest river in the U.S. west of the Continental Divide [Ebel *et*
76 *al.*, 1989; Moser *et al.*, 2003], to examine these mechanisms in the context of spatial variation in
77 terrestrial C deposition. Consistent with the priming paradigm, we hypothesize that (a) C
78 deposition associated with riparian vegetation increases total aerobic metabolism and enhances
79 oxidation of bound-OC stocks, while (b) areas without riparian vegetation foster lower rates of
80 aerobic metabolism with minimal oxidation of bound-OC.

81

82 **2. Materials and Methods**

83 *2.1. Site Description*

84 This study was conducted along the Columbia River shoreline within the Hanford Site
85 300 Area (approximately 46° 22' 15.80"N, 119° 16' 31.52"W) in eastern Washington State
86 [Graham *et al.*, 2016a; 2017; Slater *et al.*, 2010; Zachara *et al.*, 2013]. The Columbia River
87 experiences shoreline geographic variation in vegetation patterns, substrate geochemistry, and
88 microbiome composition [Arntzen *et al.*, 2006; Lin *et al.*, 2012; Peterson and Connelly, 2004;
89 Slater *et al.*, 2010; James C Stegen *et al.*, 2016; James C. Stegen *et al.*, 2012; Zachara *et al.*,

90 2013]. Accordingly, the Hanford Reach of the Columbia River embodies an ideal natural system
91 in which to examine heterogeneity of terrestrial OC inputs and subsequent OC oxidation
92 mechanisms.

93 Liquid N₂-frozen sediment profiles (0-60 cm) were collected along two shoreline
94 transects with or without riparian vegetation (hereafter, V and NV for ‘vegetated’ and ‘not
95 vegetated’, Table 1, Figure 1) perpendicular to the Columbia River in March 2015, separated by
96 a distance of ~170m. V was characterized by a moderately sloping scour zone, small boulders,
97 and a closed canopy of woody perennials *Morus rubra* (Red Mulberry) and *Ulmus rubra*
98 (Slippery Elm). The highest elevation (upper bank) samples were collected within the rooting
99 zone. In contrast, NV was characterized by a gradually sloping scour zone, cobbled armor layer,
100 and minimal vegetation. We collected profiles at three locations in each transect with 5m spacing
101 within a spatial domain of ~175 x 10m. In each transect, the lower bank profile was located at
102 ~0.5m (vertical distance) below the midpoint, which was in turn ~0.5m below the upper bank
103 profile (approximately 10m horizontal distance from upper to lower profiles). All cores were
104 collected (see below) during conditions in which the sediments were fully saturated, either being
105 underwater at the time of coring or recently underwater prior to coring. Each profile was
106 sectioned into 10-cm intervals from 0-60cm. Because OC composition (see below for Methods)
107 did not differ across upper (0-10cm) to lower (50-60cm) sections in each profile, each 10-cm
108 section was used as a replicate sample to provide sufficient sample size (n >15 at each transect)
109 for comparisons between the vegetated and non-vegetated transect.

110

111 *2.2. Sample Collection*

112 Liquid N₂-frozen sediment profiles were collected as outlined in *Moser et al.* [2003]
113 using a method developed by *Lotspeich and Reed* [1980] and modified by *Rood and Church*
114 [1994]. A pointed stainless steel tube (152 cm length, 3.3 cm outside diameter, 2.4 cm inside
115 diameter) was driven into the river bed to a depth of ~60cm. Liquid N₂ was poured down the
116 tube for ~15 minutes, until a sufficient quantity of material had frozen to the outside of the rod.
117 The rod and attached material were removed from the riverbed with a chain hoist suspended
118 beneath a tripod. Profiles were placed over an aluminum foil lined cooler containing dry ice.
119 Frozen material was removed with a mallet. The material was then wrapped in the foil and
120 transported on dry ice to storage at -80°C. In the lab, profiles were sectioned into 10cm depth
121 intervals from 0-60cm in an anaerobic glove bag (n = 6 per profile, except for the lowest
122 elevation in the non-vegetated transect which was sectioned only from 30-60cm due to low
123 recovery yield from 0-30cm; total n = 33).

124

125 2.3. Physicochemistry

126 Details concerning physicochemical assays are provided in the Supporting Information.
127 Briefly, we determined the particle distribution of sediments by separating size fractions via
128 sieving; total nitrogen, sulfur, and carbon content were determined using an Elementar vario EL
129 cube (Elementar Co.Germany); NH₄⁺ was extracted with KCl and measured with Hach Kit
130 2604545 (Hach, Loveland, Co); Fe(II) content was measured with a ferrozine assay; and all other
131 ion concentrations were measured by inductively coupled plasma mass spectrometry (ICP-MS)
132 on HCl extractions. Aerobic metabolism was determined with a resazurin reduction assay,
133 modified from *Haggerty et al.* [2009].

134

135 2.4. *Fourier transform ion cyclotron resonance mass spectrometry solvent extraction and data*
136 *acquisition*

137 We leveraged state of science chemical extraction protocols combined with Electrospray
138 ionization (ESI) and Fourier transform ion cyclotron resonance (FTICR) mass spectrometry
139 (MS) to infer differences in OC character among our samples. Previously, *Tfaily et al.* [2015;
140 2017] demonstrated the optimization of OC characterization from soils and sediments by
141 sequential extraction with polar and non-polar solvents tailored to the sample set of interest.
142 *Tfaily et al.*'s extraction procedures have been coupled to ESI FTICR-MS to distinguish OC
143 pools among ecosystems and soil types [*Tfaily et al.*, 2015; *Tfaily et al.*, 2017] as well as to
144 provide information on the metabolism of distinct OC pools among samples within a single
145 environment [*Bailey et al.*, 2017]. Other common OC characterization methods such as nuclear
146 magnetic resonance spectroscopy (NMR), Fourier transform infrared spectroscopy (FTIR), and
147 gas chromatography MS only analyze a limited number of compound classes [*Kögel-Knabner*,
148 2002; *Kögel-Knabner*, 2000]. In contrast, ESI FTICR-MS introduces intact organic molecules
149 into the MS without fragmentation and allows for the detection of a wide range of chemical
150 compounds [*Tfaily et al.*, 2015; *Tfaily et al.*, 2017]. The use of 12 Tesla (T) FTICR-MS offers
151 high mass resolving power (>1M) and mass measurement accuracy (<1 ppm), and while nascent
152 in its application within complex environmental systems, it has emerged as a robust method for
153 determining OC chemistry of natural organic matter [*Kim et al.*, 2003; *Koch et al.*, 2005; *Tfaily*
154 *et al.*, 2011; *Tremblay et al.*, 2007]. Moreover, *Tfaily et al.* [2015; 2017] demonstrated that
155 sequential extraction with targeted solvents can preferentially select OC pools with differing
156 chemical character (*e.g.*, lipid-like vs. carbohydrate-like).

157 Here, we used three solvents with different polarities —water (H₂O), methanol (CH₃OH,
158 hereafter “MeOH”) and chloroform (CHCl₃)—to sequentially extract a large diversity of organic
159 compounds from samples, according to *Tfaily et al.* [2015; 2017]. Water extractions were
160 performed first, followed by MeOH and then CHCl₃. Previous work has shown that each solvent
161 is selective towards specific types of compounds [*Tfaily et al.*, 2015]. Water is a polar solvent
162 with a selection bias for carbohydrates with high O/C ratios, amino-sugars, and other labile polar
163 compounds [*Tfaily et al.*, 2015]; and, as nearshore environments frequently experience wetting,
164 water extractions represent an estimation of readily accessible OC compounds in these
165 environments. Conversely, CHCl₃ is selective for non-polar lipids associated with mineral
166 interactions and cellular membranes (*i.e.*, physically-bound OC) [*Tfaily et al.*, 2015]. Because
167 MeOH has a polarity in between that of water and CHCl₃, it extracts both water-soluble and
168 bound-OC pools (*i.e.*, a mix of compounds that water and CHCl₃ extract), and *Tfaily et al.* [2015]
169 demonstrated compositional overlap between water-soluble and MeOH extracted OC pools. In
170 this study, we are interested in the differences in OC composition between pure water-soluble
171 and bound-OC pools, and we focus our discussion on H₂O- and CHCl₃-extractions only. We use
172 H₂O- and CHCl₃-extracted OC as proxies for readily bioavailable (*i.e.*, weakly bound) vs. less
173 bioavailable (*i.e.*, mineral- and microbial-bound) pools, respectively.

174 Extracts were prepared by adding 1 ml of solvent to 100 mg bulk sediment and shaking in
175 2 mL capped glass vials for two hours on an Eppendorf Thermomixer. Samples were removed
176 from the shaker and left to stand before spinning down and pulling off the supernatant to stop the
177 extraction. The residual sediment was dried with nitrogen gas to remove any remaining solvent,
178 and then the next solvent was added. The CHCl₃ and H₂O extracts were diluted in MeOH to
179 improve ESI efficiency. *Tfaily et al.* [2015] estimated the OC extraction efficiency to be ~15%.

180 *Tfaily et al.* [2015] have previously demonstrated extraction efficiencies as low as 2% to be
181 representative of OC pool composition. We further note that numerous studies have established
182 FTICR-MS as a robust method for distinguishing compositional differences among OC pools
183 [*Herzprung et al.*, 2017; *Kellerman et al.*, 2015; *Rossel et al.*, 2016; *Ward and Cory*, 2015;
184 *Zhang et al.*, 2016].

185 Ultra-high resolution mass spectrometry of the three different extracts from each sample
186 was carried out using a 12 Tesla Bruker Solarix FTICR-MS located at the Environmental
187 Molecular Sciences Laboratory (EMSL) in Richland, WA, USA. As per *Tfaily et al.* [2017], we
188 performed weekly calibration using a tuning solution containing C₂F₃O₂, C₆HF₉N₃O,
189 C₁₂HF₂₁N₃O, C₂₀H₁₈F₂₇N₃O₈P₃, and C₂₆H₁₈F₃₉N₃O₈P₃ with *m/z* ranging from 112 to 1333
190 (Agilent Technologies, Santa Clara, CA USA). Instrument settings were optimized using
191 Suwannee River Fulvic Acid (IHSS). The instrument was flushed between samples using a
192 mixture of water and methanol. Blanks were analyzed at the beginning and the end of the day to
193 monitor for background contaminants.

194 The extracts were injected directly into the mass spectrometer, and the ion accumulation
195 time was optimized for all samples to account for differences in OC concentration. The ion
196 accumulation time ranged between 0.5 and 1s. A standard Bruker electrospray ionization (ESI)
197 source was used to generate negatively charged molecular ions. Samples were introduced to the
198 ESI source equipped with a fused silica tube (30 μm i.d.) through an Agilent 1200 series pump
199 (Agilent Technologies) at a flow rate of 3.0 μL min⁻¹. Experimental conditions were as follows:
200 needle voltage, +4.4 kV; Q1 set to 50 *m/z*; and the heated resistively coated glass capillary
201 operated at 180 °C. One hundred forty-four individual scans were averaged for each sample and
202 internally calibrated using an organic matter homologous series separated by 14 Da (-CH₂

203 groups). The mass measurement accuracy was less than 1 ppm for singly charged ions across a
204 broad m/z range (100-1200 m/z). The mass resolution was $\sim 350K$ at 339 m/z .

205

206 *2.5 FTICR-MS data processing*

207 A depiction of our FTICR-MS data processing pipeline is presented in Figure 2 and
208 described in the following sections.

209

210 *2.5.1 Chemical formulae assignment*

211 BrukerDaltonik software version 4.2 was used to convert raw spectra to a list of m/z
212 values applying FTMS peak picker module with a signal-to-noise ratio (S/N) threshold set to 7
213 and absolute intensity threshold to the default value of 100. Putative chemical formulae were
214 then assigned using in-house software following the Compound Identification Algorithm (CIA),
215 proposed by *Kujawinski and Behn* [2006], modified by *Minor et al.* [2012], and previously
216 described in *Tfaily et al.* [2017]. Chemical formulae were assigned based on the following
217 criteria: $S/N > 7$, and mass measurement error < 1 ppm, taking into consideration the presence of
218 C, H, O, N, S and P and excluding other elements. To ensure consistent formula assignment, we
219 aligned all sample peak lists for the entire dataset to each other in order to facilitate consistent
220 peak assignments and eliminate possible mass shifts that would impact formula assignment. We
221 implemented the following rules to further ensure consistent formula assignment: (1)
222 consistently pick the formula with the lowest error and with the lowest number of heteroatoms
223 and (2) assignment of one phosphorus atom requires the presence of at least four oxygen atoms.

224

225 *2.5.2 Identification of putative biochemical transformations using FTICR-MS*

226 To identify potential biochemical transformations, we followed the procedure detailed by
227 *Breitling et al.* [2006] and employed by *Bailey et al.* [2017]. In essence, the mass difference
228 between m/z peaks extracted from each spectrum with S/N>7 were compared to commonly
229 observed mass differences associated with biochemical transformations. All possible pairwise
230 mass differences were calculated within each extraction type for each sample, and differences
231 (within 1ppm) were matched to a list of 92 common biochemical transformations (*e.g.*, gain or
232 loss of amino groups or sugars, Table S1). For example, a mass difference of 99.07 corresponds
233 to a gain or loss of the amino acid valine, while a difference of 179.06 corresponds to the gain or
234 loss of a glucose molecule. Pairs of peaks with a mass difference within 1 ppm of our
235 transformation list were considered to be related by the corresponding compound. This approach
236 is feasible with FTICR-MS data because the set of peaks in each sample are related by
237 measureable and clearly defined mass differences corresponding to gains and losses of
238 compounds. It has been previously used by *Bailey et al.* [2017] to demonstrate differences in
239 biochemical transformations among soils incubated with different microbial inocula and among
240 pore size classes in complex soil matrices.

241

242 *2.5.3 Identification of putative microbial metabolic pathways using FTICR-MS*

243 Additionally, a set of putative microbial metabolic pathways in each sample can be
244 identified by locating chemical formulae assigned to m/z's within metabolic pathways defined in
245 the Kyoto Encyclopedia of Genes and Genomes (KEGG, Release, 80.0, <http://www.kegg.jp>)
246 [*Kanehisa and Goto*, 2000]. Chemical formulae were mapped to KEGG pathways using an in-
247 house software to detect all KEGG pathways containing a giving formula. For example, a peak
248 with a mass of 400.3356 was assigned formula C₂₀H₁₆O₉ and mapped to KEGG pathway

249 ‘map00254’ (Aflatoxin biosynthesis) which contains C₂₀H₁₆O₉ as an intermediate. While only a
250 subset of compounds detected by FTICR-MS are defined within the KEGG database (*i.e.*, peaks
251 must be assigned a chemical formula and that chemical formula must be present in a KEGG
252 pathway), we found 415 unique peaks that were assigned putative molecular formulae *and* that
253 corresponded to compounds present in KEGG pathways. Additionally, we defined assignments
254 at the pathway level (*i.e.*, by “map” number) instead of using enzyme level classification (*i.e.*,
255 EC number) in order to aggregate compounds found within the same pathways. This was done to
256 facilitate functional interpretation.

257 Although we acknowledge our results do not represent a comprehensive analysis of all
258 microbial metabolic pathways present in a sample, we assume that KEGG pathways containing
259 more peaks detected by FTICR-MS within a sample are more likely to be active than those with
260 fewer mapped peaks. We further reduced possible random matches by assessing correlations
261 with aerobic metabolisms as described in the ‘Statistical Methods’ section below, and we
262 compare results across samples to yield insight into microbial pathways in each sample beyond
263 that which can be garnered from biochemical transformations. The results are, however,
264 conceptually congruent with those derived from the biochemical transformation analyses
265 described in the preceding sub-section. The KEGG pathway and transformation analyses are
266 independent of each other, yet provided consistent insights, and thus together they provide
267 greater confidence in our interpretations.

268

269 *2.6 Statistical Methods*

270 All statistical analyses were conducted using R software (<https://www.r-project.org/>).
271 FTICR m/z intensities were converted into presence/absence data prior to analysis because

272 differences in m/z intensity are influenced by ionization efficiency as well as relative abundance
273 [Kujawinski and Behn, 2006; Minor et al., 2012; Tfaily et al., 2015; Tfaily et al., 2017].

274 To examine differences in OC composition between transects, we used the ‘vegan’
275 package to construct a Sorenson dissimilarity matrix for all m/z’s identified (*i.e.*, we included
276 peaks with or without assigned formula) within each OC pool (water-soluble or
277 physicallybound). Differences between vegetation states (*i.e.*, V vs. NV) were tested with
278 PERMANOVA with permutations stratified by depth to account for non-independence among
279 samples from the same core (999 permutations, ‘vegan’) and visualized using Non-metric
280 Multidimensional Scaling (NMDS, ‘vegan’). One sample (NV, upper profile, depth 30-40cm)
281 was removed due to peak interference during FTICR-MS, and three samples (NV, middle
282 profile, depths 00-10cm, 10-20cm, 20-30cm) were excluded, because we were unable to collect
283 sufficient sample mass for all analyses.

284 To reveal transformations associated with aerobic metabolism and to study differences in
285 those transformations across vegetation states, we determined the number of times a given
286 transformation occurred within each OC pool in each sample. Specifically, for each of the 92
287 compounds in our set of biochemical transformations, we counted the number of times in each
288 sample that transformation was observed to yield an estimate of the prevalence or ‘abundance’ of
289 each transformation in each sample. We correlated these abundance estimates to rates of
290 metabolism using linear mixed models in which transformations were a fixed effect and depth
291 was a random effect. Mixed models were compared to null expectations (*i.e.*, models including
292 only random effects) with ANOVA to determine significance. Significant relationships were
293 interpreted as biochemical transformations possibly associated with biotic OC oxidation. To
294 evaluate how transformations associated with OC oxidation varied across vegetation states, we

295 used the abundances of those transformations across all samples to calculate Bray-Curtis
296 dissimilarity. Resulting Bray-Curtis dissimilarities were used to visualize multivariate
297 differences among samples using non-metric Multidimensional Scaling (NMDS, ‘vegan’), and
298 we statistically evaluated separation between vegetation states with PERMANOVA (999
299 permutations stratified by depth, ‘vegan’). We refer to H₂O- and CHCl₃-soluble OC pools at V
300 and NV, respectively, as V-W (‘vegetated water’), V-B (‘vegetated bound’), NV-W (‘not
301 vegetated water’), and NV-B (‘not vegetated bound’) for the remainder of the manuscript.

302 Similar to our analyses of biochemical transformations, we found the number of m/z’s
303 that mapped to a given KEGG pathway. We make the assumption that pathways with more m/z’s
304 mapped to them have a higher probability of actively contributing to biogeochemical function.
305 To identify which pathways were most likely to contribute to aerobic metabolism, we tested the
306 relationship between the number of m/z’s mapped to a given KEGG pathway within each sample
307 and aerobic metabolism using linear mixed models with depth as a random variable as per our
308 analyses of biochemical transformations. Pathways with significant relationships were
309 interpreted as influencing OC oxidation, and the following analysis was conducted only with
310 those pathways. To estimate relative abundances, the number of peaks mapping to each KEGG
311 pathway in a sample was normalized by the total number of peaks mapping to any KEGG
312 pathway (within the sample) that was significantly related to aerobic metabolism. To reveal
313 groups of pathways co-varying with each other across vegetation states and OC pools, we
314 statistically clustered pathways that correlated with aerobic metabolism. Clustering was based on
315 pathway relative abundances in each vegetation state and pool type. Clusters were determined
316 using the ‘hclust’ algorithm in R with the ‘complete linkage’ clustering method and visualized
317 using the ‘pheatmap’ package.

318 Finally, we examined associations between aerobic metabolism and OC thermodynamics
319 by calculating the Gibbs Free Energy of OC oxidation under standard conditions ($\Delta G^{\circ}_{\text{Cox}}$) from
320 the Nominal Oxidation State of Carbon (NOSC) as per *La Rowe and Van Cappellen* [2011].
321 NOSC was calculated from the number of electrons transferred in OC oxidation half reactions
322 and is defined by the equation:

$$323 \quad (1) \text{NOSC} = -((-Z + 4a + b - 3c - 2d + 5e - 2f)/a) + 4$$

324 , where a, b, c, d, e, and f are, respectively, the numbers of C, H, N, O, P, S atoms in a given
325 organic molecule and Z is net charge of the organic molecule (assumed to be 1). In turn, $\Delta G^{\circ}_{\text{Cox}}$
326 was estimated from NOSC following *La Rowe and Van Cappellen* [2011]:

$$327 \quad (2) \Delta G^{\circ}_{\text{Cox}} = 60.3 - 28.5(\text{NOSC})$$

328 Values of $\Delta G^{\circ}_{\text{Cox}}$ are generally positive, indicating that OC oxidation must be coupled to the
329 reduction of a terminal electron acceptor. While $\Delta G^{\circ}_{\text{Cox}}$ varies according to the availability of
330 terminal electron acceptors, our system is primarily oxic, allowing us to infer oxygen as the
331 primary electron acceptor in most reactions and make direct comparisons across samples.
332 Additionally, though the exact calculation of $\Delta G^{\circ}_{\text{Cox}}$ necessitates an accurate quantification of all
333 species involved in every chemical reaction in a sample, the use of NOSC as a practical basis for
334 determining $\Delta G^{\circ}_{\text{Cox}}$ has been validated [*Arndt et al.*, 2013; *LaRowe and Van Cappellen*, 2011].

335 Here, we assessed relationships between aerobic metabolism and $\Delta G^{\circ}_{\text{Cox}}$ of OC
336 compounds identified in each OC pool (determined by FTICR-MS analysis) using linear mixed
337 models in each vegetation state, in which aerobic metabolism was a fixed effect, depth was a
338 random effect, and average $\Delta G^{\circ}_{\text{Cox}}$ of all m/z's with assigned formula was the dependent
339 variable.

340 We note that we could not account for possible non-independence between the 3 cores
341 obtained within each transect. Given one replicate per depth per core within each transect, core
342 identity cannot be accounted for with mixed models, due to identifiability issues in model
343 estimation, once accounting for depth. However, we emphasize that the locations within each
344 transect were selected to represent a broad range of conditions within each vegetation state. That
345 is, higher elevation cores were located near vegetation (vegetated state) or in a mostly barren
346 cobbled bank (unvegetated state) in locations that submerge seasonally; lower elevation cores
347 were in locations that are usually submerged below river water and the middle elevation
348 represented a transitional location (Figure 1). The three positions within each transect therefore
349 sampled a broad range of conditions with the potential to cause variation in OC composition and
350 cycling. This sampling design enabled a regression-based analysis.

351 We originally hypothesized that OC composition/cycling would vary continuously across the
352 broad range of sampled conditions. Our analyses, however, showed that for some key aspects of
353 the data, the two sites grouped together instead of showing continuous variation. That is, despite
354 within-transect variation in the physical environment, we observed coherence in OC
355 composition/cycling within each transect. In turn, we provisionally infer that the presence or
356 absence of riparian vegetation strongly constrained OC composition/cycling within the hyporheic
357 zone of our system. The inference is made possible by the two transects being near each other
358 such that most other features of the study system were similar between the two transects (e.g.,
359 aqueous chemistries, hydrology, mineralogy, temperature, light cycles, etc.). The major feature
360 distinguishing the two transects was the presence or absence of riparian vegetation. However, as
361 noted above, we consider our inferences to be foundational for future studies because we did not
362 sampling multiple plots within each vegetated state.

363

364 **3. Results and Discussion**

365 *3.1. Shifts in physicochemical, metabolic, and OC character between vegetation states*

366 Differences in vegetation states corresponded to differences in physicochemistry, aerobic
367 metabolism, and OC pool composition. V was characterized by mature trees near the water line
368 and was nutrient-rich relative to NV (Figure S1-3). V displayed comparatively high
369 concentrations of total C and rates of aerobic metabolism (Figure S1-3). In contrast, NV
370 consisted of vegetation-free, cobble-ridden shoreline with sandier soils, low total C, and low
371 aerobic metabolism (Figure S1-3).

372 Compositional difference in OC pools indicated a possibility for distinct OC oxidation
373 processes between the vegetation states (Figure 3, PERMANOVA, $P = 0.001$ in both water-
374 soluble and bound-OC pools). If each pool contained similar OC compounds, one might infer
375 similar OC inputs and biochemical processing in each vegetation state. Differences in OC pool
376 composition within each vegetation state therefore putatively denotes either variation in OC
377 inputs and/or variation in OC processing. Additionally, total OC content explained only 42% of
378 aerobic metabolic rates across both sites ($R^2 = 0.42$, $P = 0.0009$) and had no significant influence
379 on metabolism within each site (V: $P = 0.69$, NV: $P = 0.10$, Figure S4). This suggests that factors
380 beyond bulk OC content influenced aerobic metabolism. We hypothesized that variation in OC
381 composition was responsible for variation in aerobic metabolism within each vegetation state.
382 The following sections explore this possibility.

383

384 *3.2. Associations between C transformations and aerobic metabolism*

385 Given compositional differences in OC between vegetation states and known impacts of
386 C chemistry on metabolic functioning in other systems [*Castle et al.*, 2016; *Graham et al.*,
387 2016b], we hypothesized that biochemical transformations that were related to rates of aerobic
388 metabolism would be unique to each vegetation state.

389 Consistent with this hypothesis, transformation analysis indicated that the biochemical
390 processes associated with OC oxidation were significantly different between the vegetation
391 states. Specifically, OC transformations correlated to aerobic metabolism were significantly
392 different at each vegetation state (PERMANOVA, H₂O P = 0.005 and CHCl₃ P = 0.001, Figure 4
393 a-b, Table 2). In comparing differences in transformations occurring within the water-soluble OC
394 pool, we observed higher abundances of amino- and sugar-associated transformations for V-W
395 relative to NV-W. Twenty-six and four of these transformations co-varied significantly with
396 aerobic metabolism in V-W and NV-W, respectively. These V-W transformations were primarily
397 associated with simple C molecules (*e.g.*, glucose, Table 2). Conversely, within the bound-OC
398 pool, 35 transformations co-varied with aerobic metabolism in NV-B, compared to only three in
399 V-B. In both cases, transformations associated with bound-OC consisted of a greater proportion
400 of complex C molecules (*e.g.*, biotinylation, palmitoylation, and glyoxylate, Table 2) than in water-
401 soluble pools.

402 The larger number of transformations associated with aerobic metabolism in V-W vs. V-
403 B, and the larger number in NV-B vs. NV-W, suggests that aerobic metabolism in vegetated and
404 unvegetated areas depend on water-soluble and bound-OC pools, respectively. We note some
405 oxidation of the bound-OC pool under vegetated conditions, but only three correlations were
406 observed between V-B transformations and aerobic metabolism, suggesting a relatively minor
407 role, especially considering that there were 35 significant correlations for NV-B.

408 These differences suggest that an increased supply of bioavailable compounds in
409 vegetated areas leads to bound-OC being less involved in aerobic metabolism, relative to
410 unvegetated areas where bound-OC appears to be heavily involved in aerobic metabolism. The
411 concept of priming [Kuzuyakov, 2010] would predict the opposite pattern; a greater supply of
412 bioavailable OC should increase the contributions of less bioavailable OC (here, bound-OC) to
413 aerobic metabolism. Our results run counter to a priming mechanism and indicate that the supply
414 of bioavailable compounds diminishes the contribution of bound-OC to aerobic metabolism and,
415 in turn, protects bound-OC pools. Conversely, metabolism of bound-OC pools in locations
416 without riparian vegetation possibly denotes microbial adaptations to a low-C environment that
417 supports the oxidation of mineral-bound OC. Mineral-bound OC therefore has greater potential
418 to remain sequestered along river corridors with spatially and temporally consistent inputs of
419 bioavailable OC, potentially derived from riparian vegetation.

420

421 *3.3. Associations between microbial metabolic pathways and aerobic metabolism*

422 Because we observed stark differences in the identity of OC transformations that
423 correlated with aerobic metabolism across vegetation states, we hypothesized that the microbial
424 metabolic pathways associated with OC transformations were also dependent on vegetation state.
425 Indeed, pathways associated with OC oxidation were distinct at V vs. NV, supporting our
426 hypothesis that there were differences in the metabolic processing of OC in the presence or
427 absence of riparian vegetation. Specifically, while the metabolism of plant-derived compounds
428 appeared to be a major driver of aerobic respiration at both vegetation states, metabolism at V
429 mostly involved readily bioavailable plant derivatives in the water-soluble OC pool, and
430 metabolism at NV was associated with plant derivatives in the bound-OC pool (Figure 5).

431 In V-W, pathways that corresponded to aerobic metabolism included map01110
432 biosynthesis of secondary metabolites, map000941 flavonoid biosynthesis, and map00944
433 flavone and flavonol biosynthesis, among others (see Figure 5a). Each of these pathways denotes
434 an association between plant-derived compounds and OC oxidation in sediments. Secondary
435 metabolites (map01110) are largely comprised of plant-derived compounds such as flavonoids
436 [Agati *et al.*, 2012], terpenoids [Tholl, 2015], and nitrogen-containing alkaloids [Willaman and
437 Schubert, 1961], while flavonoids [Agati *et al.*, 2012] are one of the most abundant plant-derived
438 compounds. Associations with flavone/flavonol [Agati *et al.*, 2012] and phenylpropanoids
439 [Hahlbrock and Scheel, 1989] bolster a possible relationship between plant-associated
440 metabolites and aerobic metabolism in V-W.

441 Although correlations between plant-associated KEGG pathways and aerobic metabolism
442 could indicate the persistence of plant secondary metabolites rather than microbial metabolism,
443 our results indicate a central role for vegetation in water-soluble OC oxidation in either case. For
444 example, if KEGG associations were attributable to plant metabolism instead of microbial
445 metabolism, correlations between plant-associated pathways and aerobic metabolism in V-W
446 would indicate an indirect relationship between plant growth and microbial oxidation of OC,
447 whereby plant byproducts support microbial communities in oxidizing other portions of the OC
448 pool.

449 Pathways correlated with aerobic metabolism in NV-W or V-W had some overlap, but
450 were mostly distinct from each other (Figure 5a). Although some correlations in NV-W indicated
451 associations between plant-associated metabolic pathways and OC oxidation (*e.g.*, map01110),
452 relationships also indicated the involvement of broad metabolic processes including membrane
453 transport (map02010 ABC transporters), carbohydrate metabolism (map00010

454 Glycolysis/Gluconeogenesis), and lipid metabolism (map00061 Fatty acid biosynthesis and
455 map01040 Biosynthesis of unsaturated fatty acids). We propose that these associations indicate
456 utilization of alternative resources in comparison to V-W, in which relationships almost
457 exclusively denoted the metabolism of readily bioavailable plant byproducts. In particular, we
458 note correlations in NV-W between OC oxidation and lipid-based molecules (map00061 and
459 map01040) that are typically poorly-represented in water-soluble OC pools.

460 We also observed largely distinct relationships between metabolisms and OC oxidation in
461 V-B versus NV-B. The lipid-based and plant-associated pathway map00073 (Cutin, suberine,
462 and wax biosynthesis [*King et al.*, 2007; *Raffaele et al.*, 2009; *Shepherd and Wynne Griffiths*,
463 2006]) comprised almost 35% of relationships in NV-B, and more generic plant-associated
464 pathways such as map01060 (Biosynthesis of plant secondary metabolites) and map01110
465 (Biosynthesis of secondary metabolites) were also prevalent (Figure 5b). Conversely in V-B, the
466 most abundant associations were non-lipid metabolisms map00590 (Arachidonic acid
467 metabolism) and map01100 (Metabolic pathways).

468 Together, our results indicate largely separate metabolic processes occurring in each of
469 the water-soluble and bound-OC pools in V and NV. Because of the abundance of lipid-based
470 metabolisms correlated to OC oxidation at NV (in both OC pools), we further hypothesize that
471 these metabolisms comprise the primary KEGG-identifiable pathways associated with OC
472 oxidation in areas without riparian vegetation. This contrasts to the cycling of a more diverse
473 range of secondary metabolite classes in areas with more significant OC inputs from riparian
474 vegetation. We note that correlations between KEGG pathways and aerobic metabolism in V-B
475 indicates less oxidation of lipid plant material in the bound-OC pool under vegetated conditions
476 than in NV-B. We therefore propose that plant-derived lipid compounds serve as a secondary

477 substrate for OC oxidation in shorelines with riparian vegetation, given that more correlations at
478 V were detected in the water-soluble pool.

479

480 *3.4. Thermodynamics of carbon oxidation*

481 Finally, we hypothesized that microbes would preferentially oxidize more
482 thermodynamically favorable compounds at both sites, consistent with common thermodynamic
483 constraints on biogeochemical cycles [*Burgin et al.*, 2011; *Hedin et al.*, 1998; *Helton et al.*,
484 2015]. Because we observed evidence for preferential OC oxidation of the water-soluble OC
485 pool at V and of the bound-OC pool at NV, we further hypothesized that thermodynamic-based
486 preference of OC oxidation would be observable only in the preferred substrate pool within each
487 vegetation state. Consistent with this hypothesis, aerobic metabolism was positively correlated to
488 average $\Delta G^{\circ}_{\text{Cox}}$ in V-W ($R^2 = 0.37$, $P = 0.008$, Figure 6a) and NV-B ($R^2 = 0.59$, $P = 0.0005$
489 Figure 6b), but these variables were not correlated in V-B or NV-W. In both cases, aerobic
490 metabolism corresponded to a depletion of more thermodynamically favorable OC (*i.e.*, OC
491 became less favorable as aerobic metabolism increased), resulting in progressively less favorable
492 thermodynamic conditions.

493 The priming conceptual framework would predict that terrestrial inputs associated with
494 riparian vegetation should condition microbial communities to oxidize less thermodynamically
495 favorable C, such as that found in the bound-OC pool. In such a scenario, inputs of
496 thermodynamically favorable carbon should—by minimizing community-level energy
497 constraints—allow for the rise of microbial physiologies that can oxidize less favorable C
498 [*Kuzuyakov*, 2010]. In this case, a significant relationship between thermodynamic favorability
499 and aerobic metabolism in the V-W pool should lead to a similar relationship within the V-B

500 pool. Our results reveal a strong relationship within the V-W pool, but not in the V-B pool,
501 thereby rejecting an influence of priming. Instead, our results suggest that bound-OC pools are
502 protected by thermodynamically favorable compounds that serve as preferred substrate.

503 In contrast to our expectation that water-soluble OC associated with riparian vegetation
504 would increase oxidation of bound-OC pools, we observed evidence consistent with inhibition of
505 bound-OC oxidation by thermodynamically favorable water-soluble compounds. Priming has
506 been actively debated in aquatic research [*Bengtsson et al.*, 2014; *Bianchi*, 2011; *Guenet et al.*,
507 2010], and a number of other studies have been unable to detect a priming effect in sediment and
508 aqueous habitats [*Bengtsson et al.*, 2014; *Catalán et al.*, 2015].

509

510 *3.5. Conceptual model for OC oxidation at terrestrial-aquatic interfaces*

511 Based on our work, we propose a new view of OC oxidation along terrestrial-aquatic
512 interfaces in which the oxidation of bound-OC is limited by terrestrial inputs from riparian
513 vegetation (Figure 7). While our work represents a single system, our conceptual model is
514 derived from ultra-high resolution measurements that provide greater mechanistic insight into
515 OC cycles than traditional measures. Our model is therefore meant to provide a basis for more
516 spatially extensive experiments that will allow for broader transferability.

517 We hypothesize that riparian vegetation sustains inputs of water-soluble compounds to
518 nearshore OC pools, resulting in a larger thermodynamically favorable, water-soluble OC pool
519 (Figure 7b). This leads to higher overall C content in nearshore sediments and elevated rates of
520 aerobic respiration relative to areas with less riparian vegetation. Our data also suggest that in the
521 presence of riparian vegetation, microbial carbon oxidation primarily uses the water-soluble OC
522 pool with minimal oxidation of bound-OC due to physical and/or thermodynamic protection of

523 this pool. For instance, $\Delta G^{\circ}_{\text{Cox}}$ was lower in water-soluble OC pools than in bound-OC, and a
524 large presence of this thermodynamically favorable pool may provide adequate substrate to
525 sustain metabolic functioning, limiting the need to metabolize less thermodynamically favorable
526 OC. Additionally, organomineral interactions can protect bound-OC from extracellular enzyme
527 activity [Hunter *et al.*, 2016], inhibiting the bioavailability of OC.

528 In contrast, non-vegetated riparian zones provide little input into water-soluble OC pools
529 (Figure 7a), and rates of metabolism and C pool sizes are lower in these environments. Carbon
530 oxidation in these non-vegetated zones occurs primarily within the bound-OC pool, albeit more
531 slowly and as product of different biochemical and metabolic pathways than in vegetated
532 environments (*e.g.*, complex C transformations and lipid-based metabolism of plant derivatives
533 at NV). We posit that water-soluble pools in non-vegetated sediments are sufficiently small that
534 investing in enzymes needed to metabolize this OC pool results in a net energy loss. Instead,
535 microbes in unvegetated areas must invest in cellular machinery to access bound-OC, and our
536 results imply that the cellular machinery needed to access bound-OC is distinct from the
537 machinery needed to access water-soluble OC.

538 Interestingly, aerobic metabolism within both types of sediments is related to a depletion
539 of thermodynamically favorable compounds; however, this occurs in water-soluble OC pools in
540 vegetated zones and bound-OC pools in non-vegetated zones. That is, microorganisms in both
541 environments are constrained to the metabolism of their primary substrate pool but preferentially
542 oxidize more thermodynamically favorable compounds within that pool. This suggests that
543 microorganisms are conditioned to metabolize a subset of compounds within sediment OC,
544 possibly defined by thermodynamic or physical protection mechanisms, but operate under
545 common thermodynamic constraints once adapted to oxidize a certain OC pool.

546

547 *3.6. Broader Implications*

548 Our work is particularly relevant to CO₂ emissions in light of changes in land cover and
549 precipitation, and the influence of these perturbations to on C fluxes across terrestrial-aquatic
550 interfaces. Carbon moving across terrestrial-aquatic interfaces has been examined primarily for
551 its propensity to be oxidized along land-to-sea continuums [*Battin et al.*, 2008; *Battin et al.*,
552 2009; *Regnier et al.*, 2013], but we suggest that this terrestrially-derived OC also has important
553 influences on stabilizing mineral-bound OC within nearshore environments. The magnitude,
554 distribution, and chemical quality of terrestrial C fluxes into aquatic environments are perturbed
555 by shifts in land cover [*Fang et al.*, 2005; *Knapp et al.*, 2008]. Furthermore, vegetation
556 distributions in natural ecosystems are predicted to shift in response to altered precipitation
557 regimes. Associated changes in plant phenology, morphology, and establishment will impact the
558 quantity, quality, and distribution of terrestrial material entering aquatic systems [*Knapp et al.*,
559 2008], and we currently have an incomplete understanding of how these patterns will vary across
560 ecosystems and precipitation patterns [*Fang et al.*, 2005; *Knapp et al.*, 2008]. A mechanistic
561 framework for C oxidation that captures impacts of heterogeneity in vegetation in river corridors
562 will therefore aid in predicting how terrestrial-aquatic interfaces respond to ongoing
563 perturbations. Here, we demonstrate that increases in the flux of terrestrial C to aquatic systems
564 may lead to larger mineral-bound C pools by physically and thermodynamically protecting these
565 pools. Conversely, we demonstrate a potential for oxidation of mineral-bound C pools in areas
566 with diminished terrestrial C inputs.

567 Earth System Models (ESMs) depend on mathematical representations of C cycling, and
568 the continued development of these models is tightly coupled to conceptual advances drawn

569 from field-based observations [*Burd et al.*, 2016; *Six et al.*, 2002]. Despite recent progress, these
570 models are still missing key regulatory processes [*Todd-Brown et al.*, 2013; *Wieder et al.*, 2013;
571 *Wieder et al.*, 2015]. In particular, terrestrial-aquatic interfaces are now considered to be active
572 zones of OC transformation as opposed to their representation as non-reactive ‘pipes’ in most
573 ESMs. To address this knowledge gap, we propose a new conceptual framework of OC
574 dynamics based on analysis of *in situ* observational data that explicitly considers a central
575 challenge in model improvement—biochemical, metabolic, and thermodynamic mechanisms
576 governing OC oxidation along terrestrial-aquatic interfaces. Our results directly contrast those
577 expected within a ‘priming’ framework, and we advance that water-soluble thermodynamically
578 favorable OC associated with riparian vegetation protects thermodynamically less favorable
579 bound-OC from oxidation. We also demonstrate differences in biochemical and metabolic
580 pathways associated with metabolism of water-soluble and bound-OC pools in the presence or
581 absence of riparian vegetation, furthering a process-based understanding of terrestrial-aquatic
582 interfaces. Our research provides an opportunity to enhance the mechanistic underpinning of OC
583 oxidation process representations within ESMs and proposes interactions between OC
584 thermodynamics and mineral-inhibition of OC oxidation as a key future research need.

585

586 **Author Contributions.**

587 EBG was responsible for conceptual development and data analysis and was the primary writer
588 with guidance from JCS and MT. ARC, AEG, CTR, ECR, DWK, and JCS were responsible for
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591

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601

602 **References.**

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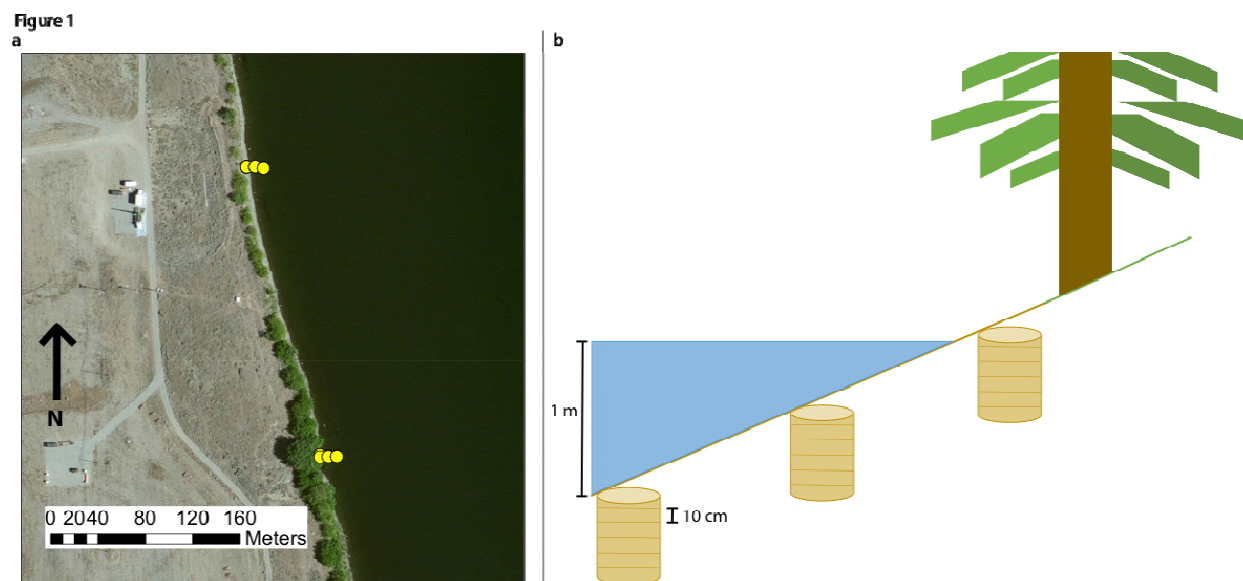
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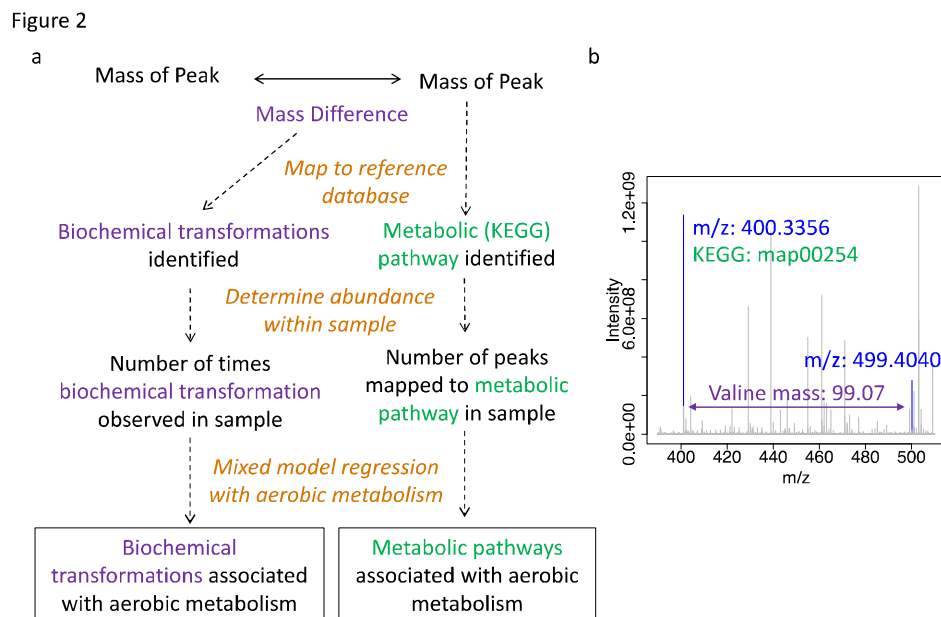
811 **Figures.**



812

813 **Figure 1. Schematic of Sampling Design.** Samples were taken from three cores across an
814 elevation gradient in a transect with or without dense vegetation. Panel (a) displays a satellite
815 image of core locations. Cores without dense vegetation were obtained from the northern
816 transect. Panel (b) shows a schematic of our sampling design. Within each core, samples were
817 taken in 10-cm intervals from 0 to 60 cm. We extracted OC pools from each sample using a
818 sequential extraction protocol (H_2O , MeOH, and $CHCl_3$) to target water-soluble and bound-OC
819 pools.

820



821

822 **Figure 2. Methodology for inferring biochemical transformations and metabolic pathways.**

823 Panel (a) depicts our workflow for analyzing biochemical transformations and metabolic

824 pathways. Biochemical OC transformations (purple) were identified by mapping mass

825 differences in pairwise m/z peak comparisons to a set of 92 known masses transferred in

826 common biochemical transformations (*e.g.*, glucose, amines). Metabolic pathways (green) were

827 identified by mapping all chemical formula assigned to m/z peaks to the KEGG database. Within

828 each sample, the abundance of each biochemical transformation and the number of peaks

829 mapping to each metabolic pathway were then correlated to aerobic metabolism to garner

830 insights into OC oxidation processes. Panel (b) displays an example portion of our FTICR-MS

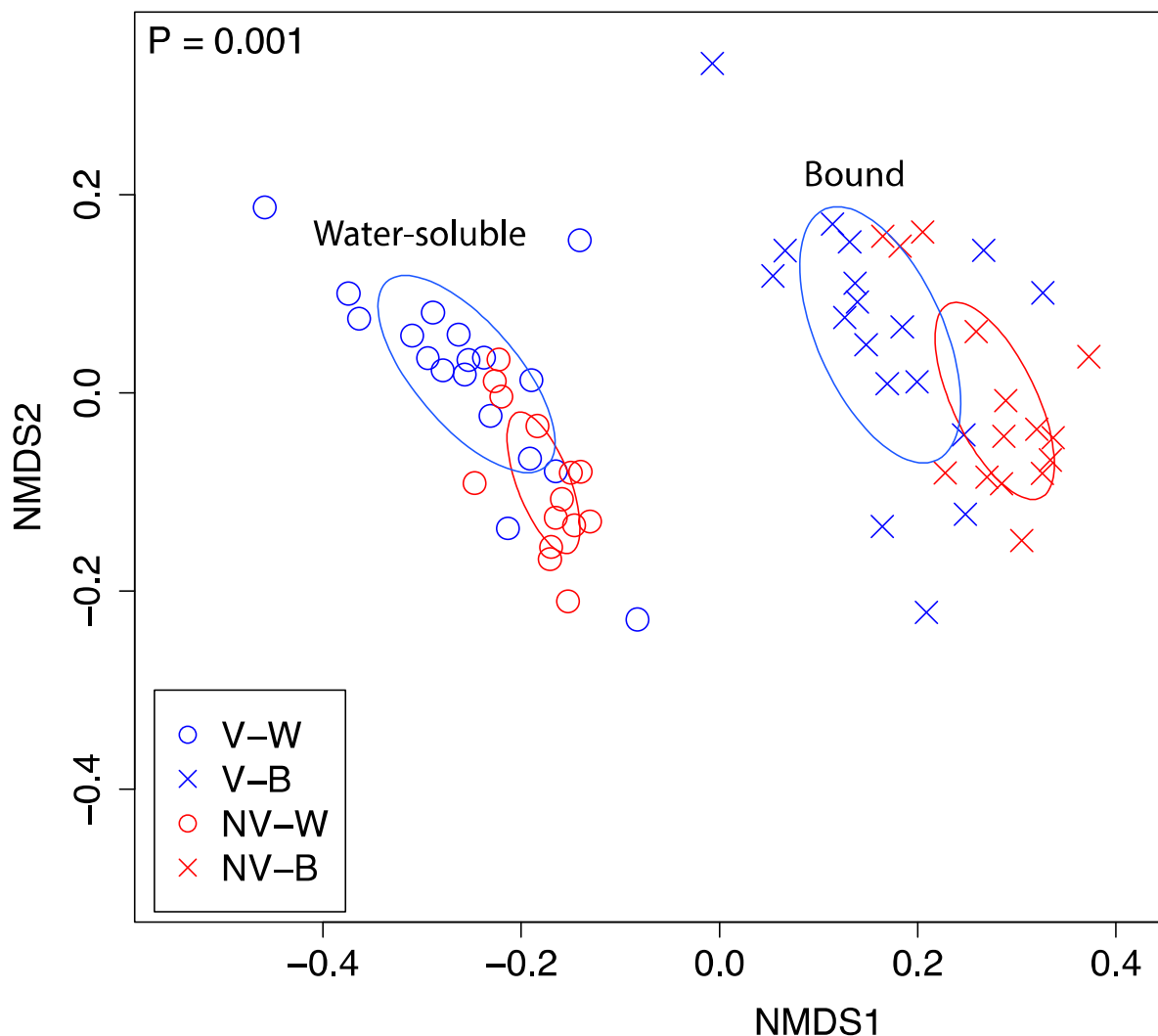
831 spectra overlain with peak assignments (blue), a biochemical transformation (mass difference

832 between peaks, denoted in purple), and a metabolic pathway (associated with the left-hand peak,

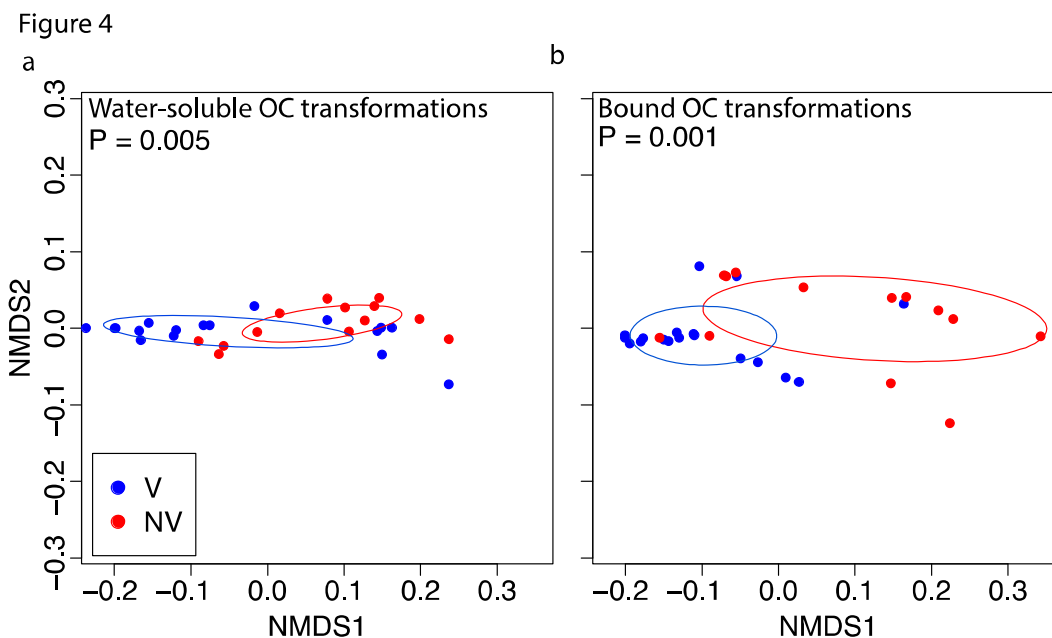
833 denoted in green).

834

Figure 3



835
836 **Figure 3. NMDS visualization of dissimilarity in OC pool composition.** Water-soluble and
837 bound-OC pools are represented by open circles and x's, respectively. Samples associated with
838 riparian vegetation are blue, and those in areas without vegetation are red. The P-value reflects
839 differences among all groups, as assessed by PERMANOVA. Ellipses represent the standard
840 deviation of the average axis scores for each group, generated using the 'ordiellipse' function in
841 the 'vegan' package. Within each extraction, the composition of OC pools was significantly
842 different across vegetation states (both $P = 0.001$).
843



844

845 **Figure 4. NMDS visualization of biochemical transformation partitioning among vegetation**

846 **states.** Biochemical transformations that were correlated to aerobic metabolism were

847 significantly different among vegetation states in both the (a) water-soluble and (b) bound-OC

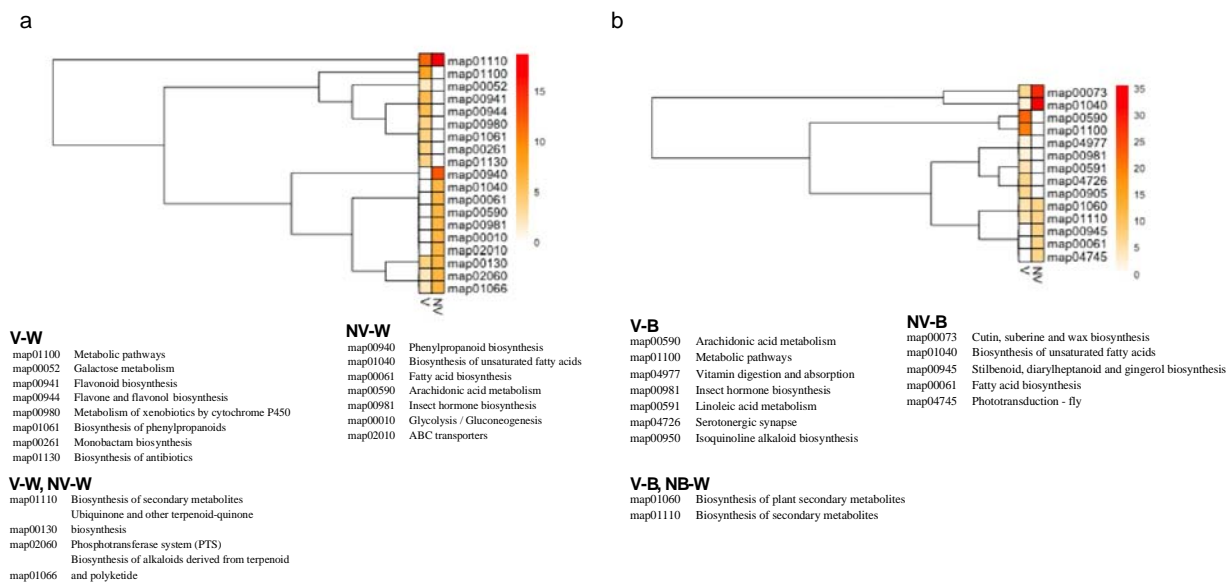
848 pools. V and NV are denoted in blue and red, respectively, and significance values are derived

849 from PERMANOVA. Ellipses represent the standard deviation of the average axis scores for

850 each group, generated using the 'ordiellipse' function in the 'vegan' package.

851

852



853

854 **Figure 5. KEGG pathways associated with aerobic metabolism.** A hierarchical clustering

855 heatmap shows KEGG pathways associated with aerobic metabolism. (a) and (b) show water-

856 soluble and bound-OC pools, respectively. Colors move from white to red on a scale of 0% to

857 20% in (a) and 0% to 35% in (b), showing percent relative abundance of each pathway in each

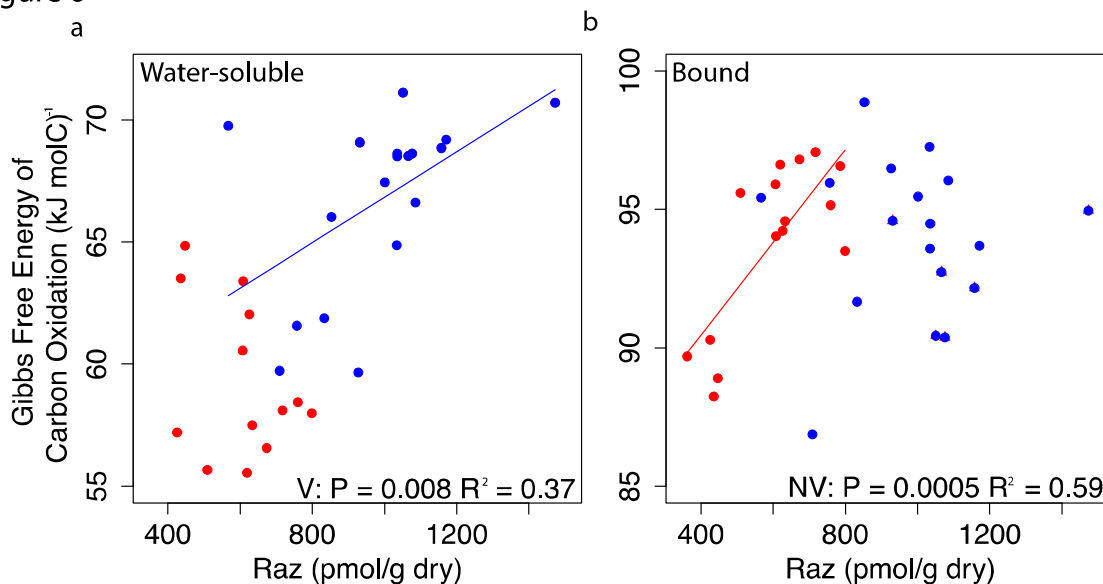
858 group. Pathways are described in the legends. Pathways listed under each condition (e.g., V-W,

859 etc.) had significant correlations to aerobic metabolism within that condition.

860

861

Figure 6



862

863 **Figure 6. Correlations between Gibbs free energy of carbon oxidation ($\Delta G^{\circ}_{\text{Cox}}$) and aerobic**

864 **metabolism.** (a) and (b) display mixed linear regressions relating $\Delta G^{\circ}_{\text{Cox}}$ to aerobic metabolism

865 in water-soluble and bound-OC pools, respectively. Aerobic metabolism is expressed as pmoles

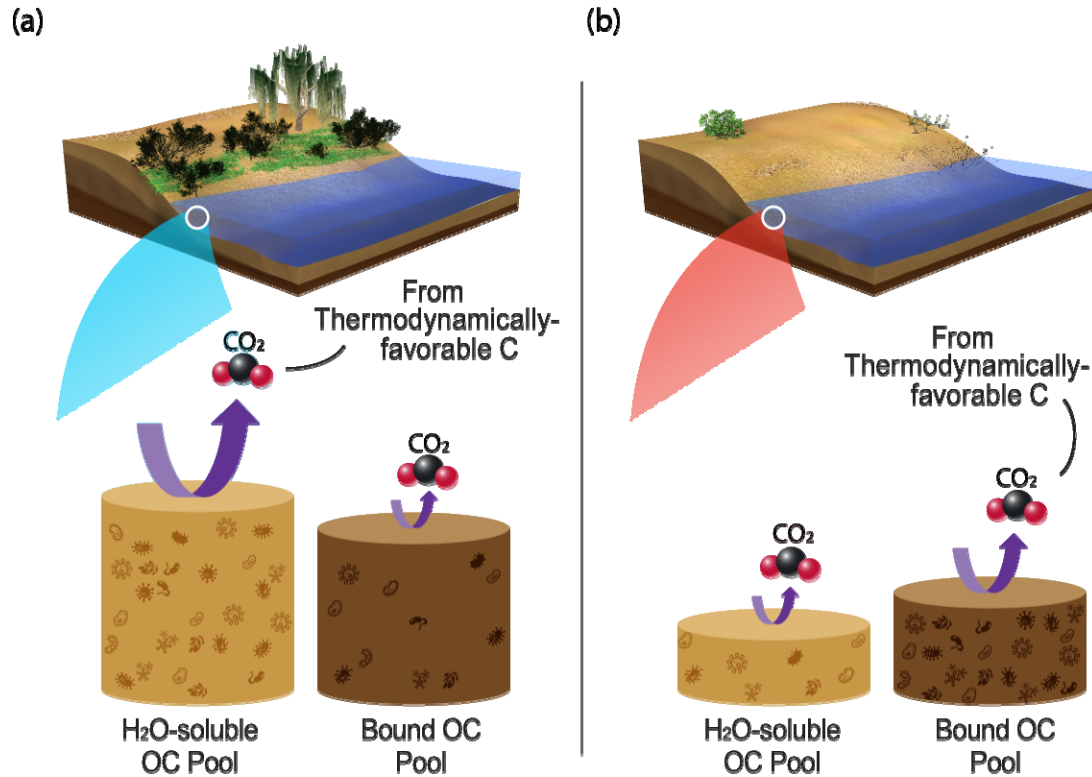
866 of resazurin reduced to resorufin per gram of sediment dry weight over a 48hr incubation period

867 (Raz, see Supplemental Information). V and NV are denoted in blue and red, and solid lines

868 show significant relationships within each vegetation state.

869

870



871

872 **Figure 7. Conceptualization of relationship between riparian vegetation and OC oxidation.**

873 We propose a conceptualization of OC oxidation at terrestrial-aquatic interfaces whereby (a)
874 more riparian vegetation results in greater terrestrial C deposition and larger water-soluble and
875 bound-OC pools. However, water-soluble OC is preferentially oxidized, which protects the
876 bound-OC pool. Conversely, (b) areas deplete in riparian vegetation experience lower inputs of
877 water-soluble OC and show lower rates of OC oxidation. This results in smaller OC pools
878 (water-soluble and bound) and microbial adaptation for oxidation of the bound-OC pool. In both
879 cases, the most thermodynamically favorable portions of the OC pool being metabolized are
880 preferentially oxidized. Height of the cylinders denotes pool sizes, and arrow thickness denotes
881 flux magnitude.

882

883

884 **Table 1. Acronyms and abbreviations used in this paper.**

Abbreviation/Acronym	Description
V	Transect with dense riparian vegetation (i.e., 'vegetated')
NV	Transect with sparse riparian vegetation (i.e., 'not vegetated')
V-W	Transect V, water extraction (water-soluble OC)
V-B	Transect V, chloroform extraction (bound-OC)
NV-W	Transect NV, water extraction (water-soluble OC)
NV-B	Transect NV, chloroform extraction (bound-OC)
C	Carbon
OC	Organic carbon
FTICR-MS	Fourier transform ion cyclotron resonance mass spectrometry
KEGG	Kyoto Encyclopedia of Genes and Genomes
$\Delta G^{\circ}_{\text{Cox}}$	Gibbs free energy of C oxidation

885

886

887 **Table 2. Biochemical transformations and the amount of variation in aerobic metabolism**
 888 **explained by each (R^2), within each OC pool and vegetation state.**

	R^2
V-W	
uridine_5_diphosphate_(-H2O)_C9H12N2O11P2	0.49
cytosine_(-H)_C4H4N3O	0.44
uridine_5_monophosphate_(-H2O)_C9H11N2O8P	0.41
guanine_(-H)_C5H4N5O	0.38
guanosine_(-H2O)_C10H11N5O4	0.37
adenine_(-H)_C5H4N5	0.36
glutathione_(-H2O)_C10H15N3O5S	0.34
uracil_(-H)_C4H3N2O2	0.33
C6H10O6	0.30
glucose_C6H12O6	0.29
Aspartic_Acid_C4H5NO3	0.28
Glucuronic_Acid_(-H2O)	0.28
D-Ribose_(-H2O)_ribosylation)	0.27
secondary_amine	0.26
C6H10O5	0.25
monosaccharide_(-H2O)	0.25
Glutamic_Acid_C5H7NO3	0.24
diphosphate_H3O6P2	0.23
thymine_(-H)_C5H5N2O2	0.23
glyoxylate_(-H2O)_C2O2	0.23
malonyl_group_(-H2O)_C3H2O3	0.22
adenosine_(-H2O)_C10H11N5O3	0.22
primary_amine	0.21
thymidine_(-H2O)_C10H12N2O4	0.21
carbamoyl_P_transfer_(-H2PO4)_CH2ON	0.20
erythrose_(-H2O)	0.20
V-B	
isoprene_addition_(-H)_C5H7	0.39
hydrogenation_dehydrogenation_H2	0.24
biotinyl_(-H2O)_C10H14N2O2S	0.16
NV-W	
Glucuronic_Acid_(-H2O)	0.49
C6H10O6	0.49
acetone_(-H)_C3H5O	0.45
adenine_(-H)_C5H4N5	0.45

889