Controls on the isotopic composition of microbial methane

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Abstract

Microbial methane production (methanogenesis) is responsible for more than half of the annual emission of this major greenhouse gas to the atmosphere. Though the stable isotopic composition of methane is often used to characterize its sources and sinks, strictly empirical descriptions of the isotopic signature of methanogenesis currently limit such attempts. We developed a biochemical-isotopic model of methanogenesis by CO_2 reduction, which predicts carbon and hydrogen isotopic fractionations, and clumped isotopologue distributions, as functions of the cell's environment. We mechanistically explain multiple-isotopic patterns in laboratory and natural settings and show that such patterns constrain the in-situ energetics of methanogenesis. Combining our model with environmental data, we infer that in almost all marine environments and gas deposits, energy-limited methanogenesis operates close to chemical and isotopic equilibrium.

Introduction

Methane (CH_4) is a major greenhouse gas, with both natural and anthropogenic sources (1). The primary natural source of biogenic methane emissions is archaeal methanogenesis in anoxic environments (2), about a third of which is hydrogenotrophic (reduction of CO_2 with dihydrogen, H_2; Ref. 3). Strong isotopic discrimination during biological and abiotic methane formation has motivated the use of methane hydrogen and carbon isotopes to trace its production and consumption processes, construct global methane budgets and evaluate its climatic impacts (1, 4, 5). Current organism-level models that rely on isotopic mass balance can explain part of the observed range of microbial isotopic discrimination (6–9), but to date, such models have prescribed rather than resolved the microbial biochemistry. It has been difficult, therefore, to distinguish between different methane sources, different modes and extents of environmental methane cycling, and different environmental controls on the microbial isotope discrimination as drivers of observed variations in the isotopic composition of methane. To constrain the microbial component of such variations, we developed and analyzed a full metabolic-isotopic model of hydrogenotrophic methanogenesis with the net reaction:

$$4H_2 + CO_2 \rightleftarrows CH_4 + 2H_2O \ . \tag{1}$$
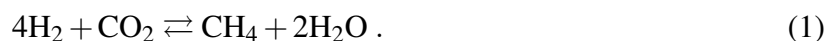
The model predicts the isotopic discrimination and its relation to the thermodynamic drive of this pathway (the Gibbs free energy of the net reaction, ΔG_{net}) and to cell-specific methanogenesis rates in laboratory cultures. Extending our analysis to energy-limited conditions, which are prevalent in natural environments, our model reveals the environmental and metabolic controls on the isotopic composition of methane.

Results and Discussion

A metabolic model of hydrogenotrophic methanogenesis

Accounting for the kinetics and thermodynamics of enzymatically-catalyzed reactions in hydrogenotrophic methanogenesis (details in the Methods and Supplementary Materials, SM), we constructed mass balance equations for the intracellular metabolites in the pathway (Fig. 1A). Given extracellular concentrations of CO_2, H_2 and CH_4, and pH, which define ΔG_{net}, these equations are solved for the steady-state concentrations of the intracellular metabolites (fig. S1). At this steady state, the model links metabolite concentrations to the net rate of methanogenesis, and to the gross forward and reverse rates of the individual reactions. We calibrated our metabolic model to available measurements of specific methanogenesis rates and their relation to H_2 concentrations in the μM to mM range (fig. S2; SM). The metabolic model accounts for electron transfer from H_2 to the intermediate metabolites through three electron carriers, cofactor F_{420}, coenzyme B (HS-CoB), and ferredoxin (Fd), thereby providing insight into the oxidation state of the cell. We validated the model's predictions for the oxidation state of these electron carriers against available measurements. Fmd catalyzes the first step in the pathway, CO_2 fixation to organic carbon with two electrons from an iron-sulfur-containing Fd (10). We found that to drive this reaction in the direction of net methanogenesis, Fd needs to be >90% reduced, and that with decreasing CO_2 concentrations, Fd must approach 100% reduced to compensate for the decreasing thermodynamic drive (fig. S1J). This is in line with previous estimates of the oxidation state of Fd for this reaction (11). F_{420} is 0.1% reduced at low H_2 concentrations, and up to 100% reduced at high H_2 concentrations (fig. S1K), in agreement with observations from lab cultures (12). Coenzyme B (HS-CoB) is 10% reduced at low H_2 concentrations, and up to 100% reduced at high H_2 concentrations (fig. S1L). The lower range is in agreement with observations of the oxidation state of HS-CoB at μM to 1 mM H_2 concentrations (13). Our observation-validated model predictions suggest that the intracellular oxidation state of electron carriers generally does not reflect electrochemical equilibrium, and is instead maintained dynamically, by oxidation and reduction fluxes that depend on the cell's metabolic activity. This

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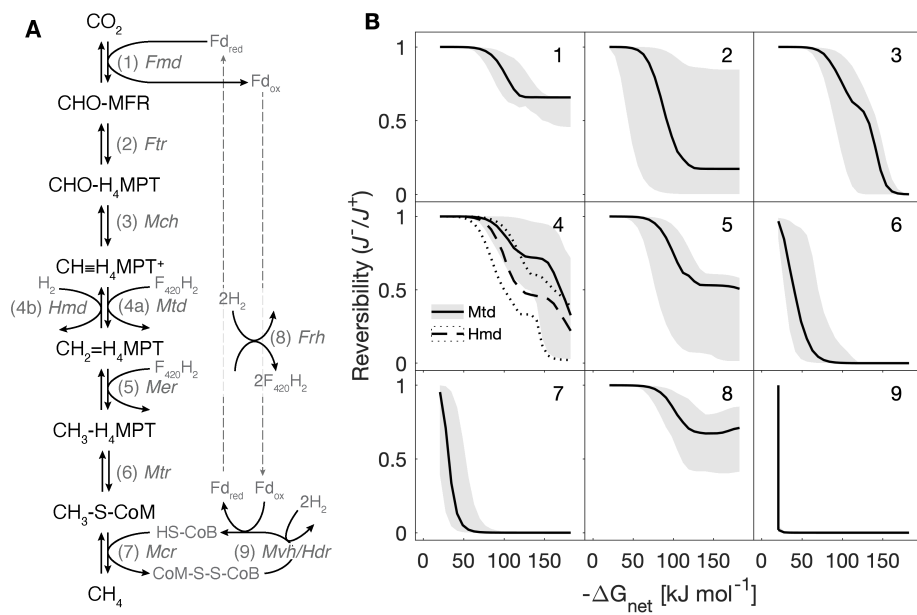
Figure 1: A metabolic model of hydrogenotrophic methanogenesis. (**A**) Schematic pathway. Abbreviated enzyme names are in italics (table S1). (**B**) Individual reaction reversibility (J_i^-/J_i^+) against ΔG_{net}. The subplots in panel B are numbered in accordance with panel A and table S1. The black lines are the median of 10^3 model simulations, and envelopes represent 95% of model results. The uncertainty originates from the enzyme kinetic parameters. Simulations were carried out at [H_2] of 1 nM to 10 mM, and typical experimental [CO_2] and [CH_4] of 10 mM and 10 μM, respectively.

dynamic control means that measurements of the oxidation state of specific reduction-oxidation couples does not necessarily bear on the oxidation state of other reduction-oxidation couples (both inorganic and organic) and on the overall oxidation state of the cell.

The calibrated and validated biochemical model forms the basis for the metabolic-isotopic coupling. With increasing H_2 concentrations, ΔG_{net} becomes increasingly negative, and the individual reactions in the pathway depart from equilibrium to various extents (Fig. 1B). We quantify the departure from equilibrium by the individual reaction reversibility, defined as the ratio of the reverse to forward gross rates of that reaction and related to its actual transformed Gibbs free energy ($\Delta G_i'$) (14–16):

$$J_i^-/J_i^+ = \exp(\Delta G_i'/RT) , \tag{2}$$

where R is the gas constant (J mol^{-1} K^{-1}), T is the temperature (K) and $\Delta G_i' = \Delta G_r'^0 + RT \ln Q$ (J mol^{-1}, where Q is the reaction quotient). For net forward methanogenesis rates, the reversibility

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76 of individual reactions varies between zero (a near-unidirectional, kinetically controlled reaction)

77 and unity (a near-equilibrium reaction). Some reactions do not fully depart from reversibility in

78 the explored ΔG_{net} space (e.g., the Frh-catalyzed reduction of F_{420}), whereas others depart from

79 reversibility at ΔG_{net} as modestly negative as -15 kJ mol^{-1} (e.g., the Mvh/Hdr, Mcr and Mtr-

80 catalyzed reactions). This differential response is a combined function of the reaction thermo-

81 dynamics, expressed as the standard-state transformed Gibbs free energy ($\Delta G_i'^0$), and the enzyme

82 kinetics, specifically the metabolic rate capacity (V^+) and Michaelis constants (K_M). Neither $\Delta G_i'^0$

83 nor V^+ and K_M in isolation predict the pattern of differential departure from equilibrium—coupled

84 thermodynamics and kinetics must be considered.

An isotopic model calibrated to laboratory cultures

86 The isotopic discrimination between substrate (s) and product (p) is described by the isotopic frac-

87 tionation factor $^r\alpha_{s/p} = {}^rR_s/{}^rR_p$, where $^{13}R = {}^{13}C/^{12}C$, $^2R = D/H$ (with $D \equiv {}^2H$ and $H \equiv {}^1H$), and

88 "r" denotes the rare isotope. The net isotopic fractionation expressed in an individual (bio)chemical

89 reaction may vary between thermodynamic equilibrium and kinetic end-members, associated re-

90 spectively with a reversible reaction and unidirectional forward reaction to form the reaction prod-

91 uct. With the reversibilities calculated in the metabolic model and with values assigned to the

92 equilibrium and kinetic isotopic fractionation factors (EFFs and KFFs, respectively) of the indi-

93 vidual reactions, the net isotopic fractionation between the pathway substrates and products may

94 be calculated (17, 18). We calibrated the relations between ΔG_{net} and the resulting net carbon and

95 hydrogen isotopic fractionations against experimental data, then used the calibrated model to de-

96 termine the relation between ΔG_{net} and the abundance of doubly substituted ("clumped") methane

97 isotopologues, which has yet to be systematically explored in experiments.

98 **Bulk carbon and hydrogen isotopic fractionation.** In laboratory cultures, the CO_2–CH_4 carbon

99 isotope fractionation ($^{13}\varepsilon_{CO_2\text{-}CH_4}$, where $\varepsilon = 1 - \alpha$ [‰]) is inversely related to ΔG_{net} (8, 19–23,

100 fig. S3A). At near-zero ΔG_{net}, the individual reactions in the pathway operate close to equilibrium

101    $(J_i^-/J_i^+ \rightarrow 1)$, and our model predicts $^{13}\varepsilon_{CO_2\text{-}CH_4}$ close to the temperature-dependent isotopic equi-

102   librium fractionation ($^{13}\varepsilon_{CO_2\text{-}CH_4}^{eq}$). When ΔG_{net} becomes slightly negative, $^{13}\varepsilon_{CO_2\text{-}CH_4}$ peaks to

103   larger-than-equilibrium values of 80-100‰ at $\approx$–45 kJ mol^{-1}, followed by a gradual decline to

104    $\approx$30‰ reached at ΔG_{net} of $\approx$–120 kJ mol^{-1} (Fig. 2A). Such larger-than-equilibrium $^{13}\varepsilon_{CO_2\text{-}CH_4}$

105   values have been observed in several experimental and environmental datasets and have yet to be

106   explained mechanistically. Our model reveals that this ΔG_{net}–$^{13}\varepsilon_{CO_2\text{-}CH_4}$ relation is controlled by

107   the landscape of departure from equilibrium of the individual reactions in the pathway. The car-

108   bon reaction network in methanogenesis is linear. In such networks, near-unidirectionality of an

109   individual reaction ($J_i^-/J_i^+ \rightarrow 0$) leads to expression of that reaction's KFF and suppresses iso-

110   topic fractionation associated with downstream reactions (18). The Mcr-catalyzed reaction's KFF

111   ($\approx$40‰, Ref. 24) is larger than its EFF ($\approx$1‰, Ref. 25). As this reaction departs from equilibrium,

112   the peak in $^{13}\varepsilon_{CO_2\text{-}CH_4}$ reflects a sum of its KFF and the EFFs of upstream reactions, with the ex-

113   ception of the Mtr-catalyzed reaction, which also partially departs from equilibrium (Fig. 1B). The

114   result is a larger-than-equilibrium $^{13}\varepsilon_{CO_2\text{-}CH_4}$ at modestly negative ΔG_{net} values. The $^{13}\varepsilon_{CO_2\text{-}CH_4}$

115   floor at ΔG_{net} more negative than $\approx$–120 kJ mol^{-1} (Fig. 2A) is defined by partial expression of the

116   KFFs of Ftr and Fmd, and suppression of the isotopic fractionations associated with downstream

117   reactions.

118   At large-negative ΔG_{net} values $^{13}\varepsilon_{CO_2\text{-}CH_4}$ is sensitive also to the extracellular partial pressure

119   of CO_2 ($p CO_2$). At a steady state, intracellular CO_2 utilization is exactly matched by net CO_2

120   diffusion across the membrane. This net diffusive flux is the difference between large gross fluxes

121   (into and out of the cell) when $p CO_2$ is high, and smaller gross fluxes at lower $p CO_2$. Thus, the

122   reversibility of net diffusion is low at low $p CO_2$, and suppression of downstream net carbon isotopic

123   fractionation results in small $^{13}\varepsilon_{CO_2\text{-}CH_4}$ (also referred to as a "reservoir effect"). The dependence of

124    $^{13}\varepsilon_{CO_2\text{-}CH_4}$ on $p CO_2$ explains the smallest net fractionations observed in laboratory cultures (lower

125   bound of red envelope in Fig. 2B, Ref. 21), as well as the dependence of $^{13}\varepsilon_{CO_2\text{-}CH_4}$ on pH in

126   hyperalkaline settings (26, 27).

127   The CH$_4$–H$_2$O hydrogen isotopic fractionation ($^2\varepsilon_{CH_4\text{-}H_2O}$) in laboratory cultures is $\approx$200‰
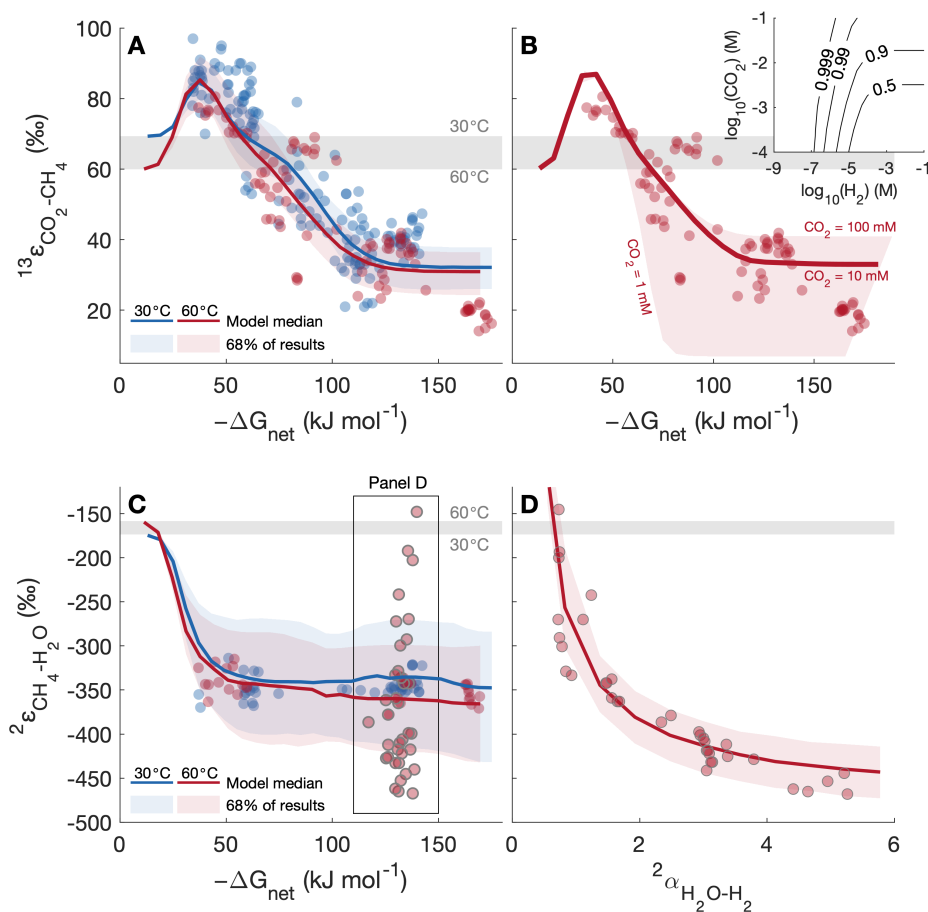
Figure 2: Model-laboratory culture comparison of bulk carbon and hydrogen isotopic fractionation. Mesophilic (30–40°C, blue circles) and thermophilic ($\geq$55°C, red circles) experimental data and model results (lines and envelopes) at 30°C (blue) and 60°C (red), and the equilibrium isotopic fractionation (gray envelopes). (**A, C**) $^{13}\varepsilon_{CO_2\text{-}CH_4}$ and $^2\varepsilon_{CH_4\text{-}H_2O}$ against ΔG_{net}. (**B**) $^{13}\varepsilon_{CO_2\text{-}CH_4}$ against extracellular CO_2 concentrations ([$CO_{2(out)}$]). Contours in the inset show the reversibility of cross-membrane CO_2 diffusion. (**D**) Mixing effects on $^2\varepsilon_{CH_4\text{-}H_2O}$ as a function of the H_2O–H_2 isotopic fractionation ($^2\alpha_{H_2O\text{-}H_2}$), compared to laboratory culture data (8).

more negative than the temperature-dependent isotopic equilibrium fractionation ($^2\varepsilon^{eq}_{CH_4\text{-}H_2O}$), and existing observations suggest that it does not display a clear dependence on ΔG_{net} (8, 19, 22, 28, 29, fig. S3B). Unlike the linear carbon reaction network, the hydrogen reaction network has four branches, each of which has the potential for hydrogen atom exchange between pathway intermediates and H_2O. Therefore, departure from equilibrium of one of the hydrogen atom exchange reactions does not preclude CH_4-H_2O hydrogen isotopic equilibrium. Specifically, the Mvh/Hdr-

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134   catalyzed reaction is near-irreversible at ΔG_{net} values as high as –25 kJ mol^{-1} (Fig. 1B), yet at this

135   ΔG_{net} value $^2\varepsilon_{CH_4\text{-}H_2O}$ is approximately equal to $^2\varepsilon^{eq}_{CH_4\text{-}H_2O}$ (Fig. 2C), and this arises from the high

136   reversibility of the other hydrogen atom exchange reactions in the pathway. Only when the Mcr-

137   and Mtr-catalyzed reactions sufficiently depart from reversibility (at $\Delta G_{net} \leq -30$ kJ mol^{-1}), cut-

138   ting off CH$_4$ and HS-CoB from exchange with upstream intermediates that are close to hydrogen

139   isotope equilibrium with H$_2$O, does $^2\varepsilon_{CH_4\text{-}H_2O}$ depart from $^2\varepsilon^{eq}_{CH_4\text{-}H_2O}$ (Fig. 2C). At ΔG_{net} more neg-

140   ative than -40 kJ mol^{-1} the KFFs of Mcr, Mtr and Mvh/Hdr control $^2\varepsilon_{CH_4\text{-}H_2O}$ (fig. S4). Overall,

141   our model reveals a clear $^2\varepsilon_{CH_4\text{-}H_2O}$–$\Delta G_{net}$ relation, which has not been accessed by the range of

142   ΔG_{net} explored in laboratory cultures to date. The apparent invariance of $^2\varepsilon_{CH_4\text{-}H_2O}$ at ΔG_{net} more

143   negative than –40 kJ mol^{-1} represents a complex, ΔG_{net}-dependent combination of isotope effects

144   associated with the enzymes in the pathway.

145   The hydrogen atom added to CH-H$_4$MPT may come from H$_2$O in the Mtd-catalyzed reaction,

146   or from H$_2$ in the Hmd-catalyzed reaction (Fig. 1A; 30). Thus, up to one quarter of the hydrogen

147   atoms in methane may come from H$_2$, depending on the relative activity of Mtd and Hmd. High

148   H$_2$ concentrations favor high methanogenesis rates and Hmd activity. Under these conditions, and

149   especially in the case of H$_2$-H$_2$O isotopic disequilibrium, $^2\varepsilon_{CH_4\text{-}H_2O}$ may vary in response to vari-

150   ations in the hydrogen isotopic composition of H$_2$. Our model captures this behavior, as observed

151   in culture experiments with 1 mM [H$_2$] and at 60°C (8, 29), displaying $^2\varepsilon_{CH_4\text{-}H_2O}$ values between

152   –145‰ and –480‰, which inversely covary with the H$_2$O-H$_2$ isotopic fractionation (Fig. 2D).

153   Our predicted trajectories for departure from equilibrium $^{13}\varepsilon_{CO_2\text{-}CH_4}$ and $^2\varepsilon_{CH_4\text{-}H_2O}$ values may

154   explain observations of near-equilibrium $^{13}\varepsilon_{CO_2\text{-}CH_4}$ values concurrent with clearly disequilibrium

155    $^2\varepsilon_{CH_4\text{-}H_2O}$ values, which have been previously explained by a decoupling of the carbon and hydro-

156   gen isotopic systems in methanogenesis (8, 31). We suggest instead, that the measured $^{13}\varepsilon_{CO_2\text{-}CH_4}$

157   values did not reflect isotopic equilibrium, but the descending branch from the $^{13}\varepsilon_{CO_2\text{-}CH_4}$ maxi-

158   mum (which occurs at $\Delta G_{net} \approx$–40 kJ mol^{-1}) with increasingly negative ΔG_{net}. In other words, we

159   suggest that apparent equilibrium $^{13}\varepsilon_{CO_2\text{-}CH_4}$ values may emerge by a fortuitous combination of

160   EFFs and KFFs, and not due to actual isotopic equilibrium between CO$_2$ and CH$_4$.

Clumped isotopologue distributions. The abundances of the doubly substituted isotopologues of CH_4 ($^{13}CH_3D$ and $^{12}CH_2D_2$) are expressed as a deviation from their concentrations at a stochastic distribution of the rare isotopes, with $\Delta^{13}CH_3D = {}^{13,2}R_{\text{sample}}/{}^{13,2}R_{\text{stochastic}} - 1$ [‰] and $\Delta^{12}CH_2D_2 = {}^{2,2}R_{\text{sample}}/{}^{2,2}R_{\text{stochastic}} - 1$ [‰]. Both $\Delta^{13}CH_3D$ and $\Delta^{12}CH_2D_2$ values depend on the methane formation temperature (6), yet applications of methane clumped isotopes to constrain its formation temperature and mechanism are complicated by source mixing and disequilibrium effects (7, 32). The dependence of $\Delta^{13}CH_3D$ and $\Delta^{12}CH_2D_2$ on ΔG_{net} has not been determined experimentally, and similar to bulk carbon isotopes, our model predicts non-monotonic departure from clumped isotopic equilibrium (fig. S5), unlike previous estimates of these relations (7, 33). As ΔG_{net} becomes negative, both $\Delta^{13}CH_3D$ and $\Delta^{12}CH_2D_2$ values decrease from the expected equilibrium compositions, and $\Delta^{12}CH_2D_2$ becomes anti-clumped (i.e., $< 0‰$) due to expression of the KFFs of the Mcr- and Mtr-catalyzed reactions (fig. S5). After the initial decrease in $\Delta^{13}CH_3D$ and $\Delta^{12}CH_2D_2$, both increase with increasingly negative ΔG_{net}, and $\Delta^{13}CH_3D$ increases to values almost as high as the equilibrium values. This behavior has two implications. First, there is a range of ΔG_{net} ($\approx$–75 to –55 kJ mol^{-1}) over which $\Delta^{13}CH_3D$ values may give the false appearance of proximity to isotopic equilibrium. Second, there is a range of ΔG_{net} ($\approx$–100 to –20 kJ mol^{-1}) over which $\Delta^{13}CH_3D$ and $\Delta^{12}CH_2D_2$ cannot uniquely constrain the energetic state of the cell (e.g., $\Delta^{13}CH_3D$ is $\approx$4 ‰ at both ΔG_{net} of $\approx$–20 and $\approx$–70 kJ mol^{-1}). However, in combination with $^{13}\varepsilon_{CO_2\text{-}CH_4}$ and $^{2}\varepsilon_{CH_4\text{-}H_2O}$ data, the position in the ΔG_{net} landscape and the degree of departure from equilibrium may be uniquely constrained (Fig. 3). The multiple-isotope composition of methane (i.e., bulk carbon and hydrogen, and clumped isotopes) may thus be a useful proxy for ΔG_{net} in natural environments where measurements of H_2, CO_2 and CH_4 concentrations are not easily obtainable. In addition to departure from equilibrium of the Mcr- and Mtr-catalyzed reactions, which may cause anti-clumped $\Delta^{12}CH_2D_2$ compositions, a testable prediction of our model is that the Hmd-catalyzed reaction may also cause anti-clumping. This arises from combinatorial effects (34–36) when Hmd activity is high, and especially during H_2-H_2O hydrogen isotopic disequilibrium (fig. S6).

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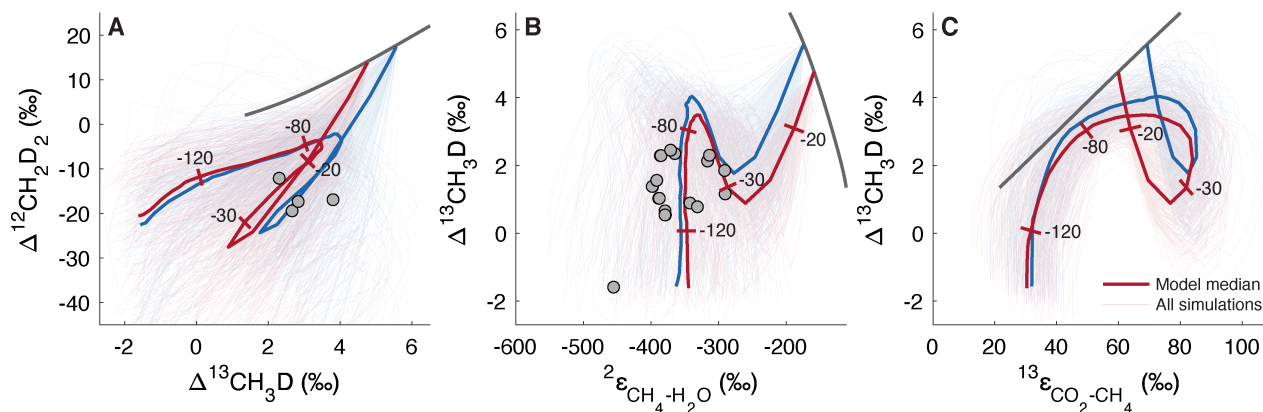
Figure 3: Model-laboratory culture comparison of clumped isotopologue abundances. Experimental data (gray circles) and model results (thin lines) of 200 simulations at 30°C (blue) and 60°C (red). The dark gray lines represent temperature-dependent isotopic equilibrium at 0-350°C, and the thick red and blue lines show the median of the individual simulations, with tickmarks at ΔG_{net} values of –20, –30, –80, and –120 kJ mol^{-1}. (**A**) $\Delta^{13}CH_3D$ against $\Delta^{12}CH_2D_2$. (**B**) $\Delta^{13}CH_3D$ against $^2\varepsilon_{CH_4-H_2O}$. (**C**) $\Delta^{13}CH_3D$ against $^{13}\varepsilon_{CO_2-CH_4}$. Laboratory culture samples are from hydrogenotrophic methanogens that do not have membrane-associated methanophenazines.

Isotopic fractionation in energy-limited environments

The metabolic-isotopic model may be used to examine the controls on the isotopic fractionation of methanogenesis not only in laboratory cultures, but in natural environments as well. Laboratory cultures operate far from equilibrium, usually at H_2 concentrations and cell-specific methanogenesis rates (csMR) much higher than those in natural environments (table S5; fig. S9). Recent reevaluations of slowly forming biogenic methane sources such as marine sediments, coalbeds or shale gas deposits, revealed that apparent CH_4–CO_2 and CH_4–H_2O isotopic equilibrium is common (8, 25, 37–40). There are currently no laboratory cultures that reproduce the isotopic effects associated with these conditions. We modified the model to use enzyme activities that were measured in H_2-limited laboratory cultures (Methods) and assessed the resulting ΔG_{net}–csMR–isotopic relations at 0-60°C.

Existing estimates of environmental ΔG_{net} are more positive than –30 kJ mol^{-1} (table S5). However, the determination of ΔG_{net} in natural environments is often difficult because of low and spatially heterogeneous in-situ H_2 concentrations (41–43), and the actual range of ΔG_{net} likely reflects this heterogeneity. As ΔG_{net} determines the csMR (fig. S9), which is easier to measure,

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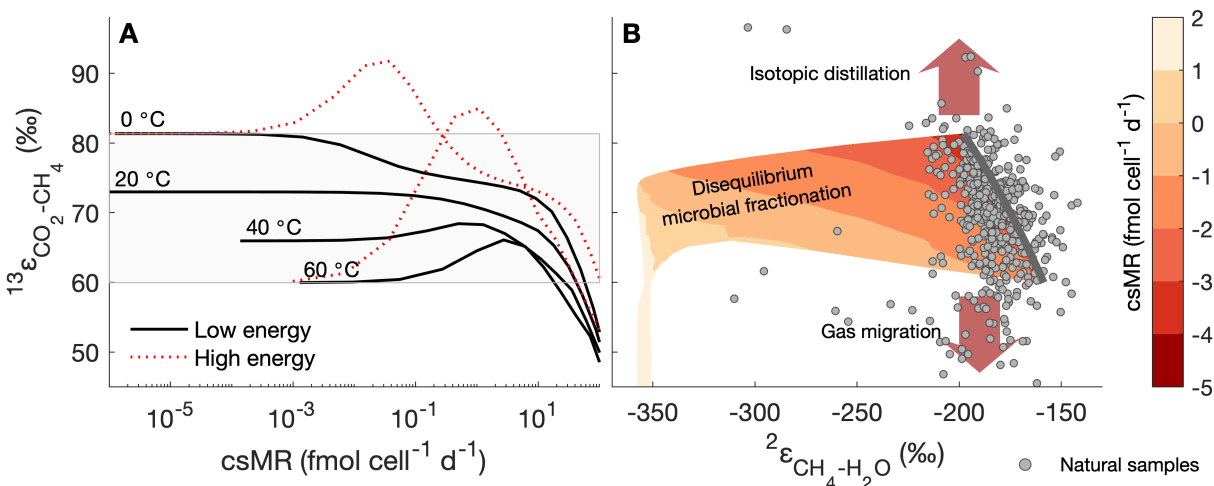
Figure 4: Isotopic fractionation during methanogenesis in energy-limited conditions. (A) $^{13}\varepsilon_{CO_2\text{-}CH_4}$ against cell-specific methanogenesis rate (csMR). Dotted red lines show the laboratory-calibrated model (high energy) results for the same temperatures. (**B**) $^{13}\varepsilon_{CO_2\text{-}CH_4}$ against $^2\varepsilon_{CH_4\text{-}H_2O}$. Contours are $\log_{10}$(csMR) predicted by our model, between 0°C and 60°C. The calculations are for [H_2] between 1 nM and 5 μM, [CO_2] and [CH_4] of 1 mM, and a cell volume of 1 μm^3. The circles are biogenic environmental samples from marine sediments, coalbed methane and natural gas deposits ($n = 491$).

we henceforth discuss csMR–isotopic fractionation relations. Our model predictions of csMR in energy-limited environments (10^{-5} to 10^0 fmol cell^{-1} d^{-1}) are considerably lower than the typical range in laboratory cultures (10^1 to 10^4 fmol cell^{-1} d^{-1}, fig. S9, Refs. 23, 29, 44). This csMR difference of several orders of magnitude reflects the model calibration to reproduce H_2–csMR relations in low-H_2 laboratory culture experiments (fig. S2). Although this calibration was limited to available experimental H_2 concentrations higher than 1 μM, our predicted csMR values at ∼nM H_2 concentrations (10^{-5} to 10^0 fmol cell^{-1} d^{-1}) overlap with the range of csMR values from shallow and deep marine sediments (10^{-4} to 10^0 fmol cell^{-1} d^{-1}), which we calculated from reported bulk methanogenesis rates and cell densities (table S6). We note that there is large uncertainty regarding these environmental csMR values, which stems in part from uncertainty on the net rate of methanogenesis. The methanogenesis rate in natural environments is often determined by radiotracer assays, which may lead to an overestimation of the net methanogenesis rate by orders of magnitude if the reversibility of net methanogenesis (i.e., between CO_2 and CH_4) is higher than 0.9 (SM). An additional source of uncertainty is the number of active methanogenic cells in the sedi-

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216 ments, and there are currently very limited estimates of these values. Despite these uncertainties,

217 our model predictions of csMR match not only environmental csMR estimates, but also estimates

218 of cell-specific power utilization, $P = -\Delta G_{net} \times csMR$ [W cell^{-1}]. We predict P between 10^{-22}

219 and 10^{-18} W cell^{-1} for ΔG_{net} between –10 to –30 kJ mol^{-1} respectively, at 10°C, in agreement

220 with previously estimated cell-specific power utilization in marine sediments (45).

221   The model reveals that in contrast to methanogenesis in lab cultures, under energy-limited

222 conditions the Mtr-catalyzed reaction departs from equilibrium before the Mcr-catalyzed reac-

223 tion (fig. S7). As a consequence, different csMR–fractionation relations emerge, most notably

224 for $^{13}\varepsilon_{CO_2\text{-}CH_4}$ (Fig. 4A). Instead of an increase in $^{13}\varepsilon_{CO_2\text{-}CH_4}$ to larger-than-equilibrium values,

225 $^{13}\varepsilon_{CO_2\text{-}CH_4}$ remains approximately constant up to csMR of ∼10 fmol cell^{-1} d^{-1}. The reason for

226 this apparent carbon isotopic equilibrium is the similar magnitude of the EFF and KFF of Mtr

227 (17‰ and 16‰ at 60°C, respectively, fig. S4). Thus, as the Mtr-catalyzed reaction departs from

228 equilibrium, the net carbon isotopic fractionation changes little. An environmental prediction of

229 the above is that $^{13}\varepsilon_{CO_2\text{-}CH_4}$ that differs measurably from $^{13}\varepsilon_{CO_2\text{-}CH_4}^{eq}$ indicates csMR higher than

230 ∼10 fmol cell^{-1} d^{-1}.

231   Analytically distinguishable departure of $^{2}\varepsilon_{CH_4\text{-}H_2O}$ from equilibrium occurs at csMR between

232 ∼0.001 and 0.1 fmol cell^{-1} d^{-1} (at 0°C and 60°C, respectively; fig. S8B), and departure from

233 clumped isotope equilibrium occurs over a similar range of csMR (fig. S8C, D). Thus, our model

234 reveals that at csMR between 0.1 and 10 fmol cell^{-1} d^{-1}, one might expect near-equilibrium

235 $^{13}\varepsilon_{CO_2\text{-}CH_4}$ concurrent with disequilibrium $^{2}\varepsilon_{CH_4\text{-}H_2O}$. The opposite situation (i.e., near-equilibrium

236 $^{2}\varepsilon_{CH_4\text{-}H_2O}$ and disequilibrium $^{13}\varepsilon_{CO_2\text{-}CH_4}$) is common in natural environments, though csMR is

237 usually unknown (fig. S9). As suggested in previous studies, this apparent hydrogen isotopic equi-

238 librium concurrent with carbon isotopic disequilibrium may be explained by diffusive mixing of

239 CO_2 and CH_4, isotopic (Rayleigh) distillation, or diagenetic isotope exchange without net methane

240 production (Fig. 4B; 8, 38, 46, 47). If disequilibrium $^{13}\varepsilon_{CO_2\text{-}CH_4}$ is explained as above rather

241 than by microbial expression of KFFs, then the scarcity of data within the field representing dis-

242 equilibrium microbial fractionation in Fig. 4B may reflect near-equilibrium isotopic fractionation

during methanogenesis in energy-limited environments. This implies csMR lower than $\sim$0.001–0.1 fmol cell^{-1} d^{-1}(depending on temperature), consistent with almost all csMR estimated from cell abundances and volumetric methanogenesis rates (SM, table S6). In energy-limited environments with higher csMR, near-equilibrium $^2\varepsilon_{CH_4\text{-}H_2O}$ concurrent with disequilibrium $^{13}\varepsilon_{CO_2\text{-}CH_4}$ may be explained by methane cycling or anaerobic oxidation of methane that operates close to the thermodynamic limit of this metabolism (48).

Conclusions

By accounting for the metabolites and reactions in the hydrogenotrophic methanogenesis pathway, we link environmental substrate and product concentrations, pH and temperature to the energetics and net rate of methanogenesis, and to the associated fractionation of carbon, hydrogen and clumped isotopes. The landscape of departure of individual reactions in the pathway from reversibility controls these fractionations, explaining rate–fractionation relations in laboratory cultures. We suggest that a combination of $^{13}\varepsilon_{CO_2\text{-}CH_4}$, $^2\varepsilon_{CH_4\text{-}H_2O}$, $\Delta^{13}CH_3D$ and $\Delta^{12}CH_2D_2$ data, interpreted within a metabolic-isotopic model framework, can uniquely constrain the in-situ energetics of methanogenic activity (i.e., the actual ΔG_{net} of methanogenesis), which may be difficult to constrain by other means. However, application of such an approach is limited to natural and artificial environments in which methane production dominates the processes that may affect apparent CH_4–CO_2 and CH_4–H_2O isotopic fractionations. Such processes include CO_2 and CH_4 diffusion, isotopic distillation and methane cycling, which are prevalent in natural environments. Nevertheless, the mechanistic understanding of isotopic effects during microbial methanogenesis provided by our model allows an inference that this metabolism operates close to chemical and isotopic equilibrium in a wide range of natural environments, including marine sediments, coal beds and natural gas deposits.

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Methods

Relating net isotopic fractionation to thermodynamic drive

The metabolic-isotopic model that we developed is based on a general framework, previously used to investigate sulfur isotopes in microbial sulfate-reduction (17, 18). This approach is applicable to any metabolic network with some modifications that we present below. In general, the net flux of a reversible enzymatically catalyzed reaction (J) is defined by the difference between the forward (J^+) and backward (J^-) gross fluxes, and can be expressed as (16):

$$J = J^+ - J^- = V^+ \times \left(\frac{\prod_j \left([r_j] / K_{M_j} \right)^{n_j}}{1 + \prod_j \left([r_j] / K_{M_j} \right)^{n_j} + \prod_i \left([p_i] / K_{M_i} \right)^{m_i}} \right) \times \left(1 - e^{\Delta G_i' / RT} \right) , \quad (3)$$

where V^+ is the maximal rate capacity of the enzyme (defined by $[E] \times k_{cat}^+$, where k_{cat}^+ is the catalytic rate constant for forward reaction), $[r_j]$ and $[p_i]$ are concentrations of the ith reactant and jth product, respectively, and K_{M_j} and K_{M_i} are the Michaelis constants for the forward and reverse directions, respectively. The n_i and m_j stand for the stoichiometry of the reactant and product, respectively. A net flux, as described by Eq. 3, is calculated for each of the reactions in the hydrogenotrophic methanogenesis pathway (table S1). These fluxes are used to construct a mass balance, which is solved at a steady-state to find the concentrations of the metabolites. The steady-state solution was obtained by forward integration using ode15s, an ordinary differential equation solver in MATLAB®. As metabolite concentrations span 6 to 7 orders of magnitude, the resulting set of differential equations is stiff, and required the use of this solver with an absolute tolerance of 10^{-20} M. We checked that the duration of integration was sufficient to reach a steady-state solution (i.e., no change in concentrations and fluxes) in all cases.

The model's inputs include the environmental temperature, pH and extracellular aqueous concentrations of H_2, CO_2 and CH_4. We assume that the intracellular concentrations of H_2 and CH_4 are equal to the extracellular concentrations due to their rapid diffusion through the membrane. For CO_2 we apply a simple diffusion model to relate intracellular to extracellular CO_2 concentrations. The model's tunable parameters include enzyme kinetic constants (K_M and V^+), thermodynamic constants ($\Delta G_i'^0$) and cellular parameters such as cell size, and concentrations of some of

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291   the metabolites. We elaborate on the parameters and the choice of their values in the SM. With

292   the inputs prescribed and values for the tunable parameters chosen, the model outputs are the con-

293   centrations of all intracellular metabolites and the gross fluxes among these metabolites, which are

294   related to the reactions in the methanogenesis pathway.

295   We used the forward and backward gross fluxes from the metabolic model to calculate the

296   net isotopic fractionations in hydrogenotrophic methanogenesis. To this end, we constructed a

297   mass balance for each isotopic system (for both bulk and clumped isotopes), which is based on

298   calculating an isotopic flux associated with each of the chemical fluxes. For the schematic reaction

299   $r \to p$, this isotopic flux can be approximated:

$$F_{rp} = J_{rp}^{+} R_r \alpha_{rp,r} \ , \tag{4}$$

300   where J_{rp}^{+} is the forward chemical flux between metabolites pools r and p, R_r is the abundance

301   ratio of heavy to light isotopes in pool r, and $\alpha_{rp,r}$ is the KFF of the reaction. A term similar to

302   Eq. 4 was assigned to account for the consumption and production of each isotopologue in the

303   pathway while considering the stoichiometry and the symmetry coefficients, where relevant, in a

304   similar manner to previous model derivations (7, 9). As described above, these terms were used

305   to construct isotopic mass balances in the form of a set of coupled differential equations, which

306   were solved numerically using the ode15s solver in MATLAB$^{\circledR}$ to obtain the steady-state bulk and

307   clumped isotopic composition of all intracellular metabolites, given bulk extracellular D/H and

308   $^{13}C/^{12}C$ of H_2O and CO_2, respectively.

309   The net isotopic fractionation in linear metabolic networks such as the carbon reaction network

310   in hydrogenotrophic methanogenesis can be solved analytically. For example, in the case of the

311   reaction $r \rightleftarrows p$, the net isotopic fractionation between the metabolite pools of r and p ($\alpha_{r,p}^{net}$) can be

312   calculated by:

$$\alpha_{r,p}^{net} = \left(\alpha_{r,p}^{eq} - \alpha_{r,rp} \right) \times (J_{pr}^{-}/J_{rp}^{+}) + \alpha_{r,rp} \ , \tag{5}$$

313   where $\alpha_{r,p}^{eq}$ is the EFF between pools r and p, $\alpha_{r,rp}$ is 1/KFF of the forward reaction. A full deriva-

314   tion for this term is presented in Wing and Halevy, 2014 (18). The reversibility J_{pr}^{-}/J_{rp}^{+} is directly

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related to $\Delta G'_i$ through the flux-force relationship (Eq. 2). In the case of longer linear reaction networks such as $s \rightleftarrows r \rightleftarrows p$, the net isotopic fractionation can be described by:

$$\alpha^{net}_{s,p} = \left(\alpha^{net}_{r,p} \times \alpha^{eq}_{s,r} - \alpha_{s,sr}\right) \times (J^-_{rs}/J^+_{sr}) + \alpha_{s,sr} . \tag{6}$$

This expression can be expanded recursively to describe the net isotopic fractionation in a linear reaction network of arbitrary length (18). Such a recursive expression has the advantage of a significantly reduced computational time compared to the numerical solution that is used for nonlinear reaction networks. We verified that this analytical solution yields identical results to the numerical solution for the net carbon isotopic fractionation.

Isotopic model parameters

To calculate the net isotopic fractionations associated with each of the steps in the methanogenesis pathway we used the temperature-dependent EFFs which were calculated for the hydrogenotrophic pathway (25, 49). The KFFs were determined experimentally only for Mcr (24). To account for missing KFFs we randomly sampled their values from prior uniform distributions 10^6 times. We included both primary and secondary isotopic effects (α_p and α_s, respectively). Primary isotopic effects are due to the breaking or formation of a bond directly with the atom of interest, and secondary isotopic effects are due to breaking or formation of an adjacent bond. Secondary KFFs are usually smaller than primary KFFs. We weighted the combinations of KFFs drawn from the prior distributions by the model-experimental mismatch, expressed as the inverse of the square of sum of the square errors ($1/SSE^2$), to generate posterior distributions of the KFFs, which were then used in the model. The selection of prior KFF value distributions is described in the SM.

The posterior distributions of the KFFs (fig. S4) serve as a sensitivity analysis of the model. Posterior distributions that are similar to the prior distributions indicate insensitivity to the model parameter, in this case, the KFF value. For carbon isotopes, the model shows sensitivity to the KFFs of Fmd, Ftr and Mtr with smallest-mismatch values of 30.8‰, 25.8‰, and 15.8‰ at mesophilic conditions, respectively. This sensitivity is in line with our observations of the reactions that determine the trajectory of $^{13}\varepsilon_{CO_2\text{-}CH_4}$ departure from equilibrium. For hydrogen isotopes, the model

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shows sensitivity to most of the primary KFFs, and to the secondary KFF of Mtr. Notably, the primary KFFs of Mtd, Mer and Frh tend to be small (i.e., closer to unity), with median values between $\approx$–450‰ and $\approx$–300‰, but larger than the secondary KFFs with median values between $\approx$–330‰ and $\approx$–220‰.

Simulating methanogenesis in energy-limited conditions

The metabolic parameters that we used to explore isotopic fractionation in methanogenesis were curated from cell cultures grown under optimal conditions. However, methanogens tightly regulate gene expression under energy-limiting conditions, resulting in shifts in enzyme specific activities (30, 50–57). Specifically, prolonged H_2 limitation promotes higher activities of Frh, Mvh/Hdr and the Mcr I isoenzyme, in parallel to a decrease in the activity of Mcr II (52, 53, 58). Mcr I has a higher affinity to the substrates CH_3-SCoM and HS-CoB relative to Mcr II, allowing the cells to increase the cell-specific respiration rate when H_2 becomes limiting (59–61). Therefore, in the simulations of methanogenesis in natural, energy-limited environments we used enzyme activities measured in crude extracts from cells that were grown under conditions favoring Mcr I activity (53) with K_M values of Mcr I (table S2). In laboratory cultures, cells under energy limitation decrease their volume in a matter of days or weeks, and a similar phenomenon was observed when comparing cell communities in energy-replete *vs.* energy-deprived environmental settings (62). Thus, we use a cellular volume smaller by a factor of two from the default value (2 fL cell^{-1} *vs.* 1 fL cell^{-1}).

Acknowledgments

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Author contributions

J.G. and I.H. conceived the study, developed and analyzed the metabolic-isotopic model, and wrote the initial draft. Q.J. developed the metabolic model. All authors contributed to the writing.

References

[1] M. Saunois, *et al.*, The global methane budget 2000–2012. *Earth Syst. Sci. Data* **8**, 697–751 (2016).

[2] S. Schwietzke, O. A. Sherwood, L. M. P. Bruhwiler, J. B. Miller, G. Etiope, E. J. Dlugokencky, S. E. Michel, V. A. Arling, B. H. Vaughn, J. W. C. White, P. P. Tans, Upward revision of global fossil fuel methane emissions based on isotope database. *Nature* **538**, 88–91 (2016).

[3] Z. Lyu, N. Shao, T. Akinyemi, W. B. Whitman, Methanogenesis. *Current Biology* **28**, R727-R732 (2018).

[4] M. J. Whiticar, Carbon and hydrogen isotope systematics of bacterial formation and oxidation of methane. *Chem. Geol.* **161**, 291–314 (1999).

[5] E. G. Nisbet, E. J. Dlugokencky, M. R. Manning, D. Lowry, R. E. Fisher, J. L. France, S. E. Michel, J. B. Miller, J. W. C. White, B. Vaughn, P. Bousquet, J. A. Pyle, N. J. Warwick, M. Cain, R. Brownlow, G. Zazzeri, M. Lanoisellé, A. C. Manning, E. Gloor, D. E. J. Worthy, E.-G. Brunke, C. Labuschagne, E. W. Wolff, A. L. Ganesan, Rising atmospheric methane: 2007-2014 growth and isotopic shift. *Glob. Biogeochem. Cycles* **30**, 1356–1370 (2016).

[6] D. A. Stolper, A. Sessions, A. Ferreira, E. Santos Neto, A. Schimmelmann, S. Shusta, D. Valentine, J. Eiler, Combined 13C–D and D–D clumping in methane: Methods and preliminary results. *Geochim. Cosmochim. Acta* **126**, 169–191 (2014).

[7] D. T. Wang, D. S. Gruen, B. S. Lollar, K.-U. Hinrichs, L. C. Stewart, J. F. Holden, A. N. Hristov, J. W. Pohlman, P. L. Morrill, M. Könneke, K. B. Delwiche, E. P. Reeves, C. N. Sutcliffe,

D. J. Ritter, J. S. Seewald, J. C. McIntosh, H. F. Hemond, M. D. Kubo, D. Cardace, T. M. Hoehler, S. Ono, Nonequilibrium clumped isotope signals in microbial methane. *Science* **348**, 428–431 (2015).

[8] T. Okumura, S. Kawagucci, Y. Saito, Y. Matsui, K. Takai, H. Imachi, Hydrogen and carbon isotope systematics in hydrogenotrophic methanogenesis under H2-limited and H2-enriched conditions: Implications for the origin of methane and its isotopic diagnosis. *Prog. Earth Planet. Sci.* **3**, 14 (2016).

[9] X. Cao, H. Bao, Y. Peng, A kinetic model for isotopologue signatures of methane generated by biotic and abiotic CO2 methanation. *Geochim. Cosmochim. Acta* **249**, 59–75 (2019).

[10] T. Wagner, U. Ermler, S. Shima, The methanogenic CO2 reducing-and-fixing enzyme is bi-functional and contains 46 [4Fe-4S] clusters. *Science* **354**, 114–117 (2016).

[11] A.-K. Kaster, J. Moll, K. Parey, R. K. Thauer, Coupling of ferredoxin and heterodisulfide reduction via electron bifurcation in hydrogenotrophic methanogenic archaea. *PNAS* **108**, 2981–6 (2011).

[12] L. M. I. de Poorter, W. J. Geerts, J. T. Keltjens, Hydrogen concentrations in methane-forming cells probed by the ratios of reduced and oxidized coenzyme F420. *Microbiology* **151**, 1697–1705 (2005).

[13] L. M. I. de Poorter, W. G. Geerts, A. P. R. Theuvenet, J. T. Keltjens, Bioenergetics of the formyl-methanofuran dehydrogenase and heterodisulfide reductase reactions in Methanothermobacter thermautotrophicus. *Eur. J. Biochem.* **270**, 66–75 (2002).

[14] J. Bassham, G. Krause, Free energy changes and metabolic regulation in steady-state photosynthetic carbon reduction. *Biochim. Biophys. Acta BBA - Bioenerg.* **189**, 207–221 (1969).

[15] D. A. Beard, H. Qian, Relationship between thermodynamic driving force and one-way fluxes in reversible processes. *PloS One* **2**, e144 (2007).

19

[16] E. Noor, A. Flamholz, W. Liebermeister, A. Bar-Even, R. Milo, A note on the kinetics of enzyme action: A decomposition that highlights thermodynamic effects. *FEBS Lett.* **587**, 2772–7 (2013).

[17] C. Rees, A steady-state model for sulphur isotope fractionation in bacterial reduction processes. *Geochim. Cosmochim. Acta* **37**, 1141–1162 (1973).

[18] B. A. Wing, I. Halevy, Intracellular metabolite levels shape sulfur isotope fractionation during microbial sulfate respiration. *Proc. Natl. Acad. Sci. U. S. A.* **111**, 18116–25 (2014).

[19] D. L. Valentine, A. Chidthaisong, A. Rice, W. S. Reeburgh, S. C. Tyler, Carbon and hydrogen isotope fractionation by moderately thermophilic methanogens. *Geochim. Cosmochim. Acta* **68**, 1571–1590 (2004).

[20] H. Penning, C. M. Plugge, P. E. Galand, R. Conrad, Variation of carbon isotope fractionation in hydrogenotrophic methanogenic microbial cultures and environmental samples at different energy status. *Glob. Change Biol.* **11**, 2103–2113 (2005).

[21] K. Takai, K. Nakamura, T. Toki, U. Tsunogai, M. Miyazaki, J. Miyazaki, H. Hirayama, S. Nakagawa, T. Nunoura, K. Horikoshi, Cell proliferation at 122 degrees C and isotopically heavy CH4 production by a hyperthermophilic methanogen under high-pressure cultivation. *Proc. Natl. Acad. Sci. U. S. A.* **105**, 10949–54 (2008).

[22] S. Hattori, H. Nashimoto, H. Kimura, K. Koba, K. Yamada, M. Shimizu, H. Watanabe, M. Yoh, N. Yoshida, Hydrogen and carbon isotope fractionation by thermophilic hydrogenotrophic methanogens from a deep aquifer under coculture with fermenters. *Geochem. J.* **46**, 193–200 (2012).

[23] B. D. Topçuoğlu, C. Meydan, T. B. Nguyen, S. Q. Lang, J. F. Holden, Growth Kinetics, Carbon Isotope Fractionation, and Gene Expression in the Hyperthermophile Methanocaldococcus jannaschii during Hydrogen-Limited Growth and Interspecies Hydrogen Transfer. *Appl. Environ. Microbiol.* **85**, 1–14 (2019).

[24] S. Scheller, M. Goenrich, R. K. Thauer, B. Jaun, Methyl-coenzyme M reductase from methanogenic archaea: Isotope effects on the formation and anaerobic oxidation of methane. *J. Am. Chem. Soc.* **135**, 14975–84 (2013).

[25] J. Gropp, M. A. Iron, I. Halevy, Theoretical estimates of equilibrium carbon and hydrogen isotope effects in microbial methane production and anaerobic oxidation of methane. *Geochimica et Cosmochimica Acta* **295**, 237–264 (2021).

[26] H. M. Miller, N. Chaudhry, M. E. Conrad, M. Bill, S. H. Kopf, A. S. Templeton, Large carbon isotope variability during methanogenesis under alkaline conditions. *Geochim. Cosmochim. Acta* **237**, 18–31 (2018).

[27] D. B. Nothaft, A. S. Templeton, J. H. Rhim, D. T. Wang, J. Labidi, H. M. Miller, E. S. Boyd, J. M. Matter, S. Ono, E. D. Young, S. H. Kopf, P. B. Kelemen, M. E. Conrad, Geochemical, biological and clumped isotopologue evidence for substantial microbial methane production under carbon limitation in serpentinites of the Samail Ophiolite, Oman. *J. Geophys. Res. Biogeosciences* **NA**, e2020JG006025 (2021).

[28] H. Yoshioka, S. Sakata, Y. Kamagata, Hydrogen isotope fractionation by Methanothermobacter thermoautotrophicus in coculture and pure culture conditions. *Geochim. Cosmochim. Acta* **72**, 2687–2694 (2008).

[29] S. Kawagucci, M. Kobayashi, S. Hattori, K. Yamada, Y. Ueno, K. Takai, N. Yoshida, Hydrogen isotope systematics among H2–H2O–CH4 during the growth of the hydrogenotrophic methanogen Methanothermobacter thermautotrophicus strain ΔH. *Geochim. Cosmochim. Acta* **142**, 601–614 (2014).

[30] C. Afting, E. Kremmer, C. Brucker, A. Hochheimer, R. K. Thauer, Regulation of the synthesis of H2-forming methylenetetrahydromethanopterin dehydrogenase (Hmd) and of HmdII and HmdIII in Methanothermobacter marburgensis. *Arch. Microbiol.* **174**, 225–232 (2000).

[31] R. Botz, H. D. Pokojski, M. Schmitt, M. Thomm, Carbon isotope fractionation during bacterial methanogenesis by CO2 reduction. *Org. Geochem.* **25**, 255–262 (1996).

[32] P. M. Douglas, D. A. Stolper, J. M. Eiler, A. L. Sessions, M. Lawson, Y. Shuai, A. Bishop, O. G. Podlaha, A. A. Ferreira, E. V. Santos Neto, M. Niemann, A. S. Steen, L. Huang, L. Chimiak, D. L. Valentine, J. Fiebig, A. J. Luhmann, W. E. Seyfried, G. Etiope, M. Schoell, W. P. Inskeep, J. J. Moran, N. Kitchen, Methane clumped isotopes: Progress and potential for a new isotopic tracer. *Org. Geochem.* **113**, 262–282 (2017).

[33] D. A. Stolper, A. Martini, M. Clog, P. Douglas, S. Shusta, D. Valentine, A. Sessions, J. Eiler, Distinguishing and understanding thermogenic and biogenic sources of methane using multiply substituted isotopologues. *Geochim. Cosmochim. Acta* **161**, 219–247 (2015).

[34] L. Y. Yeung, Combinatorial effects on clumped isotopes and their significance in biogeochemistry. *Geochim. Cosmochim. Acta* **172**, 22–38 (2016).

[35] T. Röckmann, M. E. Popa, M. C. Krol, M. E. G. Hofmann, Statistical clumped isotope signatures. *Sci. Rep.* **6**, 31947 (2016).

[36] L. Taenzer, J. Labidi, A. L. Masterson, X. Feng, D. Rumble, E. D. Young, W. D. Leavitt, Low $\Delta 12CH_2D_2$ values in microbialgenic methane result from combinatorial isotope effects. *Geochimica et Cosmochimica Acta* **285**, 225–236 (2020).

[37] D. S. Vinson, N. E. Blair, A. M. Martini, S. Larter, W. H. Orem, J. C. McIntosh, Microbial methane from in situ biodegradation of coal and shale: A review and reevaluation of hydrogen and carbon isotope signatures. *Chem. Geol.* **453**, 128–145 (2017).

[38] A. C. Turner, R. Korol, D. L. Eldridge, M. Bill, M. E. Conrad, T. F. Miller, D. A. Stolper, Experimental and theoretical determinations of hydrogen isotopic equilibrium in the system CH4-H2-H2O from 3 to 200°C. *Geochimica et Cosmochimica Acta* **NA**, NA (2021).

22

[39] J. J. Jautzy, P. M. J. Douglas, H. Xie, J. M. Eiler, I. D. Clark, CH4 isotopic ordering records ultra-slow hydrocarbon biodegradation in the deep subsurface. *Earth and Planetary Science Letters* **562**, 116841 (2021).

[40] N. Zhang, G. T. Snyder, M. Lin, M. Nakagawa, A. Gilbert, N. Yoshida, R. Matsumoto, Y. Sekine, Doubly substituted isotopologues of methane hydrate (13CH3D and 12CH2D2): Implications for methane clumped isotope effects, source apportionments and global hydrate reservoirs. *Geochimica et Cosmochimica Acta* **NA**, NA (2021).

[41] R. Conrad, T. J. Phelps, J. G. Zeikus, Gas metabolism evidence in support of the juxtaposition of hydrogen-producing and methanogenic bacteria in sewage sludge and lake sediments. *Appl. Environ. Microbiol.* **50**, 595–601 (1985).

[42] E. Giraldo-Gomez, S. Goodwin, M. S. Switzenbaum, Influence of mass transfer limitations on determination of the half saturation constant for hydrogen uptake in a mixed-culture CH4-producing enrichment. *Biotechnol. Bioeng.* **40**, 768–76 (1992).

[43] Y.-S. Lin, V. B. Heuer, T. Goldhammer, M. Y. Kellermann, M. Zabel, K.-U. Hinrichs, Towards constraining H2 concentration in subseafloor sediment: A proposal for combined analysis by two distinct approaches. *Geochimica et Cosmochimica Acta* **77**, 186–201 (2012).

[44] A. M. Zyakun, Potential of 13C/12C Variations in Bacterial Methane in Assessing Origin of Environmental Methane. *Hydrocarbon Migration And Its Near-Surface Expression*, D. Schumacher, M. A. Abrams, eds. (AAPG Special Volumes, 1996), pp. 341–352.

[45] J. A. Bradley, S. Arndt, J. P. Amend, E. Burwicz, A. W. Dale, M. Egger, D. E. LaRowe, Widespread energy limitation to life in global subseafloor sediments. *Sci. Adv.* **6**, eaba0697 (2020).

[46] C. K. Paull, T. D. Lorenson, W. S. Borowski, W. Ussler III, K. Olsen, N. M. Rodriguez, Isotopic Composition of CH4, CO2 Species, and Sedimentary Organic Matter Within Samples

504 From the Blake Ridge: Gas Source Implications. *Proc. Ocean Drill. Program Sci. Results*
505 **164**, 67–78 (2000).

506 [47] J. W. Pohlman, M. Kaneko, V. B. Heuer, R. B. Coffin, M. Whiticar, Methane sources and
507 production in the northern Cascadia margin gas hydrate system. *Earth and Planetary Science*
508 *Letters* **287**, 504–512 (2009).

509 [48] G. Wegener, J. Gropp, H. Taubner, I. Halevy, M. Elvert, Sulfate-dependent reversibility of
510 intracellular reactions explains the opposing isotope effects in the anaerobic oxidation of
511 methane. *Sci. Adv.* **7**, eabe4939 (2021).

512 [49] J. Gropp, M. A. Iron, I. Halevy, Corrigendum to "Theoretical estimates of equilibrium carbon
513 and hydrogen isotope effects in microbial methane production and anaerobic oxidation of
514 methane" [Geochim. Cosmochim. Acta 295 (2021) 237–264]. *Geochimica et Cosmochimica*
515 *Acta* **306**, 386–389 (2021).

516 [50] R. M. Morgan, T. D. Pihl, J. Nölling, J. N. Reeve, Hydrogen regulation of growth, growth
517 yields, and methane gene transcription in Methanobacterium thermoautotrophicum deltaH. *J.*
518 *Bacteriol.* **179**, 889–98 (1997).

519 [51] J. N. Reeve, J. Nölling, R. M. Morgan, D. R. Smith, Methanogenesis: Genes, genomes, and
520 who's on first? *J. Bacteriol.* **179**, 5975–86 (1997).

521 [52] P. Vermeij, J. L. Pennings, S. M. Maassen, J. T. Keltjens, G. D. Vogels, Cellular levels of factor
522 390 and methanogenic enzymes during growth of Methanobacterium thermoautotrophicum
523 deltaH. *J. Bacteriol.* **179**, 6640–8 (1997).

524 [53] J. L. Pennings, P. Vermeij, L. M. de Poorter, J. T. Keltjens, G. D. Vogels, Adaptation of
525 methane formation and enzyme contents during growth of Methanobacterium thermoau-
526 totrophicum (strain ΔH) in a fed-batch fermentor. *Antonie Van Leeuwenhoek* **77**, 281–291
527 (2000).

[54] E. L. Hendrickson, A. K. Haydock, B. C. Moore, W. B. Whitman, J. A. Leigh, Functionally distinct genes regulated by hydrogen limitation and growth rate in methanogenic Archaea. *Proc. Natl. Acad. Sci. U. S. A.* **104**, 8930–4 (2007).

[55] S. Kato, T. Kosaka, K. Watanabe, Comparative transcriptome analysis of responses of Methanothermobacter thermautotrophicus to different environmental stimuli. *Environ. Microbiol.* **10**, 893–905 (2008).

[56] Q. Xia, T. Wang, E. L. Hendrickson, T. J. Lie, M. Hackett, J. A. Leigh, Quantitative proteomics of nutrient limitation in the hydrogenotrophic methanogen Methanococcus maripaludis. *BMC Microbiol.* **9**, 149 (2009).

[57] M. Enoki, N. Shinzato, H. Sato, K. Nakamura, Y. Kamagata, Comparative Proteomic Analysis of Methanothermobacter themautotrophicus ΔH in Pure Culture and in Co-Culture with a Butyrate-Oxidizing Bacterium. *PLoS ONE* **6**, e24309 (2011).

[58] T. D. Pihl, S. Sharma, J. N. Reeve, Growth phase-dependent transcription of the genes that encode the two methyl coenzyme M reductase isoenzymes and N5-methyltetrahydromethanopterin:coenzyme M methyltransferase in Methanobacterium thermoautotrophicum delta H. *J. Bacteriol.* **176**, 6384–6391 (1994).

[59] L. G. Bonacker, S. Baudner, R. K. Thauer, Differential expression of the two methyl-coenzyme M reductases in Methanobacterium thermoautotrophicum as determined immunochemically via isoenzyme-specific antisera. *Eur. J. Biochem.* **206**, 87–92 (1992).

[60] L. G. Bonacker, S. Baudner, E. Mörschel, R. Böcher, R. K. Thauer, Properties of the two isoenzymes of methyl-coenzyme M reductase in Methanobacterium thermoautotrophicum. *Eur. J. Biochem.* **217**, 587–95 (1993).

[61] M. Dey, X. Li, R. C. Kunz, S. W. Ragsdale, Detection of Organometallic and Radical Intermediates in the Catalytic Mechanism of Methyl-Coenzyme M Reductase Using the Natural

552   Substrate Methyl-Coenzyme M and a Coenzyme B Substrate Analogue. *Biochemistry* **49**,

553   10902–10911 (2010).

554 [62] M. A. Lever, K. L. Rogers, K. G. Lloyd, J. Overmann, B. Schink, R. K. Thauer, T. M. Hoehler,

555   B. B. Jørgensen, Life under extreme energy limitation: A synthesis of laboratory- and field-

556   based investigations. *FEMS Microbiol Rev* **39**, 688–728 (2015).