Bolstering fitness via opportunistic CO_2 fixation: mixotroph dominance in modern groundwater

Martin Taubert[1,+], Will A. Overholt[1], Beatrix M. Heinze[1], Georgette Azemtsop Matanfack[2,5], Rola Houhou[2,5], Nico Jehmlich[3], Martin von Bergen[3,4], Petra Rösch[2], Jürgen Popp[2,5], Kirsten Küsel[1,6]

[1]Aquatic Geomicrobiology, Institute of Biodiversity, Friedrich Schiller University Jena, Dornburger Str. 159, 07743 Jena, Germany

[2]Institute of Physical Chemistry and Abbe Center of Photonics, Friedrich Schiller University Jena, Helmholtzweg 4, 07743 Jena, Germany

[3]Department of Molecular Systems Biology, Helmholtz Centre for Environmental Research – UFZ, Permoserstrasse 15, 04318 Leipzig, Germany

[4]Institute of Biochemistry, Faculty of Biosciences, Pharmacy and Psychology, University of Leipzig, Brüderstraße 32, 04103 Leipzig, Germany

[5]Leibniz-Institute of Photonic Technology, Albert-Einstein-Straße 9, 07745 Jena, Germany

[6]German Centre for Integrative Biodiversity Research (iDiv) Halle-Jena-Leipzig, Deutscher Platz 5E, 04103 Leipzig, Germany

[+]Corresponding author: Martin Taubert; Tel.: +49 3641 949459; Fax: +49 3641 949402; E-mail: martin.taubert@uni-jena.de

Keywords: Chemolithoautotrophy, Mixotrophy, groundwater, $^{13}CO_2$ stable isotope probing, genome-resolved metaproteomics, metagenomics, Raman microspectroscopy

Short title: Chemolithoautotrophy in shallow groundwater

Abstract

The current understanding of organic carbon inputs into ecosystems lacking photosynthetic primary

production is predicated on data and inferences derived almost entirely from metagenomic analyses.

The elevated abundances of putative chemolithoautotrophs in groundwaters suggest that dark CO_2

fixation is an integral component of subsurface trophic webs. To understand the impact of

autotrophically-fixed carbon, the flux of CO_2-derived carbon through various subpopulations of

subsurface microbiota must first be resolved, both quantitatively and temporally. Here, we

implement novel stable isotope cluster analysis to render a time-resolved and quantitative

evaluation of $^{13}CO_2$-derived carbon flow through a groundwater microbiome stimulated with reduced

sulfur compounds. We demonstrate that mixotrophs, not obligate chemolithoautotrophs, were the

most abundant active organisms in groundwater microcosms. Species of *Hydrogenophaga*,

Polaromonas, Dechloromonas, and other metabolically versatile mixotrophs drove the recycling of

organic carbon and, when chance afforded, supplemented their carbon requirements via

chemolithoautotrophy and uptake of available organic compounds. Mixotrophic activity facilitated

the replacement of 43 and 80% of total microbial carbon stores with ^{13}C in just 21 and 70 days,

respectively. This opportunistic "utilize whatever pathways net the greatest advantage in fitness"

strategy may explain the great abundances of mixotrophs in other oligotrophic habitats, like the

upper ocean and boreal lakes.

From soils to deep-sea sediments, the vast majority of cells on Earth must find a way to thrive in

environments devoid of photosynthesis[1]. To truly appreciate the global carbon cycle in all its

grandeur, it is important to understand the extent to which various types of cells rely upon

allochthonous or autochthonous carbon input. This dependence invokes selective pressures that

favor heterotrophic or chemolithoautotrophic lifestyles and provides the foundation upon which

trophic webs linking the entire subsurface biome are structured. Accurately gauging CO_2 fixation

rates and turnover in these habitats is remarkably challenging despite the invaluable utility afforded

by metagenomics to shed light on the metabolic capabilities of thousands of the organisms present[2-

5].

Modern groundwater, i.e., water having ingressed into the subsurface within the past 50 years[6], is a

transitionary ecosystem that connects surface habitats dominated by recently photosynthetically

fixed carbon with the subsurface, which is devoid of this carbon source entirely[7-9]. Here, inorganic

electron donors like reduced nitrogen, iron, and sulfur fuel chemolithoautotrophic primary

production[9-12]. Metagenomic-based studies have elucidated a diverse array of microorganisms

bearing the metabolic potential for chemolithoautotrophy[4,13-16], accounting for 12 to 47% of the

microbial population detected in groundwater[17-20]. Discoveries like these have cast doubt on

paradigms portraying modern groundwater as being dominated by heterotrophic microbes fueled by

organic material from the surface. We hypothesize that chemolithoautotrophic primary production

dictates the rates by which carbon is cycled in the modern groundwater microbiome.

To validate this hypothesis, we implemented a novel approach - stable isotope cluster analysis

(SIsCA), to render a time-resolved, quantitative assessment of CO_2-derived carbon flow through the

groundwater food web. By coupling stable isotope probing (SIP) with genome-resolved

metaproteomics[21,22], we leveraged the high sensitivity of Orbitrap mass spectrometry in SIP-

metaproteomics to acquire exceedingly accurate quantitative data on ^{13}C incorporation[23,24]. SIsCA

then employs a dimensionality reduction approach to incorporate the molecule-specific temporal

dynamics of the acquired isotopologue patterns resulting from ^{13}C-SIP time-series experiments –

ultimately discerning trophic interactions between individual members of the microbial community.

To examine the role of chemolithoautotrophy in the groundwater microbiome, we amended

groundwater microcosms with $^{13}CO_2$. Thiosulfate was used as an electron donor, as it is regularly

released into groundwater via rock weathering[25-28]. While organisms bearing the genetic potential to

oxidize reduced sulfur compounds are widespread in groundwater, little is known about their

lifestyles[15,16,29]. Under conditions favoring lithotrophic growth, we expected chemolithoautotrophy to

be the primary source of organic carbon, and a unidirectional carbon flux from autotrophs to

heterotrophs. By mapping the quantitative information derived from SIsCA to MAGs, we were able to

characterize carbon utilization and trophic interactions between active autotrophs and heterotrophs

in the groundwater microbiome over a period of 70 days. High-resolution monitoring of carbon

cycling and taxon-specific activities demonstrated that metabolically versatile mixotrophs, not strict

autotrophs, drove carbon flux in the groundwater, supplying up to 80% of the entire microbial

carbon. Insights into these microbes' lifestyles, as well as a discussion on how a metabolically flexible

mixotrophic lifestyle is optimally fit to flourish in an oligotrophic ecosystem, ensues.

Materials and Methods

Groundwater sampling and microcosms setup

Groundwater was collected from Hainich Critical Zone Exploratory (CZE) well H41 (51.1150842N,

10.4479713E) in June 2018. Well H41 provides access to an aquifer assemblage at 48 m depth in a

trochite limestone stratum. Sourced by a beech forest (*Fagus sylvatica*) recharge area, the oxic

groundwater in this well maintains mean dissolved oxygen concentrations of 5.0 ± 1.5 mg L^{-1}, < 0.1

mg L^{-1} ammonium, 1.9 ± 1.5 mg L^{-1} dissolved organic carbon, 70.8 ± 12.7 mg L^{-1} total inorganic

carbon, and a pH of 7.2 [27,30]. A total of 120 L of groundwater was sampled using a submersible pump

(Grundfos MP1, Grundfos, Bjerringbro, Denmark). To collect biomass from the groundwater, 5-liter

fractions were filtered through each of twenty 0.2-µm Supor filters (Pall Corporation, Port

Washington, NY, USA). The natural background of inorganic carbon in the groundwater was then

replaced with defined concentrations of ^{12}C or ^{13}C. Two 3-liter volumes of filtered groundwater were

acidified to pH 4 in 5-liter bottles to eliminate any bicarbonate. Following that, ^{12}C- or ^{13}C-

bicarbonate was dissolved in the groundwater to a final concentration of 400 mg L^{-1}, corresponding

to a near *in situ* concentration of 79 mg C L^{-1}. The pH of groundwater samples was then adjusted to

7.2 by addition of ^{12}C- or ^{13}C-CO_2.

Eighteen distinct microcosms were initiated for the ^{13}C-SIP experiment. For each microcosm, one

sample-laden 0.2-µm filter was placed into a 500-mL bottle containing 300 mL of treated

groundwater (as described above). Nine microcosms were sourced with water containing ^{12}C-

bicarbonate and the other nine with water containing ^{13}C-bicarbonate. Two additional microcosms

were prepared, each by transferring one 0.2-µm filter into a 1-liter bottle containing 350 mL of

untreated groundwater. One of these bottles was supplemented with 150 mL sterile D_2O (final

concentration 30%, v:v) and the other with 150 mL sterile milliQ H_2O. Sodium thiosulfate and

ammonium chloride were then was added to all microcosms to a final concentration of 2.5 mM and

105   15 µM, respectively. Finally, all microcosms were incubated with shaking (100 rpm) at 15 °C in the

106   dark.

Hydrochemical analyses

While incubating the 18 microcosms supplemented with [12]C- or ^{13}C-bicarbonate, concentrations of

oxygen, thiosulfate, and sulfate were determined at regular intervals. Oxygen concentrations were

determined using a contactless fiber-optic oxygen sensor (Fibox 4 trace with SP-PSt3-SA23-D5-YOP-

US dots [PreSens Precision Sensing GmbH, Regensburg, Germany]). Measurements were collected

from three ^{12}C microcosms and three ^{13}C microcosms every two days for the first three weeks, and

once weekly thereafter. Thiosulfate concentrations were determined via colorimetric titration assays

with iodine[31]. Samples from all microcosms were evaluated every four to seven days. For each

measurement, 2 mg potassium iodide was mixed into 1 mL of sample, followed by the addition of 10

µL of zinc iodide-starch solution (4 g L^{-1} starch, 20 g L^{-1} zinc chloride and 2 g L^{-1} zinc iodide) and 10 µL

of 17% (v:v) phosphoric acid. Titration was performed by adding 5 µL of 0.005 N iodine at a time until

the solution turned faint blue. Thiosulfate concentrations ($c_{thiosulfate}$ in mg L^{-1}) were then calculated

according to equation (1), where V_{iodine} is the volume of iodine solution added and V_{sample} is the

sample volume:

$$c_{thiosulfate} = \frac{V_{iodine} \times 561}{V_{sample}}$$ (1)

Sulfate concentrations were determined via a turbidimetric assay[32] from all microcosms every four to

seven days. For each measurement, 1 mL of either microcosm sample, standard (50 µM to 1000 µM

potassium sulfate) or blank (dH$_2$O) was mixed with 0.4 mL 0.5 M HCl and 0.2 mL BaCl$_2$-gelatin reagent

(0.5 g gelatin and 8 g BaCl$_2$ in 200 mL dH$_2$O). Following 1 h incubation in the dark, absorbances were

measured at 420 nm in a DR3900 spectrophotometer (HACH, Düsseldorf, Germany).

Detection of cellular activity by Raman microspectroscopy

Microcosms supplemented with D_2O or H_2O were sampled regularly during the first seven weeks of

incubation to quantify the incorporation of deuterium into the biomolecules of active cells (i.e.,

carbon-deuterium [C-D] bonds) via single-cell Raman microspectroscopy analysis. In preparation for

Raman microspectroscopy, 1 mL of sample was pre-filtered through a 5-μm filter, and then the cells

contained in the filtrate were washed three times with ddH_2O via centrifugation (10,000g, 2 min).

Pellets were then resuspended in 50 μL ddH_2O, and 10 μL of the final suspension was placed on

nickel foil (Raman substrate) and allowed to air dry at RT. Microbial cells were located via dark field

microscopy, and measurements were collected using a Raman microscope (BioParticleExplorer 0.5,

rap.ID Particle Systems GmbH) with an excitation wavelength of 532 nm (solid-state frequency-

doubled Nd:YAG module [Cobolt Samba 25 mW]; laser power = 13 mW at sample). The laser was

focused with an x100 objective (Olympus MPLFLN 100xBD) across a lateral spot of < 1 μm.

Backscattered light (180°) was diffracted using a single-stage monochromator (Horiba Jobin Yvon HE

532) with a 920 line mm^{-1} grating. Spectra were then registered with a thermoelectrically cooled CCD

camera (Andor DV401-BV), resulting in a resolution of ~ 8 cm^{-1}. A 5 s integration period was applied

per Raman spectrum (-57 to 3203 cm^{-1}).

Processing and analysis of Raman data

Processing and statistical analysis of raw Raman data were achieved with GNU R software[33]. Cosmic

spikes were removed from the spectra[34]. A wavenumber calibration was then applied using 4-

acetamidophenol standard spectra[35], while an intensity calibration was performed using the

SRM2242 standard[36,37]. The contribution of fluorescence was removed from spectra using the

asymmetric least-squares baseline correction method[38]. Finally, spectra were vector-normalized and

subjected to dimensionality reduction via principal component analysis (PCA). Five principal

components were used to build a linear discriminant analysis classification model, which was applied

to differentiate between deuterium-labeled and unlabeled bacterial cells. Deuterium uptake was

expressed as the C-D ratio, i.e., A(C-D) / [A(C-D) + A(C-H)], which was calculated by integrating the

areas of the C-H (2800 - 3100 cm^{-1}) and C-D (2040 - 2300 cm^{-1}) stretching vibration bands. Monitoring

deuterium incorporation into microbial cells helped gauge metabolic activity, as well as determine

optimal time points to sample microcosms.

Sampling and biomolecule extraction

After 21, 43, and 70 days of incubation, biomass was recovered from microcosms by filtering

aqueous phases through 0.2-µm Supor filters (Pall Corporation). Filters used for pre-incubation

biomass enrichment were combined with the filters used to remove the aqueous phases. A

combined DNA and protein extraction was performed using a phenol/chloroform/isoamylalcohol-

based protocol, as previously described[39]. Details regarding 16S rRNA gene amplicon sequencing and

quantitative SIP of DNA are provided in Supp. Info.

Metagenomic analysis

Metagenomic sequencing was performed on DNA samples selected from four ^{12}C microcosms: one

replicate each following 21 and 43 days of incubation, and two replicates following 70 days of

incubation. Samples were selected with the aim of covering greatest taxonomic diversity, as per the

results of PCA of 16S rRNA gene amplicon sequencing data. DNA fragment sizing, quantitation,

integrity, and purity were determined using an Agilent 2100 Bioanalyzer (Santa Clara, CA, USA).

Library preparation was achieved with a NEBNext Ultra II DNA Lib Prep Kit (New England Biolabs,

Ipswich, MA, USA) in accordance with protocols provided by the manufacturer. Multiplexed

sequencing in one flow cell of an Illumina NextSeq 500 system (300 cycles) ensued to generate 150-

bp paired-end reads.

Raw sequencing data was quality filtered using BBDuk[40] and subjected to assembly with metaSPAdes

v3.13.0[41]. Applying only contigs greater than 1,000 bp in length, three different algorithms facilitated

genomic binning: MaxBin 2.0 v2.2.7[42], MetaBAT 2 v2.12.1[43], and BinSanity v0.2.7[44]. Bin refinement

was accomplished using the MetaWRAP pipeline v1.1.3[45]. Only bins that were both more than 50%

complete and contained less than 10% contamination were considered. Bins were classified with

GTDB-Tk v0.3.2[46], and completeness parameters were appraised with CheckM v1.0.12[47]. Bins from

different samples were dereplicated using FastANI v1.0[48]. The Prokka pipeline v1.13.3[49] was used to

assign functional annotations to gene sequences and to translate into amino acid sequences for

metaproteomics analysis. Metagenomic bins of particular interest (per metaproteomics analysis)

were manually refined with Anvi'o v6.1[50], rendering the completed MAGs. Normalized coverage

values for all MAGs were calculated by dividing raw coverage values by the relative abundance of

rpoB genes in each metagenome. Gene abundances of *rpoB* were determined using ROCker[51]. Table

S1 provides an overview of the curated MAGs referred to in the study.

Metaproteomics analysis

Proteins extracted from microcosms were first subjected to SDS polyacrylamide gel electrophoresis,

followed by in-gel tryptic cleavage as previously described[39]. After reconstitution in 0.1% formic acid

(v:v), LC-MS/MS analysis was performed in LC chip coupling mode on a Q Exactive HF instrument

(Thermo Fisher Scientific, Waltham, MA, USA) equipped with a TriVersa NanoMate source (Advion

Ltd., Ithaca, NY, USA). Raw data files were analyzed using the Sequest HT search algorithm in

Proteome Discoverer (v1.4.1.14, Thermo Fisher Scientific, Waltham, MA, USA). Amino acid sequences

derived from the translation of genes present in metagenomes were compiled into a reference

database to facilitate protein identification. The following parameters were applied: enzyme

specificity was set to trypsin, two missed cleavages were allowed, oxidation (methionine) and

carbamidomethylation (cysteine) were selected as modifications, and peptide ion and Da MS/MS

tolerances were set to 5 ppm and 0.05, respectively. Peptides were considered identified upon

scoring a q-value < 1% based on a decoy database and obtaining a peptide rank of 1. Functional

classification of peptides was achieved in accordance with gene annotations generated by Prokka[49],

and taxonomic classification was based on the dereplicated and refined MAGs described above.

Stable Isotope Cluster Analysis

Peptide identifications from ^{12}C microcosms samples were mapped to mass spectra of corresponding ^{13}C-labeled samples, and the incorporation of ^{13}C was quantified by comparing expected peptide masses, chromatographic retention times, and MS/MS fragmentation patterns. Molecular masses of peptides were calculated based on amino acid sequences, isotopologue patterns of labeled peptides were extracted manually from mass spectral data using the Xcalibur Qual Browser (v3.0.63, Thermo Fisher Scientific, Waltham, MA, USA), and ^{13}C incorporation was calculated as previously described[24].

The conventional approach of calculating the most probable ^{13}C relative isotope abundance (RIA) of a peptide does not take into account the information contained in isotopologue patterns, which provide detailed information about the carbon utilization of an organism. To include this information in the analysis, we developed Stable Isotope Cluster Analysis (SIsCA). Stable Isotope Cluster Analysis (SIsCA) was performed using R[33], with scripts being available on github (https://github.com/m-taubert/SIsCA). Measured isotopologue patterns were compared to 21 predicted isotopologue patterns varying in ^{13}C RIA (5% intervals from 0 to 100% ^{13}C RIA), and coefficients of determination (R^2) were calculated for each comparison. With this approach, all information from isotopologue patterns is retained, while still data from different peptides is comparable, time series can be integrated, and the dataset can easily be used for downstream statistical analysis. To differentiate microbial lifestyles, R^2 values were averaged from samples obtained from replicate microcosms and peptides assigned to the same MAG. Resulting datasets of 21 R^2 values per time point per MAG were compared via PCA with the vegan software package[52], and clusters of MAGs were defined manually and validated by testing for overlapping confidence intervals.

Generation times of individual taxa were calculated by comparing the relative intensity of unlabeled and labeled peptide signals in mass spectrometric data, as previously described[23]. The number of doublings, n, was calculated according to equation (2) where I_{12_C} and I_{13_C} are the signal intensities of the unlabeled peptide and labeled peptide, respectively:

$$n = \log_2 \frac{I_{12_C} + I_{13_C}}{I_{13_C}} \qquad (2)$$

If the mass spectrometric signals of unlabeled and labeled and peptides overlapped, the

monoisotopic peak was used to determine the total abundance of unlabeled peptide based on the

natural distribution of heavy isotopes, as previously described[24]. Generation time, t_d, was calculated

with equation (3), where Δt is incubation time:

$$t_d = \frac{\Delta t}{n} \qquad (3)$$

Results

Sulfur oxidation by active groundwater microbes

Groundwater microbiota responded immediately to the addition of thiosulfate, yielding increasing

rates of sulfur oxidation. During the first three weeks of incubation, thiosulfate and oxygen

consumption rates remained relatively low (1.7 ± 1.9 and 5.5 ± 2.0 µmol d^{-1} [mean ± SD],

respectively; Fig. S1). Raman microspectroscopic analyses suggested that > 95% of cells were active

within the first 12 days of incubation. A distinct C-D band was observed at wavelength positions

between 2,100 and 2,300 cm^{-1} in the single-cell Raman spectra of the microcosm amended with D_2O

(Fig. 1, Fig. S2), which demonstrated new biomolecules were being synthesized by incorporating

deuterium from D_2O into carbon-deuterium bonds. The relative intensity of the C-D band increased

from 18.3% after 12 days to 25.7% after 47 days of incubation (median values; $p < 2.2 \times 10^{-16}$, $t = -$

14.038, df = 141.68, two-sided Welch's t-test), indicative of continued microbial proliferation and

cross-feeding on deuterium-labeled organic carbon.

After 70 days of incubation, consumption rates of thiosulfate (7.2 ± 2.0 µmol d^{-1}) and oxygen (12.8 ±

3.2 µmol d^{-1}) had increased significantly ($p = 6.48 \times 10^{-4}$, $t = 5.4332$, df = 7.8999 [thiosulfate] and $p =$

1.27 × 10^{-3}, $t = 4.7692$, df = 8.3019 [oxygen], two-sided Welch's t-test, Fig. S1). Sulfate was produced

at a consistent rate ranging between 8.1 and 9.6 µmol d^{-1} (no significant changes) throughout the

duration of the experiment. Recorded stoichiometry for oxygen:thiosulfate:sulfate was roughly 2.8 :

1 : 2.6 over the course of incubation, very near the theoretical ratio of 2 : 1 : 2 for oxygen-dependent

thiosulfate oxidation.

Organism-specific ^{13}C incorporation reveals distinct lifestyles

To address the carbon utilization schemes of key microbes, we conducted genome-resolved SIP-

metaproteomic analyses after 21, 43, and 70 days of incubation. SIsCA then clustered the 31 most

abundant MAGs into five distinct groupings, based on carbon utilization (Fig. 2, Fig. S3, Dataset S1).

Organisms represented by MAGs in cluster I were related to *Thiobacillus* (*Burkholderiales*) and

exhibited a stable ^{13}C RIA of 95% over throughout the 70 day experiment (Fig. 2). Such a high (>90%)

^{13}C RIA indicated exclusive CO_2 fixation. However, these strict autotrophs accounted for only 11% of

the total number of MAGs across the five clusters and 3.2 ± 3.1% (mean ± sd) of the total biodiversity

enveloped by the community, based on normalized coverages of the metagenomics dataset and 16S

rRNA gene profiles, respectively (Fig. S4, Fig. S5, Supp. Info.). By comparing the signal intensities of

^{12}C- and ^{13}C-enriched peptides, the generation time of these autotrophs was determined to be less

than 2 days (Fig. 3), highlighting the rapid production of new ^{13}C-labeled biomass from ^{13}CO$_2$.

Organisms represented by MAGs in cluster II were most closely related to species of

Methyloversatilis, *Polaromonas* and *Dechloromonas* (all *Burkholderiales*). These microbes exhibited a

moderate 65% ^{13}C RIA after 21 days of incubation (Fig. 2), which suggested the utilization of labeled

organic carbon from primary production as well as unlabeled organic carbon from the groundwater.

After incubating for 43 and 70 days, however, ^{13}C RIA increased to 91% (p = 1.573 × 10^{-3}, t = -3.5225,

df = 26.464, two-sided Welch's t-test), indicative of a switch to chemolithoautotrophic growth as

organic carbon became limited. Exhibiting generation times between 2 and 4 days (Fig. 3), MAGs

representing these mixotrophs were more than twice as abundant as those of cluster I, accounting

for 26% of the total normalized coverage.

Over the first 21 days of incubation, mean ^{13}C RIAs of cluster III and IV microbes increased from 65 to

76% and from 18 to 53%, respectively (p = 2.211 × 10^{-13}, t = -8.4984, df = 97.694 [cluster III] and p <

2.2 × 10^{-16}, t = -11.626, df = 58.764 [cluster IV], two-sided Welch's t-test; Fig. 2). This increasing trend

of ^{13}C RIA demonstrated two important points: First, it clearly indicated heterotrophic growth, based

in part on organic carbon produced by chemolithoautotrophic organisms of clusters I and II. Second,

it illustrated the increased labeling of available organic carbon, through the fixation of ^{13}CO$_2$.

Variations observed in ^{13}C RIAs between species suggested different extents of cross-feeding on

chemolithoautotrophically produced organic carbon, potentially due to preferences for different

organic carbon compounds. Cluster III housed the largest fraction of the MAG population, accounting

for 28% of the total normalized coverage, while cluster IV accounted for 20% of this total. The vast

majority of organisms represented by MAGs in these clusters exhibited generation times between 3

and 4 days (Fig. 3). However, cluster III microbes most closely related to species of *Hydrogenophaga*,

Vitreoscilla, and *Rubrivivax* exhibited growth rates as fast as their cluster I counterparts.

In cluster V, average ^{13}C RIAs reached 6% after 21 days of incubation and did not change thereafter,

which hinted at active heterotrophic lifestyles early on in the experiment. Nonetheless, these

organisms represented 15% of the total normalized coverage of all clusters. Generation times for

cluster V microbes were slightly longer and more variable, ranging from 3.5 days for species of

Acidovorax to eight days for *Aquabacterium* spp. (Fig. 3).

Analyses of corresponding peptide RIAs of all analyzed MAGs showed that 43, 68, and 80% of all

carbon available to the microbial population was replaced with ^{13}C following 21, 43, and 70 days of

incubation, respectively. Quantitative DNA-SIP confirmed this labeling pattern via increases in the

number of, and buoyant density shifts associated with, ^{13}C-labeled OTUs (Fig. S6; Supp. Info.). SIsCA

revealed carbon transfer from autotrophic cluster I to mixotrophic cluster II, and from these two

further to the heterotrophs of cluster III through V through cross-feeding on ^{13}CO$_2$-derived organic

carbon.

Functional characterization of MAGs reveals putative mixotrophs

All of the putative autotrophs detected employed the Calvin-Benson-Bassham (CBB) cycle for CO_2

fixation (Fig. 4). Subunits of the enzyme ribulose-1,5-bisphosphate carboxylase/oxygenase (RuBisCO)

were detected in the proteomes of 15 of 31 MAGs, and 14 of these contained additional enzymes of

the CBB cycle. No other complete CO_2 fixation pathways were identified. Proteins of the CBB cycle

were present not only in strict or facultative autotrophs of cluster I (i.e., relatives of *Thiobacillus* spp.)

or cluster II (e.g., relatives of *Methyloversatilis, Polaromonas,* and *Dechloromonas* spp.), but also in

heterotrophic organisms most closely related to species of *Hydrogenophaga, Rhodoferax,*

Paucibacter, and *Rubrivivax* of clusters III and IV. Mixotrophs comprised > 50% of all microbial taxa

represented across all clusters, which underscored the immense importance of their contributions to

carbon cycling in modern groundwater. The mixotrophic lifestyle, by no means a rare or insignificant

trait in these groundwater microcosms, appeared to bestow fitness on the microbes.

MAGs express pathways for the utilization of reduced sulfur compounds

Sixteen MAGs expressed proteins for sulfur oxidation via the Sox or Dsr enzyme system (Fig. 4).

Cluster II, III, and IV microbes phylogenetically affiliated with species of *Methyloversatilis,*

Dechloromonas, Hydrogenophaga, Rhodoferax and other *Betaproteobacteriales* utilized the Sox

system exclusively. MAGs harbored gene clusters of the conserved *soxCDYZAXB* gene order (Fig. S7),

featuring the core components of the Kelly-Friedrich pathway[53,54]. This pathway facilitates the

complete oxidation of thiosulfate to sulfate, without free intermediates[29]. Accessory genes *soxVW,*

soxEF, soxTRS, and *soxH* were randomly distributed through the MAGs disconnected from the main

operon.

Cluster I microbes most closely related to *Thiobacillus* spp. produced enzymes for both the Sox and

Dsr system, and corresponding MAGs housed a truncated *soxXYZAB* gene cluster that lacked genes

soxCD required to oxidize the sulfane group of thiosulfate. As such, these organisms likely used the

branched thiosulfate oxidation pathway typical for *Thiobacillus* spp.[55], whereby Dsr operating in

reverse oxidizes the sulfane-derived sulfur atom to sulfite, with elemental sulfur as intermediate[29].

Cluster I MAGs maintained the conserved operon structure *dsrABEFHCMKLJOPNR*, including genes

dsrEFH and *dsrL* typical for sulfur oxidizers but lacking gene *dsrD* for sulfate reduction[28]. These

organisms also expressed *aprAB* and *sat*, which encode Adenosine-5'-phosphosulfate reductase and

ATP sulfurylase, respectively, each of which can function in reverse to oxidize sulfite to sulfate [56].

Hence, groundwater facultative chemolithoautotrophs employed the Sox system to oxidize

thiosulfate to sulfate, while obligate chemolithoautotrophs utilized an incomplete version of this

system to oxidize the sulfone group and the Dsr/Apr/Sat system to oxidize the sulfane group of

thiosulfate.

Use of alternative electron acceptors and donors in sulfur oxidizers

Cytochrome c oxidase and other enzymes of the respiratory chain were detected in 15 sulfur oxidizer

MAGs, 12 of which also harbored enzymes for nitrate reduction (i.e., nitrate reductase, nitrite

reductase, nitric oxide reductase; Fig. 4). Several sulfur oxidizers related to species of *Dechloromonas*

and *Rhodoferax* expressed both pathways concurrently. Proteins for ammonia oxidation (*i.e.*,

ammonia monooxygenase, hydroxylamine oxidoreductase) were produced by a variety of cluster I

and IV microbes, such as *Thiobacillus* and *Methyloversatilis* species. MAG_77 (*Thiobacillus*), MAG_55

(*Dechloromonas*) and MAG_7 (*Hydrogenophaga*) even expressed [NiFe]-hydrogenase genes.

Utilization of organic carbon in oligotrophic groundwater

Cluster I, II, and III MAGs exhibited a gradient of increased versatility in utilizing various organic

carbon compounds. While cluster I's strict autotrophs only expressed pathways for sugar

degradation, MAGS of clusters II through V produced proteins germane to the breakdown and

transport of simple sugars (*e.g.*, glycolysis, pentose phosphate pathway), amino acids (TCA cycle),

fatty acids (beta-oxidation), C_1 compounds, and aromatics (Fig. 4). The TCA cycle was one of the most

abundant metabolic modules observed in MAGs of cluster II to V. Degradation pathways for toluene

and ethylbenzene were expressed by organisms most closely related to species of *Dechloromonas*

and *Rhizobacter* (*Betaproteobacteriales*), respectively. Enzymes for naphthalene and catechol

catabolism were detected in MAGs representing organisms related to *Hydrogenophaga* and

Pseudomonas spp., while gene products germane to the degradation of complex carbohydrates (*e.g.*,

starch, chitin) were produced by MAGs representing relatives of *Microbacterium* and

Sediminibacterium species. The metabolic machinery required to metabolize C_1 compounds was

detected primarily in microbes related to *Methyloversatilis* spp., which typically possessed methanol

dehydrogenase, formate dehydrogenase, and other enzymes involved in tetrahydromethanopterin-

dependent C_1-cycling.

Gene products relevant to import systems for amino acids and carboxylic acids (*e.g.*, alpha-keto

acids, C_4-dicarboxylates, lactate) were overly abundant in mixotrophs and heterotrophs of clusters II

to V (Fig. 4). Cluster III to V microorganisms that had grown exclusively heterotrophically exhibited

the greatest diversity of import-related proteins, including those for the transport of carbohydrates

and nucleotides. Only transporters targeting cations (predominantly iron) and phosphate were

detected in MAGS representing cluster I obligate autotrophs.

Discussion

Despite conditions strongly favoring autotrophic sulfur oxidizers, mixotrophs – not obligate

chemolithoautotrophs, were the most abundant active microorganisms in the groundwater

microcosms. While a diverse microbial consortium was detected, strict chemolithoautotrophs

accounted for only 3% of the total groundwater biodiversity. This is astonishing since thiosulfate and

oxygen were readily available throughout the experiment, which should have selectively promoted

the proliferation of chemolithoautotrophic microbes. Genome-resolved SIP-Metaproteomics

combined with our novel SIsCA approach facilitated identification of active microbes,

characterization of their expressed gene products (and linked pathways), and quantification of

carbon uptake and transfer within a diverse community over time (Fig. 5). Furthermore, highly

sensitive Raman spectroscopy showed that microbes were active at the outset of the incubation (no

discernable lag phase), despite low sulfur oxidation rates. Shedding new light on the mechanisms by

which CO_2-derived carbon is assimilated and cycled by groundwater microflora, this approach far

surpasses others that have implicated the importance of chemolithoautotrophy in groundwater

based solely on functional gene and metagenomics data[13,14,17,57]. Within 21 and 70 days of

incubation, 43 and 80% of the total groundwater biomass consisted of CO_2-derived carbon,

respectively. It is convincingly clear that this rapid enrichment of CO_2-derived carbon did not occur in

fixed, linear progression from chemolithoautotrophs to heterotrophs, but through a highly complex

and reticulated web of trophic interactions dominated by mixotrophs – the experts of organic carbon

recycling.

These mixotrophs strongly preferred heterotrophic growth to the fixation of CO_2, presumably a

consequence of the greater metabolic cost of carbon assimilation via the CBB cycle[58,59]. The ability to

fix CO_2 affords these microbes the luxury of an opportunistically selective lifestyle, which lends itself

to bolstered fitness (and rapid dominance) when organic carbon becomes limited in oligotrophic

systems. Cluster II mixotrophs, for example, transitioned from heterotrophy to CO_2 fixation late in

the incubation, likely due to such limitations. In a similar vein, cluster III mixotrophs expressed

pathways for autotrophic growth but were never required to fix CO_2. These microbes were able to

access a more diverse repertoire of carbon sources due to a greater metabolic versatility in organic

carbon utilization.

In support of the higher fitness associated with the opportunistic CO_2 fixation, mixotrophs grew

considerably faster (generation times of two days or less) than cluster IV and V organisms restricted

to an exclusively heterotrophic lifestyle (generation times up to 8 days). Surprisingly, however, these

heterotrophs were also able to oxidize reduced sulfur compounds, suggestive of a

chemolithoheterotrophic lifestyle. With respect to energy conservation, the constant influx of

reduced sulfur via weathering of interspersed pyrite minerals[25-27] renders sulfur oxidation an

attractive alternative to the oxidation of organic compounds, both in Hainich CZE groundwaters and

beyond. As the energetic requirement for CO_2 fixation is greater than the potential gain from organic

carbon oxidation, the most efficient strategy for both mixotrophs and heterotrophs is to net the

greatest amount of possible from sulfur oxidation and preserve precious organic carbon for anabolic

demands.

The diversity of organic carbon utilization motifs shifted gradually, and inversely, with CO_2 fixation. At

one end of the transition were the strict autotrophs of cluster I, relying exclusively on CO_2 as carbon

source. No organic carbon transporters were detected for any of these organisms. Their limited

metabolic breadth restricted growth to that from simple sugars, likely to utilize carbon assimilated

via the CBB cycle[58]. At the other end of the transition were organisms from clusters IV and V that

assimilated organic carbon exclusively. To endure the groundwater environment *sans* autotrophic

CO_2 fixation machinery, these organisms had to maintain and express a wide variety of organic

carbon transport and assimilation pathways. The most fit organisms in this modern groundwater

ecosystem, however, were the mixotrophs of clusters II and III. Establishing dominance by

opportunistically exploiting their physiological flexibility, these organisms rapidly outcompeted their

strictly autotrophic brethren (5-fold greater abundance).

Organisms most closely related to *Burkholderiales* spp., the key mixotrophic taxa in our groundwater

microcosms, gave rise to the greatest number of RuBisCO-encoding transcripts in a previous study at

our groundwater site[17]. For taxa like *Polaromonas*, *Dechloromonas*, *Hydrogenophaga*, and

Rhodoferax spp., the ability to oxidize sulfur has been posited based solely on genomic evidence[60-63].

Hitherto, chemolithoautotrophic growth on reduced sulfur compounds has not been observed from

any of these genera in pure culture. Our study demonstrates that these organisms can use reduced

sulfur as an energy source, and species of *Polaromonas*, *Dechloromonas,* and potentially

Hydrogenophaga used it to fuel autotrophic growth. These sulfur oxidizers expressed pathways for

both aerobic respiration and denitrification, despite the fact that no nitrate was added and nitrate

concentrations in the groundwater of this well never exceeded 10 mg/L[27]. Constitutive maintenance

and expression of denitrification enzymes is likely more energetically cost effective than regulating

gene expression[64]. This strategy also affords these microbes the advantage of utilizing different

electron acceptors when oxygen becomes limited.

Utilizing an incomplete TCA cycle that precludes heterotrophic growth, the *Thiobacillus*-related

organisms of cluster I are known to be obligately autotrophic[65]. Previously, by carrying out

thiosulfate- and hydrogen-driven denitrification, *Thiobacillus* spp. grew up to represent upwards of

50% of an enrichment culture obtained from Hainich CZE groundwater[66]. *In situ*, however,

Thiobacillus spp. are typically found in lower numbers[17], and most commonly appear in deeper, more

CO_2-rich subsurface systems[13]. This suggests diminished fitness and inability to compete with more

physiologically fit mixotrophs in oligotrophic modern groundwater. *Thiobacillus* can store the

elemental sulfur produced as intermediate by the Dsr enzyme system in periplasmic granules[65,67].

This storage might allow the organism to withstand times where no reduced sulfur compounds in the

groundwater are available.

There are two key advantages to being a mixotrophic sulfur oxidizer in the groundwater habitat. First

and foremost, these cells exist completely independent of surface carbon input dynamics. The energy

sources they rely on is produced autochthonously in the geological setting. Second, their diverse

breadth of physiological capabilities allows these microbes to modulate the means by which they

satisfy their anabolic requirements and energy demands based on the types of carbon available. This

includes carbohydrate degradation pathways for surface-derived plant polymers[8,68], amino acid and

nucleotide uptake systems for microbially-derived carbon[69,70], C_1 metabolic functions for C_1 carbon

compounds from biomass degradation[71], and hydrocarbon degradation pathways for rock-derived

carbon[72,73]. We hypothesize that similar strategies exploiting a myriad of carbon assimilation

pathways and versatile energy acquisition motifs benefit microbes dominating other oligotrophic

systems, such as boreal lakes or the upper ocean[74,75].

Conclusions

Our novel SIsCA-based approach facilitated the quantitative and temporal resolution of carbon flux through key subpopulations of a modern groundwater microbiome. Mixotrophs dominated this oligotrophic environment by fulfilling and supplementing their organic carbon requirements via opportunistic fixation of CO_2. This CO_2-derived organic carbon was rapidly incorporated into, and recycled throughout, microbial biomass through a highly efficient and complex trophic network. To mitigate low levels of organic carbon, autotrophic, mixotrophic, and heterotrophic microorganisms utilized reduced sulfur compounds as energy sources and preserved what organic carbon was available for anabolic demands. A wide variety of carbon assimilation pathways enabled mixotrophs and heterotrophs to make optimal use of the scarce amounts of organic carbon characteristic of oligotrophic environments. We posit that the concerted, opportunistic deployment of a wide variety of highly versatile pathways for assimilating carbon and generating energy from inorganic sources is key to microbial success in oligotrophic environments. The findings of this investigation significantly enhance our understanding of microbial survival strategies and their role in ecosystem functioning while demonstrating the powerful utility of next-generation physiology approaches like SIsCA in testing hypotheses established in metagenomics-based endeavors.

Declarations

Ethics approval and consent to participate

Not applicable

Consent for publication

Not applicable

Availability of data and materials

Metagenomic and amplicon sequencing data that support the findings of this study have been deposited into NCBI under the BioProject accession PRJNA633367.

Mass spectrometry proteomics data have been deposited into the ProteomeXchange Consortium via the PRIDE[76] partner repository with the dataset identifier PXD024889.

Competing interests

The authors declare no competing interests.

Funding

This work was supported financially by the Deutsche Forschungsgemeinschaft via the Collaborative Research Centre AquaDiva (CRC 1076 AquaDiva - Project-ID 218627073) of the Friedrich-Schiller-University Jena. Martin Taubert gratefully acknowledges funding from the DFG under Germany's Excellence Strategy - EXC 2051 - Project-ID 390713860. Climate chambers to conduct experiments under controlled temperature conditions and the infrastructure for Illumina MiSeq sequencing were financially supported by the Thüringer Ministerium für Wirtschaft, Wissenschaft und Digitale Gesellschaft (TMWWDG; project B 715-09075 and project 2016 FGI 0024 "BIODIV"). Martin von Bergen and Nico Jehmlich are grateful for support from the UFZ-funded platform for metabolomics and proteomics (MetaPro). Funding bodies played no role in the design of the study, collection, processing, and analysis of samples, interpretation of data, or writing of this manuscript.

Authors' contributions

MT and KK conceived and designed the study. MT conducted the microcosm experiments and molecular biology work. GAM, RH, PR, and JP conducted the Raman microspectroscopic analyses. NJ and MvB conducted mass spectrometric analysis for metaproteomics. MT analyzed the

490 metagenomics data with the assistance of WAO and BMH, and analyzed the SIP-metaproteomics

491 data. MT wrote the manuscript with contributions from all authors.

Acknowledgments

We are grateful to Robert Lehmann, Falko Gutmann, Heiko Minkmar, Jens Wurlitzer, and Lena

Carstens for assistance with field and lab work, and sampling of groundwater. We thank Julian

Hniopek for help with Raman measurements, and Daniel Desirò and Martin Hölzer for assistance

with sequencing. We also extend gratitude to Ivonne Görlich and Marco Groth at the Core Facility

DNA sequencing of the Leibniz Institute on Aging - Fritz Lipmann Institute in Jena for help with

Illumina sequencing.

References

1   Flemming, H. C. & Wuertz, S. Bacteria and archaea on Earth and their abundance in biofilms.

Nat. Rev. Microbiol. **17,** 247-260, doi:10.1038/s41579-019-0158-9 (2019).

2   Wrighton, K. C. *et al.* Fermentation, hydrogen, and sulfur metabolism in multiple

uncultivated bacterial phyla. *Science* **337,** 1661-1665, doi:10.1126/science.1224041 (2012).

3   Albertsen, M. *et al.* Genome sequences of rare, uncultured bacteria obtained by differential

coverage binning of multiple metagenomes. *Nat. Biotechnol.* **31,** 533-538,

doi:10.1038/nbt.2579 (2013).

4   Anantharaman, K. *et al.* Thousands of microbial genomes shed light on interconnected

biogeochemical processes in an aquifer system. *Nat. Commun.* **7,** 1-11,

doi:10.1038/Ncomms13219 (2016).

5   Parks, D. H. *et al.* Recovery of nearly 8,000 metagenome-assembled genomes substantially

expands the tree of life. *Nat. Microbiol.* **2,** 1533-1542, doi:10.1038/s41564-017-0012-7

(2017).

6   Gleeson, T., Befus, K. M., Jasechko, S., Luijendijk, E. & Cardenas, M. B. The global volume and

distribution of modern groundwater. *Nat. Geosci.* **9,** 161-167, doi:10.1038/NGEO2590 (2016).

7   Akob, D. M. & Küsel, K. Where microorganisms meet rocks in the Earth's Critical Zone.

Biogeosciences **8,** 3531-3543, doi:10.5194/bg-8-3531-2011 (2011).

8   Griebler, C. & Lueders, T. Microbial biodiversity in groundwater ecosystems. *Freshwater Biol.*

54, 649-677, doi:10.1111/j.1365-2427.2008.02013.x (2009).

9   Bell, E. *et al.* Active sulfur cycling in the terrestrial deep subsurface. *ISME J.* **14,** 1260-1272,

doi:10.1038/s41396-020-0602-x (2020).

10   Einsiedl, F. & Mayer, B. Hydrodynamic and microbial processes controlling nitrate in a

fissured-porous karst aquifer of the Franconian Alb, Southern Germany. *Environ. Sci. Technol.*

40, 6697-6702, doi:10.1021/es061129x (2006).

11   Schlesinger, W. H. On the fate of anthropogenic nitrogen. *P. Natl. Acad. Sci. USA* **106,** 203-

208, doi:10.1073/pnas.0810193105 (2009).

12   McCollom, T. M. & Seewald, J. S. Serpentinites, hydrogen, and life. *Elements* **9,** 129-134,

doi:10.2113/gselements.9.2.129 (2013).

13   Emerson, J. B., Thomas, B. C., Alvarez, W. & Banfield, J. F. Metagenomic analysis of a high

carbon dioxide subsurface microbial community populated by chemolithoautotrophs and

bacteria and archaea from candidate phyla. *Environ. Microbiol.* **18,** 1686-1703,

doi:10.1111/1462-2920.12817 (2016).

14   Probst, A. J. *et al.* Differential depth distribution of microbial function and putative symbionts

through sediment- hosted aquifers in the deep terrestrial subsurface. *Nat. Microbiol.* **3,** 328-

336, doi:10.1038/s41564-017-0098-y (2018).

15   Anantharaman, K. *et al.* Expanded diversity of microbial groups that shape the dissimilatory

sulfur cycle. *ISME J.* **12,** 1715-1728, doi:10.1038/s41396-018-0078-0 (2018).

16   Wegner, C. E. *et al.* Biogeochemical regimes in shallow aquifers reflect the metabolic

coupling of the elements nitrogen, sulfur, and carbon. *Appl. Environ. Microb.* **85,** 1-18,

doi:10.1128/AEM.02346-18 (2019).

17   Herrmann, M. *et al.* Large fractions of CO_2-fixing microorganisms in pristine limestone

aquifers appear to be involved in the oxidation of reduced sulfur and nitrogen compounds.

Appl. Environ. Microb. **81,** 2384-2394, doi:10.1128/Aem.03269-14 (2015).

18   Probst, A. J. *et al.* Genomic resolution of a cold subsurface aquifer community provides

metabolic insights for novel microbes adapted to high CO_2 concentrations. *Environ.*

Microbiol. **19,** 459-474, doi:10.1111/1462-2920.13362 (2017).

19   Jewell, T. N. M., Karaoz, U., Brodie, E. L., Williams, K. H. & Beller, H. R. Metatranscriptomic

evidence of pervasive and diverse chemolithoautotrophy relevant to C, S, N and Fe cycling in

a shallow alluvial aquifer. *ISME J.* **10,** 2106-2117, doi:10.1038/ismej.2016.25 (2016).

20   Handley, K. M. *et al.* The complete genome sequence for putative H_2- and S-oxidizer

Candidatus Sulfuricurvum sp., assembled de novo from an aquifer-derived metagenome.

Environ. Microbiol. **16,** 3443-3462, doi:10.1111/1462-2920.12453 (2014).

21   Neufeld, J. D. *et al.* DNA stable-isotope probing. *Nat. Protoc.* **2,** 860-866,

doi:10.1038/nprot.2007.109 (2007).

22   von Bergen, M. *et al.* Insights from quantitative metaproteomics and protein-stable isotope

probing into microbial ecology. *ISME J.* **7,** 1877-1885, doi:10.1038/ismej.2013.78 (2013).

23   Taubert, M. *et al.* Protein-SIP enables time-resolved analysis of the carbon flux in a sulfate-

reducing, benzene-degrading microbial consortium. *ISME J.* **6,** 2291-2301,

doi:10.1038/ismej.2012.68 (2012).

24   Taubert, M., Baumann, S., von Bergen, M. & Seifert, J. Exploring the limits of robust detection

of incorporation of ^{13}C by mass spectrometry in protein-based stable isotope probing

(protein-SIP). *Anal. Bioanal. Chem.* **401,** 1975-1982, doi:10.1007/s00216-011-5289-4 (2011).

25   Rimstidt, J. D. & Vaughan, D. J. Pyrite oxidation: A state-of-the-art assessment of the reaction

mechanism. *Geochim. Cosmochim. Ac.* **67,** 873-880, doi:10.1016/S0016-7037(02)01165-1

(2003).

26   Schippers, A., Jozsa, P. G. & Sand, W. Sulfur chemistry in bacterial leaching of pyrite. *Appl.*

Environ. Microb. **62,** 3424-3431, doi:10.1128/Aem.62.9.3424-3431.1996 (1996).

27   Kohlhepp, B. *et al.* Aquifer configuration and geostructural links control the groundwater

quality in thin-bedded carbonate-siliciclastic alternations of the Hainich CZE, central

Germany. *Hydrol. Earth Syst. Sc.* **21,** 6091-6116, doi:10.5194/hess-21-6091-2017 (2017).

28   Grimm, F., Franz, B. & Dahl, C. in *Microbial Sulfur Metabolism* (eds C. Dahl & C.G. Friedrich)

101-116 (Springer, 2008).

29   Ghosh, W. & Dam, B. Biochemistry and molecular biology of lithotrophic sulfur oxidation by

taxonomically and ecologically diverse Bacteria and Archaea. *FEMS Microbiol. Rev.* **33,** 999-

1043, doi:10.1111/j.1574-6976.2009.00187.x (2009).

30   Schwab, V. F. *et al.* Functional diversity of microbial communities in pristine aquifers inferred

by PLFA- and sequencing-based approaches. *Biogeosciences* **14,** 2697-2714, doi:10.5194/bg-

14-2697-2017 (2017).

31   DEV-D15. *Deutsche Einheitsverfahren zur Wasser-, Abwasser- und Schlammuntersuchung.*

Physikalische, chemische, biologische und bakteriologische Verfahren., Vol. DEV-D15 (VCH

Verlagsgesellschaft, 1975).

32   Tabatabai, M. A rapid method for determination of sulfate in water samples. *Environ. Lett.* **7,**

237-243 (1974).

33   R Core Team. in *R Foundation for Statistical Computing*   (R Core Team, Vienna, Austria,

2019).

34   Ryabchykov, O. *et al.* Automatization of spike correction in Raman spectra of biological

samples. *Chemometr. Intell. Lab.* **155,** 1-6, doi:10.1016/j.chemolab.2016.03.024 (2016).

35   Dörfer, T., Bocklitz, T., Tarcea, N., Schmitt, M. & Popp, J. Checking and improving calibration

of Raman spectra using chemometric approaches. *Z. Phys. Chem.* **225,** 753-764,

doi:10.1524/zpch.2011.0077 (2011).

36   Bocklitz, T. W., Dörfer, T., Heinke, R., Schmitt, M. & Popp, J. Spectrometer calibration

protocol for Raman spectra recorded with different excitation wavelengths. *Spectrochim.*

Acta A. **149,** 544-549, doi:10.1016/j.saa.2015.04.079 (2015).

37   Guo, S. X. *et al.* Towards an improvement of model transferability for Raman spectroscopy in

biological applications. *Vib. Spectrosc.* **91,** 111-118, doi:10.1016/j.vibspec.2016.06.010

(2017).

38   Baseline correction of spectra v. 1.3-0 (2020).

39   Taubert, M. *et al.* Tracking active groundwater microbes with D_2O labelling to understand

their ecosystem function. *Environ. Microbiol.* **20,** 369-384, doi:10.1111/1462-2920.14010

(2018).

40   Bushnell, B. BBTools software package. *URL* http://sourceforge.net/projects/bbmap (2014).

41   Nurk, S., Meleshko, D., Korobeynikov, A. & Pevzner, P. A. metaSPAdes: a new versatile

metagenomic assembler. *Genome Res.* **27,** 824-834, doi:10.1101/gr.213959.116 (2017).

42   Wu, Y. W., Simmons, B. A. & Singer, S. W. MaxBin 2.0: an automated binning algorithm to

recover genomes from multiple metagenomic datasets. *Bioinformatics* **32,** 605-607,

doi:10.1093/bioinformatics/btv638 (2016).

43   Kang, D. W. D. *et al.* MetaBAT 2: an adaptive binning algorithm for robust and efficient

genome reconstruction from metagenome assemblies. *Peerj* **7,** 1-13, doi:10.7717/Peerj.7359

(2019).

44   Graham, E. D., Heidelberg, J. F. & Tully, B. J. BinSanity: unsupervised clustering of

environmental microbial assemblies using coverage and affinity propagation. *Peerj* **5,** 1-19,

doi:10.7717/Peerj.3035 (2017).

45   Uritskiy, G. V., DiRuggiero, J. & Taylor, J. MetaWRAP - a flexible pipeline for genome-resolved

metagenomic data analysis. *Microbiome* **6,** 1-13, doi:10.1186/S40168-018-0541-1 (2018).

46   Chaumeil, P. A., Mussig, A. J., Hugenholtz, P. & Parks, D. H. GTDB-Tk: a toolkit to classify

genomes with the Genome Taxonomy Database. *Bioinformatics* **36,** 1925-1927,

doi:10.1093/bioinformatics/btz848 (2019).

47   Parks, D. H., Imelfort, M., Skennerton, C. T., Hugenholtz, P. & Tyson, G. W. CheckM: assessing

the quality of microbial genomes recovered from isolates, single cells, and metagenomes.

Genome Res. **25,** 1043-1055, doi:10.1101/gr.186072.114 (2015).

48   Jain, C., Rodriguez, R. L., Phillippy, A. M., Konstantinidis, K. T. & Aluru, S. High throughput ANI

analysis of 90K prokaryotic genomes reveals clear species boundaries. *Nat. Commun.* **9,** 1-8,

doi:10.1038/s41467-018-07641-9 (2018).

49   Seemann, T. Prokka: rapid prokaryotic genome annotation. *Bioinformatics* **30,** 2068-2069,

doi:10.1093/bioinformatics/btu153 (2014).

50   Eren, A. M. *et al.* Anvi'o: an advanced analysis and visualization platform for 'omics data.

Peerj **3,** 1-29, doi:10.7717/peerj.1319 (2015).

51   Orellana, L. H., Rodriguez-R, L. M. & Konstantinidis, K. T. ROCker: accurate detection and

quantification of target genes in short-read metagenomic data sets by modeling sliding-

window bitscores. *Nucleic Acids Res.* **45,** e14, doi:10.1093/nar/gkw900 (2017).

52   Dixon, P. VEGAN, a package of R functions for community ecology. *J. Veg. Sci.* **14,** 927-930,

doi:10.1111/j.1654-1103.2003.tb02228.x (2003).

53   Friedrich, C. G., Rother, D., Bardischewsky, F., Quentmeier, A. & Fischer, J. Oxidation of

reduced inorganic sulfur compounds by bacteria: Emergence of a common mechanism? *Appl.*

Environ. Microb. **67,** 2873-2882, doi:10.1128/Aem.67.7.2873-2882.2001 (2001).

54   Kelly, D. P., Shergill, J. K., Lu, W. P. & Wood, A. P. Oxidative metabolism of inorganic sulfur

compounds by bacteria. *Anton. Leeuw. Int. J. G.* **71,** 95-107, doi:10.1023/A:1000135707181

(1997).

55   Beller, H. R. *et al.* Whole-genome transcriptional analysis of chemolithoautotrophic

thiosulfate oxidation by *Thiobacillus denitrificans* under aerobic versus denitrifying

conditions. *J. Bacteriol.* **188,** 7005-7015, doi:10.1128/Jb.00568-06 (2006).

56   Beller, H. R. *et al.* The genome sequence of the obligately chemolithoautotrophic,

facultatively anaerobic bacterium *Thiobacillus denitfificans. J. Bacteriol.* **188,** 1473-1488,

doi:10.1128/Jb.188.4.1473-1488.2006 (2006).

57 Alfreider, A., Vogt, C., Geiger-Kaiser, M. & Psenner, R. Distribution and diversity of

autotrophic bacteria in groundwater systems based on the analysis of RubisCO genotypes.

Syst. Appl. Microbiol. **32,** 140-150, doi:10.1016/j.syapm.2008.11.005 (2009).

58 Berg, I. A. Ecological aspects of the distribution of different autotrophic CO_2 fixation

pathways. *Appl. Environ. Microb.* **77,** 1925-1936, doi:10.1128/Aem.02473-10 (2011).

59 Quayle, J. R. & Ferenci, T. Evolutionary aspects of autotrophy. *Microbiol. Rev.* **42,** 251-273,

doi:10.1128/Mmbr.42.2.251-273.1978 (1978).

60 Mattes, T. E. *et al.* The genome of *Polaromonas* sp. strain JS666: Insights into the evolution of

a hydrocarbon- and xenobiotic-degrading bacterium, and features of relevance to

biotechnology. *Appl. Environ. Microb.* **74,** 6405-6416, doi:10.1128/Aem.00197-08 (2008).

61 Salinero, K. K. *et al.* Metabolic analysis of the soil microbe *Dechloromonas aromatica* str.

RCB: indications of a surprisingly complex life-style and cryptic anaerobic pathways for

aromatic degradation. *BMC Genomics* **10,** 1-23, doi:10.1186/1471-2164-10-351 (2009).

62 Kämpfer, P. *et al. Hydrogenophaga defluvii* sp. nov. and *Hydrogenophaga atypica* sp. nov.,

isolated from activated sludge. *Int. J. Syst. Evol. Micr.* **55,** 341-344, doi:10.1099/ijs.0.03041-0

(2005).

63 Jin, C. Z. *et al.* Genomic and metabolic insights into denitrification, sulfur oxidation, and

multidrug efflux pump mechanisms in the bacterium *Rhodoferax sediminis* sp. nov.

Microorganisms **8,** 262, doi:10.3390/microorganisms8020262 (2020).

64 Geisel, N. Constitutive versus responsive gene expression strategies for growth in changing

environments. *Plos One* **6,** e27033, doi:10.1371/journal.pone.0027033 (2011).

65 Boden, R., Hutt, L. P. & Rae, A. W. Reclassification of *Thiobacillus aquaesulis* (Wood & Kelly,

1995) as *Annwoodia aquaesulis* gen. nov., comb. nov., transfer of *Thiobacillus* (Beijerinck,

1904) from the *Hydrogenophilales* to the *Nitrosomonadales*, proposal of *Hydrogenophilalia*

class. nov within the '*Proteobacteria*', and four new families within the orders

Nitrosomonadales and *Rhodocyclales*. *Int. J. Syst. Evol. Micr.* **67,** 1191-1205,

doi:10.1099/ijsem.0.001927 (2017).

66 Kumar, S. *et al.* Thiosulfate- and hydrogen-driven autotrophic denitrification by a microbial

consortium enriched from groundwater of an oligotrophic limestone aquifer. *FEMS*

Microbiol. Ecol. **94,** fiy141, doi:10.1093/femsec/fiy141 (2018).

67 Katayama-Fujimura, Y., Tsuzaki, N., Hirata, A. & Kuraishi, H. Polyhedral inclusion-bodies

(Carboxysomes) in *Thiobacillus* species with reference to the taxonomy of the genus

Thiobacillus. *J. Gen. Appl. Microbiol.* **30,** 211-222, doi:10.2323/Jgam.30.211 (1984).

68 Küsel, K. *et al.* How deep can surface signals be traced in the Critical Zone? Merging

biodiversity with biogeochemistry research in a Central German Muschelkalk landscape.

Front. Earth. Sci. **4,** 32, doi:10.3389/Feart.2016.00032 (2016).

69 Roth, V. N. *et al.* Persistence of dissolved organic matter explained by molecular changes

during its passage through soil. *Nat. Geosci.* **12,** 755-761, doi:10.1038/s41561-019-0417-4

(2019).

70 Herrmann, M. *et al.* Predominance of *Cand.* Patescibacteria in groundwater is caused by their

preferential mobilization from soils and flourishing under oligotrophic conditions. *Front.*

Microbiol. **10,** 1407, doi:10.3389/Fmicb.2019.01407 (2019).

71 Gray, C. M., Monson, R. K. & Fierer, N. Emissions of volatile organic compounds during the

decomposition of plant litter. *J. Geophys. Res.-Biogeo.* **115,** G03015,

doi:10.1029/2010jg001291 (2010).

72   Benk, S. A. *et al.* Fueling diversity in the subsurface: Composition and age of dissolved organic

matter in the Critical Zone. *Front. Earth. Sci.* **7,** 296, doi:10.3389/Feart.2019.00296 (2019).

73   Schwab, V. F. *et al.* ^{14}C-free carbon is a major contributor to cellular biomass in geochemically

distinct groundwater of shallow sedimentary bedrock aquifers. *Water Resour. Res.* **55,** 2104-

2121, doi:10.1029/2017WR022067 (2019).

74   Eiler, A. Evidence for the ubiquity of mixotrophic bacteria in the upper ocean: Implications

and consequences. *Appl. Environ. Microb.* **72,** 7431-7437, doi:10.1128/Aem.01559-06 (2006).

75   Hansson, T. H., Grossart, H. P., del Giorgio, P. A., St-Gelais, N. F. & Beisner, B. E.

Environmental drivers of mixotrophs in boreal lakes. *Limnol. Oceanogr.* **64,** 1688-1705,

doi:10.1002/lno.11144 (2019).

76   Perez-Riverol, Y. *et al.* The PRIDE database and related tools and resources in 2019:

improving support for quantification data. *Nucleic Acids Res.* **47,** D442-D450,

doi:10.1093/nar/gky1106 (2019).

Figures

Figure 1

Figure 1: Quantification of deuterium incorporation by single-cell Raman microspectroscopy.

Boxplots depict the relative intensity of Raman C-D bands, determined by A(C-D) / [A(C-D) + A(C-H)],

from single-cell Raman spectra. Spectra were obtained from groundwater microcosms with 30% D_2O

(shaded) or H_2O (empty) at various time points. Boxes show median, and first and third quartile;

whiskers denote 5th and 95th percentile. Outliers are depicted as dots. A minimum of 147 spectra

were obtained at each time point.

Figure 2

Figure 2: Clustering of selected MAGs based on carbon utilization. (a) Stable isotope cluster analysis

based on PCA of ^{13}C incorporation profiles over incubation time obtained from SIP-metaproteomics

of ^{13}C-microcosm samples. Each point represents a distinct organism represented by one MAG. MAG

clusters are indicated by Latin numbers. Ellipses depict 95% confidence intervals. All MAGs shown

facilitated the acquisition of at least two replicates of ^{13}C incorporation patterns per time point. (b)

Representative ^{13}C incorporation profiles of MAGs marked with asterisks are given for each cluster.

Heatmaps depict the extent of ^{13}C incorporation in peptides of the corresponding MAG after 21 (T1),

43 (T2), and 70 days (T3) of incubation (5% intervals, ranging from 0 to 100% ^{13}C relative isotope

abundance).

   Figure 3

cluster

Figure 3: Generation times of groundwater microorganisms. Values were determined for the first 3

weeks of incubation, based on the relative abundance of ^{12}C and ^{13}C peptides. Shown are mean and

standard deviation based on n ≥ 4 replicate determinations. Colored horizontal lines indicate average

generation time for each cluster. bdl: generation time fell below the detection limit of 2 days. na:

quantification of generation time was not possible.

Figure 4

Figure 4: Metabolic functionality of selected MAGs. The sizes of the bubbles correspond to the total

number of peptides detected for each MAG and each functional category identified at any time

point. Metabolic functions are grouped into CO_2 fixation (red), sulfur cycling (yellow), nitrogen cycling

(green), aerobic respiration and ATP synthesis (blue), organic carbon utilization (black), and import

functions (brown). The taxonomic categories "other" include peptides that were assigned to multiple

MAGs affiliated with the same genus. Only MAGs considered in the stable isotope cluster analysis are

shown. RuBisCO: ribulose-1,5-bisphosphate carboxylase/oxygenase, CODH/ACS: carbon monoxide

dehydrogenase/acetyl-CoA synthase, TCA cycle: tricarboxylic acid cycle.

Figure 5

Figure 5: Carbon flux between microbial clusters. Red arrow inlays illustrate the fraction of $^{13}CO_2$-

derived carbon assimilated by each microbial cluster after 21, 43, and 70 days. Arrow width scales

with the total amount of carbon assimilated based on the relative abundance of the respective

microbial cluster in the metagenomics analysis. Fading grey arrows indicate uptake of unlabeled

organic carbon from the groundwater. Checkboxes highlight the presence and activity of metabolic

functions for CO_2 fixation, utilization of organic carbon, and sulfur oxidation.