1   Running title: Biological sedimentary iron oxides

2   Subject: Geomicrobiology and microbial contributions to geochemical cycles

3   **A biological source of marine sedimentary iron oxides**

4   Jacob P. Beam[1*], Jarrod J. Scott[1], Sean M. McAllister[2], Clara S. Chan[2], James McManus[1], Filip

5   J. R. Meysman[2], and David Emerson[1]

[1]Bigelow Laboratory for Ocean Sciences, East Boothbay, Maine, USA, 04544

[2]University of Delaware, Department of Geological Sciences, Newark, Delaware, USA, 19716

[3]Royal Netherlands Institute for Sea Research, Department of Ecosystem Studies, Korringaweg

7, 4401 NT Yerseke, The Netherlands

[*]To whom correspondence should be addressed: jbeam@bigelow.org; PO Box 380, East

Boothbay, ME, USA, 04544; phone: 207-315-2567 extension 406

Abstract

The biogeochemical cycle of iron is intricately linked to numerous element cycles.

Although reductive biological processes that bridge the iron cycle to other element cycles are

established, little is known about microbial oxidative processes on iron cycling in sedimentary

environments—resulting in the formation of iron oxides. Here, we show that a major source of

sedimentary iron oxides originates from the metabolic activity of iron-oxidizing bacteria from

the class Zetaproteobacteria, stimulated by burrowing animals in coastal sediments.

Zetaproteobacteria were estimated to be a global total of 10^{26} cells in coastal, bioturbated

sediments and would equate to an annual production of approximately 7.9×10^{15} grams of

sedimentary iron oxides—twenty-five times larger than the annual flux of iron oxides by rivers.

These data suggest that iron-oxidizing Zetaproteobacteria are keystone organisms in marine

sedimentary environments given their low numerical abundance; yet exert a profound impact via

the production of iron oxides.

Keywords: iron-oxidizing bacteria/Zetaproteobacteria/bioturbation/iron

biogeochemistry/keystone organism

Main

Iron oxides are important components of coastal and continental shelf sediments, and are

thought to originate primarily by river deposition (Poulton and Raiswell, 2002). Authigenic iron

oxides can also be formed by the oxidation of ferrous iron [Fe(II)], which has been largely

attributed to chemical oxidation of Fe(II) in pore waters in areas with significant sediment

mixing and irrigation by animals—bioturbation and bioirrigation—and subsequent reaction with

oxygen in anoxic sediments (Aller, 1982; Canfield, 1989). Although sedimentary chemical

oxidation of iron is important under saturated oxygen conditions at neutral pH (Canfield, 1989),

bioirrigated sediments contain microenvironments—formed by burrowing animals—that are well

below saturation (Kristensen and Kostka, 2005) where the biological contribution to iron

oxidation is quantitatively more significant (Emerson *et al.*, 2010). Under low-oxygen conditions

(< 100 µM O_2) and without rapid replenishment of highly reactive, poorly-crystalline iron oxides

(i.e., ferrihydrite and lepidocrocite), they would be quickly exhausted by hydrogen sulfide and by

bacterial iron reduction, and form iron sulfides or be released to the water column. Collectively,

these findings suggest that biology is involved and important in marine sedimentary iron

oxidation under ferruginous conditions commonly observed in bioturbated sediments.

The Zetaproteobacteria represent a class of iron-oxidizing bacteria (FeOB) that are

exclusively found in marine or saline-influenced environments that contain high ferrous iron

[Fe(II)] concentrations (McAllister *et al.*, 2011; McBeth *et al.*, 2013; Scott *et al.*, 2015, 2017).

Coastal marine sediments can have Fe(II) pore water concentrations ranging from ~1-2,000 µmol

L^{-1} that are capable of supporting lithoautotrophic populations of Zetaproteobacteria (Emerson,

2016). Recent studies have identified Zetaproteobacteria in surface openings of benthic

macrofauna in the Mediterranean Sea (Rubin-Blum *et al.*, 2014), worm burrows in sub-marine

groundwater discharge into sands in Delaware (McAllister *et al.*, 2015), coastal sediments in

Denmark (Laufer *et al.*, 2016), and Baltic and North Sea sediments (Reyes *et al.*, 2016). These

recent studies suggest that Zetaproteobacteria may play a significant role in iron oxidation in

marine sediments—a quantitative estimate of their abundance is necessary to determine their

biogeochemical role on a global scale.

We analyzed the coastal sediment microbial communities to determine the extent of

Zetaproteobacteria from geographically diverse sites (n=90; Supplementary Table 1) utilizing

16S rRNA gene sequencing, which highlights their importance on a global scale. Iron oxidation

appears ubiquitous in the Zetaproteobacteria (Field *et al.*, 2015; Scott *et al.*, 2015; Barco *et al.*,

2015), thus the 16S rRNA gene can be used to infer this specific metabolism. We also enriched

for environmentally relevant iron-oxidizing bacteria from coastal sediments, which provided

further metabolic evidence of the importance of iron oxidation in marine sediments. A meta-

analysis of 16S rRNA gene studies revealed the extent and importance of Zetaproteobacteria on

the global sedimentary iron biogeochemical cycle.

We identified Zetaproteobacteria in 60 % of our samples (Supplementary Table S1), and

their median relative abundance was 1.1 percent of the total microbial community (range = 0.04-

15 %) in worm (e.g., polychaetes) burrows in coastal marine sediments (Fig. 1a).

Zetaproteobacteria were ten times less abundant in bulk sediments (Fig. 1a) with a median near

zero percent (range = 0-1 %), and were statistically different from worm burrows (p-value=9.2 x

10^{-7}, Wilcoxon test). The large ranges (0.04-15 %) and non-normal distribution of

Zetaproteobacteria relative abundance in worm burrows (see Supplementary Materials and

Methods) was most likely a combination of differences in burrow ventilation rates and

efficiencies (Kristensen and Kostka, 2005), differences in sediment physicochemical conditions

(Supplementary Table S1), and sampling bias (for example, residual sediment on worm

burrows). Bioirrigation by benthic animals increases the extent of oxidative processes in these

sediments, thus biological iron oxidation can occur at greater depths (10s of centimeters) than

typical oxygen penetration of a few millimeters into coastal surface sediments. The abundance of

Zetaproteobacteria at the burrow walls correlated with the concentration of pore water ferrous

iron [Fe(II)] (Fig. 1b), which is their main energy source—resulting in the production of solid

phase iron oxides around worm burrows (Fig. 1c). Quantitatively, highly reactive, iron oxides—

operationally extractable by sodium dithionite (Poulton and Canfield, 2005)—were 3 times

higher at burrow walls, which accounted for 20-40 % of the iron oxides with depth

(Supplementary Figure S1 and S2). These freshly-produced iron oxides are important substrates

for iron-reducing microorganisms that release Fe(II) into pore waters, and are essential to the

supply of dissolved iron (dFe) to phytoplankton in the water column in coastal communities and

continental shelves (Severmann *et al.*, 2010). Iron oxides are also important to the mineralization

of organic matter in marine sediments by iron reducers (Canfield, 1989) substrates for early

pyrite diagenesis (Berner, 1984), enhance organic matter burial (Lalonde *et al.*, 2012), and

inhibit accumulation of pore water hydrogen sulfide, preventing conditions detrimental to

benthic animals (Kristensen and Kostka, 2005), perhaps functioning as a local firewall against

euxinic conditions(Seitaj *et al.*, 2015).

The relative abundance of Zetaproteobacteria in worm burrows resembles the Fe(II)

concentration profile (Supplementary Figure S3), and both were at their maximum values around

2-3 cm. The high Fe(II) (~40-140 μM) and low oxygen (~20-60 μM) conditions present in

bioturbated sediment pore waters (Supplementary Figure S3) were ideal habitats for

microaerophilic Zetaproteobacteria to thrive (Emerson *et al.*, 2010). Zetaproteobacteria relative

abundance decreased with depth in both burrows and sediments (Supplementary Fig. S3), likely

due to the decrease in Fe(II) with depth and increase in hydrogen sulfide production by sulfate-

reducing bacteria. Although there is oxygen in these sediments at depth, hydrogen sulfide may

inhibit oxygen respiratory machinery under these conditions. The formation of iron sulfide

minerals with increasing depth by biogenic sulfide may also compete with Zetaproteobacteria for

access to Fe(II). Under these ferruginous settings, biotic rates of Fe(II) oxidation exceed abiotic

chemical oxidation (Emerson *et al.*, 2010).

Two Zetaproteobacteria Operational Taxonomic Units (herein, referred to as ZetaOTUs)

dominated the Zetaproteobacterial diversity in worm burrows (Supplementary Table S2). The

dominant ZetaOTU across all samples was ZetaOTU14, which comprised 32 % of all ZetaOTUs

(Supplementary Table S2), and is represented by four single cell amplified genomes (SAGs)

from diffuse flow vent systems (Field *et al.*, 2015; Scott *et al.*, 2015, 2017). We isolated the first

member of ZetaOTU14, strain CSS-1 from iron oxide surface flocculent in a laboratory

bioturbation microcosm (Supplementary Figure S4). This strain grew best under low oxygen

(~60 μM O_2) and high Fe(II) concentrations similar to those measured from sediment pore

waters (Supplementary Figure S3). Strain CSS-1 produced stalks encrusted with poorly-

crystalline iron oxides under laboratory conditions (Supplementary Figure S4). These iron oxides

are consistent with those produced by other Zetaproteobacteria (Chan *et al.*, 2010), as well as in

naturally occurring iron mats associated with hydrothermal vents, which are highly reactive, and

resistant to undergoing diagenesis to more crystalline oxides (e.g., goethite) (Picard *et al.*, 2015).

Single cell genomes from ZetaOTU14 representatives contained genes essential for growth on

iron and low oxygen conditions (Supplementary Table S2). The second most abundant OTU was

ZetaOTU9 (22 %) and is represented by two cultured isolates (*Ghiorsea bivora* strains TAG-1

and SV108) (Mori *et al.*,, 2017), as well as 5 SAGs from deep-sea vents (Supplementary Table

S2). ZetaOTU9 isolates also had genes necessary for growth on iron and low oxygen

(Supplementary Table S2), and have also been shown to oxidize hydrogen, which may explain

the ubiquity of this OTU in sediments and other environments (see below). There was no clear

distribution of ZetaOTUs 14 and 9 with respect to depth (Supplementary Figure S3) in worm

burrows and sediments—although it is likely that they are adapted for specific Fe(II) and O_2

concentrations, which were hypothesized for other Zetaproteobacteria (Field *et al.*, 2015).

We searched for Zetaproteobacterial 16S rRNA gene sequences in marine sediment

datasets (Supplementary Table S3), and identified them in numerous sediments on a global scale

(Figure 2). We found a pattern consistent with our samples—ZetaOTUs 14 and 9 were present

and generally the most abundant ZetaOTUs in coastal and shelf sediments (Figure 2).

Zetaproteobacteria relative abundance was not found to exceed one percent in other studies, as

microenvironments were not considered, which are abundant in bioturbated sediments

(Kristensen and Kostka, 2005). Accordingly, we hypothesize that when the abundance of

Zetaproteobacteria exceeds ~0.1 % in sediments, there is active growth and iron oxidation

associated with bioturbating and bioirrigating animals. We estimated a median global population

size of Zetaproteobacteria to be 1.05×10^{26} cells (range = 3.83×10^{24}-1.44×10^{27} cells) from our

measurements (Fig. 1a) and from other studies (Fig. 2) in continental shelf sediments using

cellular abundance in the upper 10 centimeters—the worldwide average depth of bioturbation

(Boudreau, 1998)—of continental shelf environments (total cells =10^{29}; <150 m water depth

(Kallmeyer *et al.*, 2012)). The global Zetaproteobacteria abundance estimate was then used with

recent iron oxide production rate measurements from diffuse flow vents (~1.3×10^{-16} mol Fe cell⁻

¹ hr⁻¹) (Emerson *et al.*, 2016), which could result in the production of ~7 petagrams of iron

oxides per year (range = 0.1-70 petagrams per year). Recent two-dimensional, sub-millimeter

Fe(II) measurements in bioturbated sediments revealed extensive Fe(II) oxidation occurring

within the immediate vicinity of worm burrows and a rapid re-oxidation rate of 3.78 ±1.4 mmol

Fe m^{-2} day^{-1} (de Chanvalon *et al.*, 2017). These chemical rate measurements combined with an

estimate of the global volume of bioturbated coastal sediments 10 cm deep (~2.1 x 10^{13} m^3)

(Teal *et al.*, 2008) would equate to an annual production of 1.6 ± 1.1 petagrams of iron oxides.

These two independent estimations of iron oxide production rates are well within the range of

one another. Based on these estimates, the annual biological oxidation of iron in sediments—

forming iron oxides—could exceed the annual flux of iron oxides from rivers to coastal

sediments (Poulton and Raiswell, 2002) up to a factor of twenty-five.

Zetaproteobacteria exert a profound impact on global sedimentary biogeochemistry via

the production of biogenic, highly-reactive iron oxides despite their low global abundance (~0.11

%)—effectively functioning as keystone organisms in coastal sediments stimulated by burrowing

animals. Zetaproteobacteria contribute significantly to the rapid rates of Fe(II) re-oxidation

measured and observed in coastal sediments (Chanvalon *et al.*, 2017). Climate change outcomes

such as coastal hypoxia may have positive or negative effects on the sedimentary iron

biogeochemical cycle—either stimulating microaerobic bacterial iron oxidation resulting in an

increase in iron oxide production, thus enhancing dFe release or inhibiting oxidation by the

increase in hydrogen sulfide production, precipitating Fe as iron sulfides. The result of an

increase or decrease in dFe flux would be enhanced or reduced primary productivity by

phytoplankton, respectively. Thus, sedimentary iron oxide formation by Zetaproteobacteria may

have a direct impact on important water column processes such as carbon and nitrogen fixation.

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References

Aller RC. (1982). The Effects of Macrobenthos on Chemical Properties of Marine Sediment and

Overlying Water. In: McCall PL, Tevesz MJS (eds). *Animal-Sediment Relations: The Biogenic*

Alteration of Sediments. Springer US: Boston, MA, pp 53–102.

Barco RA, Emerson D, Sylvan JB, Orcutt BN, Jacobson Meyers ME, Ramírez GA, *et al.* (2015).

New insight into microbial iron oxidation as revealed by the proteomic profile of an obligate

iron-oxidizing chemolithoautotroph. *Appl Environ Microbiol* **81**: 5927–5937.

Berner A. (1984). Sedimentary pyrite formation: An update. *Geochim Cosmochim Acta*

Boudreau BP. (1998). Mean mixed depth of sediments: The wherefore and the why. *Limnol*

Oceanogr **43**: 524–526.

Canfield DE. (1989). Reactive iron in maine sediments. *Geochim Cosmochim Acta* **53**: 619–632.

Chan CS, Fakra SC, Emerson D, Fleming EJ, Edwards KJ. (2010). Lithotrophic iron-oxidizing

bacteria produce organic stalks to control mineral growth: implications for biosignature

formation. *ISME J* **5**: 717–727.

de Chanvalon AT, Metzger E, Mouret A, Knoery J, Geslin E, Meysman FJR. (2017). Two

dimensional mapping of iron release in marine sediments at submillimetre scale. *Mar Chem*.

191: 34-49.

Emerson D. (2016). The irony of iron - Biogenic iron oxides as an iron source to the ocean.

Front Microbiol **6**: 1–6.

Emerson D, Fleming EJ, McBeth JM. (2010). Iron-oxidizing bacteria: an environmental and

genomic perspective. *Annu Rev Microbiol* **64**: 561–583.

Emerson D, Scott JJ, Leavitt AH, Fleming E, Moyer CL. (2016). In situ estimates of iron-

oxidation and accretion rates for iron-oxidizing bacterial mats at Loihi Seamount. *bioRxiv*.

Field EK, Sczyrba A, Lyman AE, Harris CC, Woyke T, Stepanauskas R, *et al.* (2015). Genomic

insights into the uncultivated marine Zetaproteobacteria at Loihi Seamount. *ISME J* **9**: 857–870.

Kallmeyer J, Pockalny R, Adhikari RR, Smith DC, D'Hondt S. (2012). Global distribution of

microbial abundance and biomass in subseafloor sediment. *Proc Natl Acad Sci* **109**: 16213–

16216.

Kristensen E, Kostka JE. (2005). Macrofaunal Burrows and Irrigation in Marine Sediment:

Microbiological and Biogeochemical Interactions. In: Kristensen E (ed) *Interact Between*

Macro- Microorg Mar Sediments. American Geophysical Union Washington, DC, pp. 125–157.

Lalonde K, Mucci A, Ouellet A, Gelinas Y. (2012). Preservation of organic matter in sediments

promoted by iron. *Nature* **483**: 198–200.

Laufer K, Nordhoff M, Schmidt C, Behrens S, Jørgensen BB, Kappler A. (2016). Co-existence

of microaerophilic, nitrate-reducing, and phototrophic Fe(II)-oxidizers and Fe(III)-reducers in

coastal marine sediment. *Appl Environ Microbiol* **82**: 1433–1447.

McAllister SM, Barnett JM, Heiss JW, Findlay AJ, MacDonald DJ, Dow CL, *et al.* (2015).

Dynamic hydrologic and biogeochemical processes drive microbially enhanced iron and sulfur

cycling within the intertidal mixing zone of a beach aquifer. *Limnol Oceanogr* **60**: 329–345.

McAllister SM, Davis RE, McBeth JM, Tebo BM, Emerson D, Moyer CL. (2011). Biodiversity

and emerging biogeography of the neutrophilic iron-oxidizing Zetaproteobacteria. *Appl Environ*

Microbiol **77**: 5445–5457.

McBeth JM, Fleming EJ, Emerson D. (2013). The transition from freshwater to marine iron-

oxidizing bacterial lineages along a salinity gradient on the Sheepscot River, Maine, USA.

Environ Microbiol Rep **5**: 453–463.

Mori JF, Scott JJ, Hager KW, Moyer CL, Küsel K, Emerson D (2017). Physiological and

ecological implications of an iron- and hydrogen-oxidizing member of the Zetaproteobacteria,

Ghiorsea bivora, gen. nov. sp. nov. Accepted.

Picard A, Kappler A, Schmid G, Quaroni L, Obst M. (2015). Experimental diagenesis of organo-

mineral structures formed by microaerophilic Fe(II)-oxidizing bacteria. *Nat Commun* **6**: 6277.

Poulton SW, Canfield DE. (2005). Development of a sequential extraction procedure for iron:

Implications for iron partitioning in continentally derived particulates. *Chem Geol* **214**: 209–221.

Poulton SW, Raiswell R. (2002). The low-temperature geochemical cycle of iron: From

continental fluxes to marine sediment deposition. *Am J Sci* **302**: 774–805.

Reyes C, Dellwig O, Dahnke K, Gehre M, Noriega-Ortega BE, Bottcher ME, *et al.* (2016).

Bacterial communities potentially involved in iron-cycling in Baltic Sea and North Sea

sediments revealed by pyrosequencing. *FEMS Microbiol Ecol* **92**.

Rubin-Blum M, Antler G, Tsadok R, Shemesh E, Austin JA, Coleman DF, *et al.* (2014). First

evidence for the presence of iron oxidizing zetaproteobacteria at the levantine continental

margins. *PLoS One* **9**: 1–10.

Scott JJ, Breier JA, Luther GW, Emerson D. (2015). Microbial iron mats at the mid-atlantic ridge

and evidence that zetaproteobacteria may be restricted to iron-oxidizing marine systems. *PLoS*

One **10**: 1–19.

Scott JJ, Glazer BT, Emerson D. (2017). Bringing microbial diversity into focus: high-resolution

analysis of iron mats from the Lōʻihi Seamount. *Environ Microbiol* **19**: 301-316.

Seitaj D, Schauer R, Sulu-Gambari F, Hidalgo-Martinez S, Malkin SY, Burdorf LDW, *et al.*

(2015). Cable bacteria generate a firewall against euxinia in seasonally hypoxic basins. *Proc Natl*

Acad Sci U S A **112**: 13278-13283.

Severmann S, McManus J, Berelson WM, Hammond DE. (2010). The continental shelf benthic

iron flux and its isotope composition. *Geochim Cosmochim Acta* **74**: 3984–4004.

Teal LR, Bulling MT, Parker ER, Solan M. (2008). Global patterns of bioturbation intensity and

mixed depth of marine soft sediments. *Aquat Biol* **2**: 207–218.

266   **Figure Legends**

Figure 1. Boxplots of the relative abundance of Zetaproteobacteria (**a**) in iron-oxide lined worm

burrow walls (n=29) and surrounding sediments (n=61). Notches are representative of 95 %

confidence interval and the medians (solid black lines) between worm burrows and sediments

(1.1 % and 0 %, respectively) are statistically different (p-value=9.2 x 10^{-7}, Wilcoxon test).

Filled circles represent individual data points and open circles indicate outliers.

Zetaproteobacteria relative abundance (%) as a function of pore water ferrous iron [Fe(II)]

concentration (μmol L^{-1}) (**b**) from worm burrows (blue circles, fitted blue line, orange fill = 95 %

confidence interval) and sediments (black circles, fitted black line, grey fill = 95 % confidence

interval) (see Methods for details on line fits). Characteristic iron oxide lined worm burrow walls

(c) from "The Eddy", Sheepscot River, Maine, USA (image from 27 August 2015). Burrow

walls are likely created by the polychaete, *Nereis diversicolor* or hemichordate, *Saccoglossus*

kowalevskii, which are both common to these intertidal sediments in Maine.

Figure 2. Global distribution of Zetaproteobacteria in marine sediments (circles) and non-

sediment sites (triangles) such as hydrothermal vents. The relative abundance of

Zetaproteobacteria in sediments from other 16S rRNA gene studies was never above 1 % and

was typically within the range measured from bulk marine sediments from Maine. Sequences are

from numerous studies (Supplementary Table S3) that include Sanger, 454, and Illumina

sequencing technologies.

Figure 1

Figure 2