Metabolic imprints in the hydrogen isotopes of *Archaeoglobus fulgidus* tetraether lipids

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Abstract

The stable hydrogen isotope composition of archaeal lipids is emerging as a potential paleoenvironmental proxy, adding to the well-established application of plant leaf wax-derived *n*-alkanes in paleohydrological reconstruction. A handful of studies reported relatively invariant and depleted hydrogen isotope compositions for archaeal lipids despite the range of different organisms and growth conditions explored. However, how modes of metabolism and physiological state (growth phase) affect the hydrogen isotope signatures of archaeal lipids remains poorly understood, limiting our ability to interpret archaeal lipid biomarker records from the environment. Here we conducted water isotope label experiments with a metabolically flexible and well-studied model archaeon *Archaeoglobus fulgidus* and quantified the hydrogen isotope fractionation between lipids and water in response to different carbon substrates and electron donor-acceptor pairs. The $^2H/^1H$ fractionation between lipids and water ($\varepsilon_{L/W}$) was overall negative, ranging from –280 to –226 ‰, and overlapped with the ranges observed for other archaea in previous studies. Isotope flux-balance model results suggest that ≥80 % and ≥50 % of lipid-bound H in *A. fulgidus* cultures directly reflect water isotope compositions (i.e., not via organic substrate or H_2) during autotrophy and heterotrophy, respectively. The model results also suggest the final saturation during isoprenoid lipid biosynthesis catalyzed by a flavoenzyme geranylgeranyl reductase likely contributes to the large $^2H/^1H$ fractionation observed in this study, consistent with previous isotope flux-balance model results for a different archaeon. Finally, we synthesized available data to compare $\varepsilon_{L/W}$ patterns across all three domains of life: Bacteria, Archaea and Eukarya. Emerging patterns between archaeal and eukaryotic lipids are consistent with the notion of highly

42 fractionating geranylgeranyl reductase, and the patterns between archaeal and bacterial lipids
43 suggest that the general state of energy limitation may also contribute to large, negative values of
44 $\varepsilon_{L/W}$ observed in prokaryotic lipids. Altogether, these findings lend further support for the potential
45 of archaeal lipid $\varepsilon_{L/W}$ as a paleohydrological proxy and provide a broader insight into the $^2H/^1H$
46 fractionation mechanisms potentially shared among prokaryotic and eukaryotic lipid biomarkers.

1.   Introduction

Hydrogen (H) plays a crucial role in most metabolic processes and is incorporated into biological

molecules, often with distinct isotope fractionation patterns. Consequently, the ratio of stable

hydrogen isotopes ($^2H/^1H$) of certain biomolecules can be used to track the H isotope compositions

of source water, or the operation of specific biosynthetic pathways (Leavitt et al., 2023). Lipid-

bound H in hydrocarbons may resist isotope exchange over geologic time scales for up to 10^4 to

10^8 years (Sessions et al., 2004). Thus, sedimentary records of lipid δ^2H can be applied to

reconstruct past ecological or hydrological states (Sachse et al., 2012; Sessions, 2016; McFarlin et

al., 2019). The δ^2H of lipids produced by photoautotrophic eukaryotes as well as chemoautotrophic

and heterotrophic bacteria have been studied in detail for this reason (e.g., Sessions et al., 1999;

Chikaraishi and Naraoka, 2003; Zhang and Sachs, 2007; Zhang et al., 2009; Sachse et al., 2012;

Sachs, 2014; Dawson et al., 2015; Osburn et al., 2016; Leavitt et al., 2016; Wijker et al., 2019;

McFarlin et al., 2019). In contrast, there are only a handful of studies on the H isotope compositions

of archaeal lipids (Kaneko et al., 2011; Dirghangi and Pagani, 2013; Wu et al., 2020; Lengger et

al., 2021; Leavitt, Kopf et al., 2023). As a consequence, the factors controlling the lipid/water

fractionation ($\varepsilon_{L/W}$), as well as the extent and limits of proxy applications remain poorly understood

for archaeal lipids.

The $\varepsilon_{L/W}$ capture a combination of source water δ^2H and the kinetic isotope effect (KIE) of enzymes

involved in lipid biosynthesis. For lipids produced by photoautotrophic eukaryotes (e.g., plants

and algae), KIE is relatively invariant but can be altered by some environmental factors (e.g.,

temperature, salinity, irradiance) and some factors intrinsic to an organism (e.g. growth rate in

response to different stimuli and underlying genetics) (Sachs, 2014; Van Der Meer et al., 2015;

Sachs and Kawka, 2015; Sachs et al., 2016, 2017; Maloney et al., 2016; Bender, 2017; Ladd and

Sachs, 2017; Wolfshorndl et al., 2019). In general, these are well studied in eukaryotes, and have

been accounted for in hydroclimate proxy studies (Sachse et al., 2012; Sessions, 2016; McFarlin

et al., 2019).

Bacterial lipids show a significantly wider range in $\varepsilon_{L/W}$ than eukaryotic lipids. Depending on the

central metabolism a given bacteria operates, $\varepsilon_{L/W}$ can range from –400 to –200 ‰ (exclusively

negative values) for (an)aerobic chemoautotrophs to –154 to +380 ‰, associated with aerobic

heterotrophs (Sessions et al., 2002; Campbell et al., 2009; Zhang et al., 2009; Dawson et al., 2015;

Osburn et al., 2016; Leavitt et al., 2016; Wijker et al., 2019). For lipids produced by aerobic

heterotrophs, H fluxes through $NADP^+$-reducing and NADPH-balancing reactions provide a

quantitative explanation for the full range of $\varepsilon_{L/W}$ values observed (Wijker et al., 2019). Bacterial

lipids found in sediments, especially those produced by aerobic heterotrophs, may have more

useful applications as proxies for microbial metabolism and ecology rather than past hydroclimate

(Zhang et al., 2009; Wijker et al., 2019).

Less is known about what information archaeal $\varepsilon_{L/W}$ record. Studies to-date indicate that archaeal

lipids are generally depleted in ^2H, with $\varepsilon_{L/W}$ values similar to those observed in eukaryotic

photoautotrophs and bacterial chemoautotrophs (Kaneko et al., 2011; Dirghangi and Pagani, 2013;

Wu et al., 2020; Lengger et al., 2021; Leavitt, Kopf et al., 2023). In one recent study, continuous

cultures of the obligate autotroph *Nitrosopumilus maritimus* revealed that $\varepsilon_{L/W}$ remains nearly

constant across a 3-fold range in growth rate (Leavitt, Kopf et al., 2023). Other studies have

reported similarly large, negative values of $\varepsilon_{L/W}$ for archaea grown under heterotrophic conditions

(Kaneko et al., 2011; Dirghangi and Pagani, 2013; Wu et al., 2020; Lengger et al., 2021). However,

it remains to be determined whether and how carbon metabolism affects $\varepsilon_{L/W}$ within a single

organism. Testing this is important, considering that mixotrophy (using both autotrophic and

heterotrophic carbon sources) is a widespread mode of metabolism in nature and that many archaea

are adapted to fluctuating environments. For bacterial lipids, changes in carbon metabolism

resulted in significantly different $\varepsilon_{L/W}$ values within a single organism for some species (e.g.,

Cupriavidus oxalaticus and *Paracoccus denitrificans*; Zhang et al., 2009; Osburn et al., 2016), but

not in other species (sulfate-reducing bacteria; Dawson et al., 2015; Osburn et al., 2016; Leavitt et

al., 2016). As such, understanding the role of carbon metabolism in setting archaeal lipid H isotope

compositions is important for interpreting archaeal lipid records from the environment.

Furthermore, our current understanding of broader patterns of $\varepsilon_{L/W}$ that emerge from comparison

across lipids produced by all three domains of life—Archaea, Bacteria, and Eukarya—is limited,

despite the similarities observed among certain subgroups of different domains.

In this study, we conducted pure culture experiments to better understand factors influencing $\varepsilon_{L/W}$

in archaeal lipids. To test the effect of metabolism on $\varepsilon_{L/W}$ within a single organism, we cultivate

a metabolically flexible archaeon *Archaeoglobus fulgidus* in batch cultures on different carbon

substrates and electron donor-acceptor pairs (lactate-thiosulfate, lactate-sulfate, and H_2/CO_2-

thiosulfate). We also probe the effect of growth phase by comparing biomass harvested at

exponential and stationary growth phases and track the fraction of water incorporated into the final

product lipid by using different water isotope compositions of growth media. We then interpret

the regression parameters derived from water isotope label experiments with an isotope flux-

balance model informed by the enzymatic steps involved in the Wood-Ljungdahl (reductive acetyl-

CoA) pathway used by *A. fulgidus* as well as archaeal isoprenoid lipid biosynthesis pathways used

by all tetraether-producing archaea. In addition to *A. fulgidus*, we examine two thermoacidophilic

archaea (*Acidianus* sp. DS80 and *Metallosphaera sedula* DSM 5348T) for additional $\varepsilon_{L/W}$

measurements. We synthesize our new findings in the domain Archaea along with other available

archaeal lipid $\varepsilon_{L/W}$ data as well as representative $\varepsilon_{L/W}$ values from domains Bacteria and Eukarya,

and highlight factors that underlie patterns in $\varepsilon_{L/W}$ across the three domains of life.

2.   Materials and Methods

2.1.   Culture Strains and Growth

Archaeoglobus fulgidus str. VC-16 (DSM4304) was obtained from Deutsche Sammlung von

Mikroorganismen und Zellkulturen (Braunschweig, Germany) and grown in batch cultures on two

different carbon substrates and electron acceptors (Table 1). Culture media was prepared following

the DSMZ 399 recipe with modifications as follows (per liter): 0.34 g KCl, 4.00 g $MgCl_2 \cdot 6 \ H_2O$,

3.45 g $MgSO_4 \cdot 7 \ H_2O$, 0.25 g NH_4Cl, 0.14 g $CaCl_2 \cdot 2 \ H_2O$, 0.11 g $K_2HPO_4 \cdot 3 \ H_2O$, 0.20 g KH_2PO_4,

18.00 g NaCl, 2.52 g $NaHCO_3$, 2 mL $Fe(NH_4)_2(SO_4)_2 \cdot 7 \ H_2O$ (0.1% w/v), and 10 mL trace elements

(DSMZ 141). After autoclaving, sterile anoxic solutions were added to the medium to a final

concentration of 3.2 mM L-cysteine-HCl and 1% (v/v) vitamin solution (DSMZ 141). When

thiosulfate was used as an electron acceptor, 4.00 g/L $MgCl_2 \cdot 6 \ H_2O$ and 3.45 g/L $MgSO_4 \cdot 7 \ H_2O$

were replaced with 4.70 g/L MgCl$_2$·6 H$_2$O and 3.48 g/L Na$_2$S$_2$O$_3$·5 H$_2$O. For heterotrophic growth,

lactate was used as an electron donor and carbon source, 20 mM of Na-lactate solution was added

and incubated with approximately 1 bar of N$_2$/CO$_2$ (80:20, v/v) headspace; for autotrophic growth,

lactate was omitted and H$_2$ and CO$_2$ were used as an electron donor and carbon source,

respectively, with 21 psig H$_2$/CO$_2$ (80:20, v/v) in the headspace. The headspaces of autotrophic

cultures were flushed and re-pressurized with H$_2$/CO$_2$ once per day after each OD measurement.

The δ^2H of culture media was manipulated by volumetrically diluting 99.9 % purity ^{2}H$_2$O.

All cultures of *A. fulgidus* were incubated at 82 °C. Heterotrophic cultures were incubated

statically, and autotrophic cultures were shaken at 200 rpm to increase gas exchange between the

H$_2$/CO$_2$ headspace and culture media. The optical density at 600 nm (OD$_{600}$) of samples was used

for monitoring growth, and linearity against direct cell counts was confirmed (depth 0.02 mm;

Petroff-Hausser Counter 3900) throughout the absorbance range. Biomass samples were harvested

twice per experiment, one during mid-exponential phase and the other during early stationary

phase. To harvest, 50 mL of culture was centrifuged for 15 minutes at 4 °C and 4500 *g* and resulting

wet cell pellets were stored at –80 °C until lipid extraction.

Batch cultures of *Acidianus* sp. DS80 and *M. sedula* DSM 5348T were grown for additional

archaeal lipid $\epsilon_{L/W}$ measurements (not for water isotope labeling experiments or isotope flux-

balance model). *Acidianus* sp. DS80 was grown as previously described in (Rhim, Zhou et al.,

2024), and *M. sedula* was grown in Brock basal salts medium with Allen's trace element mixture

(Brock et al., 1972). See Supporting Information for additional details for growth conditions for

Acidianus sp. DS80 and *M. sedula*.

2.2. Lipid Extraction, Ether Cleavage, and Quantification

Frozen cell pellets were lyophilized overnight, and dry cell pellets were extracted using acidic

hydrolysis-methanolysis. Briefly, dry cell pellets were physically disrupted in 1.5-mL

microcentrifuge tubes by vortexing with methanol (MeOH) and *ca*. 250 µL of 100 µm combusted

glass beads for 10 minutes at 3000 rpm using a Disruptor Genie (Scientific 126 Industries, SI-

DD38). After excess MeOH wasevaporated, octacosanol (C_{28}) or lignocerol (C_{24}) and C_{46} GTGT

were added to all samples as internal extraction standards. Samples were incubated at 65 °C for 90

minutes with 3 N methanolic HCl (33 % water content) to hydrolyze tetraether headgroups. Methyl

tert-butyl ether (MTBE) was added to the cooled samples at a ratio of 3:2 (acid:MTBE, v/v), and

samples were sonicated for 5 minutes (Qsonica Q500, Newtown, CT, USA). After sonication, n-

hexane was added at a 1:1 ratio (MTBE:hexane, v/v), vortexed, and centrifuged at 15,000 g for 3

minutes. The upper organic phase was collected three times with n-hexane for the two subsequent

rounds. The resulting total lipid extracts (TLEs) were dried under N_2 and proceeded to ether

cleavage or stored at –20 °C until ether cleavage.

Ether bonds were cleaved in 57% hydroiodic acid (HI) at 125 °C for 4 hours. Then the organic

phase was collected three times using n-hexane. The resulting alkyl iodides were reduced to

alkanes (biphytanes, BPs) with H_2 in the presence of $Pt^{(IV)}O_2$ (Kaneko et al., 2011). Hydrogenated

samples were passed through combusted pasteur pipettes with glass wool plugs to remove PtO_2

powder. Filtered samples were dried under N_2 and stored at –20 °C until analysis. The resulting

ether-cleaved BPs were analyzed on a single quadrupole gas chromatography-mass spectrometer

(GC-MS; Thermo ISQ LT with TRACE 1310) for compound identification and on a GC-flame

ionization detector (GC-FID; Thermo TRACE 1310) for quantification, both housed in the Organic

Geochemistry Laboratory at the University of Colorado Boulder.

2.3.   Hydrogen isotope analysis and data processing

The δ^2H of BPs was analyzed on a GC-pyrolysis-isotope ratio MS (GC-P-IRMS) on a GC IsoLink

II IRMS System (Thermo Scientific) in the Organic Geochemistry Laboratory at the University of

Colorado Boulder. The GC-P-IRMS system consists of a Trace 1310 GC fitted with a

programmable temperature vaporization (PTV) injector and either a 30 m ZB5HT column (i.d. =

0.25 mm, 0.25 μm, Phenomenex, Torrance, CA, USA) or a 60 m DB1 column (i.d. = 0.25 mm,

0.25 μm, Agilent, Santa Clara, CA, USA), ConFlo IV interface, and MAT 253 Plus mass

spectrometer (Thermo Scientific).

A. fulgidus has been observed to produce both tetraether and diether lipids, with the majority of

GDGT with no ring (GDGT-0) and a small fraction with up to two pentacyclic rings (Lai et al.,

2008). The tetraether-to-diether ratio for *A. fulgidus* is expected to be ~1:1 at the growth

temperature in this study (82 °C; Lai et al., 2008), and we focus on the tetraether fraction for the

rest of the discussion. In line with the previous observations, the majority of BPs derived from *A.*

fulgidus GDGT did not have rings (BP-0). All δ^2H values reported in this study are for BP-0.

All hydrogen isotope ratios are reported in the standard delta notation (δ^2H) in permil (‰) units,

against the Vienna Standard Mean Ocean Water (VSMOW) international standard:

$$\delta\,2H \;=\; \left[\frac{(2H/1H)_{sample}}{(2H/1H)_{VSMOW}} - 1\right] \times 1000 \qquad \text{Eq. (1)}$$

The lipid-water hydrogen isotope fractionation factor is reported in alpha notation ($^2\alpha_{L/W}$), and

fractionation is reported in epsilon notation ($^2\varepsilon_{L/W}$) in ‰:

$$2\alpha_{L/W} = \frac{(2H/1H)_{lipid}}{(2H/1H)_{water}} \qquad\qquad \text{Eq. (2)}$$

$$2\varepsilon_{L/W} = \left[2\alpha_{L/W} - 1\right] \times 1000 \qquad\qquad \text{Eq. (3)}$$

Values of δ^2H were first determined relative to H_2 reference gas (δ^2H_{raw}), and then calibrated

externally using a standard n-alkane mixture (A6, containing C_{15} through C_{30} n-alkanes spanning

–9 to –263 ‰ vs. VSMOW; A. Schimmelmann, Indiana University). The A6 standard was

combined with a C_{36} n-alkane (nC_{36}, –259.2 ‰ vs. VSMOW; A. Schimmelmann, Indiana

University) and measured at regular intervals at different concentrations. The BP δ^2H calibration

was done in R, using the packages *isoreader* (Kopf et al., 2021) and *isoprocessor* available at

github.com/isoverse. Details of the data reduction and calibration methods are described in

(Leavitt, Kopf et al., 2023). The RMSE of the calibration peaks was 4.2 ‰ (n=613) across a peak

size range from 6.2 to 59.9 Vs. The RMSE of the C_{36} n-alkane internal standard was 7.5 ‰ (n=90)

across a peak size range from 6.4 to 50.9 Vs. Calibrated δ^2H values for the biphytanes were

corrected for the H added during hydrogenation of alkyl iodides (see Leavitt, Kopf et al., 2023).

The hydrogenation correction ranged from 9.8 to 13.7 ‰ and increased analytical uncertainty by

up to 1.8 ‰.

The δ^2H of the H_2 gas in the H_2/CO_2 headspace of autotrophic cultures was analyzed on a Thermo

Gasbench II in the Stable Isotope Laboratory at Northwestern University. The δ^2H of non-

exchangeable H in lactate used for heterotrophic cultures was assumed to be –60 ‰, which is the

value previously determined for glucose (Zhang et al., 2009).

The δ^2H of natural abundance and deuterium-enriched medium water was measured at the Stable

Isotope Laboratory at Dartmouth College as previously described (Kopec et al., 2019). Briefly,

water was reduced with chromium at 850 °C using the H-Device, and the δ^2H of the resulting gas

was measured by an IRMS (Thermo Delta Plus XL). The measured δ^2H values were converted to

the water δ^2H equivalent by calibrating with water standards with known isotopic compositions

analyzed in the same way. For samples expected to be outside the range of the standards (e.g., the

deuterium-enriched medium), samples were diluted with water standards. The 1σ uncertainty for

δ^2H measurements is <0.5 ‰.

2.4.   Model implementation

An isotope flux-balance model was adapted from the conceptual framework described in Leavitt,

Kopf et al. (2023). Our model considers the Wood-Ljungdahl (reductive acetyl-CoA) pathway

used by A. fulgidus for determining the sources and fluxes of H during acetyl-CoA (Ac-CoA)

synthesis. It also considers the general archaeal lipid biosynthesis pathway for the accounting of

total H budget in archaeal lipids. Analytical solutions for the model were calculated in R. Details

of the model design and outcomes are described in section 4.

3.   Results

3.1.   Cultures and growth rates

Archaeoglobus fulgidus is a metabolically flexible strain from the phylum Euryarchaeota. We

cultivated *A. fulgidus* in batch (closed-system) cultures on three different combinations of electron

donors and acceptors (Table 1; Fig. 1A–C). In two treatments, *A. fulgidus* was grown as a

heterotroph, using lactate as the carbon source and electron donor and sulfate or thiosulfate as the

terminal electron acceptor. In the third treatment, *A. fulgidus* was grown as an autotroph, using

CO_2 as the carbon source, H_2 as the electron donor, and thiosulfate as the terminal electron

acceptor. Cultures were acclimated to each treatment by growing on respective substrate

combinations for two or more passages before inoculation. Overall, heterotrophic conditions

resulted in faster growth (T-L, $t_d = 6.2 \pm 0.6$ h; S-L, $t_d = 10.3 \pm 1.7$ h) compared to autotrophic

growth (T-H_2/CO_2, $t_d = 19.3 \pm 3.9$ h) (Table 2; Fig. 1A–C).

Table 1. Summary of culture experiments.

Experimental Condition	electron donor	electron acceptor	carbon source
Thiosulfate-Lactate (T-L)	Lactate (20 mM)	Thiosulfate (14 mM)	Lactate (20 mM)
Sulfate-Lactate (S-L)	Lactate (20 mM)	Sulfate (14 mM)	Lactate (20 mM)
Thiosulfate-H_2/CO_2 (T-H_2/CO_2)	H_2 (g)* (80% v/v)	Thiosulfate (14 mM)	CO_2 (g)* (20% v/v)

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255 * The headspace of autotrophic cultures was filled with 21 psig H_2/CO_2 (80:20, v/v) and re-pressurized with H_2/CO_2
256 once per day after each OD measurement.
257

258 3.2.   Carbon metabolism and growth phase affect lipid δ^2H and regression parameters

259 Different modes of carbon metabolism in *A. fulgidus* not only resulted in different growth rates

260 but also affected lipid δ^2H and regression parameters. The lipid isotopic composition reflects the

261 sum of distinct isotope fractionations between lipids and each external H source, resulting in the

262 isotopic mass balance (Zhang et al., 2009):

$$R_{BP} = f_W \cdot 2\alpha_{L/W} \cdot R_W + f_S \cdot 2\alpha_{L/S} \cdot R_S \qquad \text{Eq. (4)}$$

where R_{BP}, R_W, and R_S represent the $^2H/^1H$ ratios of BPs from GDGT-0 (BP-0), media water, and

substrate (lactate for the heterotrophic conditions; H_2 for the autotrophic condition), respectively.

f_W and f_S are the fractions of lipid-bound H originating from water and substrate, respectively. $^2\alpha$

represents the net isotope fractionation factor between lipid and water ($^2\alpha_{L/W}$) and between lipid

and substrate ($^2\alpha_{L/S}$), respectively. In parallel cultures where only one parameter (R_W) is varied

experimentally, measurements of R_L, R_W, and R_S provide the basis for regression of R_L on R_W.

The isotope analysis results and regression parameters are summarized in Table 2.

The δ^2H_L values were strongly correlated with δ^2H_W values in all experiments (Table 2; $R^2 \geq 0.993$

and $p \leq 0.05$ in every regression). The regression slopes varied as a function of carbon metabolism

and growth phase. Overall, the two heterotrophic conditions (T-L and S-L) and exponential growth

phase resulted in lower slopes compared to the autotrophic condition (T-H_2/CO_2) and stationary

growth phase. Slopes for the two heterotrophic conditions were similar (0.481 ± 0.017 for T-L,

0.477 ± 0.028 for S-L during exponential phase; 0.548 ± 0.015 for T-L, 0.519 ± 0.016 for S-L

during stationary phase; Table 2). The autotrophic condition resulted in higher slopes (0.552 ±

0.019, exponential phase; 0.628 ± 0.025, stationary phase; Table 2). Notably, the higher slope

observed during the stationary phase of the autotrophic culture was the closest to the slope

estimated assuming $f_W = 1$ (0.707) (Table 2; Fig. 1D).

Substrate[a]	t_d (h)[b]	Growth Phase[c]	Avg. OD[d]	$\delta^2 H_W$[e] (‰)	s.d. (‰)	$\delta^2 H_L$[f] (‰)	s.d. (‰)	$^2\varepsilon_{L/W}$ (‰)	s.d. (‰)	n[g]	Slope[h] ($^2\alpha_{L/W} \cdot f_W$)	s.e.	R^2	p
T-L	6.2 ± 0.6	Exp	0.147	362.4	0.2	−70.7	6.6	−317.9	29.4	3				
			0.135	176.0	0.2	−164.6	7.0	−289.6	12.4	3	0.481	0.017	0.996	1 × 10⁻⁷
			0.141	−64.0	0.2	−275.8	7.1	−226.2	5.9	2				
		Sta	0.346	361.3	0.2	−52.7	6.7	−304.1	38.5	2				
			0.344	175.5	0.2	−152.4	7.5	−279.0	13.7	2	0.548	0.015	0.999	3 × 10⁻⁶
			0.351	−63.7	0.2	−285.3	8.7	−236.7	7.2	2				
S-L	10.3 ± 1.7	Exp	0.096	361.8	0.2	−87.5	9.7	−329.9	36.4	1				
			0.100	174.7	0.2	−163.5	4.8	−287.9	8.5	3	0.477	0.028	0.993	7 × 10⁻⁵
			0.082	−61.7	0.2	−289.7	5.9	−243.0	5.0	2				
		Sta	0.307	361.4	0.2	−66.9	5.7	−314.6	26.9	2				
			0.295	175.4	0.2	−168.1	4.9	−292.3	8.5	2	0.519	0.016	0.998	4 × 10⁻⁷
			0.282	−61.5	0.2	−287.1	10.2	−240.3	8.6	2				
T-H₂/CO₂	19.3 ± 3.9	Exp	0.052	348.3	0.2	−72.8	5.0	−312.3	21.5	2				
			0.046	164.8	0.2	−176.8	7.0	−294.0	11.5	1	0.552	0.019	0.998	8 × 10⁻⁵
			0.054	−66.5	0.2	−302.1	4.7	−252.4	4.0	1				
		Sta	0.137	348.5	0.2	−71.7	8.7	−311.7	37.8	3				
			0.132	166.0	0.2	−173.8	10.0	−291.4	16.8	1	0.628	0.025	0.997	1 × 10⁻⁵
			0.134	−64.5	0.2	−326.4	10.2	−279.9	8.8	2				
T-H₂/CO₂ (calculated)	-	-	-	-	-	-	-	-	-	-	*0.707 (f_W = 1)*	-	-	-

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[a] Thiosulfate-lactate (T-L); sulfate-lactate (S-L); thiosulfate-H_2/CO_2 (T-H_2/CO_2)
[b] Doubling time calculated for logarithmic growth
[c] Biomass harvested during mid-exponential growth phase (Exp); biomass harvested during early stationary phase, including biomass from both exponential and
stationary phase (Sta)
[d] Optical density (OD)
[e] $\delta^2 H$ of culture medium at the time of biomass sampling. Average of biological and technical replicates ± 1 s.d.
[f] Weighted mean $\delta^2 H$ of biphytane from GDGT-0
[g] Number of biological replicates used to calculate the average values of $\delta^2 H_{Water}$, $\delta^2 H_{Lipid}$, and $^2\varepsilon_{L/W}$
[h] Slope for the regression of $\delta^2 H_{Lipid}$ against $\delta^2 H_{Water}$ (*see* Figure 1D-F)
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Figure 1. Growth curves and regression of lipid δ²H versus water δ²H for *Archaeoglobus fulgidus*. Growth curves of *A. fulgidus* grown autotrophically on (A) thiosulfate and H_2/CO_2 or heterotrophically on (B) thiosulfate and lactate or (C) sulfate and lactate. Regression of biphytane hydrogen isotopic composition (δ^2H_L) on water isotopic composition (δ^2H_W) grown on (D) thiosulfate and H_2/CO_2, (E) thiosulfate and lactate, and (F) sulfate and lactate.

4. *A. fulgidus* H-isotope flux-balance model

To further interpret these findings, we modified the isotope flux-balance model recently developed for an ammonia-oxidizing archaeon *N. maritimus* SCM1 (Leavitt, Kopf et al., 2023) to reflect the biochemistry underlying the different modes of carbon metabolism in *Archaeoglobus fulgidus* VC-16. The overall conceptual framework consists of two parts for the stoichiometric accounting of:

1) methyl-H in Ac-CoA and 2) the 80 H atoms in the C_{40} alkyl chain of BPs inherited from Ac-

CoA. The former requires metabolism-specific information considering the mode(s) of carbon

fixation and experimental conditions such as substrate availability, while the latter is informed by

more universal information about archaeal isoprenoid biosynthesis. Below, we describe the details

of the two parts of our model.

4.1.   Sources of lipid-H via the modified Wood-Ljungdahl pathway

4.1.1.   Heterotrophy

A. fulgidus can grow autotrophically or heterotrophically via the modified reductive Ac-CoA or

Wood-Ljungdahl pathway (Möller-Zinkhan et al., 1989; Möller-Zinkhan and Thauer, 1990;

Hocking et al., 2014). During chemoorganoheterotrophic growth on lactate, the type species *A.*

fulgidus VC16 couples the complete oxidation of organic substrates to CO_2 with the reduction of

sulfate or thiosulfate to sulfide. In this study, the organic substrate was lactate. *A. fulgidus* oxidizes

lactate via pyruvate and Ac-CoA using enzymes and cofactors involved in the C_1-pathway, similar

to those found in methanogens, in reverse direction (Fig. 2A; Möller-Zinkhan et al., 1989; Möller-

Zinkhan and Thauer, 1990). A D-lactate dehydrogenase in *A. fulgidus* catalyzes the first step of

lactate oxidation to pyruvate coupled with the reduction of menaquinone (MQ), resulting in

menaquinol (MQH_2) (Step 1, Fig. 2A; Kunow et al., 1993; Reed and Hartzell, 1999). The next step

of pyruvate oxidation to Ac-CoA is catalyzed by pyruvate:ferredoxin oxidoreductase (Step 2, Fig.

2A; Kunow et al., 1993). The three H atoms in the methyl group of lactate most likely remain

intact and are inherited to Ac-CoA. Some of these Ac-CoA products are used for biomass synthesis

(Step 3.2, Fig. 2A), including isoprenoid lipid biosynthesis, while some of them are degraded

further for the complete oxidation of lactate. For the latter, carbon monoxide dehydrogenase

catalyzes both the reversible carbon-carbon cleavage of Ac-CoA and the oxidation of carbonyl

group to CO_2 coupled with ferredoxin (Fd) reduction (Step 3.1, Fig. 2A; Möller-Zinkhan and

Thauer, 1990).

The coenzymes involved in the downstream oxidative branch of Ac-CoA degradation could

transfer hydrides (e.g., $F_{420}H_2$ in Steps 4 and 5, Fig. 2A). These fluxes of H are important to track

because the hydrides from Ac-CoA (ultimately from lactate) can be recycled for NADP reduction,

the product of which (NADPH) can donate hydrides to isoprenoid precursors during lipid

biosynthesis. We assign F_{420} and Fd as the electron acceptors for Steps 4–5 and Step 8, respectively

(Fig. 2A), given that the genes relating to all steps of the reductive Ac-CoA pathway are expressed

in *A. fulgidus* (Hocking et al., 2014) and considering the known mechanisms of how the enzymes

of interest operate in methanogens. In CO_2-reducing methanogens, both

methylenetetrahydromethanopterin reductase (Mer) and methylenetetrahydromethanopterin

dehyrogenase (Mtd) oxidize $F_{420}H_2$ (Steps 4 and 5, respectively, in reverse; Fig. 2A), and

formylmethanofuran dehydrogenase (Fwd) requires Fd_{red} (Step 8, in reverse; Fig. 2A) (Thauer,

2012).

4.1.2. Autotrophy

During chemolithoautotrophic growth *A. fulgidus* VC16 couples H_2 or formate oxidation with

thiosulfate or sulfite reduction (Achenbach-Richter et al., 1987; Steinsbu et al., 2010). We fed *A.*

fulgidus H_2 and thiosulfate as the electron donor and acceptor for the autotrophic condition (T-

H_2/CO_2). *A. fulgidus* reduces CO_2 to synthesize Ac-CoA via the C_1-pathway (Fig. 2B), in principle

a reversal of the Ac-CoA degradation during heterotrophy (Fig. 2A). The same enzymes involved

in H additions used by CO_2-reducing methanogens catalyze the successive addition of H during

autotrophic carbon fixation in *A. fulgidus*: Fwd (Step 1), Mtd (Step 4), and Mer (Step 5) (Fig. 2B).

During chemolithoautotrophy, the H atoms in methyl precursors in theory derive from intracellular

water. However, one of the isoenzymes for methylenetetrahydromethanopterin (CH_2-H_4MPT)

dehydrogenase is H_2-dependent (Hmd) and uses H_2 directly, as opposed to $F_{420}H_2$ used by the other

isoenzyme Mtd, to reduce methyl-H_4MPT (Hendrickson and Leigh, 2008). Given the similarities

between methanogens and *A. fulgidus*, we consider the possible flux of H from H_2 to Ac-CoA

(Step 4–6, Fig. 2B) in one of the end-member scenarios (see 4.2; Table 3).

Figure 2. Carbon metabolism by *Archaeoglobus fulgidus* via the modified reductive acetyl-CoA or Wood-Ljungdahl pathway. (A) Chemoorganoheterotrophic growth on lactate: lactate is oxidized via pyruvate and acetyl-CoA, which either serves as the building blocks for biosynthesis or is further degraded for the complete oxidation of lactate. (B) Chemolithoautotrophic growth on H_2/CO_2: CO_2 is reduced to produce acetyl-CoA via the C_1-pathway, with enzymes and cofactors similar to those involved in methanogenesis. Abbreviations: MQH_2, menaquinol or reduced menaquinone; Fd_{red}, reduced ferredoxin; Ac-CoA, acetyl-CoA; CH_3-H_4MPT, methyl-tetrahydromethanopterin (H_4MPT); CH_2=H_4MPT, methylene-H_4MPT; CH^+≡H_4MPT, methenyl-H_4MPT; CHO-H_4MPT, formyl-H_4MPT; CHO-MFR, formylmethanofuran; $F_{420}H_2$, reduced coenzyme F_{420}.

4.2.   Model schematics and stoichiometric accounting

Fractions of Ac-CoA produced via the modified Wood-Ljungdahl pathway described above are used for the synthesis of archaeal isoprenoids. Leavitt, Kopf et al. (2023) provide a detailed overview of the sources of H and mechanisms of the biochemical steps involved in archaeal tetraether biosynthesis (Appendix B). In short, the mevalonate pathway synthesizes isopentenyl pyrophosphate (IPP), which are C_5 isoprene units, from three units of Ac-CoA. Then, three units of IPP and its isomer condense to yield the C_{20} precursor, geranylgeranyl diphosphate (GGPP). The geranylgeranyl chains are transferred around to yield di-O-geranylgenranylglyceryl phosphate (DGGGP) with the glycerol-1-phosphate (G1P) backbone. The tetraether synthase (Tes) enzyme catalyzes the condensation of two DGGGP units to yield the membrane-spanning tetraether. Finally, the geranylgeranyl reductase (GGR) enzyme catalyzes the saturation of tetraether to yield GDGT. We combine the stoichiometric accounting of the methyl-H in Ac-CoA for *A. fulgidus* (see 4.1) and of the 80 H atoms in the C_{40} alkyl chain of BPs inherited from Ac-CoA during tetraether biosynthesis described above (Leavitt, Kopf et al., 2023) to build an overall schematic of the specific model scenarios explored in this study (Fig. 3).

Figure 3. Schematic diagram of the isotope flux-balance model for *Archaeoglobus fulgidus*. (A) Heterotrophic scenario for lactate oxidation; (B) autotrophic scenario for H_2 oxidation. Color codes for the different H fluxes (*f*) are shown in the legend. The subscript W stands for water as the ultimate H source; the subscript S stands for substrate as the ultimate H source, where the substrates are lactate and H_2 for heterotrophy and autotrophy, respectively. See 4.2.1 and Appendix A (Eq. A5–A8) for detailed descriptions of individual fluxes. Abbreviations: LS, lipid synthesis; AS, acetyl-CoA synthesis; Cofa, cofactor; $F_{420}H_2$, reduced coenzyme F_{420}; Fd_{red}, reduced ferredoxin; Ac-CoA, acetyl-CoA; GGR, geranylgeranyl reductase.

4.2.1. *Empirically informed isotopic mass balance*

With the fractions defined in the overall hydrogen budget (see Appendix A), the isotopic mass balance in Eq. (4) can be further refined. Replacing the f_W and f_S terms in Eq. (4) with the sums of all fractions in Eq. (A5) (Appendix A), isotopic mass balance for the heterotrophic case becomes:

$$R_{BP,het} = \left(f_{W,direct} \cdot \alpha_{W,direct} + f_{W,cofa-LS} \cdot \alpha_{W,cofa-LS}\right) \cdot R_W$$

$$+ \left(f_{S,cofa-LS} \cdot \alpha_{S,cofa-LS} + f_{S,AS} \cdot \alpha_{S,AS}\right) \cdot R_S$$

Eq. (5)

Similarly, isotopic mass balance for the autotrophic case becomes:

$$R_{BP,auto} = \left(f_{W,direct} \cdot \alpha_{W,direct} + f_{W,cofa-AS} \cdot \alpha_{W,cofa-AS} + f_{W,cofa-LS} \cdot \alpha_{W,cofa-LS}\right) \cdot R_W$$

$$+ f_{S,AS} \cdot \alpha_{S,AS} \cdot R_S$$

Eq. (6)

Now we can use the regression parameters estimated from the δ^2H_{H2O} label experiments (Fig. 1;

Table 2) to calculate analytical solutions for reasonable ranges of individual terms in Eq. (5–6). In

Eq. (5–6), the coefficient for R_W in the first term represents the slope, m, and the entire second

term represents the intercept, c. For heterotrophy:

$m_{het} = f_{W,net} \cdot \alpha_{W,net} = f_{W,direct} \cdot \alpha_{W,direct} + f_{W,cofa-LS} \cdot \alpha_{W,cofa-LS}$

$\varepsilon_{W,net} = \dfrac{m_{het}}{f_{W,net}} - 1 = \dfrac{f_{W,direct} \cdot \alpha_{W,direct} + f_{W,cofa-LS} \cdot \alpha_{W,cofa-LS}}{f_{W,net}} - 1$

Eq. (7)

$c_{het} = f_{S,net} \cdot \alpha_{S,net} \cdot R_S = f_{S,cofa-LS} \cdot \alpha_{S,cofa-LS} + f_{S,AS} \cdot \alpha_{S,AS}$

$\varepsilon_{S,net} = \dfrac{c_{het}}{f_{S,net} \cdot R_S} - 1 = \dfrac{f_{S,cofa-LS} \cdot \alpha_{S,cofa-LS} + f_{S,AS} \cdot \alpha_{S,AS}}{f_{S,net}} - 1$

Eq. (8)

where $\varepsilon_{W,net}$ and $\varepsilon_{S,net}$ refer to the net isotope fractionation between water and lipid (combining

isotope effects associated with $f_{W,direct}$ and $f_{W,cofa-LS}$) and between substrate and lipid (combining

isotope effects associated with $f_{S,cofa-LS}$ and $f_{S,AS}$), respectively.

4.3.   Model results

Table 3, Fig. 4 and Fig. 5 show the summary of model results. We first estimated $f_{W,net}$, which we

define as the total flux that directly or indirectly (e.g., via cofactors) reflects water isotope

composition (i.e., not affected by substrate isotopic composition) ($f_{W,net} = f_{W,direct} + f_{W,cofa}$), for a

range of x_{ex} from 0 % to 100 %. Overall, the autotrophic case had a higher range of $f_{W,net}$ compared

to the two heterotrophic cases (Table 3; Fig. 4A–C). Even for the minimal water exchange and

maximal substrate contribution scenario ($x_{ex} = 0$ %, $s = 100$ %), when the contribution from

substrate is expected to be the highest, $f_{W,net}$ was 80 % (Table 3; dashed line, Fig. 4A). For the two

heterotrophic cases, the $f_{W,net}$ values were 31 and 70 % for the highest ($x_{ex} = 0$ %, $s = 100$ %) and

lowest ($x_{ex} = 100$ %, $s = 0$ %) substrate contribution scenarios, respectively (Table 3; Fig. 4B–C).

For the net isotope fractionation between lipid and water ($\varepsilon_{W,net}$), the two heterotrophic cases

showed similar ranges and patterns where the $\varepsilon_{W,net}$ values decrease with increasing x_{ex} (Fig. 4E–

F). The $\varepsilon_{W,net}$ range was relatively more positive during stationary growth phase (–218 to 776 ‰

for the T-L condition; –259 to 682 ‰ for the S-L condition) compared to exponential growth phase

(–312 to 561 ‰ for the T-L condition; –319 to 546 ‰ for the S-L condition) (Table 3: Fig. 4E–F).

For the autotrophic case, the growth phase-dependent trend was also observed (more positive $\varepsilon_{W,net}$

during stationary phase; Table 3; Fig. 4D). Unlike in the heterotrophic cases, all $\varepsilon_{W,net}$ values for

the autotrophic case were negative and fell within a relatively narrow range for the entire range of

x_{ex} (–448 to –312 ‰ for exponential phase; –372 to –228 ‰ for stationary phase; Table 3). For the

autotrophic case, a theoretical end-member $\varepsilon_{W,net}$ with no substrate contribution can be estimated

by calculating the regression slope through the origin (i.e., no intercept, which represents the

substrate contribution term, $f_S \cdot {}^2\alpha_{L/W} \cdot R_S$; Eq. (4)): –293 ‰ for the autotrophic case (black dotted

line; Fig. 4D). This value falls within the range of $\varepsilon_{W,net}$ for stationary, rather than exponential,

growth phase (Fig. 4D). For the net isotope fractionation between lipid and substrate ($\varepsilon_{S,net}$), the

two heterotrophic cases showed similar ranges and patterns where $\varepsilon_{S,net}$ values increase with

increasing x_{ex} (Fig. 4H–I). The growth phase-dependent trend for $\varepsilon_{S,net}$ contrasted that of $\varepsilon_{W,net}$,

where the $\varepsilon_{S,net}$ range was relatively more negative during stationary growth phase (–688 to –282

‰ for the T-L condition; –654 to –201 ‰ for the S-L condition) compared to exponential growth

phase (–582 to –36 ‰ for the T-L condition; –592 to –59 ‰ for the S-L condition) (Table 3; Fig.

4H–I). A similar trend was observed for the autotrophic case (more negative $\varepsilon_{S,net}$ during stationary

phase; Table 3; Fig. 4G).

The net isotope fractionation factors ($\varepsilon_{W,net}$ and $\varepsilon_{S,net}$) can be further broken down into individual

fluxes associated with H isotopes with different routes of incorporation. The net isotope

fractionation factor between lipid and water ($\varepsilon_{W,net}$) can be divided into: 1) $\varepsilon_{W,direct}$ associated with

H atoms that are transferred through the $f_{W,direct}$ flux (Fig. 3) via H exchange between isoprenoid

precursors and intracellular water or addition via GGR, and 2) $\varepsilon_{W,cofa}$ associated with the $f_{W,cofa}$

flux ($f_{W,cofa-LS}$ for heterotrophy; $f_{W,cofa-LS}$ and $f_{W,cofa-AS}$ for autotrophy; Fig. 3) via cofactors (Fd_{red}

and $F_{420}H_2$ for autotrophy; Fd_{red} for heterotrophy) (Fig. 5A–C). The net isotope fractionation factor

between lipid and substrate ($\varepsilon_{S,net}$) can be divided into: 1) $\varepsilon_{S,AS}$ associated with H atoms that are

transferred through the $f_{S,AS}$ flux (Fig. 3) via Ac-CoA synthesis, and 2) $\varepsilon_{S,cofa}$ associated with the

$f_{S,cofa}$ flux ($f_{S,cofa-NADPH}$ and $f_{S,cofa-GGR}$ for heterotrophy; Fig. 3) via cofactors ($F_{420}H_2$) (Fig. 5D–E).

Interpretation of these results are discussed further in 5.2.1.

Table 3. Isotope flux-balance model results for *Archaeoglobus fulgidus*.

Growth phase	Scenario		$f_{W,net}$ (%)[c]			$\varepsilon_{W,net}$ (‰)[d]			$\varepsilon_{S,net}$ (‰)[e]		
	x_{ex} (%)[a]	s (%)[b]	T-L	S-L	T-H	T-L	S-L	T-H	T-L	S-L	T-H
Exp.	0	0	41	41	100	179	167	−448	−511	−523	n/a
	0	100	31	31	80	561	546	−312	−582	−592	−564
	100	0	70	70	100	−312	−319	−448	−36	−59	n/a
	100	100	60	60	90	−198	−206	−387	−277	−294	−139
Sta.	0	0	41	41	100	341	270	−372	−636	−595	n/a
	0	100	31	31	80	776	682	−218	−688	−654	−785
	100	0	70	70	100	−218	−259	−372	−282	−201	n/a
	100	100	60	60	90	−87	−136	−302	−461	−401	−577

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[a] the degree of water exchange during lipid synthesis, which depends on the extent of re-equilibration during
isomerization steps (Steps 2 and 6; Appendix B).
[b] the degree of H contribution to lipids from substrate (lactate for heterotrophy; H_2 for autotrophy)
[c] the net total contribution of H from water to lipids ($f_{W,net} = f_{W,direct} + f_{W,cofa}$)
[d] the net isotope fractionation between water and lipid (combining isotope effects associated with $f_{W,direct}$ and $f_{W,cofa-LS}$)
[e] the net isotope fractionation between substrate and lipid (combining isotope effects associated with $f_{S,cofa-LS}$ and $f_{S,AS}$)

Figure 4. Isotope flux-balance model results for *Archaeoglobus fulgidus*. (A–C) The total flux that directly or indirectly (e.g., via cofactors) reflects water isotope composition (i.e., not affected by substrate isotopic composition) ($f_{W,net}$); (D–F) Net isotope fractionation between lipid and water ($\varepsilon_{W,net}$); (G–I) Net isotope fractionation between lipid and substrate ($\varepsilon_{S,net}$). Each column represents the metabolic mode and growth substrates: autotrophy on thiosulfate and H_2/CO_2 (A, D and G); heterotrophy on thiosulfate and lactate (B, E and H); and heterotrophy on sulfate and lactate (C, F and I). Lighter shaded areas refer to exponential growth phase; darker shades refer to stationary growth phase. Solid lines refer to the s = 0 % (minimal substrate contribution) scenarios; dashed lines refer to the s = 100 % (maximal substrate contribution) scenarios. See 4.2.1 and Table 3 for details of the conditions for each scenario.

Figure 5. Isotope flux-balance model results for *Archaeoglobus fulgidus*. (A–C) Breakdown of the net isotope
fractionation between lipid and water ($\varepsilon_{W,net}$; Fig. 4D–F) into two different isotope fractionation factors: $\varepsilon_{W,direct}$ for H
atoms that are transferred through exchange during isomerization or addition via GGR, corresponding to $f_{W,direct}$ in Fig.
3, and $\varepsilon_{W,cofa}$ for H atoms that are transferred through cofactors (Fd_{red} and $F_{420}H_2$ for autotrophy, panel A; Fd_{red} for
heterotrophy, panels B and C), corresponding to $f_{W,cofa}$ in Fig. 3. (D–E) Breakdown of the net isotope fractionation
between lipid and substrate ($\varepsilon_{S,net}$; Fig. 4G–I) into two different isotope fractionation factors: $\varepsilon_{S,AS}$ for H atoms that are
inherited directly from substrate via Ac-CoA, corresponding to $f_{S,AS}$ in Fig. 3, and $\varepsilon_{S,cofa}$ for H atoms that are transferred
through cofactors (only applicable for heterotrophy via $F_{420}H_2$, panels D and E), corresponding to $f_{S,cofa}$ in Fig. 3. The
lighter shaded areas refer to the s = 0 % (minimal substrate contribution) scenarios; darker shaded areas refer to the s
= 100 % (maximal substrate contribution) scenarios. Solid lines and brown colors refer to heterotrophic conditions;
dashed lines and purple colors refer to autotrophic conditions. See 4.2.1 and Table 3 for details of the conditions for
each scenario.

496   5.   Discussion

5.1.   Effects of carbon metabolism and growth phase on lipid H isotopes composition in Archaea

The isotope flux-balance model results show a large fraction of lipid-H directly reflects water H

isotope compositions, rather than substrate (lactate or H_2) isotope compositions in *A. fulgidus*. The

$f_{W,net}$ is the net fraction of total biphytane H that directly records water H isotopes and is explicitly

not derived from organic substrate or H_2. We observe larger $f_{W,net}$ for autotrophic growth ($f_{W,net}$ =

80–100 %), whereas $f_{W,net}$ for heterotrophic growth was lower ($f_{W,net}$ = 31–70 %; Table 3; Fig. 4A–

C). Given that $\varepsilon_{W,net}$ is likely negative (i.e., a normal kinetic isotope fractionation), we can

constrain the $f_{W,net}$ range that corresponds to $\varepsilon_{W,net} < 0$ (*see* Fig. 4D–E). This yields $f_{W,net}$ ranges of:

38–52 % for T-L and 42–52 % for S-L during exponential growth phase, versus 50–56 % for T-L

and 52–54 % S-L during stationary growth phase. The lowest $f_{W,net}$ recorded were during

heterotrophy in exponential phase, independent of the terminal electron acceptor. For stationary

phase cultures of *A. fulgidus*, the $f_{W,net}$ was ≥80 % for autotrophy versus ≥50 % for heterotrophy.

The $f_{W,net}$ range for lipids produced during stationary phase alone is likely larger, because lipids

collected during stationary phase are a mixture of those produced during exponential phase and

those newly produced during stationary phase. The higher $f_{W,net}$ range observed for autotrophy is

reasonable, given that lactate is an additional source of H that contributes directly to Ac-CoA

synthesis during heterotrophy. A similar observation was reported for bacterial lipids where, for

instance, the mole fraction of lipid H derived from water ($\equiv X_W$ in Zhang et al., 2009; herein

equivalent to $f_{W,net}$) was larger for *Escherichia coli* grown on glucose compared to the same

organism grown on complex medium (Zhang et al., 2009). Though the lipid synthesis pathways

are different for Bacteria versus Archaea, it appears that the fractional contribution of substrates

that are more direct precursors for lipid biosynthesis generally increases when organic substrates

are used or H_2 is supplied in excess as an electron donor.

The effects of carbon metabolism and growth phase on lipid isotope composition have not been

investigated in detail for archaeal lipids. Our results indicate that both the mode of metabolism and

growth phase affect the lipid isotope composition (Table 3; Fig. 4). A recent study reported $\varepsilon_{L/W}$

of BPs produced by *N. maritimus* in continuous cultures at a range of doubling time and observed

little change in $\varepsilon_{W/L}$ (0.18 ± 0.09 ‰ increase in $\varepsilon_{L/W}$ per hour increase in doubling time; Leavitt,

Kopf et al., 2023). The relative variation in doubling time we observe for *A. fulgidus* grown on

different substrates here in batch cultures (Table 2) is similar (*ca*. 3-fold variation) to the variation

observed for *N. maritimus* in Leavitt, Kopf et al. (2023). A key difference here is that we do not

strictly control growth and metabolic rates, as was the case for chemostat cultivation of *N.*

maritimus (Leavitt, Kopf et al., 2023), and for *Sulfolobus acidocaldarius* (Harris et al., 2022).

Nonetheless, the range of $\varepsilon_{L/W}$ we observe across different metabolic modes by *A. fulgidus* (–231

± 6 ‰ for T-L, –241 ± 4 ‰ for S-L, and –271 ± 18 ‰ for T-H_2/CO_2; Table 2) is significant

compared to the entire range of $\varepsilon_{L/W}$ observed across a similar variation in growth rate by *N.*

maritimus (–272 ± 20 to –279 ± 8 ‰ for BP-0; Leavitt, Kopf et al., 2023). *N. maritimus* is an

obligate chemoautotroph, such that all BP-bound H is strictly biosynthetic, whereas we show here

that patterns of $\varepsilon_{W/L}$ can change between heterotrophic and autotrophic growth conditions (Fig. 4

and Fig. 6). Our results support this hypothesis and show variation in $\varepsilon_{L/W}$ up to ~40 ‰ within the

same metabolically versatile organism, in response to different electron donor-acceptor pairs and

modes of carbon metabolism.

Even though the range of $\varepsilon_{L/W}$ in *A. fulgidus* is relatively larger compared to *N. maritimus*, the

overall variation of $\varepsilon_{L/W}$ across all archaeal lipid studies to-date is consistently small compared to

the variation observed across bacterial lipids (Fig. 6).

5.2.   Factors that affect lipid-water H isotope fractionation across the three domains of life

Here we present a synthesis of the broader $\varepsilon_{L/W}$ patterns observed in lipids from all three domains

of life. Archaeal lipids are characterized by a relatively narrow range of large and negative values

of $\varepsilon_{L/W}$ across a wide range of organisms, metabolisms, and growth conditions (Fig. 6; Kaneko et

al., 2011; Wu et al., 2020; Leavitt, Kopf et al., 2023; this study). This is in contrast with the domain

Bacteria, where fatty acid $\varepsilon_{L/W}$ ranges a large span of >600 ‰ with both highly positive and

negative values (Fig. 6; Campbell et al., 2009; Zhang et al., 2009; Dawson et al., 2015; Osburn et

al., 2016; Wijker et al., 2019). Archaeal lipid $\varepsilon_{L/W}$ largely overlaps with that of eukaryotic lipid

$\varepsilon_{L/W}$, although eukaryotic lipid $\varepsilon_{L/W}$ spans a significant range depending on the types of organism

and classes of compound (Fig. 6; Sessions et al., 1999; Chikaraishi et al., 2004, 2009; Zhang and

Sachs, 2007; Sachs et al., 2016). Notably, there are subgroups of both domains Bacteria and

Eukarya that share similarities with Archaea in their $\varepsilon_{L/W}$. The shared similarities are interesting,

given the difference in lipid biosynthesis mechanism among the three domains. Below, we take a

closer look at the specific subsets of $\varepsilon_{L/W}$ values from the domains Euakrya and Bacteria that

resemble those from Archaea and discuss factors that may contribute to the large, negative values

of $\varepsilon_{L/W}$ observed across the three domains.

Figure 6. Summary of hydrogen isotope fractionation ($\varepsilon_{L/W}$) between lipids and water observed in culture studies. $\varepsilon_{L/W}$
values observed in all three domains of life: Archaea (circle), Bacteria (square), and Eukarya (diamond). Color codes
represent different types of metabolism: aerobic heterotrophy (red), anaerobic heterotrophy (purple), oxygenic
photoautotrophy (green), aerobic chemoautotrophy (orange), and anaerobic chemoautotrophy (blue). Electron donors
used for biosynthesis are shown on the horizontal axis: acetate (ac); succinate (su); pyruvate (py); lactate (la); glucose
(gl); methane (me); H_2+CO_2 (h2); formate (fo); oxalate (ox); H_2+Fe^{3+} (hf); H_2+S^0 (hs); ammonium (aa); yeast extract
(ye); H_2, acetate or methanol for methanogenesis (mg); water for photoautotrophy (pa). Tested organisms for bacterial
cultures include *C. oxalaticus*, *C. necator*, *R. palustris* (Zhang et al., 2009); *E. coli* (Zhang et al., 2009; Osburn et al.,
2016; Wijker et al., 2019); *D. autotrophicum* (Campbell et al., 2009; Osburn et al., 2016); *M. capsulatus* (Sessions et
al., 2002); *D. multivorans* (Dawson et al., 2015); *D. alaskensis* (Leavitt et al., 2016); *P. denitrificans*, *S. oneidensis*, *D.
hydrogenophilus*, *D. alaskensis*, *D. propionicus* (Osburn et al., 2016); *B. subtilis*, *P. fluorescens*, *R. radiobacter*, and *E.
meliloti* (Wijker et al., 2019). Tested organisms for eukaryotic cultures include *A. fundyense*, *I. galbana*, *Ascophyllum*
sp., *F. vesiculosis*, *Z. marina*, *S. alterniflora* (Sessions et al., 1999); *C. japonica* (Chikaraishi et al., 2004); *B. braunii*, *E.
unicocca*, *V. aureus* (Zhang and Sachs, 2007); *C. sativuc* (Chikaraishi et al., 2009); *E. huxleyi* (Sachs et al., 2016).
Tested organisms for archaeal cultures include *Acidianus* sp. DS80 (this study); *N. maritimus* (Leavitt, Kopf et al., 2023);
Sulfolobus sp. (Kaneko et al., 2011); *A. fulgidus* (this study); *Metallosphaera sedula* (this study); and *M. barkeri* (Wu et
al., 2020).

5.2.1. Large lipid-water fractionation upon isoprenoid saturation in Archaea and Eukarya

The consistently large, negative values of $\varepsilon_{W/L}$ for archaeal lipids are notable considering the

different types of organism and metabolic modes studied so far (Fig. 6). In Fig. 5, we present the

model solution spaces for individual H fluxes, within either net flux that reflects water isotope

composition (i.e., not affected by substrate contributions; $f_{W,net}$) or net flux that reflects isotopic

contributions from organic substrate or H_2 ($f_{S,net}$), to further investigate specific parts of archaeal

lipid biosynthesis pathway that may contribute to the observed large and negative values of $\varepsilon_{L/W}$.

Prior work yielded estimates of $\alpha_W \cong 0.9$ for isotope fractionation associated with direct water

incorporation across many bacteria (Zhang et al., 2009; Wijker et al., 2019), and $\varepsilon_{L/W}$ values

modeled for *N. maritimus* with this empirical α_W were consistent with experimental observations

(Leavitt, Kopf et al., 2023). Assuming an equivalent value of $\varepsilon_{W,direct}$ (i.e., isotope fractionation

associated with the $f_{W,direct}$ flux is –100 ‰), we constrain plausible ranges of $\varepsilon_{W,cofa}$ in *A. fulgidus*

(Fig. 5A–C). The ranges of $\varepsilon_{W,cofa}$ at $\varepsilon_{W,direct}$ = –100 ‰ are particularly well constrained for the

autotrophic case (–599 to –436 ‰ at s = 0 %; –519 to –276 ‰ at s = 100 %; Fig. 5A).

It is plausible that both $\varepsilon_{W,cofa-AS}$ and $\varepsilon_{W,cofa-LS}$ have negative values that are comparable to the

combined value of $\varepsilon_{W,cofa}$, as the less likely alternative would be a negative $\varepsilon_{W,cofa}$ that consists of

very different $\varepsilon_{W,cofa-AS}$ and $\varepsilon_{W,cofa-LS}$. Our current model cannot distinguish between isotope

fractionation factors associated with different fluxes within $f_{W,cofa}$ (i.e., $f_{W,cofa-AS}$ through Ac-CoA

vs. $f_{W,cofa-LS}$ through NADPH), both of which can inherit H from cofactors $F_{420}H_2$ and/or Fd_{red} (Fig.

3B). However it is worth noting that, for both heterotrophic and autotrophic cases, a fraction of

the $f_{W,cofa-LS}$ flux represents hydride transfer during the final saturation (double bond reduction) via

GGR (Fig. 3; Step 10, Appendix B). Results of the isotope flux-balance model for *N. maritimus*

also indicated that the GGR reduction step is strongly fractionating (ε_{GGR} = –690 to –820 ‰ if

ε_{NADPH} = –100 to –300 ‰, depending on the degree of water exchange, or x_{ex} in this study),

regardless of how other model parameters are set (Leavitt, Kopf et al., 2023). Given how *N.*

maritimus and *A. fulgidus* share little in common otherwise (metabolically, genetically, and

biochemically), large fractionation associated with the GGR reduction step provides a

parsimonious explanation for the similar $\varepsilon_{L/W}$ values observed in the two organisms, as it is the

final step of archaeal lipid biosynthesis shared across all tetraether-producing Archaea. This would

be consistent with the similarly large, negative values of $\varepsilon_{L/W}$ observed in other archaeal lipids

(Fig. 6); testing more Archaea with the combined approach of isotope labeling experiments and

biochemically informed isotope flux-balance model will help further confirm this idea.

A closer examination of the eukaryotic lipids that share similar $\varepsilon_{L/W}$ values to archaeal lipids lends

further support to the notion that GGR reduction contributes to large H isotope fractionation

between lipids and water. Fig. 7A shows $\varepsilon_{L/W}$ for different subgroups of eukaryotic lipids,

categorized by three different biosynthetic pathways. In general, the two major families of lipids

are acetogenic and isoprenoid lipids which consist of acetate and isoprene monomeric units,

respectively (Pearson, 2014). Isoprenoid lipids are further divided into two groups, depending on

whether the isoprene (C_5) building blocks are synthesized via the mevalonic acid (MVA) pathway

(Lynen et al., 1958; Bloch, 1959) or the methylerythritol phosphate (MEP) pathway (Rohmer et

al., 1993). All Archaea discussed in this study synthesize isoprenoid lipids via the MVA pathway.

The MEP pathway is generally defined as bacterial and the MVA pathway as eukaryotic and

archaeal; both pathways often simultaneously occur in plants, where the MEP pathway is

expressed in chloroplasts (relics of the cyonobacterial plastid ancestor) and the MVA pathway is

expressed in the cytosol (Lichtenthaler et al., 1997; Lichtenthaler, 1999; Pearson, 2014). An

exception to this is the case of green algae, where both the cytosolic sterols and plastidic phytol

isoprenoids are synthesized via the MEP pathway (Lichtenthaler et al., 1997; Disch et al., 1998;

Pearson, 2014).

The three different biosynthetic pathways of eukaryotic lipids (acetogenic, MVA, and MEP) are

associated with distinct ranges of $\varepsilon_{L/W}$. In general, isoprenoid lipids have more negative $\varepsilon_{L/W}$

compared to acetogenic lipids (Estep and Hoering, 1980; Sessions et al., 1999). Among isoprenoid

lipids, plastidic isoprenoids such as phytols are more depleted in ^{2}H than the cytosolic counterparts

(Sessions et al., 1999; Chikaraishi et al., 2004, 2009). This general trend is captured in Fig. 7A,

where acetogenic lipids (e.g., *n*-alkanol and *n*-alkyl lipids including fatty acids) have more positive

$\varepsilon_{L/W}$ compared to isoprenoid lipids. The C_{15} and C_{30} isoprenoids synthesized via the MVA pathway

(e.g., sterol, triterpenoids, sesquiterpenes, squalene) are relatively more enriched in ^{2}H compared

to the C_{20} isoprenoids (e.g., phytol, phytadiene, diterpenes) and botryococcenes (triterpenes)

synthesized via the MEP pathway (Fig. 7A).

While the relative difference in $\varepsilon_{L/W}$ between MVA and MEP pathways holds in general, the larger

and more negative values of $\varepsilon_{L/W}$ appear to correlate more with highly fractionating enzymatic

steps, rather than specific pathways. For example, Chikaraishi et al. (2004) observed in

Cryptomeria japonica (conifer) that $\varepsilon_{L/W}$ values for cytosolic lipids produced via the MVA

pathway, sesquiterpenes (–228 to –238 ‰) and squalene (–225 to –251 ‰), are similar to those

for plastidic diterpenoids (–238 to –251 ‰) produced via the MEP pathway (Fig. 7A). Phytol and

botryococcenes, on the other hand, have particularly negative $\varepsilon_{L/W}$ values among eukaryotic lipids

produced via the MEP pathway. For example, Zhang and Sachs (2007) found that botryococcenes

with longer chain lengths had progressively lower δ^2H compared to the C_{30} botryococcene

(precursor to all other botryococcenes) in *Botryococcus braunii* race B. They hypothesized that

significantly ^2H-depleted hydrogen from methionine is incorporated into botryococcene during the

step-wise methylation (Zhang and Sachs, 2007). Similarly, phytol in *C. japonica* had up to 65 ‰

more negative $\varepsilon_{L/W}$ compared to diterpenoids, another plastidic lipid produced via the MEP

pathway within the same organism (Chikaraishi et al., 2004). Later, Chikaraishi et al. (2009)

observed a gradual depletion in δ^2H from phytol precursors (starting from geranylgeraniol) to

phytol in cucumber, indicating that significantly ^2H-depleted hydrogen is incorporated into phytol

and its precursors during the step-wise hydrogenation catalyzed by GGR ($\varepsilon_{GGR} \cong -650$ ‰;

Chikaraishi et al., 2009).

The $\varepsilon_{L/W}$ values of phytol (and its dehydration product phytadiene) investigated in a wide range of

eukaryotes so far are all significantly large and negative (e.g., –357 to –278 ‰ for phytol in

dinoflagellate, coccolithophore, marsh grass, cucumber, and conifer; Fig. 7A). Similarly, all values

of $\varepsilon_{L/W}$ investigated in a range of different archaea are significantly large and negative (e.g., –341

to –148 ‰ for phytanes and biphytanes in *N. maritimus*, *Sulfolobus* sp., *M. sedula*, *Acidianus* sp.,

M. barkeri, and *A. fulgidus*; Fig. 7A). Taken together, empirical and model-based results suggest

that the GGR reduction is a highly fractionating step that contribute to the large fractionation

observed in both eukaryotic and archaeal lipid synthesis (Chikaraishi et al., 2009; Leavitt, Kopf et

al., 2023; this study, Fig. 5 and Fig. 6).

Figure 7. Comparison of $\varepsilon_{L/W}$ values observed in eukaryotic *vs*. archaeal lipids and in bacterial *vs*. archaeal lipids. **(A)**
$\varepsilon_{L/W}$ values observed in acetogenic and isoprenoid lipids from the domains Archaea and Eukarya. Different classes of
compounds are shown on the horizontal axis. Compounds shown for acetogenic pathway: n-alkanol, *n*-alkane, *n*-alkyl
(*n*-alkanoic acids and fatty acids), and alkadiene. For the MVA pathway: C_{15} and C_{30} isoprenoids (sterol, squalene,
triterpenoids, and sesquiterpenes) and P+BP (phytanes and biphytanes). For the MEP pathway: C_{20} isoprenoids
(diterpenes, sandaracopimarinol, ferruginol, diterpenol, and sugiol), phytol+ptd (phytol and phytadiene), and btc
(botryococcene). **(B)** $\varepsilon_{L/W}$ values observed in fatty acids from anaerobic and/or autotrophic Bacteria and isoprenoid
lipids from Archaea. Different organisms are shown on the horizontal axis. For Bacteria: *E. coli* (Ec), *Paracoccus*
denitrificans (Pd), SRB (*Desulfobacterium autotrophicum*, *Desulfococcus multivorans*, *Desulfobacter hydrogenophilus*,
Desulfovibrio alaskensis, *Desulfolobus propionicus*), and *Cupriavidus* sp. (Cu). For Archaea: *Acidianus* sp. DS80 (Ac),
Sulfolobus sp. (Su), *Archaeoglobus fulgidus* (Af), *Metallosphaera sedula* (Ms), *Nitrosopumilus maritimus* (Nm), and
Methanosarcina barkeri (Mb). A full list of references can be found in Appendix C.

5.2.2. Energy limitation and lipid-water fractionation in Archaeal versus Bacterial lipids

Besides the shared biochemical mechanism underlying negative $\varepsilon_{L/W}$ in eukaryotic and archaeal

lipids, another interesting pattern emerges from comparison of $\varepsilon_{L/W}$ across the three domains of

life: a subset of bacterial lipids resembles archaeal lipids in its consistently negative $\varepsilon_{L/W}$ values

(Fig. 6). For bacterial lipids, anaerobiosis and/or autotrophy appear to be associated with

consistently large, negative values of $\varepsilon_{L/W}$. Examples include aerobic autotrophy (*Cupriavidus*

oxalaticus, Zhang et al., 2009; *Paracoccus denitrificans*, Osburn et al., 2016), anaerobic

autotrophy (*Desulfobacterium autotrophicum*, Campbell et al., 2009; *Desulfococcus multivorans*,

Dawson et al., 2015; *P. denitrificans*, *Desulfobacterium autotrophicum*, *Desulfovibrio alaskensis*,

Osburn et al., 2016), and anaerobic heterotrophy (*D. multivorans*, Dawson et al., 2015; *P.*

denitrificans, *Escherichia coli*, *D. autotrophicum*, *Desulfobacter hydrogenophilus*, *D. alaskensis*,

Desulfolobus propionicus, Osburn et al., 2016) (Fig. 6; Fig. 7B). Of these modes of metabolism,

aerobic or anaerobic autotrophy has a more constrained range of $\varepsilon_{W/L}$ and more overlaps with $\varepsilon_{L/W}$

observed in archaeal lipids, compared to anaerobic heterotrophy (Fig. 6; Fig. 7B). We propose that

NADPH flux imbalances associated with energy limitation that broadly characterize autotrophic

and/or anaerobic modes of metabolism contribute to the apparent similarity in $\varepsilon_{L/W}$ observed

between the domains of Bacteria and Archaea.

The concept of NADPH flux imbalance set forth by Wijker et al. (2019) provides a quantitative

framework to relate metabolism-dependent patterns of $\varepsilon_{L/W}$ observed in aerobic heterotrophs to the

fluxes of H through specific dehydrogenase and transhydrogenase enzymes. For the highly

conserved fatty acid biosynthesis pathway, NADPH serves as the reducing agent for the reduction

of acyl intermediates (Wakil et al., 1983). As a result, NADPH accounts for ~50% of fatty acid H

(Sessions et al., 1999). Wijker et al. (2019) showed that $\varepsilon_{L/W}$ values of fatty acids produced by five

wild-type species of aerobic heterotrophic bacteria are positively correlated with relative NADPH

imbalance flux. Among the tested aerobic heterotrophs, two species (*E. coli* and *B. subtilis*) had

moderately negative values of NADPH flux imbalance (–32 and –11 %, respectively) (Wijker et

al., 2019). Here, negative values reflect underproduction of NADPH where catabolic production

of NADPH does not meet the anabolic requirement (Wijker et al., 2019). Furthermore, the results

from *E. coli* knockout mutants confirmed the direction of isotope fractionation for specific

transhydrogenases (e.g., ^{2}H-depletion by NADPH-producing PntAB and ^{2}H-enrichment by

NADPH-consuming UDhA) (Wijker et al., 2019).

For microbes adapted to chronic energy limitation, even more negative values of NADPH flux

imbalance would be expected where just enough NADPH is produced such that most to all of it is

used for biosynthesis (i.e., little to no flux from NADPH back to NADH). This prediction is

consistent with the general concept of energy spilling (also known as growth uncoupling or

overflow metabolism), which suggests that microbes under energy limitation tend to have more

efficient biomass production (Russell, 2007). For example, *E. coli* grown aerobically resulted in

more than 9-fold increase in ATP production per mole of glucose compared to the anaerobic

counterpart; however, the biomass yield was only 2-fold greater during aerobic respiration

(Stouthamer and Bettenhaussen, 1977; Gottschalk, 1986). Similar phenomenon of energy spilling

has also been observed in methanogenic archaea. For example, *Methanothermobacter*

thermautotrophicus and *Methanococcus maripaludis* produced less biomass per mole of methane

when H_2 (electron donor) was present in excess compared to H_2-limiting condition (Morgan et al.,

1997; De Poorter et al., 2007; Costa et al., 2013).

In line with the observed and expected outcomes of energy limitation (negative NADPH flux

imbalance and tighter coupling between catabolism and anabolism), Leavitt, Kopf et al. (2023)

found that $\varepsilon_{L/W}$ values for *N. maritimus* fall within the 95% confidence interval for the extension

of Wijker et al. (2019) regression when an extremely negative value (–100 %) of NADPH flux

imbalance was assigned. The apparent fit of *N. maritimus* $\varepsilon_{L/W}$ on the negative end of the

framework is reasonable, given that *N. maritimus* is an obligate autotroph adapted to oligotrophic

nitrification with energy-efficient CO_2 fixation pathway (Martens-Habbena et al., 2009; Könneke

et al., 2014). The fact that all archaeal lipids investigated for $\varepsilon_{L/W}$ so far fall within a relatively

narrow range of negative values (Fig. 6; Fig. 7B) may, at least partially, reflect the general concept

that Archaea are optimized for chronic energy limitation (Valentine, 2007). The results from this

study show that, while metabolic modes and growth rates have significant impacts on $\varepsilon_{L/W}$ within

a single organism (*A. fulgidus*), the observed variation in $\varepsilon_{L/W}$ still falls within the overall range

observed across archaeal lipids (e.g., heterotrophic and autotrophic *A. fulgidus* 'Af' data *vs*. other

archaeal lipids data; Fig. 7B).

The similarity between archaeal lipids and a subset of anaerobic and/or autotrophic bacterial lipids

is striking (Fig. 6; Fig. 7B) given the fundamental differences in fatty acid and isoprenoid lipid

biosynthesis pathways. This suggests that factors controlling microbial $\varepsilon_{L/W}$ may not be domain-

specific; instead, common challenges associated with anaerobic respiration and/or autotrophy (i.e.,

lower free energy yields and/or additional energetic costs for CO_2 fixation) may become primary

controlling factors of $\varepsilon_{L/W}$ for both Bacteria and Archaea. Notably, the range of $\varepsilon_{L/W}$ in fatty acids

produced by sulfate-reducing bacteria (SRB) is much narrower compared to that observed in

aerobic heterotrophs regardless of metabolic pathway or substrate (Fig. 7B).

One of the transhydrogenases that plays an important role in SRB is NfnAB, and it has been shown

that this enzyme is important for H isotope fractionation in *D. alaskensis* (Leavitt et al., 2016). The

details of the magnitude of $\varepsilon_{L/W}$ set by NfnAB under severe NADPH underproduction may be

complicated by other processes such as intracellular isotope distillation (i.e., more quantitative

transfer of NADPH hydride to lipids) as suggested by Leavitt et al. (2016). Regardless, given that

decoupling of NADPH production for anabolism from NADH for catabolism becomes more

important during energy limitation, transhydrogenases that are widely distributed among anaerobic

bacteria (namely NfnAB; Pereira et al., 2011; Leavitt et al., 2016; Buckel and Thauer, 2018) may

be key contributing factors for the negative $\varepsilon_{W/L}$ observed across these organisms (Fig. 6; Fig. 7B).

It is worth noting that *nfnAB* was not found in archaeal species investigated by Pereira et al. (2011),

including *A. fulgidus*.

A complex interplay among environmental, physiological, and enzymatic factors underlies net

isotope fractionation patterns observed in a biosynthetic product such as $\varepsilon_{L/W}$. For archaeal lipids

and a subset of bacterial lipids, however, two factors appear to play key roles in shaping the

consistently negative $\varepsilon_{L/W}$ patterns observed in the lipids from both domains: highly fractionating

enzymatic steps (e.g., GGR in archaeal isoprenoid lipid biosynthesis; transhydrogenases in

bacterial fatty acid biosynthesis) and energy limitation (i.e., NADPH flux imbalance). The

negative $\varepsilon_{L/W}$ commonly observed in archaea and a subset of bacteria is likely a convergent feature

resulting from a combination of both factors. A better mechanistic understanding of factors

conducive to similar patterns of $\varepsilon_{L/W}$ in bacterial and archaeal lipids could lead to biogeochemical

and/or ecological insights that are unique to microbial life. This would, in turn, significantly

expand the application of lipid biomarkers as paleoenvironmental or paleoecological proxy

currently possible with eukaryotic or bacterial lipids alone.

6. Conclusion

In this study, we conducted water isotope label experiments with a metabolically flexible and well-

studied model archaeon *A. fulgidus* and quantified the hydrogen isotope fractionation between

lipids and water in response to different carbon substrates and electron donor-acceptor pairs. We

observed changes in regression parameters among experimental conditions, where higher slopes

($f_W \cdot {}^2\alpha_{L/W}$) were observed in autotrophic compared to heterotrophic conditions and during

stationary compared to exponential growth phase. The bio-isotopic model further constrains

individual parameters such as fluxes that reflect water isotope composition (i.e., not affected by

substrate isotope composition) (f_W) and isotope fractionation between lipids and water ($\varepsilon_{L/W}$) that

together make up the slope. These results are consistent with observations from regression

parameters, where f_W is larger under autotrophic condition and during stationary growth phase

compared to the heterotrophic conditions and exponential growth phase. The model results also

allow us to constrain plausible ranges of isotope fractionation associated with a specific route of hydrogen flux (e.g., flux for direct water incorporation *vs*. flux from water through cofactors). The model results point toward the isotope fractionation associated with the $f_{W,cofa}$ flux, which includes the hydride transfer during the final saturation via GGR, as a highly fractionating step. This is consistent with recent bio-isotopic model results for another archaeon *N. maritimus*. Finally, we synthesized a three-domain comparison of $\varepsilon_{L/W}$ patterns which suggests that the highly fractionating double bond reduction via GGR may also explain the similarly large and negative values of $\varepsilon_{L/W}$ observed between eukaryotic and archaeal lipids. Furthermore, the shared negative $\varepsilon_{L/W}$ values between fatty acids in anaerobic and/or autotrophic bacteria and isoprenoids in archaea may be associated with the general state of energy limitation experienced by the different groups of organisms. Additional experimental work with archaea that use carbon metabolism pathways that have not been examined with the bio-isotopic model (i.e., 3HP/4HB and reductive Wood-Ljungdahl pathways) as well as more systematic comparisons between bacterial and archaeal lipids will help determine the scope of applicability for microbial lipid $\varepsilon_{L/W}$ in paleoenvironmental and paleoecological studies.

7. Data availability

Access to all data products and code for the isotope flux-balance model for this manuscript are available at: https://github.com/jrhim/Rhim_et_al_2024_GCA.

8. CRediT authorship contribution statement

Jeemin H. Rhim: Conceptualization, Methodology, Investigation, Data Curation, Writing – Original Draft, Writing – Review & Editing, Visualization. **Sebastian Kopf:** Conceptualization,

Methodology, Resources, Writing – Review & Editing, Supervision, Funding acquisition. **Jamie**

McFarlin: Investigation, Writing – Review & Editing. **Ashley E. Maloney:** Investigation,

Writing – Review & Editing. **Harpreet Batther:** Investigation, Writing – Review & Editing.

Carolynn M. Harris: Investigation, Writing – Review & Editing. **Alice Zhou:** Investigation,

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Weber: Investigation, Writing – Review & Editing. **Shelley Hoeft-McCann:** Investigation,

Writing – Review & Editing. **Ann Pearson:** Resources, Writing – Review & Editing, Supervision,

Funding acquisition. **William D. Leavitt:** Conceptualization, Resources, Writing – Review &

Editing, Supervision, Funding acquisition.

823   9.   Declaration of competing interest

824   The authors declare that they have no known competing financial interests or personal

825   relationships that could have appeared to influence the work reported in this paper.

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11. Appendix A. Supplementary material

These include the Overview of hydrogen budget, Culturing conditions for *Acidianus* sp. and *Metallosphaera sedula*, Full list of references for Figure 7, and Supplemental Figures 1–2.

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