

1 **Evidence for a growth zone for deep subsurface microbial clades in near-**
2 **surface anoxic sediments**

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11 **Abstract**

12 Global marine sediments harbor a large and highly diverse microbial biosphere, but the
13 mechanism by which this biosphere is established during sediment burial is largely
14 unknown. During burial in marine sediments, concentrations of easily-metabolized
15 organic compounds and total microbial cell abundance decrease steadily. However, it is
16 unknown whether some microbial clades increase with depth, despite the overall trend of
17 abundance decrease. We show total population increases in 38 microbial families over 3
18 cm of sediment depth in the upper 7.5 cm of White Oak River (WOR) estuary sediments.
19 Clades that increased with depth were more often anaerobic, uncultured, or common in
20 deep marine sediments relative to those that decreased. Minimum turnover times (which
21 are minimum *in situ* doubling times of growth rates) were estimated to be 2-25 years by
22 combining sedimentation rate with either quantitative PCR (qPCR) or the product of the
23 Fraction Read Abundance of 16S rRNA genes and total Cell counts (FRAXC). Turnover
24 times were within an order of magnitude of each other in two adjacent cores, as well as in
25 two laboratory enrichments of Cape Lookout Bight (CLB), NC, sediments (average
26 difference of $28 \pm 19\%$). qPCR and FRAXC in WOR cores and FRAXC in CLB

27 incubations produced similar turnover times for key deep subsurface uncultured clades
28 *Bathyarchaeota* (8.7 ± 1.9 years) and *Thermopfundales*/MBG-D (4.1 ± 0.7 years). We
29 conclude that common deep subsurface microbial clades experience a narrow zone of
30 growth in shallow sediments, offering an opportunity for natural selection of traits for
31 long-term subsistence after resuspension events.

32

33 **Significance statement:** The current dogma is that the deeply-branching uncultured
34 microbes that dominate global marine sediments do not actually increase in population
35 size as they are buried in marine sediments – rather they exist in a sort of prolonged
36 torpor for thousands of years. This is because no evidence has ever been found that these
37 clades actually increase population sizes, or grow, as they are gradually buried. We
38 discovered that they actually do increase population sizes during burial, but only in the
39 upper few centimeters. This changes our dogma about marine sediments as a vast
40 repository of non-growing microbes, to a vast repository of non-growing microbes with a
41 thin and relatively rapid area of growth in the upper 10 centimeters.

42

43 **Introduction**

44 Marine sediments harbor one of the largest microbiomes on Earth, comprising
45 $\sim 2.9 \times 10^{29}$ microbial cells (1). In most marine sediments, total microbial cells decrease
46 with depth in a log-log relationship since the energy available to support the microbial
47 community decreases as resources are depleted (1, 2). Therefore most microbes in marine
48 sediments are nutrient-limited and in a state of near-zero growth (3, 4). Bulk sedimentary

49 microbial communities have been calculated to have turnover times on the order of tens
50 to hundreds of years, based on microbial respiration rates and energetic requirements (5,
51 6). The assembly of these deep subsurface communities over many meters of burial
52 appears to occur through selective survival (7) or environmental filtering (8), rather than
53 net growth. Only the most abundant organisms, or those best suited to the changing redox
54 regime, persist while all other clades die off (7–10). However, marine sediment microbes
55 appear to have adaptations to marine sediments, since enzyme affinity (lower K_m values)
56 to carbon substrates increases with depth and specialization shifts to highly-degraded
57 substrates with depth (11). These adaptations imply that heritable natural selection has
58 occurred in a sedimentary environment, which implies that these common subsurface
59 clades experience net growth, rather than just persistence, at some point during burial.
60 This net growth may occur in the upper few centimeters of sediments, since microbes
61 have been shown to grow and increase biomass in upper sedimentary layers (12, 13).
62 However, the methods used in these growth studies could not distinguish between
63 different clades nor whether this growth resulted in net population increases with
64 increasing depth. We therefore hypothesized that growth occurs for common subsurface
65 clades in the upper few centimeters of marine sediments, over relatively short time
66 intervals of a few tens of years. This growth would have been missed by previous studies
67 that demonstrated persistence rather than growth of subsurface clades during burial (7–
68 10) because these studies were conducted in deeper sediments, after net growth had
69 largely ceased.

70 Microbial populations in marine sediments are phylogenetically diverse (14), so
71 clades may differ in their response to burial. One option is that the cells from each clade

72 die-off at similar rates, with no evidence of growth (Fig. 1). The second option is that
73 some clades experience growth as they are buried, allowing them to increase their
74 population size, even as whole-community cell numbers decrease (Fig. 1). In the second
75 option, the rate of increase in total population size, or turnover time, therefore represents
76 a minimum growth rate estimate. The maximum growth rate could be even higher since
77 growth and death rates may be in equilibrium at each depth layer. Marine sediments
78 contain many phylogenetically divergent non-cultured cells (PDNC) belonging to high-
79 level taxonomic groups with no cultured representatives (15). We hypothesized that 1)
80 some PDNC clades increase in cell abundance with depth in the upper 10 cm of coastal
81 sediments, 2) their turnover times are slower than the < 24 hours doubling times of most
82 microbial cultures, and 3) turnover times are specific to each clade, and are therefore
83 repeatable in different sediment cores and laboratory incubations. We measured cell
84 abundance changes for individual microbial clades with depth in the upper 10 cm (or 40
85 years of burial time) in the White Oak River (WOR) estuary. Like most marine
86 sediments, they contain diverse PDNCs, and have a steady sediment deposition rate
87 allowing a conversion between depth and time (16). As a second test, we measured the
88 growth of these organisms in 2.2 year-long incubations of sediments from Cape Lookout
89 Bight (CLB) without any nutrient additions so that time could be measured directly (17).

90 Measuring microbial growth rates in natural marine sediments is not
91 straightforward. Quantification methods, such as catalyzed reporter deposition
92 fluorescent in situ hybridization (CARD-FISH) and quantitative PCR (qPCR) are often
93 inaccurate in marine sediments, due to problems with permeabilizing cells and
94 amplification biases (18, 19). These methods are also limited because they only measure

95 the microbial groups targeted by the chosen probes or primers. Measuring the relative
96 abundance of 16S rRNA gene sequence reads amplified from marine sediments
97 encompasses a wider range of phylogenetic groups, but also does not provide accurate
98 quantifications since DNA extraction methods and primer sequences are biased (20).

99 However, when considering relative population increases, these biases cancel
100 each other out (Fig. 2). To quantify the relative population changes with depth, one can
101 multiply the Fraction of 16S rRNA gene Read Abundance by the total Cell counts
102 (FRAx C) at each depth or timepoint. The ratio of FRAx C at different depths or
103 timepoints has the bias term in both the numerator and the denominator, so it cancels out,
104 allowing the quantification of relative changes in population size (i.e., a 3-fold increase or
105 a doubling in the example in Fig. 2), but not the exact number of cells that have been
106 gained or lost, over the time interval (18). A second measurement for turnover time was
107 made using qPCR assays, which can also be used to calculate relative population turnover
108 times (18). We found that PDNC and clades that could be inferred to be anaerobic
109 increased with population turnover times of 2-25 years, with good replication for each
110 clade with FRAx C and qPCR, between duplicate cores from the WOR estuary, and
111 between cores from the WOR estuary and CLB sediment incubations. Close relatives of
112 cultures and aerobic organisms decreased with population half lives that were not
113 replicable through the different measurements.

114

115 **Methods**

116

117 **Sample acquisition and treatment.** Four sediment cores of ~10 cm depth each were
118 retrieved from Station H, WOR estuary (34 44.490' N, 77 07.44' W) May 16, 2016, in
119 about 1.5 m water depth, with water salinity of 11 ppt (cores 30, 21, 32, and 34). Each
120 core was sectioned into 1 cm intervals later that day at room temperature (roughly equal
121 to the measured *in situ* water temperature of 25°C) at the University of North Carolina
122 Institute for Marine Sciences. Two cores (cores 30 and 32) were sectioned for qPCR and
123 16S rRNA gene libraries, two cores (cores 34 and 31, which were taken adjacent to cores
124 30 and 32, respectively) were sectioned for sulfate concentrations and porosity. The core
125 for cell counts was taken May 28, 2013 (core 7) from the same site (34°44.482'N,
126 77°7.435'W), and was sectioned in 3 cm intervals. The top 3 cm of marine sediment from
127 CLB (34.6205°N, 76.5500°W) were collected from 10 m water depth on October 2, 2013
128 and combined into a 2L Erlenmeyer flask that was incubated anoxically for 802 days, as
129 reported previously (17), without any amendment of substrates.

130 **Geochemistry.** To measure sulfate, a 15 ml plastic tube was filled completely with
131 sediment and centrifuged at 5,000x g for 5 minutes. A syringe was used to remove the
132 supernatant below the air interface. The porewater was filtered using a 0.2 µm syringe
133 filter into 100 µl of 10% HCl to a final volume of 1 ml. Porewater sulfate concentrations
134 were determined by ion chromatography (Dionex, Sunnyvale, CA). Porosity was
135 determined by comparing dry and wet weights as described previously (21).

136 **DNA extraction, qPCR, sequencing, and data analysis.** Each 1 cm depth
137 interval of cores 30 and 32 was placed into a Whirlpak bag and frozen immediately on
138 dry ice for later DNA extraction after storage at -80°C. DNA was extracted from frozen
139 sediments (MoBio RNA Powersoil kit with DNA accessory) and enumerated on a

140 NanoDrop 3000. At the timepoints from the CLB incubations, DNA was extracted from
141 frozen samples using the FastDNA kit for Soil (MP Bio, Santa Ana, CA). For both the
142 WOR cores and the CLB incubations, the V4 region of each DNA extract was amplified
143 using the universal primers 515f and 806r (22), prepared via Nexterra kit and sequenced
144 at the Center for Environmental Biotechnology at the University of Tennessee
145 (Knoxville, TN) on an Illumina MiSeq. 16S rRNA gene reads were processed, chimera-
146 checked, and classified via Silva taxonomy (v126; (23)) in mothur (24). All calculations
147 were performed on clades at the fifth taxonomic level, which is roughly the family level,
148 only from clades with >150 total reads in the Miseq run, which left 327 clades from core
149 30 and 288 clades from core 32. 16S rRNA gene reads were deposited at the Short Read
150 Archive (SRA) at NCBI's Genbank with BioProject ID PRJNA614649 for WOR
151 sediments and PRJNA321388 for CLB incubations.

152 Quantitative PCR (qPCR) was used to determine the 16S rRNA gene copy
153 numbers of bacteria, archaea, *Bathyarchaeota*, and Marine Benthic Group – D in the
154 *Thermoplasmata* (with the latter for core 32 only) using the Quantifast SYBRGreen kit
155 (Qiagen) on a BioRad Opticon2 thermocycler. Since absolute copy numbers cannot be
156 accurately measured for DNA in marine sediments (18, 25), only the relative changes in
157 copy number were measured. Our primers were MCG528F and MCG732R for
158 *Bathyarchaeota* (26), MBG-D322F and MBG-D569R for Marine Benthic Group – D
159 (27), ARCH915F (28) and ARC1059R (29) for archaea, and BAC340F and BAC515R
160 for bacteria (30).

161 **Population turnover time calculations.** Population turnover times for a family-level
162 classification were calculated from relative 16S rRNA gene abundances and total cell
163 counts as described previously (17), with the following relationship:

$$164 \quad t_{d,i} = \frac{t \ln(2)}{\ln\left(\frac{FRA_{i,t} C_{tot,t}}{FRA_{i,0} C_{tot,0}}\right)} \quad \text{Equation 1}$$

165 where $t_{d,i}$ is the doubling time of the i th family clade, t is the elapsed time, $FRA_{i,t}$ is the
166 fraction read abundance of the i th clade at time t , and $C_{tot,t}$ is the total number of cells at
167 time t . The product of the fraction 16S rRNA gene abundance of a particular clade and
168 the total cell abundance will be referred to as the FRAxC of a particular clade. To
169 calculate the population turnover for a specific group of target organisms using qPCR, we
170 used the following relationship:

$$171 \quad t_{d,i} = \frac{t \ln(2)}{\ln\left(\frac{copies_{i,t}}{copies_{i,0}}\right)} \quad \text{Equation 2}$$

172 where *copies* is the number of copies of target 16S rRNA gene quantified via qPCR.
173 Equations 1 and 2 are equivalent to $\ln(2)$ divided by the slope time vs. the natural log of
174 $FRA_i C_{tot}$ or the natural log of $copies_i$. For incubations from CLB, time from the start of
175 the experiment was known. For samples taken from the WOR estuary, time at each depth
176 was determined based on an age model (31), using the porosity measurements for core 31
177 and a sedimentation rate of 0.26 cm/year (32).

178 **Cell counts.** Sediments were fixed in 3% paraformaldehyde for a few hours, washed
179 twice with PBS, and stored at -20°C in PBS:ethanol. Sediments were sonicated at 20%
180 power for 40 seconds, diluted to 1:40 in PBS, filtered onto a 0.2 µm polycarbonate filter
181 (Fisher Scientific, Waltham, MA), stained with SYBR Gold (Invitrogen, Carlsbad, CA),

182 and counted at 100x magnification on an epifluorescent microscope. To interpolate cell
183 counts into 1 cm intervals, an exponential curve fit was applied to the samples within the
184 0-10 cm interval with the following fit: $\text{depth} = 137.77 \exp(-1 \times 10^{-9} \times \text{cells/ml})$, $R^2=0.94$.
185 Two more cell count curves were used to test the sensitivity of turnover times to cell
186 count profile variations. In one, cell counts from a 2012 WOR core were previously
187 published (17), and resulted in the following curve fit over 0-10 cm: $\text{depth} = 14.13 \exp(-$
188 $4 \times 10^{-10} \times \text{cells/ml})$, $R^2=0.93$. In the second, cell counts were taken using
189 WebPlotDigitizer from previously published estuarine sediments from Ashleworth Quay
190 (33). For the CLB incubations, SYBR Gold direct counts of paraformaldehyde-fixed cells
191 were performed as described previously (17).

192

193 **Results and Discussion**

194 Sulfate concentrations and porosity were constant with depth in the upper 5 cm of
195 core 34 and the upper 3 cm of core 31 (Fig. S1) due to bioirrigation and bioturbation.
196 Below this, sulfate concentrations decreased due to microbial sulfate reduction as has
197 been observed in other cores at this site (21, 34). Total cell abundance decreased with
198 depth (Fig. S2). The FRAXC and qPCR of total archaea, *Bathyarchaeota* (previously
199 called MCG or MG 1.3 (35)), and *Thermopfundales* (previously called MBG-D (36))
200 increased slightly below the depth of bioirrigation in both WOR cores (Fig. S3). The
201 FRAXC and qPCR of total bacteria decreased with depth in both cores. qPCR values were
202 lower than FRAXC values, demonstrating that these methods are subject to amplification
203 biases and are best used to measure relative changes rather than absolute values (18). The
204 turnover times measured with FRAXC and qPCR methods were within a year of each

205 other for *Bathyarchaeota* and MBG-D (Table 1), suggesting that FRAXC is as good a
206 measure of relative abundance as qPCR. FRAXC, however, opens the possibility of
207 measuring turnover times for any clade represented in the 16S rRNA gene libraries,
208 without developing primer sets for each one.

209 Using FRAXC in the WOR cores, the doubling time (defined as a positive value
210 for turnover time) or half-life (defined as negative a value for turnover time) was
211 calculated for each family-level clade across the three centimeters below the bioirrigated
212 upper layer (5.5-7.5 cm for core 30 and 3.5-5.5 cm for core 32, Table S1). These three
213 depths produced the highest number of clades that increased with depth. A similar
214 increase in the relative contribution from anaerobic organisms was observed immediately
215 below the bioirrigation zone in the marine sediments of Aarhus Bay, Denmark (8). Of the
216 315 clades, 278 were present in both cores, and 60% of these had $R^2 > 0.4$ to the curve fit
217 to equation 1, meaning that a change in FRAXC over depth could be observed. The
218 lenient cut-off of $R^2 > 0.4$ was chosen in order to perform downstream analyses on the
219 largest possible selection of clades. Of these, 80% had doubling times or half-lives
220 between 2 and 25 years. This means that the sampling interval of ~8 years was sufficient
221 to resolve clades that underwent at least a third of one doubling or halving in that
222 interval. Clades that increased faster or more slowly with depth were outside the
223 detection limit. Read abundance between different clades did not correlate with turnover
224 time (Fig. S4). In other words, a clade's relative abundance in the 16S rRNA gene
225 libraries did not predict whether it would increase or decrease with depth.

226 To determine how sensitive turnover times were to differences in the cell
227 abundance curves used to calculate FRAXC, comparisons were made with cell counts

228 from a 2012 WOR core (17) and estuarine sediments from Ashleworth Quay (33). For
229 core 30, 92% and 100% of the clades with turnover times of 2-25 years still fell within
230 that range using the 2012 and Ashleworth Quay cell counts. For core 32, these numbers
231 were 98% and 100%. This suggests that this method is fairly robust to variations in cell
232 abundance curves.

233 The direction of either increase or decrease was replicable in WOR cores 30 and
234 32 for 131 out of 133 clades (Fig. 3a). Clades that increased with depth were common
235 marine subsurface inhabitants: *Chloroflexi*, *Deltaproteobacteria*, *Planctomycetes*,
236 *Bathyarchaeota*, *Lokiarchaeota*, MBG-D, etc (37). Those that decreased with depth were
237 common seawater inhabitants: *Betaproteobacteria*, *Alphaproteobacteria*,
238 *Gammaproteobacteria*, *Thaumarchaeota* (MG-I), *Bacteroidetes*, *Acidobacteria* (38).
239 Clades from uncultured phyla were more likely to increase rather than decrease with
240 depth (9/11 clades). Clades from cultured orders, families, genera, or species were more
241 likely to decrease rather than increase with depth (57/60 clades). This suggests that easily
242 cultured organisms were dying off and difficult-to-culture clades, or PDNC, were
243 growing. Aerobes and nitrate reducers were more likely to decrease with depth (23/23
244 clades where all cultures are obligate aerobes or nitrate reducers, Fig. 3a, inset). Doubling
245 times for clades that increased were well-correlated for individual clades between the two
246 WOR cores (slope = 0.87, $R^2 = 0.64$), suggesting that they may have been growing at
247 rates determined by their particular physiological traits. The mean deviation from the
248 mean between the duplicate cores was $28 \pm 19\%$. This means that turnover times were
249 accurate at least within an order of magnitude. For instance, a clade with a turnover time
250 of 10 is very likely to have an actual turnover time within at least a range of 5-15 years.

251 In contrast, half-lives for clades that decreased were poorly correlated for individual
252 clades between cores (slope = 0.27, $R^2 = 0.11$). Clades that decreased with depth may
253 have been dying or decaying because they could not meet their energetic requirements in
254 this anoxic environment, and, unlike the turnover times of the growing organisms, their
255 rates of decay/death were not specific to their physiologies. These results support the
256 conclusion that FRAx_C doubling times represent real population changes during burial
257 and not random variations.

258 As a separate check on the feasibility of these turnover times, we used the depth-
259 integrated sulfate reduction rate ($0.93 \mu\text{M cm}^{-3} \text{d}^{-1}$, Fig. S1) to estimate a turnover time
260 for sulfate reducers of 12 years, assuming that 10% of total cells were sulfate reducers
261 with growth yields of 2 g cell C/mol (3), and cell carbon contents of 23 fg/cell (39). This
262 falls well within our results from FRAx_C of 2-25 years doubling times, including one
263 clade with cultured sulfate reducers, *Desulfarculaceae*, that had doubling times of 15 and
264 18 years in the two cores. Previous work in arctic and temperate marine sediments also
265 supports generation times of over a year in marine sediment incubations where total cell
266 abundance of sulfate reducing bacteria held constant despite increasing metabolic rates
267 over a year-long incubation (40).

268 Should the population turnover times estimated here approximate the growth rate
269 of these clades, then these microbes would be considered dormant relative to pure
270 cultures (4). However, if the populations are in steady state with some resource that
271 gradually increases with depth (such as a lessening influence of bioturbation, for
272 example), then cell replication is roughly balanced by cell death, so an increase in total
273 population size does not necessarily reflect the growth rate, which may be much faster.

274 To estimate how much the growth rate is likely to differ from doubling times, imagine a
275 clade which increased in relative population from 4% to 50% of the population during a
276 drop in total community size of 1,000,000 cells to 100,000 cells over 2 years. This clade
277 would have a doubling time of 6.21 years. During those two years, the net gain was
278 therefore 10,000 cells. If no cells died, then 10,000 divisions occurred and the doubling
279 time is a growth rate. Death rates can vary greatly among different microbial groups (41),
280 but if we estimate the death rate for a clade to be equal to the total community death ratio
281 of 10:1 for this example, then 100,000 divisions would have occurred and the true growth
282 rate would have been 1.16 years. So, it is possible that the actual growth rates of the
283 clades in our study were faster than 2-25 years, but they may not have been orders of
284 magnitude faster if the decay rates are slow for cells adapted to low energy environments.

285 Two alternative possibilities could also explain an increase with depth in sediment
286 cores, without indicating population growth. First, this increase could represent a mixing
287 curve between a deep source of microbial cells diffusing up toward the surface sediments.
288 Second, the increase could be the result of depositional changes, with greater loading of
289 these subsurface clades in previous depositional events. To rule out these alternative
290 possibilities, we tracked community succession in marine sediment from Cape Lookout
291 Bight, NC, in a laboratory mesocosm under methanogenic conditions over 2.2 years (Fig.
292 3b). Doubling times or half-lives were calculated with FRAXC for each family level clade
293 using timepoints between 40 and 802 days (40, 47, 54, 61, 75, 80, 86, 94, 107, 114, 122,
294 and 802 days, Table S2). All 601 taxa were present in both incubations. R-squared was
295 not a good method of quality control for the CLB incubation data because the data were
296 not evenly distributed across the time interval. Instead, an abundance cutoff was imposed

297 such that each clade must have had at least one timepoint with a value of at least 14 for
298 $\ln(\text{FRAXC})$, which left 195 taxa remaining. Of these, 96%, or 188, had doubling times or
299 half-lives between 0.1 and 7 years (Fig. 3b, Table S2). As with the WOR cores, the
300 sampling interval of 0.1 to 2.2 years was sufficient to resolve clades that underwent a
301 least a third of one doubling or halving in that interval, with slower or faster clades
302 outside the detection limit. Also, similarly to the WOR cores, the direction of increase or
303 decrease was replicable in duplicate incubations (182/188 clades), clades that increased
304 were more similar between replicate incubations (slope = 1.05, $R^2 = 0.96$), and clades
305 that decreased were less similar between replicate cores (slope = 0.83, $R^2 = 0.53$).

306 The clades that increased over time in the CLB incubations were largely methane-
307 cycling archaea (*Methanosarcinales*, *Methanomicrobiales*, and ANME-1), which was
308 expected since the onset of methanogenesis occurred during the sampling interval. The
309 turnover times of methanogen-like archaea of 0.3 to 0.8 years, were consistent with
310 previously measured doubling times of one of these clades, ANME-1, grown in an
311 enrichment, at 0.6 years (42). Of the five clades that increased in both the WOR
312 sediments and the CLB incubations (*Bathyarchaeota*, C3, MBG-D, MG-III, and 20a-9),
313 four increased faster in the CLB incubations, but most had multi-year turnover times in
314 both types of experiments and measurement methods (Table 1). Two clades
315 (*Bathyarchaeota* and MBG-D) had both types of measurements (qPCR and FRAXC) in
316 both types of experiments (WOR cores and CLB incubations). The mean of all these
317 measurements resulted in turnover times of 8.7 ± 1.9 years for *Bathyarchaeota* and $4.1 \pm$
318 0.7 years for MBG-d; the relatively small standard deviation suggests that these turnover
319 times are accurate within a few years. Many of the clades with slower turnover times in

320 the WOR cores were present in the CLB incubations, but did not have turnover times that
321 met our quality cut-off. We may have been able to resolve their growth rates if we had
322 waited ~8 years. The lab incubation experiment therefore also supports a multi-year
323 population turnover time for anoxic uncultured marine subsurface organisms, in an
324 experimental system that rules out diffusional mixing and depositional changes as drivers
325 for the population changes.

326 These results suggest that the log-log decay rate with depth for the total microbial
327 community does not apply to some clades within the marine sediment biosphere. Some
328 clades experience periods of net growth, resulting in population increases during burial
329 (Option 2 in Fig. 1). This has important implications for microbial adaptations to the
330 marine subseafloor biosphere. Many of these clades have been shown to be dormant or
331 non-growing for most of the sediment column (3, 9). In other environments, such as
332 agricultural soils, periods of dormancy are interspersed with periods of growth (4). Our
333 results suggest that the period of growth for marine sediment microbes may occur in the
334 upper few cm, over a time period that is much shorter than the full marine sediment
335 column, which can be millions of years old.

336 This suggests a possible mechanism for adaptation to long-term dormancy in
337 marine sediments, building on the Growth Advantage in Stationary Phase (GASP), where
338 *E. coli* cells that have been starved for years outcompete fresh cells under low substrate
339 conditions (43). Our results suggest that these GASP-like adaptations to long-term
340 starvation could also be selected for after resuspension events, allowing GASP to be
341 extrapolated over much longer timescales in marine sediment, as was predicted
342 previously (44). If deep sedimentary cells with subsistence-promoting or GASP-like

343 mutations are resuspended either by a storm or sediment slumping, these organisms
344 would have a growth advantage over others as they were buried a second time through
345 the relatively brief and shallow growth zone. This means that although evolutionary
346 changes do not occur during the process of burial and near-zero cell growth over
347 thousands of years (7), the ability to subsist over thousands of years in marine sediment is
348 a trait that could have been selected for within the growth zone in shallow sediments.
349 Microbes might even be ejected from deeper sediments depths by mud volcanoes into
350 overlying seawater, where they can be carried by currents and re-deposited at the seafloor
351 (45). Our results suggest that such organisms may revive and begin growing shortly after
352 deposition.

353 If these turnover times are similar to growth rates, then these organisms grow
354 much more slowly than most laboratory cultures, which double in less than a day. This
355 agrees with previous work showing that microbes grow much more slowly in the
356 environment than in laboratory culture. *Staphylococcus aureus* grows 48-fold faster in
357 culture than it does in an infected lung (46) and *Leucothrix mucor* grows 7-fold faster in
358 culture than it does in seawater (47). The fact that so many of the clades that grew in our
359 study belong to uncultured clades suggests that slow growth rates may be a reason why
360 so many of them remain uncultured. Recent successes with culturing slow-growing
361 strains from previously-uncultured phyla *Lokiarchaeota* and *Atribacteria* attest to the
362 necessity of extremely long wait times for growth; 2000 days in one case (48, 49).

363 Even if it is possible to speed up such organisms sufficiently to culture them, their
364 traits expressed under such conditions are likely to greatly diverge from those expressed
365 in their natural environment. Therefore, even when cultures are available, it is crucial to

366 study these organisms in their natural environmental settings. Using label incorporation
367 (46, 50, 51), or tracking slow growth in a mixed population under different conditions,
368 will allow for growth experiments on the environmentally abundant, evolutionarily
369 distinct, uncultured microbes from resource-limited environments. Such experiments are
370 necessary to assess modifications to cell cycle regulation, proteins, or plasma membranes
371 that enable such extraordinarily slow growth, as well as determine how this narrow
372 growth zone in surficial sediments may select for the traits that enable long-term
373 subsistence of the deep subsurface biosphere.

374

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384

385 **Table 1.** Doubling times for all clades that increased in both cores of White Oak River
386 (WOR) estuary sediments and replicate incubations from the Cape Lookout Bight (CLB)
387 marine sediments. Data are expressed in years. FRaxC denotes values calculated by the

388 increase of the product of 16S rRNA gene read percentages and total cells over depth
 389 (WOR) or time (CLB). qPCR denotes values calculated by the increase of qPCR of 16S
 390 rRNA genes over depth. Hyphen indicates no data because qPCR was not attempted for
 391 that clade or sample.

Clade	WOR marine sediment				CLB marine sediment	
	Core 30		Core 32		Incubation 2	Incubation 3
	FRAxC	qPCR	FRAxC	qPCR	FRAxC	FRAxC
<i>Bathyarchaeota</i> , or MCG	10.1	10.3	8.9	10.0	6.4	6.3
Group C3, MCG-related	16.2	-	14.9	-	2.1	1.8
<i>Thermoprofundales</i> (MBG-D)	4.6	-	3.4	4.8	3.4	4.4
<i>Thermoprofundales</i> (MG-III)	3.7	-	2.8	-	1.4	1.5
<i>Thermoprofundales</i> (20a-9)	7.1	-	3.7	-	0.8	0.7

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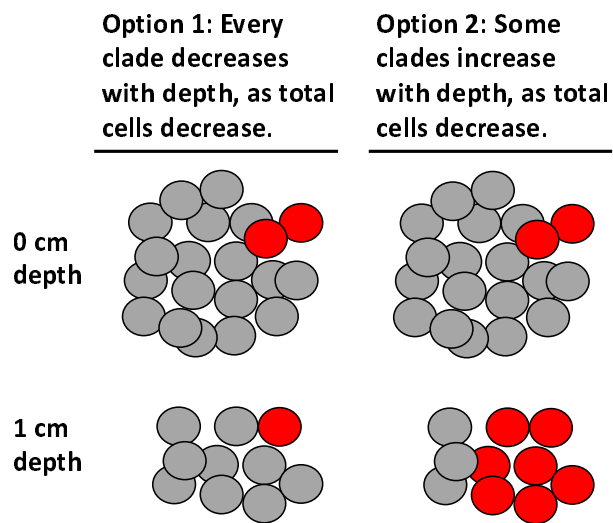


Figure 1. Options for how an individual clade of microbes (red cells) either decrease (option 1) or increase (option 2) with total population decrease with 1 cm sediment depth.

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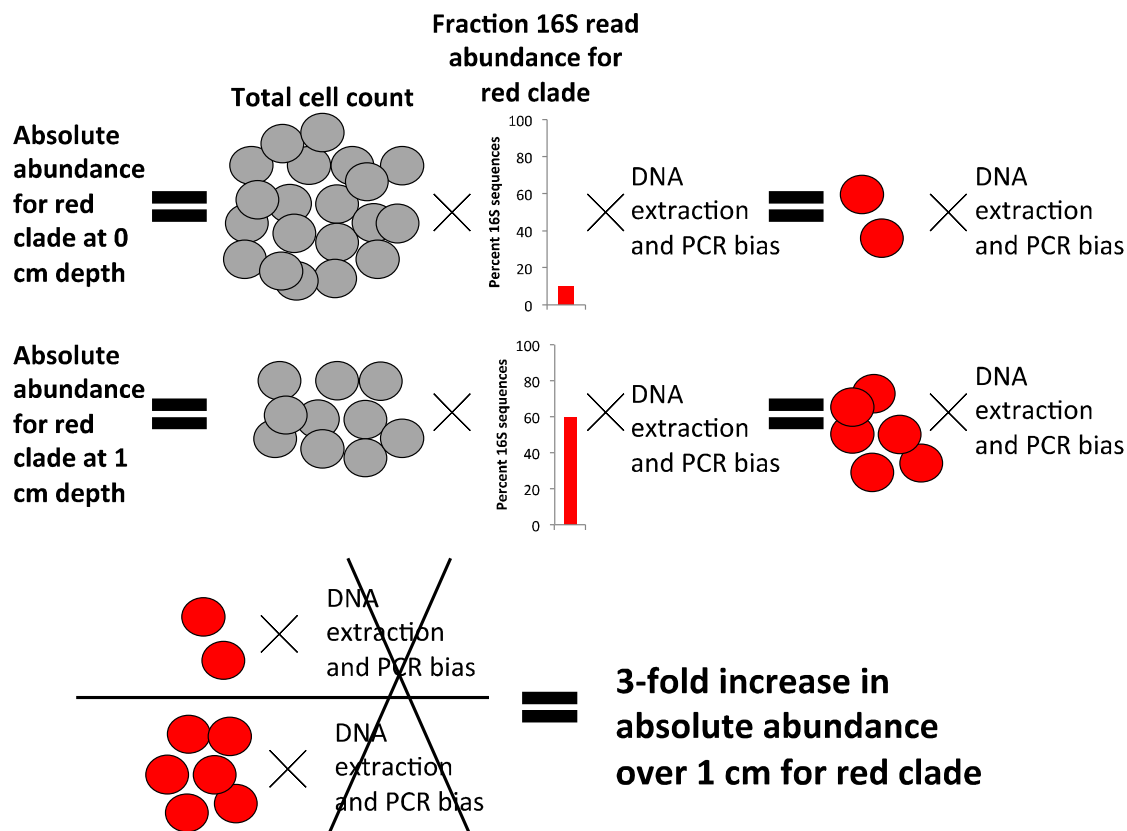


Figure 2. Example to explain how Fraction Read Abundance times Cell abundance (FRAxC) works. The total cell count (grey) decreases with depth. The absolute abundance of an imaginary clade of organisms (red) is equal to the total cell abundance multiplied by the fraction of 16S rRNA gene reads of the red clade and a bias term that cannot be measured accurately and is different for each clade of microbes. Assuming that this bias term is similar over adjacent sediment depths for the same clade of microbes, the relative increase in the red clade with depth can be measured absolutely, since the bias term cancels out.

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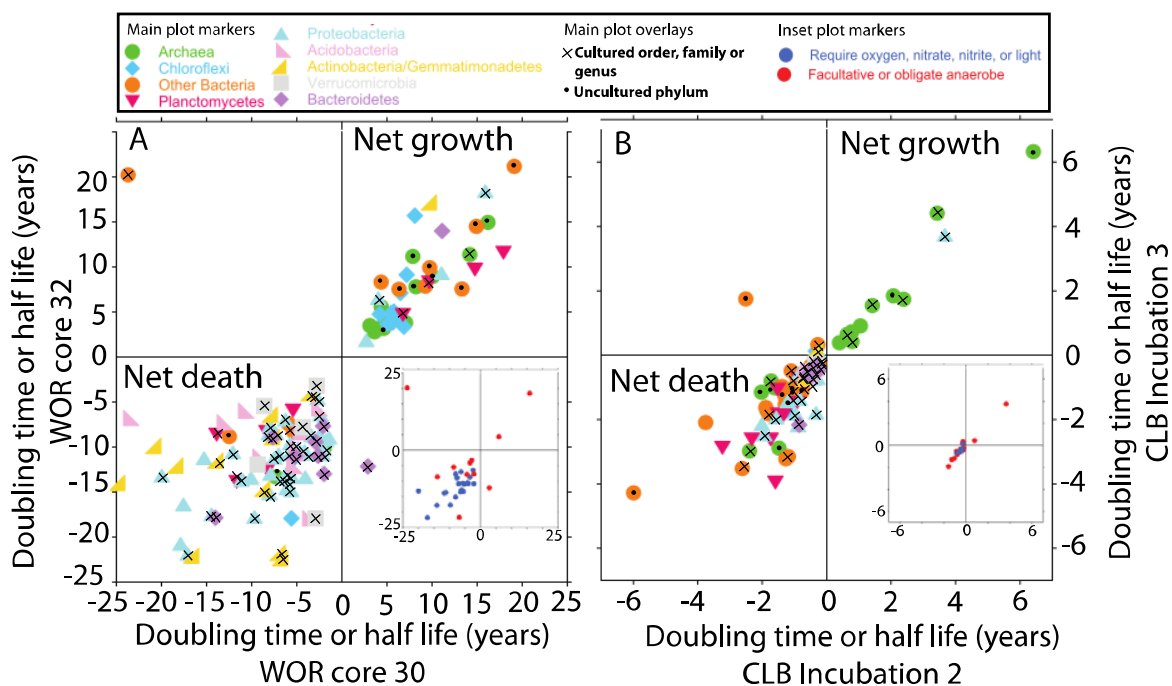


Figure 3. Multi-year doubling times for uncultured microbes in marine sediments, measured in A) down-core trends in White Oak River estuary or B) incubation time course experiments from Cape Lookout Bight. In each panel, replicates are plotted against each other to show good correlation between clades that increase (upper right quadrants) or decrease (lower left quadrants) with depth. In A) clades that grow tend to be common subsurface clades (in dark colors) and uncultured phyla (black dots); clades that decay tend to be common seawater clades (in pastels) and more closely related to cultures (x's). Insets show the subset of clades belonging to cultured families where all cultured relatives are oxygen- or nitrate-reducing (blue), or contain at least some non-nitrogen-cycling facultative anaerobes (red). Note different timescales for WOR and CLB data.

422 References

- 423 1. **Kallmeyer J, Pockalny R, Adhikari RR, Smith DC, D'Hondt S.** 2012. Global
 424 distribution of microbial abundance and biomass in subseafloor sediment. Proc
 425 Natl Acad Sci U S A **109**:16213–16216.
- 426 2. **Parkes RJ, Cragg BA, Wellsbury P.** 2000. Recent studies on bacterial
 427 populations and processes in subseafloor sediments: A review. Hydrogeol J 11–

- 428 28.
- 429 3. **Hoehler TM, Jørgensen BB.** 2013. Microbial life under extreme energy
430 limitation. *Nat Rev Microbiol* **11**:83–94.
- 431 4. **Lennon JT, Jones SE.** 2011. Microbial seed banks: the ecological and
432 evolutionary implications of dormancy. *Nat Rev Microbiol* **9**:119–30.
- 433 5. **Braun S, Mhatre SS, Jaussi M, Røy H, Kjeldsen KU, Seidenkrantz M,**
434 **Jørgensen BB, Lomstein BA.** 2017. Microbial turnover times in the deep seabed
435 studied by amino acid racemization modelling. *Sci Rep* **7**:5680.
- 436 6. **Jørgensen BB.** 2011. Deep subseafloor microbial cells on physiological standby.
437 *Proc Natl Acad Sci* **108**:18193–18194.
- 438 7. **Starnawski P, Bataillon T, Ettema TJG, Jochum LM, Schreiber L, Chen X,**
439 **Lever MA, Polz MF, Jørgensen BB, Schramm A, Kjeldsen KU.** 2017.
440 Microbial community assembly and evolution in subseafloor sediment. *Proc Natl*
441 *Acad Sci* **114**:2940–2945.
- 442 8. **Marshall IPG, Ren G, Jaussi M, Lomstein BA, Jørgensen BB, Røy H,**
443 **Kjeldsen KU.** 2019. Environmental filtering determines family-level structure of
444 sulfate-reducing microbial communities in subsurface marine sediments. *ISME J*
445 **19**:1920–1932.
- 446 9. **Kirkpatrick JB, Walsh EA, D’Hondt S.** 2016. Fossil DNA persistence and decay
447 in marine sediment over hundred-thousand-year to million-year time scales.
448 *Geology* **44**:615–618.
- 449 10. **Jochum LM, Chen X, Lever MA, Loy A, Jørgensen BB, Schramm A,**

- 450 **Kjeldsen KU**. 2017. Depth Distribution and Assembly of Sulfate-Reducing
451 Microbial Communities in Marine Sediments of Aarhus Bay. *Appl Environ*
452 *Microbiol* **83**:1–15.
- 453 11. **Steen AD, Kevorkian RT, Bird JT, Dombrowski N, Baker BJ, Hagen SM,**
454 **Mulligan KH, Schmidt JM, Webber AT, Royalty T, Alperin MJ**. 2019.
455 Kinetics and identities of extracellular peptidases in subsurface sediments of the
456 White Oak River Estuary, NC. *Appl Environ Microbiol* **85**:1–14.
- 457 12. **Karl D, Novitsky J**. 1988. Dynamics of microbial growth in surface layers of a
458 coastal marine sediment ecosystem. *Mar Ecol Prog Ser* **50**:169–176.
- 459 13. **Novitsky J, Karl D**. 1986. Characterization of microbial activity in the surface
460 layers of a coastal subtropical sediment. *Mar Ecol Prog Ser* **28**:49–55.
- 461 14. **Parkes RJ, Webster G, Cragg BA, Weightman AJ, Newberry CJ, Ferdelman**
462 **TG, Kallmeyer J, Jørgensen BB, Aiello IW, Fry JC**. 2005. Deep sub-seafloor
463 prokaryotes stimulated at interfaes over geological time. *Nature* **436**:390–394.
- 464 15. **Lloyd KG, Steen AD, Ladau J, Yin J, Crosby L**. 2018. Phylogenetically novel
465 uncultured microbial cells dominate Earth microbiomes. *mSystems* **3**:e00055-18.
- 466 16. **Kubo K, Lloyd KG, F Biddle J, Amann R, Teske A, Knittel K**. 2012. Archaea
467 of the Miscellaneous Crenarchaeotal Group are abundant, diverse and widespread
468 in marine sediments. *ISME J* **6**:1949–65.
- 469 17. **Kevorkian R, Bird JT, Shumaker A, Lloyd KG**. 2018. Estimating population
470 turnover rates by relative quantification methods reveals microbial dynamics in
471 marine sediment. *Appl Environ Microbiol* **84**:1–16.

- 472 18. **Lloyd KG, May MK, Kevorkian RT, Steen AD.** 2013. Meta-analysis of
473 quantification methods shows that archaea and bacteria have similar abundances in
474 the subseafloor. *Appl Environ Microbiol* **79**:7790–7799.
- 475 19. **Buongiorno J, Turner S, Webster G, Asai M, Shumaker AK, Roy T,**
476 **Weightman A, Schippers A, Lloyd KG.** 2017. Inter-laboratory quantification of
477 Bacteria and Archaea in deeply buried sediments of the Baltic Sea (IODP
478 Expedition 347). *FEMS Microbiol Ecol* fix007.
- 479 20. **Polz MF, Cavanaugh CM.** 1998. Bias in template-to-product ratios in
480 multitemplate PCR. *Appl Environ Microbiol* **64**:3724–3730.
- 481 21. **Lloyd KG, Alperin MJ, Teske A.** 2011. Environmental evidence for net methane
482 production and oxidation in putative ANaerobic MEthanotrophic (ANME) archaea.
483 *Environ Microbiol* **13**:2548–2564.
- 484 22. **Caporaso JG, Lauber CL, Walters WA, Berg-Lyons D, Huntley J, Fierer N,**
485 **Owens SM, Betley J, Fraser L, Bauer M, Gormley N, Gilbert JA, Smith G,**
486 **Knight R.** 2012. Ultra-high-throughput microbial community analysis on the
487 Illumina HiSeq and MiSeq platforms. *ISME J* **6**:1621–1624.
- 488 23. **Pruesse E, Quast C, Knittel K, Fuchs BM, Ludwig W, Peplies J, Glöckner FO.**
489 2007. SILVA: a comprehensive online resource for quality checked and aligned
490 ribosomal RNA sequence data compatible with ARB. *Nucleic Acids Res* **35**:7188–
491 7196.
- 492 24. **Schloss PD, Westcott SL, Ryabin T, Hall JR, Hartmann M, Hollister EB,**
493 **Lesniewski RA, Oakley BB, Parks DH, Robinson CJ, Sahl JW, Stres B,**

- 494 **Thallinger GG, Horn DJ Van, Weber CF.** 2009. Introducing mothur: Open-
495 source, platform-independent, community-supported software for describing and
496 comparing microbial communities. *Appl Environ Microbiol* **75**:7537–7541.
- 497 25. **Buonigiorno J, Turner S, Webster G, Asai M, Shumaker AK, Roy T,**
498 **Weightman A, Schippers A, Lloyd KG.** 2017. Interlaboratory quantification of
499 Bacteria and Archaea in deeply buried sediments of the Baltic Sea (IODP
500 Expedition 347). *FEMS Microbiol Ecol* **93**.
- 501 26. **Kubo K, Lloyd KG, F Biddle J, Amann R, Teske A, Knittel K.** 2012. Archaea
502 of the Miscellaneous Crenarchaeotal Group are abundant, diverse and widespread
503 in marine sediments. *ISME J* **6**.
- 504 27. **Lloyd KG, Schreiber L, Petersen DG, Kjeldsen KU, Lever MA, Steen AD,**
505 **Stepanauskas R, Richter M, Kleindienst S, Lenk S, Schramm A, Jørgensen**
506 **BB.** 2013. Predominant archaea in marine sediments degrade detrital proteins.
507 *Nature* **496**:215–218.
- 508 28. **Stahl DA, Amann R.** 1991. Development and application of nucleic acid probes.
509 John Wiley & Sons, Chichester, UK.
- 510 29. **Yu Y, Lee C, Kim J, Hwang S.** 2005. Group-specific primer and probe sets to
511 detect methanogenic communities using quantitative real-time polymerase chain
512 reaction. *Biotechnol Bioeng* **89**:670–9.
- 513 30. **Nadkarni MA, Martin FE, Jacques NA, Hunter N.** 2002. Determination of
514 bacterial load by real-time PCR using a broad-range (universal) probe and primers
515 set. *Microbiology* **148**:257–266.

- 516 31. **Larowe DE, Amend JP.** 2015. Power limits for microbial life. *Front Microbiol*
517 **6**:1–11.
- 518 32. **Benninger LK, Martens CS.** 1983. Sources and fates of sedimentary organic
519 matter in the White Oak and Neuse estuaries. *Water Resources Research Insitute of*
520 *the University of North Carolina, UNC-WRRI-83-194.*
- 521 33. **Wellsbury P, Herbert RA, Parkes RJ.** 1996. Bacterial activity and production in
522 near-surface estuarine and freshwater sediments **19**.
- 523 34. **Martens CS, Albert DB, Alperin MJ.** 1998. Biogeochemical processes
524 controlling methane in gassy coastal sediments — Part 1 . A model coupling
525 organic matter flux to gas production , oxid1. Martens CS, Albert DB, Alperin MJ
526 (1998) Biogeochemical processes controlling methane in gassy coastal sedim.
527 *Methods* **18**.
- 528 35. **Meng J, Xu J, Qin D, He Y, Xiao X, Wang F.** 2013. Genetic and functional
529 properties of uncultivated MCG archaea assessed by metagenome and gene
530 expression analyses. *ISME J* 1–10.
- 531 36. **Zhou Z, Liu Y, Lloyd KG, Pan J, Yang Y, Gu J-D, Li M.** 2018. Genomic and
532 transcriptomic insights into the ecology and metabolism of benthic archaeal
533 cosmopolitan, Thermopfundales (MBG-D archaea). *ISME J*.
- 534 37. **Teske A, Sørensen KB.** 2008. Uncultured archaea in deep marine subsurface
535 sediments: have we caught them all? *ISME J* **2**:3–18.
- 536 38. **Galand PE, Casamayor EO, Kirchman DL, Lovejoy C.** 2009. Ecology of the
537 rare microbial biosphere of the Arctic Ocean. *Proc Natl Acad Sci U S A*

- 538 **106**:22427–32.
- 539 39. **Braun S, Morono Y, Littmann S, Kuypers M, Aslan H, Dong M, Jorgensen**
540 **BB, Lomstein BA.** 2016. Size and carbon content of sub-seafloor microbial cells
541 at Landsort Deep, Baltic Sea. *Front Microbiol* **7**:1–13.
- 542 40. **Robador A, Brüchert V, Jørgensen BB.** 2009. The impact of temperature change
543 on the activity and community composition of sulfate-reducing bacteria in arctic
544 versus temperate marine sediments. *Environ Microbiol* **11**:1692–1703.
- 545 41. **Hao X, Wang Q, Zhang X, Cao Y, Mark Loosdrecht CM van.** 2009.
546 Experimental evaluation of decrease in bacterial activity due to cell death and
547 activity decay in activated sludge. *Water Res* **43**:3604–3612.
- 548 42. **Nauhaus K, Albrecht M, Elvert M, Boetius A, Widdel F.** 2007. In vitro cell
549 growth of marine archaeal-bacterial consortia during anaerobic oxidation of
550 methane with sulfate. *Environ Microbiol* **9**:187–196.
- 551 43. **Finkel SE.** 2006. Long-term survival during stationary phase: evolution and the
552 GASP phenotype. *Nat Rev Microbiol* **4**:113–20.
- 553 44. **Lever MA, Rogers KL, Lloyd KG, Schink B, Thauer RK, Hoehler TM,**
554 **Jørgensen BB.** 2015. Life under extreme energy limitation□: a synthesis of
555 laboratory- and field-based investigations 1–41.
- 556 45. **Hoshino T, Toki T, Ijiri A, Morono Y, Machiyama H, Ashi J, Okamura K,**
557 **Inagaki F.** 2017. Atribacteria from the subseafloor sedimentary biosphere disperse
558 to the hydrosphere through submarine mud volcanoes. *Front Microbiol* **8**:1–14.
- 559 46. **Kopf SH, Sessions AL, Cowley ES, Reyes C, Sambeek L Van, Hu Y, Orphan**

- 560 **VJ, Kato R, Newman DK.** 2015. Trace incorporation of heavy water reveals slow
561 and heterogeneous pathogen growth rates in cystic fibrosis sputum. *Proc Natl Acad*
562 *Sci USA* **113**:E110–E116.
- 563 47. **Brock TD.** 1967. Bacterial growth rate in the sea: direct analysis by thymidine
564 autoradiography. *Science* **155**:81–83.
- 565 48. **Imachi H, Nobu MK, Nakahara N, Morono Y, Ogawara M, Takaki Y,**
566 **Takano Y, Uematsu K, Ikuta T, Ito M, Matsui Y, Miyazaki M, Murata K,**
567 **Saito Y, Sakai S, Song C, Tasumi E, Yamanaka Y, Yamaguchi T, Kamagata**
568 **Y, Tamaki H, Takai K.** 2019. Isolation of an archaeon at the prokaryote-
569 eukaryote interface. *bioRxiv* <http://dx>.
- 570 49. **Katayama T, Nobu MK, Kusada H, Meng X-Y, Yoshioka H, Kamagata Y,**
571 **Tamaki H.** 2019. Membrane-bounded nucleoid discovered in a cultivated
572 bacterium of the candidate phylum “Atribacteria.” *bioRxiv* <http://dx>.
- 573 50. **Morono Y, Terada T, Nishizawa M, Ito M, Hillion F, Takahata N, Sano Y,**
574 **Inagaki F.** 2011. Carbon and nitrogen assimilation in deep seafloor microbial
575 cells. *Proc Natl Acad Sci U S A* **108**:18295–300.
- 576 51. **Hatzenpichler R, Scheller S, Tavormina PL, Babin BM, Tirrell DA, Orphan**
577 **VJ.** 2014. In situ visualization of newly synthesized proteins in environmental
578 microbes using amino acid tagging and click chemistry. *Environ Microbiol*
579 **16**:2568–2590.

580