Stable isotope fractionation reveals similar atomic level controls during

aerobic and anaerobic microbial Hg transformation pathways

Running title: Isotopes unify controls on bacterial Hg transformations

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ABSTRACT

Mercury (Hg) is a global pollutant and potent neurotoxin that bioaccumulates in food webs as monomethylmercury (MeHg). The production of MeHg is driven by anaerobic and Hg redox cycling pathways such as Hg reduction, which control the availability of Hg to methylators. Anaerobes play an important role in Hg reduction in methylation hotspots, yet their contributions remain underappreciated due to how challenging these pathways are to study in the absence of dedicated genetic targets and low levels of Hg^0 in anoxic environments. In this study we used Hg stable isotope fractionation to explore Hg reduction during anoxygenic photosynthesis and fermentation in the model anaerobe *Heliobacterium modesticaldum* Ice1. We show that cells preferentially reduce lighter Hg isotopes in both metabolisms leading to mass-dependent fractionation, but mass-independent fractionation commonly induced by UV-visible light is absent. Due to variability associated with replicated experiments, we could not discern whether dedicated physiological processes drive Hg reduction during photosynthesis and fermentation. However, we demonstrate that fractionation is affected by the interplay between pathways controlling Hg recruitment, accessibility, and availability alongside metabolic redox reactions. The combined contributions of these processes lead to isotopic enrichment during anoxygenic photosynthesis that is in between the values reported for anaerobic respiratory microbial Hg reduction and abiotic photoreduction. Isotope enrichment during fermentation is closer to what has been observed in aerobic bacteria that reduce Hg through dedicated detoxification pathways. Our work suggests that similar controls likely underpin diverse microbe-mediated Hg transformations that affect Hg's fate in oxic and anoxic habitats.

IMPORTANCE

Anaerobic and photosynthetic bacteria that reduce mercury affect mercury delivery to microbes in methylation sites that drive bioaccumulation in food webs. Anaerobic mercury reduction pathways remain underappreciated in the current view of the global mercury cycle because they are challenging to study, bearing no dedicated genetic targets to establish physiological mechanisms. In this study we used stable isotopes to characterize the physiological processes that control mercury reduction during photosynthesis and fermentation in the model anaerobe *Heliobacterium modesticaldum* Ice1. The sensitivity of isotope analyses highlighted the subtle contribution of mercury uptake towards the isotope signature associated with anaerobic mercury reduction. When considered alongside the isotope signatures associated with microbial pathways for which genetic determinants have been identified, our findings underscore the narrow range of isotope enrichment that is characteristic of microbial mercury transformations. This suggests that there exist common atomic-level controls for biological mercury transformations across a broad range of geochemical conditions.

INTRODUCTION

Mercury (Hg) is a global pollutant and potent neurotoxin that bioaccumulates in aquatic and terrestrial food webs as monomethylmercury (MeHg) (1). Anaerobic microbes with the *hgcAB* gene cluster, which encodes for the metabolic machinery responsible for Hg methylation (2), contribute to MeHg production in habitats such as aquatic sediments, water columns, and flooded soils (3-6). Hg methylation is thought to occur intracellularly and as such, it is ultimately controlled by the bioavailability of Hg to methylators in anoxic habitats (7).

Hg redox cycling plays a key role in determining the inorganic Hg substrates available to methylators. Its role is two-fold: anaerobic Hg^0 oxidation can supply dissolved Hg^{II} required to generate MeHg and evasion of Hg^0 due to reduction can limit MeHg production by removing inorganic Hg substrates. Although it is well established that abiotic photochemical Hg reduction dominates Hg redox cycling in oxic surface systems where light is present (8-16), Hg reduction pathways in anoxic zones where Hg methylation occurs have mostly been characterized under laboratory conditions and remain challenging to assess in the field.

Anoxic Hg reduction can occur via abiotic redox reactions with dissolved organic carbon (17) and magnetite (18, 19), but also through biotic pathways mediated by chemotrophic and phototrophic microorganisms (20-24). Unlike many aerobes that reduce Hg using dedicated enzymatic machinery encoded by the *mer* operon (25, 26), Hg reduction in anaerobes occurs through cometabolic pathways tied to anaerobic respiration, fermentation, and anoxygenic photosynthesis (20-24). In contrast to *mer*-mediated Hg reduction and demethylation, dedicated genetic determinants have yet to be identified for anaerobic Hg reduction, making these pathways challenging to study from a mechanistic standpoint. Furthermore, while some studies have shown that Hg^0 can significantly contribute to Hg speciation in anoxic habitats (27, 28), it is

possible that low levels of Hg^0 are maintained by an active cycle of reduction and oxidation (29).

In the absence of genetic targets and measurable Hg^0 accumulation, the contributions of

anaerobic Hg^0 production to global Hg cycling remain cryptic and difficult to assess.

The use of stable Hg isotope fractionation is a powerful tool for providing

biogeochemical proxies for different abiotic and biotic Hg transformations (30, 31). All abiotic

and biotic Hg transformation pathways studied to date demonstrate kinetic mass dependent

fractionation (MDF) in which the product pool becomes isotopically lighter (enrichment of light

isotopes) as the reaction proceeds (32). Some pathways, such as abiotic photochemical reduction

and demethylation, also demonstrate mass independent fractionation (MIF), which is defined as

fractionation that occurs outside of the predicted MDF pattern (30). MIF is most commonly

observed due to the magnetic isotope effect for odd-Hg isotopes ^{199}Hg and ^{201}Hg in the presence

of sunlight (33) and the extent to which MIF occurs is controlled in part by water chemistry and

the presence of photosynthetic organisms (34).

The combination of MDF and MIF tracers, and the rate of isotopic enrichment, referred

to as fractionation factor, can be used to outline the underlying components of different microbe-

mediated Hg transformation pathways. Over the last decade, Hg isotope fractionation has been

studied in microbes capable of *mer*-mediated reduction (35) and MeHg demethylation (36);

anaerobic non-*mer*-mediated Hg reduction (37, 38); Hg methylation (38-41); and oxygenic

phototrophic MeHg degradation and Hg reduction (34). These studies have yielded valuable

insights and mechanistic details that explain how microbes interact with Hg at the atomic level

making them a promising option for studying mechanisms for anaerobic Hg reduction.

Our objective in this study was to characterize Hg isotope fractionation patterns during

anoxygenic phototrophic and fermentative Hg^{II} reduction, for which there exists no data. By

providing the first isotopic characterization of these pathways we aim to better understand the

underlying physiological pathways involved in anaerobic Hg reduction and their role in

controlling Hg availability to methylators in anoxic habitats.

In this work, we test whether Hg reduction by the model anaerobe *Heliobacterium*

modesticaldum Ice1 during anoxygenic photosynthesis and fermentation leads to distinct isotope

enrichment signatures, which would suggest the different metabolic Hg reduction pathways are

controlled by separate processes. We chose to work with this model Heliobacteria because its

metabolic versatility is representative of the strategies that allow this family to thrive in the same

niche as methylators in habitats such as rice paddies where MeHg accumulation is a major health

concern (42-45). Furthermore, our previous work showed that *H. modesticaldum* has the capacity

to reduce high proportions of Hg^{II}, making this strain well-suited to producing the high levels of

Hg^0 required for measuring isotope fractionation (23). In addition to comparing isotope

fractionation associated with different metabolic pathways in *H. modesticaldum*, we compare the

isotope enrichment observed in this study to previously published isotope signatures for biotic

and abiotic Hg transformations to critically evaluate whether common controls exist for microbe-

mediated Hg transformations. Finally, we discuss whether current analytical capacity for isotope

fractionation can be used to track the contributions of such pathways in the environment.

METHODS

Cell growth conditions and bioreactor setup

All experiments with live cells employed the strain *Heliobacterium modesticaldum* Ice1 grown in PYE medium. PYE medium was prepared by dissolving the following ingredients in 1 L of ultra-filtered high purity (18.2 MΩ) water: Dipotassium phosphate (K_2HPO_4) (1 g L^{-1}); magnesium sulphate heptahydrate ($MgSO_4$ x $7H_2O$) (0.2 g L^{-1}); calcium chloride dihydrate ($CaCl_2$ x $2H_2O$) (0.02 g L^{-1}); sodium thiosulphate pentahydrate ($Na_2S_2O_3$) x $5H_2O$ (0.2 g L^{-1}); sodium pyruvate (2.20 g L^{-1}); yeast extract (YE) (4.00 g L^{-1}); and ammonium sulphate $[(NH_4)_2SO_4]$ (1.00 g L^{-1}). The pH of the medium was adjusted to 6.8, after which 1 mL of a filter-sterilized (0.2 µm pore size) trace element solution was added to the medium as a 1000X dilution. The trace element stock was prepared by adding the following ingredients to 1 L of ultra-filtered high purity water: Zinc sulphate heptahydrate ($ZnSO_4$ x $7H_2O$) (0.10 g L^{-1}); manganese chloride tetrahydrate ($MnCl_2$ x $4H_2O$) (0.03 g L^{-1}); boric acid (H_3BO_3) (0.30 g L^{-1}); cobalt chloride hexahydrate ($CoCl_2$ x $6H_2O$) (0.20 g L^{-1}); cupric chloride dihydrate ($CuCl_2$ x $2H_2O$) (0.01 g L^{-1}); nickel chloride hexahydrate ($NiCl_2$ x $6H_2O$) (0.02 g L^{-1}); sodium molybdate dihydrate (Na_2MoO_4 x $2H_2O$) (0.03 g L^{-1}). The medium was then boiled for 15 minutes under a stream of sterile nitrogen (N_2) gas before being aliquoted to serum bottles or balch tubes that were crimped shut, flushed with N_2 gas, and autoclaved at 121°C for 30 min.

Phototrophic cultures were grown at a constant visible light intensity of 80 µmol photon m^{-2} s^{-1} with a peak irradiance in the near-red to red spectrum (600 nm to 700 nm) at 50°C whereas fermentative cultures were grown in the dark at 50°C in line with our previous work (23). *H. modesticaldum* Ice1 was obtained from the Leibniz Institute Deutsche Sammlung von Mikroorganismen und Zellkulturen (DSMZ) culture collection (catalogue number DSM-792).

The bioreactor methodology employed in these experiments was similar to the one used in

previous work (23) with the following modifications: The bioreactor was periodically supplied

with sterile N_2 gas that passed through an activated carbon filter and a 0.2 µm pore size air filter

at a flow rate of 0.25 L min^{-1} and two gold traps were connected in series at the reactor outlet to

capture Hg^0 (**Fig S1**).

The bioreactor was kept in an incubator set to 50°C. Phototrophic experiments were

performed using a 60 W incandescent light bulb with a visible light intensity being maintained at

20 µmol photon m^{-2} s^{-1} and mirroring the irradiance of growth conditions. Fermentative

experiments were performed in the dark. Background Hg^0 was purged and removed from

reactors prior to phototrophically or fermentatively-grown *H. modesticaldum* cells being added

as a 10 % (v/v) inoculum after which an additional subsample was withdrawn to verify initial

cell density via O.D. 600 nm. For sterile treatments, no cells were added to the bioreactor and the

volume of growth medium was adjusted to ensure experiments were carried out at a working

volume of 500 mL. We chose to use sterile medium rather than heat-killed cultures based on our

previous work, which showed that autoclaved cultures of *H. modesticaldum* supplied with Hg^{II}

under the same conditions reduced negligible amounts of Hg^{II} (23). In all experiments, National

Institute of Standards and Technology (NIST) 3133, a standard with certified Hg isotope values,

was added to the bioreactor to approximately 10 nM Hg final concentration. Additional details

on the bioreactor setup are provided in **Supporting Methods**.

Hg⁰ measurements

The bioreactor was bubbled periodically rather than continuously in contrast to previous work because continuous bubbling led to fouling of gold traps and poor total Hg (THg) and isotope recoveries. Subsamples were collected at designated timepoints by connecting dual gold traps to the reactor and purging the reactor for 30 min to collect Hg^0. In the final 5 mins, the reactor liquid was sampled for Hg^{II} concentration, isotopic composition, and cell growth after which gas flow was stopped. All gold traps were capped to prevent Hg^0 loss and aqueous samples were conserved with 1% (v/v) trace element grade hydrochloric acid (HCl) at 4°C in the dark. Reactor washes were performed using trace element grade HCl following previously established protocols (23).

Total Hg analyses and sample preparation for stable isotope measurements

All samples for aqueous Hg analyses were oxidized with 10 % (v/v) bromine chloride (BrCl) in line with previous work (40). THg concentrations were determined by stannous chloride reduction coupled to cold vapor atomic fluorescence spectrometry at the U.S. Geological Survey Mercury Research Lab (USGS MRL, Middleton, WI) (46). During THg analysis, duplicates and matrix standard spikes were analyzed every 10 samples; analyses showed less than 10 % relative percent difference between duplicates and spike recoveries of 90-110%. The detection limit was 0.2 pM.

Total Hg mass balances were calculated in line with previous work (23, 24) with corrections to account for the mass of Hg removed for each subsample. For Hg^0 collected onto gold traps, a thermal desorption system was used to liberate Hg^0 over a 40 min desorption cycle which was then captured in 40% nitric acid (HNO_3):BrCl (3:1) oxidizing solution. An aliquot of

the trapping solution was subsequently analyzed for THg concentration following the protocol

established in previous work (47). Desorption efficiency during pre-concentration was tested by

trapping a NIST 3133 standard with each batch of samples, and recoveries for standard trapping

were 101.27 ± 4.26 % (n = 15). THg recoveries from bioreactor experiments were 94.90 ± 9.50

% (n=7) (**Fig S2**).

Hg stable isotope analyses

Using the previously determined THg concentration, an aliquot of the trapping solution was

diluted with ultra-high purity water to an acid content $< 10\%$ H^+ and a THg concentration of 3.75

to 5.00 nM (48). A Neptune Plus multicollector-inductively coupled plasma mass spectrometer

(MC-ICP-MS, Thermo Scientific) was used for Hg stable isotope ratio measurements at the

USGS MRL. A concentration and matrix-matched NIST 3133 Hg standard was used for sample

bracketing (49). To achieve detection at low Hg concentrations, Hg was reduced with stannous

chloride in-line then introduced continuously into a custom designed gas liquid separator along

with a thallium standard (NIST SRM 997, 195 nM in 3% HCl) for mass bias correction (48, 50).

MC-ICP-MS instrument conditions and analytical expectations were derived from previously

published work. All quality control metrics for Hg stable isotope analysis and raw data are

provided in **Supporting Methods, Table S1** and **Table S2**. Note that very low levels of Hg^0

were produced in abiotic controls for both phototrophic and fermentative experiments and as

such there was insufficient Hg^0 to measure isotope fractionation in the product pool (**Table S2**).

Isotope calculations

Delta calculations followed the conventions set forth by others (49). This convention calls for

mass dependent fractionation (MDF) to be expressed in terms of $\delta^{xxx}Hg$ where graphically

$\delta^{202}Hg$ is used to signify MDF. $\delta^{xxx}Hg$ is calculated as:

$\delta^{xxx}Hg(‰) = [(^{xxx}Hg/^{198}Hg_{sample})/(^{xxx}Hg/^{198}Hg_{NIST-3133}) - 1] \times 1000.$ (**Eq 1**)

XXX is used to signify the isotope of interest. Hg also undergoes MIF of both even and odd

isotopes. Here, odd-MIF is described by $\Delta^{199}Hg$ and even-MIF by $\Delta^{200}Hg$. MIF is calculated as:

$\Delta^{xxx}Hg \approx \delta^{xxx}Hg - (\delta^{202}Hg * \beta).$ (**Eq 2**)

β denotes the mass dependent scaling constant, which is determined by the laws of mass

dependence (49). Isotope enrichment effects based on the ratio of products to reactants (i.e. $\varepsilon_{p/r}$,

herein referred to as ε_{AP} for anoxygenic phototrophic reduction and ε_{FM} for fermentative

reduction) were calculated using Rayleigh fractionation models to account for the Hg^0 that was

removed following methods from previous work on microbial Hg reduction (35). This method

involved fitting a linear regression (shown in **Fig S3, Fig S4,** and **Table S3**) based on the change

in relative isotope ratios, $\ln(R/R_0)$, as a function of $\ln(f_r)$ where:

$R = (\delta^{202}Hg/1000) + 1$ for Hg^{II} (reactant) or Hg^0 (product) at a given time point and

$R_0 = (\delta^{202}Hg/1000) + 1$ at the beginning of the experiment (**Eq 3**);

$f_r = Hg^{II}$ in the reactor at a given time / Hg^{II} in the reactor at the start of the experiment (**Eq 4**);

and the slope of the linear regression, which is equivalent to ε_{AP} or ε_{FM}:

$\varepsilon = \ln (R/R_0)/ \ln(f_r)$ (**Eq 5**).

We chose to use the total Hg analyses from the bioreactor to establish f_r as this is a true

representation of the instantaneous isotope fractionation rather than a time integrated sample

following the methods outlined in previous work (35). Note that in our slope calculations we do

not have data for $^{202}Hg^0$ when f_r=1 (the 0h sampling point) because no isotope fractionation was

taking place prior to the start of Hg reduction (**Table S2** and **Table S3**). Additionally, in

experiments where negligible amounts of Hg^{II} were reduced to Hg^0, data points for $\delta^{202}Hg$ in the

Rayleigh fractionation plots tend to cluster around to $f_r = 1$ since ~100% of the original Hg

supplied remained in the reactor after 48 hours. Additional calculations for isotope mass balances

were carried out to test for deviation from the Rayleigh model and can be found in **Table S4.**

ε values for MDF were compiled from previous studies on pure microbial cultures and

abiotic Hg transformations for comparison purposes. Currently, there is no standard guideline for

how to calculate ε and studies vary widely in how they present isotope enrichment data and their

associated variability. We have provided details on how we normalized ε values from different

studies in **Table S5**. In this work we only compared enrichment factors from studies that used

MC-ICP-MS techniques and slope calculations in line with our analyses, however we have

included all of the ε values for microbe-driven Hg transformations published to date in **Table S5**.

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RESULTS AND DISCUSSION

Hg^0 production during anoxygenic photosynthesis and fermentation

In our experiments Hg^0 production relied on the presence of live cells for both fermentative and

phototrophic growth conditions. Cumulative Hg^0 production was an order of magnitude higher

for live cell treatments compared to abiotic controls (>0.60 vs <0.05 nmol) (**Fig 1**). Fermentative

cells produced double the amount of Hg^0 compared to phototrophic cells (1.33 to 1.64 nmol vs

0.60 to 0.81 nmol, respectively) despite having been supplied with slightly less Hg^{II} (**Fig 1B, E**).

In relative terms, phototrophic cells reduced 12 to 15% of the initial Hg^{II} supplied whereas

fermentative cells reduced 31 to 40% (**Fig S2**).

Relative Hg^0 production in these experiments was lower, yet consistent with our previous

work, which showed fermentatively-grown cells reduced 15% more Hg^{II} than phototrophically-

grown cells, reducing between 60 to 75% of the initial Hg^{II} supplied (23). The relative amount of

Hg^0 production did not increase with higher Hg^{II} exposure (10 nM) versus previous experiments

(250 pM) (23). It is unlikely that the lower reduction rate observed in the experiments presented

here was due to the toxicity of 10 nM Hg supplied, given that *H. modesticaldum* grows well

under the same conditions at 50 times higher Hg concentrations (23). Instead, we suspect that the

lack of constant bubbling of the reactor may be responsible for the observed results. It is possible

that the accumulation of CO_2 in the bioreactor following organic carbon oxidation may compete

with Hg^{II} as an electron sink, decreasing the amount of Hg^0 produced. Inorganic carbon is

required for anaplerotic reactions that fulfill the biosynthetic needs of *H. modesticaldum* and CO_2

could have competed with Hg^{II} for reducing power, as previously shown (51). These results

could also be due to Hg^0 oxidation allowed by the increased residence time of Hg^0 in the reactor

afforded by the intermittent sparging, as has been observed in other anaerobes (52, 53).

Despite carefully controlling the growth conditions in our experiments, fermentative cells

from replicate #1 exhibited a lag phase of 12 hours (**Fig 1F**). Although this lag did not affect

final cell density, the slow-growing culture from replicate #1 initiated Hg^0 production earlier

than the culture in replicate #2 (**Fig 1E, F**). These results may be attributable to cells in replicate

#1 maintaining a higher Hg to cell ratio in the first 12 hours of the experiment, which suggests

that Hg is being reduced intracellularly in line with our previous work on Heliobacteria (23, 54).

Indeed, even at low initial densities, cells can maintain µM levels of intracellular reducing

power, which could be used to reduce nM levels of Hg^{II} (24).

Mass dependent fractionation during photosynthetic and fermentative growth

Hg reduction in phototrophically-grown *H. modesticaldum* cultures resulted in consistent positive MDF and negligible Δ^{199}Hg values (-0.01 ± 0.05 ‰, n = 28) in the reactant pool indicating an absence of MIF (**Fig 2AB**, **Table S1,** and **Table S2**). Abiotic controls where light was supplied showed minimal change in Hg isotope values related to both MDF and MIF, confirming that Hg reduction and subsequent isotopic fractionation were predominantly driven by cellular processes (**Fig 2A**, **Table S1,** and **Table S2**).

We attribute the lack of photochemical fractionation in our experiments to the low levels of visible light and anoxic conditions employed in our setup. These conditions were designed to minimize the amount of abiotic photoreduction observed compared to high energy UV light sources (natural or controlled), such as those used to study Hg stable isotope fractionation in the oxygenic phototrophic green alga *Isochrysis galbana* (34). In the study on *I. galbana*, the authors suggested that UV-mediated production of free radical pairs, such as reactive oxygen species, during photosynthetic electron transport contributed to MIF during Hg reduction (34). Such reactions would likely be inhibited in our experimental setup because of the low energy wavelengths characteristic of the irradiance we provided and the active removal of oxygen required for the anoxygenic phototrophic growth of *H. modesticaldum*.

For phototrophically-grown cultures, the reactant pool δ^{202}Hg^{II} increased steadily over time suggesting cells preferentially reduced lighter Hg^{II} (**Fig 2A**). A mirror trend was observed for δ^{202}Hg^{0}, which was depleted in ^{202}Hg^{0} at the beginning of the experiment, became enriched with heavier isotopes as the reaction proceeded, and eventually approached the initial isotopic composition of the NIST 3133 standard (approximately 0‰) (**Fig 2B**).

Curiously, the $\delta^{202}Hg^0$ for $f_r = 0.94$ at 3h in the two phototrophic live cell replicates showed

that the product pool was initially enriched with the $^{202}Hg^0$ isotope (**Fig 2B**). A similar pattern

has been observed in iron reducing bacteria that preferentially reduced heavier Fe^{III} (55), and the

model anaerobic Hg methylator and reducer *Geobacter sulfurreducens* PCA (38, 40). The most

recent work on *G. sulfurreducens* has shown that uptake of Hg selects for lighter Hg isotopes but

cells can also access an isotopically heavier pool of Hg^{II} from the dissolved phase (38). It is

possible that a similar process is occurring for phototrophically-grown *H. modesticaldum*

wherein cells access an isotopically heavier pool of Hg during equilibrium binding of Hg^{II} to the

outside of the cell, followed by an alternative uptake process that selects for lighter pools of Hg.

Based on these results, we included all time points in isotopic enrichment calculations for Hg^{II}

but omitted the 3h time point in calculations for Hg^0 given that other cellular fractionation

processes could be occurring (**Fig S3, Fig S4,** and **Table S3**).

Replicated experiments with fermentative cultures showed more variability with respect to

their fractionation patterns compared to phototrophic cultures, however positive MDF was

discernable (**Fig 2CD**). Replicate #1 showed an initial depletion in reactant pool $\delta^{202}Hg^{II}$ and

became progressively enriched with $^{202}Hg^{II}$ relative to the NIST 3133 standard suggesting lighter

isotopes were still preferentially reduced (**Fig 2C**). Note that the second $\delta^{202}Hg^{II}$ data point

available for replicate #1 occurs at $f_r=0.72$ due to the rapid rate of Hg^{II} reduction leading to

substantially less Hg remaining in the reactor (**Fig 1E** and **Fig 2C**). Results for replicate #2

further suggested that lighter Hg^{II} isotopes were preferentially reduced with $\delta^{202}Hg^{II}$ slightly

increasing at the beginning of the experiment before rising to a maximum of 0.81 ‰ between $f_r =$

0.85 at 12h and $f_r = 0.48$ at 48h (**Fig 2C**).

The trends for $\delta^{202}Hg^0$ in fermentative cultures mirrored what was observed for $\delta^{202}Hg^{II}$.

$\delta^{202}Hg^0$ from replicate #1 showed that the product pool was initially enriched with heavy $^{202}Hg^0$,

further suggesting cells may have accessed a readily bioavailable pool of heavy Hg^{II}, before

undergoing a pronounced depletion and progressive enrichment with heavy isotopes similar to

phototrophic cultures (**Fig 2BD**). In contrast, the $\delta^{202}Hg^0$ values for cells in replicate #2 suggest

that the product pool was depleted in $^{202}Hg^0$ before undergoing enrichment for heavier Hg^0,

isotopes reaching a maximum of 1.73 ‰ (**Fig 2D**). This enrichment in $^{202}Hg^0$ product exceeded

the initial composition of the NIST 3133 standard (0 ‰) suggesting fractionation deviated from

the predicted Rayleigh fractionation for Hg reduction (**Table S4**). Our observations in the Hg^0

product pool are similar to what has been reported for anaerobic cultures of *G. sulfurreducens*,

which produced product pools $\delta^{202}Hg$ values that exceeded the initial composition of the Hg^{II}

supplied during methylation (40). As previously mentioned, this type of enrichment in the

product pool could be driven by differences in the bioaccessible pools of Hg. Indeed, more

recent work with *G. sulfurreducens* has reinforced the view that subcellular partitioning

processes can contribute to Hg isotope fractionation, although such changes were not detected at

the level of THg pool (38).

In our work, Hg reduction provided such a strong isotopic shift that we could detect

fractionation in the THg pool (**Fig 2**). However, we cannot discount the contributions of smaller

magnitude (<0.10 ‰) subcellular partitioning processes or the preferential utilization of different

intracellular Hg pools during Hg^{II} reduction that may more drastically affect the isotope values of

the Hg^0 product pool. Such processes could be driving the slight depletion of $^{202}Hg^{II}$ (-0.29 to -

0.18 ‰) early on for fermentative cells in replicate #1, where low cell densities may have

amplified uptake and adsorption-driven fractionation in comparison to experiments where cell

density increased by an order of magnitude over the same time frame (**Fig 1F** and **Fig 2C**). It is

more challenging to discuss the contributions of such processes for replicate #2, where the low

cumulative Hg^0 production at $f_r = 0.94$ and 0.88 (3h and 6h time points, respectively) precluded

our ability to measure Hg^0 isotope fractionation (**Fig 1E** and **Table S2**). Despite this limitation,

the high $\delta^{202}Hg^0$ value obtained for the final timepoint for replicate #2 reinforces that additional

processes are contributing to fractionation during fermentative growth that warrant further

exploration.

Our experimental results suggest that Hg^0 oxidation is not a major contributor at earlier

stages of reduction (i.e. 3 to 12 hours) but may be contributing to isotope fractionation at later

time points associated with longer Hg residence times in the reactor. Under phototrophic

conditions, the measured values of $\delta^{202}Hg^{II}$ at 48h were isotopically heavier than predicted

values, which suggests Hg^0 oxidation may be occurring (**Table S4**). Under fermentative

conditions the potential contribution of Hg oxidation is unclear, since a consistent enrichment

was not observed between replicates (**Table S4**). The variability in isotope results for later time

points in both fermentative replicates suggests processes other than reduction may be occurring.

Although Hg^0 oxidation has never been observed in anoxygenic phototrophs, it has been

observed in other model chemotrophic anaerobes (53). Recent work showed that thiol-bearing

molecules preferentially oxidize heavy $^{202}Hg^0$, leading to negative MDF, in addition to a

negative MIF signal (56). Although we did not observe any negative MDF or MIF in our

experiments (**Table S1** and **Table S2**) to the extent seen in previous work (56), we acknowledge

that such processes have the potential to contribute to net fractionation alongside internal

partitioning related to uptake or adsorption processes.

The results presented in this work are an important first step to resolving the physiological

processes that drive anaerobic Hg transformations in *H. modesticaldum.* We demonstrate that

while biotic reduction of Hg during anoxygenic photosynthesis and fermentation drives Hg

isotope fractionation, other processes such as adsorption, cellular uptake, and equilibrium effects

have an important impact on the net isotope values observed. Our results suggest that the extent

to which these additional processes affect isotope fractionation is tightly coupled to cellular Hg

availability and the physiological state of cell cultures. These biological controls and overlapping

Hg reactions are key to elucidating the Hg isotope signature associated with environmental

processes, such as reduction. Bearing these nuances in mind, we aim to carry out additional

experiments that assess the contributions of uptake, adsorption, and redox transformations to the

net Hg isotope fractionation associated with anaerobic Hg reduction.

Comparing enrichment factors for abiotic and biotic Hg transformations

Our initial motivation for this study was to test if phototrophic and fermentative Hg reduction led

to different isotope enrichment factors and establish whether different underlying processes

supported different metabolic Hg reduction pathways. When comparing ε values, which are used

as an index to distinguish between different isotope fractionation processes, we obtained average

enrichment factors using the ratio of product to reactant ($\varepsilon_{p/r}$) ± 2 standard deviations (2SD) for

MDF in the reactant pool Hg^{II} of -0.80 ± 0.56 ‰, for anoxygenic photosynthesis (ε_{AP}), and -1.15

± 0.48 ‰, for fermentation (ε_{FM}) (**Fig 3, Fig S3, Table S3,** and **Table S5**). The enrichment factor

could not be calculated using the product Hg^0 pool for fermentative cultures, and ε was

overestimated in phototrophic cultures (-2.86 ± 1.32 ‰), likely due to challenges in obtaining

analytical replicates of gaseous Hg^0 and quantitatively collecting this pool (57) (**Fig S4** and

Table S3). The small difference between ε_{AP} and ε_{FM} (i.e. 0.35 ‰) could suggest common

controls for Hg reduction during photosynthesis and fermentation. However, we acknowledge

that we could not effectively test our hypothesis due to the variability observed in replicated

fermentative experiments where the cell populations used were in different physiological states,

as supported by the growth curves (**Fig 1F**). Despite this limitation, our work demonstrates that

net isotopic enrichment stems from a combination of multiple processes most likely affected by

Hg transport (e.g. adsorption, uptake, and efflux) that would occur regardless of the intracellular

transformations in question.

Given that this is the first report of isotope enrichment signatures for Hg reduction during

anoxygenic photosynthesis and fermentation, there is value in comparing the signatures from this

work to previously published data. This comparison can help better understand the physiological

processes that control microbial Hg transformations and advance Hg isotope applications. For

convenience, we have presented the enrichment values for our replicated experiments separately

in **Figure 3** and provide data compiled from the literature alongside any manipulations carried

out to facilitate comparisons in **Table S5**.

The isotopic enrichment observed for both metabolisms tested in *H. modesticaldum* falls

between the values observed for abiotic and microbial Hg reduction pathways (**Fig 3** and **Table

S5**). The isotopic enrichment for MDF observed for replicated experiments with phototrophic *H.

modesticaldum* is slightly lower or equal to photoreduction in the presence of dissolved organic

carbon (-1.00 to -0.60 ‰ vs -0.60 ± 0.20 ‰, respectively) (32), but higher than photoreduction in

the presence of marine exudates and UVB light (-1.00 to -0.60 ‰ vs -1.47 ‰, respectively) (34)

(**Fig 3** and **Table S5**). When comparing Hg reduction in the anoxygenic phototroph *H.

modesticaldum* to the oxygenic phototrophic green alga *I. galbana*, the isotopic enrichment

observed is slightly lower or equal, depending on the treatment in question (-1.00 to -0.60 ‰ vs -

0.14 for high UV and -0.60 ‰ for low UV, respectively) (34) (**Fig 3** and **Table S5**). Though the

similar ε value of the low UV treatment suggests anoxygenic and oxygenic photosynthetic Hg

reduction share a common physiological pathway, the presence of MIF in *I. galbana* due to free

radical generation within the cell rules out this possibility (34). (**Table S1**). Thus, Hg reduction

during anoxygenic photosynthetic Hg reduction could be distinguished from other photo-induced

reduction pathways, which display MIF in the reactant Hg^{II} pool (58).

When compared with previous work on microbial Hg reduction in pure cultures, the isotopic

enrichment observed for phototrophically-grown *H. modesticaldum* was higher than what has

previously been reported for aerobic *mer*-mediated reduction and Hg reduction during anaerobic

respiration (-1.00 to -0.60 ‰ vs -1.80 to -1.20 ‰, respectively) (35, 37, 38) (**Fig 3** and **Table**

S5). Interestingly, the isotopic enrichment observed in fermentatively-grown *H. modesticaldum*

(-1.32 to -0.98 ‰) was closer to aerobic *mer*-mediated reduction (-1.40 to -1.20 ‰) (35) than to

Hg reduction during anaerobic respiration in *Shewanella oneidensis* MR-1 and a modified strain

of *G. sulfurreducens* PCA incapable of Hg methylation (-1.80 to -1.60 ‰) (**Fig 3** and **Table S5**)

(37, 38). The enrichment values reported for dark abiotic Hg^{II} reduction catalyzed by dissolved

organic matter (DOM) are slightly lower compared to fermentative Hg^{II} reduction (-1.32 to -0.98

‰ vs -1.52‰). Similar to what was mentioned for photoreduction previously, DOM-mediated

transformations display varying degrees of MIF such that fermentative Hg^{II} reduction could be

distinguished from DOM-mediated Hg reduction in the dark (**Fig 3** and **Table S5**) (59). An

additional line of evidence supporting that Hg^0 oxidation is unlikely to be contributing to the

isotope signature observed in our system are the ε values for thiol mediated Hg^0 oxidation in the

dark. These values are considerably larger than what we observed for *H. modesticaldum* (1.51 ±

0.20 ‰ for cysteine and 1.47 ± 0.24 ‰ for glutathione) (**Fig 3** and **Table S5**) (56)

These comparisons illustrate that isotopic enrichment for microbial Hg reduction

pathways that only display MDF fall within a narrow range (i.e. -0.60 to -1.80 ‰) (**Fig 3** and

Table S5). We find this to be a strikingly small difference given the ecological diversity of the

model organisms and variety of growth conditions used to study microbial Hg reduction. The

similar ε values reported for fermentation in this study and the *mer* operon are also noteworthy,

as they suggest nearly identical isotope enrichment can occur in aerobes harbouring dedicated

enzymatic machinery to reduce Hg and anaerobes where *mer*-based strategies are largely absent.

Our study also shows that similar isotopic enrichment can be observed for different

anaerobic Hg transformations. The anaerobic Hg methylating chemotrophs *G. sulfurreducens*

PCA and *Desulfovibrio desulfuricans* ND132 displayed ε signatures (i.e. -0.92 and -1.10 ‰,

respectively) that fall within the range of ε values we report in our study (i.e. -1.32 to -0.60 ‰)

(**Fig 3** and **Table S5**) (40). The similarity in enrichment signatures suggests that common

processes may affect the delivery of inorganic Hg substrates to intracellular sites independently

of the Hg transformations in question. Although it is outside of the scope of the current study,

identifying the nature of these processes may provide insights into similar controls that exist for

Hg uptake and transformation in anaerobes.

Our work showcases the strengths but also the challenges of using isotope fractionation-

based approaches to decipher cryptic Hg cycling pathways. To illustrate the challenges of using

isotopes to track Hg cycling pathways in the environment, we visualized our comparison of ε

values in the context of a vertically redox stratified habitat (e.g. a water column) (**Fig 3**).

Through Hg stable isotope fractionation, we can distinguish between abiotic and biotic redox

reaction pathways that control Hg's availability in methylation hotspots, such as the limit of the

photic zone, on the bases of MIF signatures associated with photochemical and DOM-mediated

reactions. In contrast, teasing apart biological pathways that contribute to the net isotope

signature of Hg^{II} in the environment is more challenging because competing MDF processes

occur in both dark and photo-mediated pathways where Hg^{II} is a reactant (e.g. reduction and

methylation) and a product (e.g. demethylation and oxidation). This overlap makes it difficult to

translate δ^{202}Hg values observed under controlled conditions that are sensitive to changes in

cellular physiology to environments where pools of bioavailable Hg are not homogenous due to

variable mixing.

CONCLUSION

In this study we provide the first reports of MDF associated with Hg reduction during

anoxygenic photosynthesis and fermentation. We show that isotopic enrichment during

anoxygenic phototrophic Hg reduction leads to an intermediate fractionation process that lies

between photochemical reactions and dark microbial Hg reduction. We also show that

fermentative Hg reduction leads to isotopic enrichment that is similar to aerobic *mer*-mediated

reduction. Although we could not discern whether dedicated physiological processes support Hg

reduction during anoxygenic photosynthesis and fermentation, our study shows that isotopes are

well-suited to teasing apart the complex interwoven processes that drive microbe-mediated Hg

transformations and highlighting common processes that control Hg bioavailability to microbes.

In future work, it will be important to isolate the different steps involved in Hg uptake

and transformation by rigorously controlling Hg availability and cellular physiology. Leveraging

molecular tools alongside isotopes will be a useful strategy to address these knowledge gaps and

the recent availability of genetically tractable *H. modesticaldum* deletion mutants offers a

promising means to do so (60). In that regard, gene deletion approaches targeting redox-active

enzymes could be used to characterize the physiological components contributing directly to Hg^{II}

reduction, whereas targeting homologues for proteins involved in Hg uptake in other model

anaerobes could be the key to identifying entry routes for Hg into *H. modesticaldum* (61, 62). By

combining molecular tools with isotope-based approaches to study anaerobic Hg reduction, we

may finally be able characterize the contributions of such pathways to the global Hg cycle in

present day environments and historical archives (e.g. sediment records). Such studies would

help better understand the ecological conditions that gave rise to the evolution of anaerobic Hg

redox cycling pathways prior to the acquisition of dedicated Hg detoxification strategies such as

the *mer* operon.

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AUTHOR CONTRIBUTIONS

507   DSG, AJP, SEJ and MTT designed all experiments; DSG and NCL carried out bioreactor

508   experiments and data analyses; SEJ and MTT carried out total Hg and stable isotope analyses;

509   DSG, AJP, SEJ and MTT wrote the manuscript.

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FIGURES AND CAPTIONS

Figure 1: Total Hg and cumulative Hg^0 production by _H. modesticaldum_ grown phototrophically and fermentatively. A, D THg for live cells and sterile controls without cells. **B, E** Cumulative Hg^0 production for live cells and sterile controls without cells. **C, F** Microbial growth as measured by O.D. 600 nm for live cells. 10 nM Hg was supplied in all experiments from the same stock (NIST 3133).

Figure 2: Mass dependent fractionation of ^{202}Hg in Hg^{II} (A, C) and Hg^{0} (B, D) during phototrophic and fermentative growth of *H. modesticaldum* and no cell controls. δ^{202}Hg values for Hg^{II} and Hg^{0} are plotted with respect to the fraction of remaining total inorganic Hg^{II} in the bioreactor. Insufficient Hg^{0} was recovered from the no cell controls and thus no results are displayed. 10 nM Hg was used in all experiments and Hg was supplied from the same stock (NIST 3133).

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Figure 3: Compilation of isotopic enrichment factors (ε) associated with mass dependent fractionation for abiotic and biotic Hg transformations. Hg transformation pathways are denoted by colours and the detection of mass-independent fractionation (MIF) is denoted by different shapes. Note that the data originating from this work is presented as separate replicates to emphasize the range of ε that observations fall within. For all other data, error bars denote 2 standard deviations for studies where this data was available. In the event that data was presented as a standard error between replicated experiments, standard deviation was calculated based on the sample size indicated in each study. Abbreviations on the y-axis denote the following: UV (ultraviolet light), DOC (dissolved organic carbon), DOM (dissolved organic matter), AP (anoxygenic photosynthesis), and FM (fermentation). The curves in this figure represent gradients for light and oxygen (O_2). The values in this figure were obtained from the following studies: Abiotic – Hg(II) photoreduction (32, 34); abiotic – Hg(II) reduction (59); abiotic – thiol Hg(0) oxidation (56); aerobic respiration - MerA Hg(II) reduction (35, 37); anaerobic respiration – Hg(II) methylation (40); anaerobic respiration – Hg(II) reduction (37, 38); anoxygenic photosynthesis – Hg(II) reduction (this study); fermentation – Hg(II) reduction (this study); oxygenic photosynthesis – Hg(II) reduction (34). The raw data in this figure are in **Table S5**.