Title

Ancient tropical extinctions contributed to the latitudinal diversity gradient

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Running title

Asymmetric gradient of tropical extinction

Data accessibility statement

All the data used in this manuscript are presented in the manuscript and its supplementary

material or have been published or archived elsewhere.

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Abstract

Biodiversity currently peaks at the equator, decreasing towards the poles. Growing fossil evidence suggest that this hump-shaped latitudinal diversity gradient (LDG) has not been persistent through time, with similar species diversity across latitudes flattening out the LDG during past "greenhouse" periods. This provides a new starting point for LDG research. Most studies assume the processes shaping the LDG have acted constantly through time and seek to understand why diversity accumulated in the Holarctic at lower levels than at the equator, *e.g.* as the result of limited dispersal, or higher turnover in Holarctic regions. However, fossil investigations suggest that we need to explain when and why diversity was lost at high latitudes to generate the LDG. Unfortunately, diversity lost scenarios in the Holarctic have been repeatedly proposed but not yet clearly demonstrated. Here, we outline the 'asymmetric gradient of extinction' (AGE) framework, which contextualize previous ideas behind the LDG in the frame of a time-variable scenario. We suggest the current steep LDG may be explained by the extinction of clades adapted to warmer conditions from the new temperate regions formed in the Neogene, together with the equator-ward dispersal of organisms tracking their own climatic preferences, when tropical biomes became restricted to the equator. Conversely, high rates of speciation and pole-ward dispersal can account for the formation of an ancient flat LDG during the Cretaceous–Paleogene greenhouse period. Phylogenies and fossils of the Testudines, Crocodilia and Lepidosauria support the AGE scenario and showed the LDG to have varied over time, with high latitudes serving as a source of tropical diversity but suffering disproportionate extinction during transitional periods to cold climate. Our results demonstrate that the inclusion of fossils in macroevolutionary studies allows detecting extinction events less detectable in analyses restricted to present-day data only.

50   **Keywords:** climate change; fossils; biodiversity; dispersal; extinction; Holarctic; niche;

51   phylogeny; speciation; tropics.

Introduction

The current increase in species richness from the poles towards the equator, known as the latitudinal diversity gradient (LDG), is one of the most conspicuous patterns in ecology and evolution. This pattern has been described for microbes, insects, vertebrates, and plants, and for marine, freshwater, and terrestrial ecosystems[1–6].

For decades, it has been thought that the modern-type steep LDG (with higher diversity at the equator) persisted throughout the Phanerozoic (the last 540 million years), even if the gradient was sometimes shallower[7], based on published fossil record studies[8,9]. However, the methodological limitations of fossil sampling have called this conclusion into question. Analyses controlling for sampling bias have suggested that, for many groups, the LDG was less marked in the past than it is today, flat (*i.e.* with similar species diversity across latitudes) or even developed a paleotemperate peak during some periods in the past (see Mannion *et al.*, (2014) for a review). This sampling-corrected flatter LDG in deep time has been demonstrated for non-avian dinosaurs[11], mammals[12,13], tetrapods[14], insects[15–17], brachiopods[18–20], bivalves[21], coral reefs[22], crocodiles[23], turtles[24,25], and plants[26–28]. The pattern emerging from fossil studies also suggests that steep LDGs, such as that currently observed, have been restricted to the relatively small number of short "coldhouse" periods during the history of the Earth: the Ordovician/Silurian, the Carboniferous/Permian, the end of the Jurassic, and the Neogene. Most of the Phanerozoic has instead been characterized by warm greenhouse climates associated with a flatter LDG1 (**Fig. 1**).

This recent fossil evidence now provides a new starting point for LDG research. Most hypotheses on the origin of the LDG are based on the assumptions that *(i)* the processes shaping the LDG have acted constantly through time and *(ii)* equatorial regions are the source of world diversity[29,30], with the LDG resulting from lower levels of diversity accumulation in the Holarctic than at the equator through time[7,31,32]. Previous studies have explained the LDG

Figure 1 | Changes in global temperature and extension of the tropical belt during the Cenozoic, in relation with the shape of the LDG. Early Cenozoic global temperatures were higher than today and paratropical conditions extended over northern and southern latitudes. From the early Eocene climatic optimum (EECO; ca. 53-51 Ma), a global cooling trend intensified on Earth and culminated with the Pleistocene glaciations. Global cooling was punctuated by sharp declines of temperatures, such as the terminal Eocene Event (TEE, ca. 33.9 Ma) and periods of relative warmth. Warm-equable regimes got then restricted to the equator. The LDG evolved following these global changes; during greenhouse periods diversity was uniform across latitudes, such that the LDG flattened, while in cold periods diversity peaked at the equator (a steep LDG) (Mannion et al., 2014). The question mark denotes the focus of this study, which is to unveil the processes that mediated the transition between a flat and steep LDG. The relative temperature curve of the Cenozoic is adapted from Zachos et al., (2008). Maps represent the extension of the tropical belt and Earth tectonic changes as derived from Ziegler et al., (2003) and Morley (2007). Abbrev.: P=Pleistocene, Pli=Pliocene

as a result of greater tropical diversification and limited dispersal out of the equatorial

region[7,31,33], or by high rates of turnover in the Holarctic (*i.e.* similar high speciation (λ) and

extinction (μ) rates; $\lambda \approx \mu$), all keeping diversity levels in this region low over time (**Table 1**),

for amphibians[34,35], birds[32,36,37], butterflies[38], conifers[39], fishes[40], mammals[33,37], and

lepidosaurs[41], for example. However, the recent fossil investigations showing, for many

lineages, similar high diversity levels in the Holarctic and at the equator in the past suggest

we do not necessarily need to explain why diversity accumulated at slower rates in the

Holarctic through time, but the question being how and when diversity was lost at high

latitudes, giving rise to the current shape of the LDG[10]?

Diversity losses in the Holarctic have been traditionally considered to underlie the

LDG[37]. They were initially attributed to Pleistocene glaciations[42], but this hypothesis has been

called into question by the finding that the LDG substantially predates the Pleistocene[7]. More

ancient extinctions have also been considered[43–49]. For example, Hawkins *et al.* (2006)[46]

suggested that the avian LDG resulted from the differential extirpation of older warm-adapted

clades from the temperate regions newly formed in the Neogene. Pyron (2014)[41] suggested

that higher temperate extinction represents a dominant force for the origin and maintenance of

latitudinal gradients. More recently, Pulido-Santacruz & Weir (2016)[49] proposed the

terrestrial LDG is largely the effect of a post-Eocene increase in extinction rates at high

latitudes resulting from the cooling Cenozoic trend. Unfortunately, using phylogenies alone,

these and other studies on the LDG have not clearly demonstrated diversity losses in the

Holarctic but instead high regional turnover[33,35–39,50]. Nonetheless, high turnover can only

explain a slow accumulation of lineages, with one fauna being replaced by another, but does

not explain diversity decline (*i.e.* a reduction in the net number of species). Diversity declines

occur when extinction exceeds speciation, resulting in negative net diversification rates

($r = \lambda - \mu$; $r < 0$). Accordingly, 'diversity loss' hypotheses need to be differentiated in the

Table 1. Predictions of the most common LDG hypotheses, including the *Asymmetric gradient of tropical extinction* (AGE) hypothesis proposed in this study. The main evolutionary hypotheses published to date can be classified according to three criteria: **(1)** the mechanisms behind regional differences in species richness[7], including explanations based on evolutionary time, on dispersal *(d),* and on diversification *(r=λ, μ)*. Explanations based on evolutionary time assume that most groups originated in the tropics and had less time to diversify in the temperate regions[90], but regional differences in evolutionary rates are not invoked. Hypotheses focussing on the role of geographic movements *(d),* include the "*tropical niche conservatism*" model, assuming that most groups originated in the tropics and the LDG results from limited dispersal to the temperate regions, as only few of tropical species succeeded to adapt to the temperate regimes[30]. The "*Into the tropics*" model assumes instead the LDG results from dispersals towards the equator[38,50]. Hypotheses that emphasize the LDG is generated by regional differences in net diversification rates (the composite value $r = \lambda - \mu$), being higher in the tropics[32,34] assume that the outstanding tropical diversity could be the outcome of higher rates of speciation in the tropics than in the extra-tropical regions ($\lambda_t > \lambda_e$) under the "*cradle of diversity*" hypothesis, and/or could result from lower rates of extinction ($\mu_t < \mu_e$) under the "*museum of diversity*" hypothesis[66]. The LDG could also result from higher turnover rates (*i.e.* higher λ and μ) in the Holarctic[37]. Diversification and dispersal hypotheses are not mutually exclusive. The "*out of the tropics*" model synthesizes this idea that the tropics are regarded as both a cradle and a museum, with lineages preferentially originating in the tropics and expanding over time into high latitudes[29,65]. In addition, evolutionary hypotheses could be classified according to **(2)** the rate at which processes acted through time; most studies assumed evolutionary processes acted constantly through time to explain the LDG. The AGE model, conversely, includes various diversification and dispersal parameters for each temporal interval (the transition from coldhouse to greenhouse, and *viceversa*). **(3)** Finally, the hypotheses can be classified according to the source of tropical diversity: "*Lower Holarctic species accumulation*" hypotheses assume that the equator is the source of world diversity and species accumulated at slower rates on the higher latitudes. Conversely, "*Holarctic diversity loss*" hypothesis assumes the Holarctic was also a source of diversity but this diversity was lost at some periods of the evolutionary history (e.g. climate change from greenhouse to coldhouse).

1. mechanism	Time (*t*)	Differences in *r*			Differences in *d*			Both	
2. rate	Time-constant models							Time-variable models	
3. source	Lower Holarctic species accumulation							Holarctic diversity loss	
Hypothesis (reference)	*Time for speciation*[90]	*Cradle of diversity*[66]	*Museum of diversity*[66]	*Holarctic turnover*[37]	*Into the tropics*[38,50]	*Tropical niche conservatism*[32]	*Out of the tropics*[29,65]	*Asymmetric gradient of extinction*	
								Coldhouse to greenhouse transition	**Greenhouse to coldhouse transition**
Predictions	$t_{trop} > t_{temp}$ $(r_{trop} = r_{temp})$ $(d_{trtp} = d_{tptr})$	$r_e > r_h$ $(\lambda_e > \lambda_h)$ $d_{eh} = d_{he}$	$r_e > r_h$ $(\mu_e < \mu_h)$ $d_{eh} = d_{he}$	$r_h = 0$ $(\mu_h = \lambda_h)$ $d_{eh} = d_{he}$	$r_e = r_h$ $d_{eh} < d_{he}$	$r_e = r_h$ $d_{eh} > d_{he}$	$r_e > r_h$ $(\lambda_e > \lambda_h)$ $(\mu_e < \mu_h)$ $d_{eh} > d_{he}$	$r_e < r_h$ $(\mu_h < \lambda_h; r_h > 0)$ $d_{eh} > d_{he}$	$r_e > r_h$ $(\mu_h > \lambda_h; r_h < 0)$ $d_{eh} < d_{he}$

Abbreviations: λ = speciation, μ = extinction, d = dispersal, t = time, e = equatorial, h = Holarctic, trop = Tropical biome, temp = Temperate biome.

literature from 'high turnover' scenarios. The perceived difficulty for inferring negative

diversification rates from present-day phylogenetic data[51,52] and the assumption that diversity

levels were always lower in the Holarctic than at the equator have resulted in 'diversity loss'

hypotheses being repeatedly proposed but seldom demonstrated. Meanwhile, numerous fossil

investigations have detected signatures of extinction and diversity loss in the Northern

Hemisphere. For instance, Archibald *et al.*[15,16] sampled insect diversity at an Eocene site in

Canada, and in present-day temperate Massachusetts (USA) and tropical sites of Costa Rica.

Insect diversity was higher at the Eocene paleotropical site than the modern temperate

locality, and comparable to the modern-day tropical locality, suggesting that post-Eocene

insects have thus suffered greater levels of extinction in the Nearctic regions than around the

equator. This pattern is consistent with results of other studies on various taxonomic groups,

including birds[53], invertebrates[15,16,54], mammals[12,13,55] and plants [56–58]. However, fossil studies

are generally restricted to a geographic and temporal scale, which makes difficult to

extrapolate local inferences of extinction in the context of the LDG.

Here we capitalize on the current knowledge and we aim at: *(1)* outlining a new

framework for studying the LDG, the 'asymmetric gradient of extinction' (AGE) hypothesis,

which formalizes and expands previous hypotheses about the contribution of extinctions to

the LDG in the context of a time-variable pattern; and *(2)* testing the predictions of the AGE

hypothesis with three tetrapod groups.

From a flat to a hump-shaped LDG: an asymmetric gradient of extinction

Over the Earth's history, the geographic extent of the tropical biome around the equator has

fluctuated, with periods of pole-ward expansion during which warm paratropical conditions

appeared at high latitudes, followed by periods of equator-ward contractions[59–63] (**Fig. 1**). To

account for these biodiversity changes, we propose to include a temporal component to study

the LDG in which prevailing speciation, extinction and dispersal dynamics change between

warm and cold intervals. The AGE hypothesis makes the following main postulates (P) during

climatic transitions towards cooler climates (**Fig. 2; Table 1**): (**P1**) extinction rate exceeds the

speciation rate at high latitudes ($_H$), *i.e.* declining diversity ($r_H < 0$), while the equatorial ($_E$)

regions remain stable, and (**P2**) the rate of dispersal '*into the equator*' is greater than that '*out*

of the equator' ($d_{HE} > d_{EH}$), thus triggering the formation of a steep LDG. Conversely,

during the transition to greenhouse periods, different processes drove the formation of a flat

gradient, and AGE postulates: (**P3**) diversification rates increase in the new paratropical

regions ($r_E < r_H$), and (**P4**) '*out of the tropical*' dispersals increase over '*into the equator*'

migrations ($d_{HE} < d_{EH}$).

For instance, the last 100 million have witnessed the contraction of tropical conditions

towards the equator, due to the global cooling since the latest Cretaceous–early Cenozoic

period (the most recent greenhouse period), culminating in the Pleistocene glaciations[64].

According to the AGE hypothesis, the expansion of tropical-like conditions to higher latitudes

induced species diversification in the new paratropical areas (**Fig. 2a**) and facilitated

movements within the broad 'tropical belt', such that tropical equatorial clades were able to

disperse '*out of the equator*' into high-latitude warm regions[29,65]. An equable Cretaceous-

early Cenozoic greenhouse climate thus triggered the formation of a flat LDG (**Fig. 2a**). By

contrast, tropical biome retractions following the climate cooling induced periods of declining

diversity at high latitudes (where climate change was more intensively felt), and initiated

biotic movements '*into the equator*' (**Fig. 2b**). Extinction rates were high for tropical-adapted

lineages at high latitudes, but lower for low-latitude tropical lineages. Climate change would

thus have driven the development of an asymmetric gradient of extinction (AGE) and

dispersal within the tropical biome, and mediated the formation of a steep LDG (**Fig. 2c**).

Figure 2 | Prevalent evolutionary processes behind the latitudinal diversity gradient under the AGE time-variable framework. The graphic shows the hypothetic change in evolutionary dynamics (diversification "r", speciation "λ", extinction "μ" and dispersal "d") between Holarctic and equatorial regions (denoted with "h" and "e" subscripts, respectively) and across main climatic intervals: **a**, the greenhouse late Cretaceous-early Cenozoic period. **b**, the late Eocene-Oligocene climatic transition. **c**, the Neogene coldhouse interval. **d**, the past and future transition to greenhouse climates. For each period, inset Figs. represent the distribution of species richness across latitudes (the LDG shape) and the hypothetical change in global evolutionary dynamics under the AGE hypothesis.

The AGE hypothesis attempts to reconcile previous contending hypotheses on the

origin of the LDG by placing them in a temporal scenario (**Table 1**). For instance, there is

controversial support around the tropics being '*cradle*' or '*museum of diversity*' [66], and

dispersal prevailing '*out of* [29,65] or '*into the tropics*'[38,41,50]. The AGE hypothesis alternatively

invokes the '*museum of diversity*' regarding the equatorial tropics as refuge during coldhouse

transitions, but also the '*cradle of diversity*' during greenhouse periods. Similarly, the

hypothesis invokes '*out of the equator*' dispersals during greenhouse transitions and '*into the*

equator' dispersals during coldhouse transitions. The AGE hypothesis also assumes tropical

niche conservatism due to physiological limits[30,31] as we postulate that most of the tropical-

adapted taxa at high latitudes were unable to adapt and either went extinct or suffered

restrictions of their distributions when the tropical biome retreated towards the equator.

Accordingly, the AGE hypothesis focuses on the fate of tropical-adapted clades under

variable climate conditions but does not apply for groups having evolved the appropriate

adaptations to cope with climate change, e.g.[67,68].

Here, we used comparative methods for both phylogeny-based and fossil-based

datasets to test the predictions of the AGE hypothesis for the Testudines, Crocodilia and

Lepidosauria. The modern-day Crocodilia and Lepidosauria comprise mostly tropical-adapted

species with a classic LDG pattern as shown by an accumulation of diversity at equatorial

latitudes[41,44]. We evaluated the applicability of our framework to subtropical taxa as well, by

extending the study to Testudines, a group displaying a hump-shaped gradient of diversity

today centred on subtropical latitudes (10°S–30ºN)[69]. By contrast, the paleolatitudinal

distribution of turtles was concentrated in the Holarctic (30–60ºN) during the Cretaceous[24,25].

All these lineages are ancient and likely experienced climatic transitions during the early

Cenozoic[23,25,41,44,69]. They display contrasting patterns of species richness: turtles and

crocodiles are species-poor (330 and 25 species, respectively), while lepidosaurs include a

large number of species (9500+ species) and have a rich fossil record extending back to the

Triassic (Early Cretaceous for crocodiles), providing information about the variation of

latitudinal species richness accumulation during their evolutionary history.

Results

Phylogeny-based diversification analyses: are diversification rates higher at the equator?

According to current distribution data, the species richness of turtles, lepidosaurs and

crocodiles peaks near the equator, with 84% of all extant species living in the tropics, only

15% living in temperate regions and 1% spanning both biomes. We classified each species

reported in the phylogeny (Supplementary Tables 1-3) as living close to the equator (the

modern-day tropical biome) or the Holarctic and Southern Hemisphere (the modern-day

temperate biome). For turtles, there were 239 tropical species, 84 temperate and 6 spanning

both biomes (7 were marine species). For lepidosaurs, there were 7955 tropical species, 1337

temperate and 124 spanning both biomes. The species-poor crocodile clade had only 23

tropical and two temperate species.

We analyzed differences in diversification rates between the Holarctic and equatorial

regions, with the binary state change speciation and extinction model 'BiSSE'[70,71], see

Methods. We did not use the geographic state change speciation and extinction model[72],

which is appropriate for dealing with widespread species, because most of the species in our

datasets were endemic to the Holarctic or equatorial regions, and, for a character state to be

considered in SSE models, it must account for at least 10% of the total diversity[73]. We did not

apply the BiSSE model to crocodiles, because simulation studies have shown that trees

containing fewer than 300 species may have to weak a phylogenetic signal to generate

sufficient statistical power[73].

We first used the time-constant BiSSE model, which is generally used in studies of the

LDG. For turtles, net diversification rates were higher in the Holarctic than at the equator

(**Table 2**, Supplementary Fig. 1a), but this difference was not significant, and rates of

dispersal '*into the equator*' were ten times higher than those '*out of the equator*'. For

lepidosaurs, a similar dispersal pattern was recovered, but net diversification rates were

significantly higher towards the equator (Supplementary Fig. 1b). We tested the AGE

hypothesis by introducing two shift times, at 51 and 23 million years ago (Ma), to detect

differences in diversification dynamics between greenhouse, transitional, and coldhouse

periods. This model indicated that the net diversification of turtles was similar in the Holarctic

and at the equator, whereas it was lower in the Holarctic for lepidosaurs until the coldhouse

period, when Holarctic diversification increased (**Table 2**, Supplementary Fig. 2). Dispersal

was considered to be symmetric between regions (*into the equator = out of the equator*)

during greenhouse periods, and asymmetric (*into the equator > out of the equator*) during the

climatic transition and coldhouse period. The same patterns were obtained for analyses with

the same model but with different combinations of shift times (51/66 Ma and 34/23 Ma;

Supplementary Fig. 3).

Fossil-based diversification analyses: evidence for ancient tropical extinctions?

We also analyzed differences in diversification rates between the Holarctic and equatorial

regions based exclusively on fossil data. The turtle fossil dataset comprised 4084 occurrences

for 420 genera (65 extant and 355 extinct; Supplementary Table 4). The lepidosaur fossil

dataset comprised 4798 occurrences for 638 genera (120 extant and 518 extinct;

Supplementary Table 5). The crocodile fossil dataset comprised 1596 occurrences for 121

genera (9 extant and 112 extinct; Supplementary Table 6). We first inferred global

diversification dynamics by analyzing the fossil datasets as a whole with a Bayesian approach

Table 2. Results of the diversification and biogeographic analyses performed in this study for the transition from greenhouse to coldhouse climates. Abbreviations: λ = speciation, μ = extinction, d = dispersal, $R\mu$ = range extirpations, e = equatorial, h = Holarctic.

	Data source	Model	Turtles	Squamates	Crocodiles
			Green- to coldhouse transition	**Green- to coldhouse transition**	**Green- to coldhouse transition**
Diversification analyses	Fossil	PyRate	$\lambda_e > \lambda_h$ $\mu_e < \mu_h$ $r_e > r_h$	$\lambda_e = \lambda_h$ $\mu_e > \mu_h$ $r_e < r_h$	$\lambda_e > \lambda_h$ $\mu_e < \mu_h$ $r_e > r_h$
	Present	BiSSE (time-variable)	$\lambda_e = \lambda_h$ $\mu_e = \mu_h$ $d_{he} > d_{eh}$	$\lambda_e > \lambda_h$ $\mu_e < \mu_h$ $d_{he} > d_{eh}$	-
	Present	BiSSE (constant)	$\lambda_e < \lambda_h$ $\mu_e = \mu_{h,}$ $d_{he} > d_{eh}$	$\lambda_e < \lambda_h$ $\mu_e < \mu_{h,}$ $d_{he} > d_{eh}$	-
Biogeographic analyses	Present	DEC	$R\mu_e < R\mu_h \,(R\mu_e < R\mu_h)^*$ $d_{he} < d_{eh} \,(d_{he} < d_{eh})^*$	$R\mu_e > R\mu_h \,(R\mu_e = R\mu_h)^*$ $d_{he} < d_{eh} \,(d_{he} = d_{eh})^*$	$R\mu_e = R\mu_h \,(R\mu_e = R\mu_h)^*$ $d_{he} = d_{eh} \,(d_{he} = d_{eh})^*$
	Present + fossil	DEC fossil	$R\mu_e < R\mu_h \,(R\mu_e < R\mu_h)^*$ $d_{he} > d_{eh} \,(d_{he} > d_{eh})^*$	$R\mu_e > R\mu_h \,(R\mu_e < R\mu_h)^*$ $d_{he} < d_{eh} \,(d_{he} > d_{eh})^*$	$R\mu_e < R\mu_h \,(R\mu_e < R\mu_h)^*$ $d_{he} > d_{eh} \,(d_{he} > d_{eh})^*$

* Prevalent dynamics when the number of events is calculated relative to the number of taxa currently distributed in each region

to inferring the temporal dynamics of origination and extinction rates based on fossil

occurrences[74] (see *Methods*). For turtles, origination rates peaked during the Jurassic,

subsequently decreasing until the present day. Extinction rates were generally low and

constant during the Mesozoic, but increased during the coldhouse periods of the Jurassic and

Paleogene, resulting in negative net diversification during the Paleogene (**Fig. 3**, **Table 2**,

Supplementary Figs. 4, 5). For lepidosaurs, origination rates peaked in the Jurassic and Late

Cretaceous, whereas extinction increased steadily until the Late Cretaceous. In the Paleogene,

net diversification approached zero, suggesting a high rate of turnover (**Fig. 3**, Supplementary

Figs. 6, 7). Crocodile origination peaked in the Early Cretaceous, subsequently decreasing

towards the present day, and extinction rates were generally low and constant. We also

identified diversity losses in the Paleogene extending to the present, suggesting that

crocodiles are still in a phase of declining diversity (**Fig. 3**, Supplementary Figs. 8, 9).

We performed additional analyses with different subsets of the three fossil datasets, to

separate the speciation and extinction signals between geographic regions (equator or

Holarctic) and ecological conditions (temperate or tropical, see *Methods*). These analyses

showed that the diversity losses experienced by turtles and crocodiles during the Paleogene

were mostly attributable to species living in the Holarctic and under tropical conditions (**Figs.**

4, **5, Table 2**). The global diversity loss inferred for crocodiles during the Neogene was

attributed to taxa living in both the Holarctic and equatorial regions (adapted to temperate and

tropical conditions respectively), providing further support for the hypothesis that this whole

group is in decline. For all groups, temperate taxa have been estimated to have high rates of

diversification during the Oligocene, but lower rates during the Neogene. For the equatorial

datasets, extinction and speciation rates decreased over time, resulting in constant net

diversification rates (except for lepidosaurs, which displayed a decrease in diversification

during the Paleogene, followed by an increase in diversification during the Neogene).

Figure 3 | Global pattern of turtles, squamates and crocodiles diversification through time based on the fossil record, and analysed with a Bayesian model. Origination (blue) and extinction (red) rates were estimated using time bins as defined by epochs of the geological timescale (on the top, main climatic periods are shown as follows: Greenhouse, Tran. = climatic transition, and C. = coldhouse). Solid lines indicate mean posterior rates, whereas the shaded areas show 95% credibility intervals. Net diversification rates (black) are the difference between origination and extinction. The vertical lines indicate the boundaries between geological periods. Tr, Triassic; J, Jurassic; K, Cretaceous; Pg, Paleogene; and Ng, Neogene.

Holarctic *vs.* equatorial

Figure 4 | Global pattern of turtle, squamate and crocodile diversification between Holarctic and equatorial regions, based on the fossil record. Diversification dynamics are compared between fossils distributed in Holarctic and equatorial regions. Origination (blue) and extinction (red) rates were estimated using time bins as defined by the main climatic intervals since the Mesozoic (on the top, climatic periods are shown as follows: Greenhouse, Tran. = climatic transition, and C. = coldhouse). Solid lines indicate mean posterior rates, whereas the shaded areas show 95% credibility intervals. Net diversification rates (black) are the difference between origination and extinction. The vertical lines indicate the boundaries between climatic intervals. Tr, Triassic; J, Jurassic; K, Cretaceous; Pg, Paleogene; and Ng, Neogene

Temperate *vs.* tropical

Figure 5 | Global pattern of turtle, squamate and crocodile diversification across temperate and tropical climates, based on the fossil record. Diversification dynamics are compared between fossils inhabiting under temperate and tropical macroclimates. Origination (blue) and extinction (red) rates were estimated using time bins as defined by the main climatic intervals since the Mesozoic (on the top, climatic periods are shown as follows: Greenhouse, Trans. = climatic transition, and C. = coldhouse). Solid lines indicate mean posterior rates, whereas the shaded areas show 95% credibility intervals. Net diversification rates (black) are the difference between origination and extinction. The vertical lines indicate the boundaries between climatic intervals. Tr, Triassic; J, Jurassic; K, Cretaceous; Pg, Paleogene; and Ng, Neogene.

Estimations of ancestral origins: did groups preferentially originate close to the equator?

We tested the predictions of the AGE hypothesis further, by performing biogeographic analyses with the dispersal-extinction-cladogenesis (DEC) model[75] and dated phylogenies (see *Methods*). We first analyzed the data in an unconstrained DEC analysis in which all ranges covering three areas could be in an ancestral state. We inferred an equatorial distribution for the deepest nodes for the turtles and lepidosaurs, whence these lineages colonized the other regions (**Fig. 6a**, **Table 2**, Supplementary Fig. 10). Crocodile ancestors were found to have been widespread during the Cretaceous, with an early vicariant speciation event separating *Alligator* in the Holarctic from the other Alligatoridae genera in equatorial regions (Supplementary Fig. 11).

Our biogeographic estimates based exclusively on extant data conflict with the fossil record[23,24,76]. We overcame this bias by introducing information about the distribution of fossils into DEC, in the form of hard (HFC) and soft (SFC) geographic fossil constraints at specific nodes (see *Methods*; Supplementary Tables 7–9). The inclusion of fossil information yielded very different biogeographic histories for the three groups (**Table 2**; turtles: **Fig. 6b**, Supplementary Fig. 12; lepidosaurs: Supplementary Figs. 13, 14; and crocodiles: Supplementary Figs. 15, 16). Under the SFC model, turtles were found to have originated in the Northern Hemisphere (under the HFC model they were spread over both regions), whence lineages migrated towards the equator and southern regions (**Fig. 6b**, Supplementary Fig. 12). Most dispersal therefore occurred '*into the equator*' (Supplementary Fig. 17, Supplementary Table 10). We also detected a larger number of geographic extinctions when fossil ranges were considered, predominantly for turtle lineages in the Holarctic (53 and 11 lineages disappeared from this region under the HFC and SFC models, respectively) and in southern temperate regions (9 in the HFC model; Supplementary Fig. 17, Supplementary

(a) Unconstrained biogeographical model

(b) Fossil-informed biogeographical model

Pleurodira

Cryptodira

Chelidae
Podoc.
Trionychidae
Cheloniidae
Emydidae
Testudinidae
Geoemydidae

Range dispersal event
Range extinction event
Range inheritance scenario
Ancestral range reconstruction:
Northern temperate regions
Equatorial and tropical regions
Southern temperate regions

Tr | J | K | Pg | Ng

Figure 6 | Biogeographic reconstruction of Testudines showing the effects of the incorporation of fossil information into biogeographic inference. **a**, Biogeographic reconstruction inferred with DEC based on the distribution of extant taxa. **b**, Biogeographic reconstruction under the fossil-informed HFC (*hard fossil constraint*) model. Coloured circles at tips and nodes represent current and ancestral ranges, respectively, while squares represent range inheritance scenarios. Colours correspond with the discrete areas in the legend. Black circles indicate fossil range constraints included in the analysis, with numbers corresponding with taxa in Table Supplementary S7. The reconstruction under the *soft fossil constraint* (SFC, see text) model is presented in Supplementary Fig. 12.

Table 11). The same trend was observed when the number of extinction/dispersal events was

controlled for the number of lineages currently distributed in each region (**Fig. 7**).

Lepidosaurs originated in both regions in both SFC and HFC analyses (Supplementary

Figs. 13, 14). During the greenhouse period, dispersal '*into the equator*' occurred at the same

rate (or at a higher rate in the HFC model) than dispersal '*out of the equator*', and dispersal

'*out of the equator*' prevailed thereafter (Supplementary Fig. 17, Supplementary Table 10).

Estimated range extinction rates were high in this group under the unconstrained model, with

30 lineages extirpated from the Holarctic, two from southern temperate regions and 152 from

the equator (Supplementary Fig. 17, Supplementary Table 11). Under fossil-informed models,

the number of Holarctic extinctions was higher (109 and 66 lineages in the HFC and SFC

models, respectively), whereas the number of lineages extirpated from the equator was similar

(144 and 109 in the HFC and SFC models, respectively; Supplementary Fig. 17). When the

number of events was controlled for the actual number of lineages distributed in each region,

the number of Holarctic extinctions and dispersals '*into the equator*' increased dramatically,

exceeding equatorial dispersal/extinctions (**Fig. 7**). For crocodiles, analyses including fossil

ranges showed that all the early nodes were distributed in the Holarctic (Supplementary Figs.

15, 16), and range extinctions were detected: four lineages disappeared from the Holarctic,

three from southern temperate regions, and two from the equator (HFC model; Supplementary

Fig. 17, Supplementary Tables 11, 12). Only two lineages disappeared from the Holarctic in

the SFC model. The same trends were observed after controlling the number of events for the

current number of lineages in each region (**Fig. 7**).

Discussion

Generation of the current LDG

Figure 7 | Estimated number of range-extinction and dispersal events through time. Analyses were performed for Testudines, Squamata and Crocodiles under the *unconstrained model* (Unc.), based on present evidence only, and the fossil-based *hard* (HFC) and *soft fossil constraint* (SFC) biogeographic models. **a**, Inferred number of range extinction events through time and across regions relative to the number of lineages currently distributed in each region. The global mean temperature curve is modified from Zachos et al., (2008). **b**, Inferred number of dispersal events from the Holarctic into the equator (IntoEq) and out of the equatorial zone (OutEq), relative to the current number of lineages distributed in the Holarctic and equatorial zones, respectively. Abbreviations: Tr, Triassic; J, Jurassic; K, Cretaceous; Pg, Paleogene; and Ng, Neogene, Trans. = climatic transition, and C. = coldhouse.

Fossil investigations have shown that at certain times during the Phanerozoic, the LDG has

weakened, flattened, or developed a palaeotemperate peak, with diversity at high latitudes

being greater in some periods of the past than currently for many groups[10,13]. Hypotheses

relating to 'slow Holarctic diversity accumulation', such as limited dispersal to the

Holarctic[31], high Holarctic turnover[37,41,49], or high rates of equatorial diversification[32–35,77],

cannot themselves account for the formation of a flatten LDG, or the transition from higher to

lower diversity in the Holarctic observed in many groups. Furthermore, although the

processes shaping biodiversity vary over time and space, this has been largely overlooked in

the context of the LDG, which has been generally explained in terms of the actions of time-

constant process.

We account for temporal changes in the global distribution of biodiversity by

proposing a scenario involving losses/gains of tropical diversity at high latitudes during

transitional periods from warm to cool conditions. The AGE hypothesis captures components

of previous studies[10,31,35,50,65] in the context of a time-variable LDG to disentangle the relative

contributions of speciation, extinction and dispersal for each particular geological period in

the formation of the LDG. Below, we evaluate the support of our analyses to the postulates of

the time-variable AGE hypothesis (**Fig. 2; Table 1**):

P1: Extinction exceeds speciation at high latitudes during cool transitions, *i.e.*

declining diversity ($r_H < 0$). Our diversification analyses based on extant species (time-

constant and time-variable BiSSE analyses) do not support this postulate, suggesting instead

higher levels of Holarctic diversification for turtles, and of equatorial diversification for

lepidosaurs. By contrast, results for fossil-only (PyRate) diversification analyses were

consistent with this prediction for turtles and crocodiles. We found that diversification rates of

turtles and crocodiles decreased in all regions during the transition to colder climates, but the

slowing of diversification was much stronger in the Holarctic than at the equator, with

extinction exceeding speciation in this region (**Fig. 3, 4**). This suggests that Holarctic

diversity loss ($r < 0$) during the Paleogene could explain the formation of a steep LDG for

these groups. For lepidosaurs, **P1** is not supported by fossil data; diversity losses occurred

during Cenozoic cooling in the equator but not at higher latitudes (**Fig. 4**). Diversity dynamics

for the species distributed at the equator, however, may not be entirely reliable, due to the

poverty of the equatorial dataset in terms of the number of fossil lineages and the small

number of records per lineage (Supplementary Table 12). Uncertainties therefore remain on

these estimates, which have wide credibility intervals, probably due to geographic biases in

the fossil record[78]. Turnover rates were very high in the Holarctic during the transitional

period, indicating that species did disappear from high latitudes, but that a new lepidosaur

community replaced them. This result suggests the number of lepidosaur species may always

have been unbalanced between regions. The high Holarctic turnover would contribute to the

maintenance of this pattern, together with the inferred temporal increases in diversification at

the equator (**Fig. 4**), as previously hypothesized [41].

P2: Higher '*into the equator*' dispersal than '*out of the equator*' dispersal

($d_{HE} > d_{EH}$) during cool transitions. The DEC biogeographic analyses based on extant

species do not support this postulate but instead the '*tropical niche conservatism*' hypothesis

for turtles, with an equatorial origin and recent invasion of high-latitude regions, resulting in

less time for lineages to diversify[31] (**Figs. 6, 7**, **Table 1, 2**). This result is consistent with the

findings of recent investigations[79,80]. For lepidosaurs, they support the '*out of the tropics*'

(**Fig. 7**, Supplementary Figs. 1-3, 10), and for crocodiles the diversification hypothesis, with

higher origination rates close to the equator and no effect on dispersal (**Fig. 7**, **Table 2**,

Supplementary Fig. 11). In contrast, time-constant and time-variable BiSSE analyses are

consistent with **P2** and the '*into the equator*' hypothesis to explain the LDG (Supplementary

Figs. 1a, 2a, 3a), as so they are the results for fossil-informed DEC biogeographic analyses;

all groups are inferred to have had a widespread ancestral distribution that subsequently contracted towards the equator due to both higher levels of range extirpations at higher latitudes and '*into the equator*' dispersals during Cenozoic cooling (**Figs. 6, 7**). This result is also in agreement with previous fossil investigations on turtles[24,25,76] and crocodiles[23,44]. For lepidosaurs, in absolute terms, more species migrated "*out of*" than "*into the equator*" (Supplementary Fig. 17), but the number of species in the equatorial region today is four times the number of lineages elsewhere. After controlling for the imbalance in species sampling in our tree, we found that a higher proportion of lepidosaur species lost their ancestral Holarctic distribution and emigrated '*into the equator*' [41] than the other way around (**Fig. 7**). Although the number of fossil constraints in the biogeographic analysis of lepidosaurs was relatively low given the size of the tree (30 Holarctic and equatorial fossils for 4161 nodes), these constraints significantly increased the absolute number of Holarctic range extinctions (from 30 to 109) and '*into the equator*' dispersals (from 40 to 124) relative to estimates without such constraints (Supplementary Tables 10, 11). Meanwhile, the inclusion of fossil data did not alter the number of events estimated for equatorial taxa. This finding suggests that a deeper understanding of lepidosaur fossil taxonomy might facilitate the assignment of fossils on the tree, and the detection of additional high-latitude extinctions not detected here.

Unfortunately, the age of the taxa evaluated here prevents us for testing the predictions associated with the transition from coldhouse to greenhouse conditions (**P3-4**). Nonetheless, our fossil-based analyses show similar diversification rates in the Holarctic and equatorial regions during the greenhouse period of the Cretaceous-early Cenozoic for all groups (overlapping credibility intervals; **Fig. 4**; Supplementary Figs. 1-3, 4-9), consistent with the idea of the existence of a flattened LDG during this phase [10]. Similarly, the AGE hypothesis focuses essentially on the Northern Hemisphere, but diversity losses and dispersals '*into the*

equator' may have also occurred in the temperate regions of the Southern Hemisphere.

Indeed, we found high rates of range extinctions in this region for all groups (**P2**; red lines on

Fig. 7). Unfortunately, the scarce fossil record prevents any diversification estimates for this

region (see *Methods*).

Overall, the general pattern that could be extracted from our study is that the AGE

hypothesis was supported for crocodiles and turtles using fossil and fossil-informed

phylogenetic investigations (**Table 1, 2**; **P1–P2** of AGE). However, if we rely only on

analyses based on data for extant species using biogeographic and constant and time-variable

BiSSE diversification models, this evolutionary scenario was poorly supported (although

BiSSE analyses support **P2** of AGE**)**. Support for the AGE hypothesis is mixed in

lepidosaurs; on the one hand the detected Holarctic range contractions are in agreement with

"higher Holarctic diversity loss" scenarios (**P2** of AGE), on the other hand, the evidence for

high Holarctic turnover are more in line with previous "slower Holarctic diversity

accumulation" hypotheses.

The timing and effect of the last greenhouse to coldhouse climatic transition

Recent fossil investigations suggest that changes in the shape of the LDG have been

associated with major climatic oscillations[10,13]. Accordingly, we hypothesized the

impoverishment of the Holarctic resulted from the contraction of the tropical biome during

the last greenhouse to coldhouse transition. But when this coldhouse transition took place?

The transitional period to cold was here defined between 51 and 23 Ma, after the early Eocene

Climatic Optimum (EECO), based on paleontological evidence showing that paratropical

conditions and the associated warm-adapted taxa disappeared from high latitudes between the

mid-late Eocene and the Neogene[59,61,62]. Our diversification results with time intervals defined

by the main climatic periods are consistent with these observations, and detect Holarctic

diversity loss during the late Paleogene. We cannot exclude, however, that diversity losses at

high latitudes occurred since the Cretaceous, as suggested by our fossil-based diversification

analyses with time intervals defined by the main geological periods (Supplementary Figs. 4,

6, 8) and by our fossil-based biogeographic analyses suggesting the prevalence of '*into the*

equator' dispersals since the Cretaceous (**Fig. 7**, Supplementary Table 11). These findings

could suggest that other processes different to climate change mediated the extinction and

range contraction of Holarctic lineages in the Cretaceous, or alternatively, that a transition

phase to cold started before the relatively short interval considered here. Some studies

consider the EECO only represented a transient temperature peak within an otherwise cooling

trend that started in the Cretaceous[64,81]. This trend was intensified by the Cretaceous-

Paleogene (K-Pg) mass extinction, and the drop in temperatures caused by the impact-

associated winter[82,83]. In our study, lineage extinctions, range extinctions and southward

dispersals increased between the K-Pg and Neogene (**Fig. 4, 7**), suggesting an additive effect

of K-Pg and Neogene cooling on depopulation of the Holarctic.

The ancestors of turtles, lepidosaurs and crocodiles were adapted to tropical conditions

during the Late Cretaceous[44,84]. Our results indicate that extinction events were not random,

instead preferentially affecting taxa living in tropical-like climates at high latitudes[48] (**Figs. 4,**

5). This suggests that many species adapted to warm conditions living in the Holarctic were

unable to adapt to the new temperate regimes and either went extinct or escaped extinction by

contracting their ranges in a southerly direction (**Fig. 7**). Meanwhile, we found that the

diversification rates of turtles, crocodiles and lepidosaurs living in temperate climatic

conditions were significantly higher than those of tropical-adapted taxa living in Holarctic and

equatorial regions after the transition to temperate climates in the late Eocene (**Fig. 5**, **Table**

2). The new temperate habitats could have constituted an opportunity for diversification

because they increased geographic ranges and ecological niches[34], and may have driven an

inverse LDG for some groups[39,68]. Several radiations following the appearance of the

temperate biome have been identified in other groups of organisms, such as plants [85,86],

mammals[87,88] or insects[89]. After this period, speciation decreased dramatically in the

temperate lineages of our focal groups, possibly due to the effect of the Pleistocene

glaciations, and no difference in diversification between tropical and temperate lineages is

currently evident (**Fig. 5**). In summary, our study suggests that differences in species richness

between geographic regions may be explained by differences in diversification and dispersal

rates. Differences in species richness between ecological types may be explained by the

longer time available for tropical-adapted clades to diversify in tropical areas[90] rather than

higher rate of speciation under warm tropical environments, as previously postulated [91].

Reconciling fossil and phylogenetic evidence

Our results unequivocally demonstrate that the inclusion of fossils in macroevolutionary

studies makes it possible to detect signals of ancient high-latitude extinctions and range

extirpations (**Figs. 4-7**), otherwise hardly detectable with analyses based exclusively on

present-day data. This conflict between extant and fossil evidence may extend beyond our

study, pervading the LDG literature. High extinction rates have occasionally been inferred in

tropical lineages[67,92–94], with hypotheses relating to extinction focused on temperate taxa and

recent time scales, such as the effects of recent Pleistocene glaciations, for example[33,37,41,42].

In reported cases of extinction, origination rates were also found to be elevated in high-

latitude groups (high turnover)[33,37–39,41,50], while diversity losses ($r < 0$) have to our

knowledge never been inferred in phylogenetic studies of the LDG (with the exception of the

recent Pulido-Santacruz & Weir (2016) study using time constant BiSSE models, but see

below). On the other hand, ancient tropical extinction at high latitudes is supported by fossil

studies on various taxonomic groups[15,16,28,53–56].

The last decade has seen many efforts to reconcile fossil and phylogenetic evidence.

Birth-death diversification models have been developed to detect negative diversification

rates on reconstructed phylogenies, or total-evidence trees[95–97]. Still, their use in the literature

is limited and these models are difficult to implement in a trait evolution context, such as in

the study of the LDG. LDG studies are often based on state-dependent speciation and

extinction models[35,38,40,41,49,50]. These models are designed to test differential diversification

and asymmetric transition scenarios, such as that suggested here, but LDG studies often

assume that diversification parameters remain constant over time. If the evolutionary

processes shaping the LDG have varied across latitudes and time, then time-constant models

are not appropriate for testing more complex scenarios underlying the LDG. Moreover, the

potential of time-constant models for detecting negative diversification rates is questionable,

since inferring negative diversification for the entire history of lineages conflicts with the fact

that these groups are still extant. Testing our hypothesis thus requires the implementation of

time-variable models. When applied to the study of diversity patterns, these models have

revealed marked extinction signatures in ancestral tropical plant clades[98]. The incorporation of

time-shifts into our BiSSE analyses improves but not completely reconciles the fossil

evidence with extant diversity. Identifying the causes of this problem and finding solutions

are beyond the scope of this study, but this artifact highlights the importance of fossils in

macroevolutionary inferences[99]. Fossil records remain incomplete, but they nevertheless

provide the only direct evidence of the diversity that existed in the past. By contrast to

molecular phylogenies, the incompleteness of the fossil record has a less problematic effect

on the estimation of speciation and extinction rates, because removing a random set of taxa

does not affect the observed occurrences of other lineages[74]. Indeed, simulations have shown

that PyRate correctly estimates the dynamics of speciation and extinction rates under low

levels of preservation or severely incomplete taxon sampling.

Conclusion

After decades of research, the processes shaping the LDG remain among the most hotly

debated topics in ecology and evolutionary biology. We propose here the AGE hypothesis,

which explains the origin of the current LDG through the changes in global diversification

and dispersal dynamics imposed by large-scale climatic transitions. Our analyses for turtles

and crocodilians indicated that the processes shaping the LDG have changed over time, the

current form of this gradient being the result of ancient high-latitude tropical diversity loss

and range contractions as a consequence of the retraction of the tropical biome and due to

climate cooling. The AGE hypothesis might account for the LDG of tropical-adapted groups

that were once diverse at high latitudes, but might not be fully applicable to all organisms

currently displaying a LDG, as shown here for lepidosaurs.

Methods

Time-calibrated phylogenies and the fossil record. We compared the predictions of the AGE

hypothesis with the LDG of three vertebrate groups: turtles (order Testudines), crocodiles

(order Crocodilia), and scaled lizards (order Lepidosauria). A time-calibrated phylogeny for

each group was obtained from published data. For turtles, we used the phylogeny of Jaffe *et*

al. (2011), including 233 species. We preferred this phylogeny over other more recent and

slightly better sampled trees[101] because the divergence time estimates of Jaffe *et al.* (2011) are

more consistent with recent estimates based on genomic datasets[79,102]. For lepidosaurs, we

retrieved the most comprehensive dated tree available, including 4161 species[41], and a

complete phylogeny was obtained for crocodiles[103].

Fossil occurrences were downloaded from the *Paleobiology Database*

(https://paleobiodb.org/#/, last accessed October 25th 2017). We reduced potential biases in

the taxonomic assignation of turtle, crocodile and lepidosaur fossils, by compiling occurrence

data at the genus level. The fossil datasets were cleaned by checking for synonymies between

taxa and for assignment to a particular genus or family on the basis of published results

(Supplementary Table 4–6).

Estimation of origination and extinction rates with phylogenies. We investigated possible

differences between Holarctic and equatorial regions, by combining the turtle and lepidosaur

phylogenies with distributional data (Supplementary Tables 1, 2) to fit trait-dependent

diversification models in BiSSE[70]. We accounted for incomplete taxon sampling in the form

of trait-specific global sampling fraction of extant species[104].

We ensured comparability with previous LDG studies, by initially using a constant-

rate trait-dependent diversification model. The constant-rate BiSSE model has six parameters:

two speciation rates (without range shift, or *in situ* speciation), one associated with the

Holarctic (hereafter 'H', λ_H) and the other with other equatorial and subtropical regions

(hereafter 'equator' or 'E', λ_E), two extinction rates associated with the Holarctic (μ_H) and the

equator (μ_E), and two transition rates (dispersal or range shift), one for the Holarctic to

equator direction (q_{H-E}), and the other for the equator to Holarctic direction (q_{E-H}).

We then assessed the effect of species distribution on diversification, allowing for rate

changes at specific time points. This approach is associated with a lower bias than the use of

constant rates. We used the time-dependent BiSSE (BiSSE.td) model, in which speciation,

extinction, and dispersal rates are allowed to vary between regions and to change after the

shift times. We introduced two shift times to model different diversification dynamics

between greenhouse, transitional, and coldhouse periods. We assumed that a global warm

tropical-like climate dominated the world from the origin of the clades until 51 Ma

(corresponding to the temperature peak in the Cenozoic). Thereafter, the climate

progressively cooled until 23 Ma (the transitional period), when the climate definitively

shifted to a temperate-like biome in the Holarctic[61,62,64]. The shift times at 51 Ma and at 23

Ma are initial values that are re-estimated by the model during the likelihood calculation. The

climatic transition in the Cenozoic may have different temporal boundaries, with potential

effects on the results. We thus applied the same model but with different combinations of shift

times (we tested 51/66 Ma and 34/23 Ma for the upper and lower bounds of the climatic

transition).

Analyses were performed with the R package *diversitree* 0.9-7[71], using the *make.bisse*

function to construct likelihood functions for each model from the data, and the functions

constrain and *find.mle* to apply different diversification scenarios. Finally, we used a Markov

Chain Monte Carlo (MCMC) approach to investigate the credibility intervals of the parameter

estimates. Following previous recommendations[71], we used an exponential prior $1/(2r)$ and

initiated the chain with the parameters obtained by maximum likelihood methods. We ran

10,000 MCMC steps, with a burn-in of 10%.

Estimation of origination and extinction rates with fossils. We also used fossil data to

estimate diversification rates over time. We analyzed the three fossil records, using a

Bayesian model for simultaneous inference of the temporal dynamics of origination and

extinction, and of preservation rates[74]. This approach, implemented in PyRate[105], uses fossil

occurrences that can be assigned to a taxon, in this case fossil genera. The preservation

process is used to infer the individual origination and extinction times of each taxon from all

fossil occurrences and an estimated preservation rate; it is expressed as expected occurrences

per taxon per million years.

We followed a birth-death shift approach[106], which focuses on the variation of

origination and extinction at a global scale and over large temporal ranges. We used a

homogeneous Poisson process of preservation (-mHPP option). We also accounted for the variation of preservation rates across taxa, using a Gamma model with gamma-distributed rate heterogeneity (-mG option). We used four rate categories to discretize the gamma distribution, to allow for a greater variability of preservation rates across taxa.

Given the large number of occurrences analyzed and the vast timescale considered, we dissected the birth–death process into time intervals, and estimated origination and extinction rates within these intervals. In one set of analyses we defined the time intervals using the geological epochs of the stratigraphic timescale[107] (Supplementary Figs. 4, 6, 8). In another set of analyses, we defined the intervals according to the different climatic periods characterizing the Cenozoic (Supplementary Figs. 5, 7, 9), as discussed above: the greenhouse world (Cretaceous), the climatic transition (Paleogene), and the coldhouse world (Neogene until the present). We adopted this solution as an alternative to the algorithms implemented in the original PyRate software for joint estimation of the number of rate shifts and the times at which origination and extinction shift[74]. The estimation of origination and extinction rates within fixed time intervals improved the mixing of the MCMC and made it possible to obtain an overview of the general trends in rate variation over a long timescale [106]. Both the preservation and birth–death processes were modeled in continuous time but without being based on boundary crossings. Thus, the origination and extinction rates were measured as the expected number of origination and extinction events per lineage per million years. One potential problem when fixing the number of rate shifts *a priori* is over-parameterization. We overcame this problem by assuming that the rates of origination and extinction belonged to two families of parameters following a common prior distribution, with parameters estimated from the data with hyper-priors[108].

We ran PyRate for 10 million MCMC generations on each of the 10 randomly replicated datasets. We monitored chain mixing and effective sample sizes by examining the

log files in Tracer 1.6[109]. After excluding the first 20% of the samples as a burn-in, we

combined the posterior estimates of the origination and extinction rates across all replicates to

generate plots of the change in rate over time. The rates of two adjacent intervals were

considered significantly different if the mean of one lay outside the 95% credibility interval of

the other, and vice versa. We looked at the marginal posterior distributions of origination and

extinction rates through the evolutionary history of the three groups and assessed the effect of

different environments.

In the context of the LDG, we performed additional analyses with different subsets of

fossils, to separate the speciation and extinction signals of different geographic regions

(equator or Holarctic) and ecological conditions (temperate or tropical). For example, for

turtles, we split the global fossil dataset into four subsets: one for the fossil genera occurring

at the equator (429 occurrences), one for the fossils occurring in the Holarctic (3568

occurrences), one for the fossil genera considered to be adapted to temperate conditions (993

occurrences), and one for the fossils considered to be adapted to tropical conditions (2996

occurrences). We excluded the few fossil occurrences for the southern regions of the South

Hemisphere (about 180) only in subset analyses, as they were poorly represented in our

dataset. Note that a given fossil can be present in both the 'Holarctic' and 'tropical' datasets.

We encoded tropical/temperate preferences by considering macroconditions in the Holarctic

to be paratropical until the end of the Eocene, as previously reported [61,62] (and references

therein). We also assumed that taxa inhabiting the warm Holarctic were adapted to tropical-

like conditions (*i.e.* a high global temperature, indicating probable adaptation to tropical

climates). This is, of course, an oversimplification that may introduce bias into the analysis,

but general patterns may nevertheless emerge from such analyses[110]. For turtles, this

assumption is supported by a recent study modeling the climatic niche of this group during

the Late Cretaceous, which found that the Holarctic ancestors of turtles were adapted to

tropical conditions[84]. After the late Eocene, we categorized each species as living in the

temperate biome or the tropical biome, according to the threshold latitudes defining the

tropics (23.4°N and 23.4°S) suggested in a previous study[33]. This delineation is also

consistent overall with the Köppen climate classification. With these datasets, we reproduced

the same PyRate analyses as for the whole dataset (see above). In general, the fossil datasets

included mostly Holarctic fossils, with a smaller number of occurrences for the equator.

Caution is therefore required when drawing conclusions from the equatorial datasets.

Inferring ancestral geographic distribution with phylogenies and fossils. We performed

biogeographic analyses with the parametric likelihood method DEC[75] using the fast C++

version[111] (https://github.com/rhr/lagrange-cpp). Turtle, lepidosaur, and crocodile species

distributions were obtained from online databases (www.iucnredlist.org and www.reptile-

database.org). We chose 23.4°N and 23.4°S as the threshold latitudes defining the tropics, and

categorized each species as living in the Holarctic, in the southern temperate regions, or in the

equatorial tropics and subtropical regions. We considered that all ranges comprising three

areas could be considered an ancestral state (*maxareas* =3).

We set up three different DEC analyses. We first ran DEC with no particular

constraints, using only the distribution of extant species. We then performed DEC analyses

including fossil information in the form of 'fossil constraints' at certain nodes, according to

the range of distribution of fossil occurrences assigned to a particular taxon during the

relevant time frame. For example the crown age of Carettochelyidae (Testudines) dates back

to the Late Jurassic (150 Ma, *node 5*, **Fig. 3**; Supplementary Table 7), and we set a constraint

on this node reflecting the distribution of all the Late Jurassic fossils attributed to

Carettochelyidae. Similarly, for the origin of turtles (210 Ma, *node 1*), distribution constraints

represent the range of Late Triassic fossils assigned to turtles. For the crown of Trionychidae,

in the Early Cretaceous (123 Ma, *node 2*), the early fossils assigned to the clade were used to

constrain the geographic origin of Trionychidae. In total, we implemented 23 fossil

constraints for turtles (Supplementary Table 7), 30 fossil constraints for lepidosaurs

(Supplementary Table 8), and 8 for crocodiles (Supplementary Table 9).

We included the fossil distribution in two different approaches: *(i)* a soft (SFC), and

(ii) hard fossil constraints (HFC). For the SFC approach, fossil data were incorporated into the

anagenetic component of the likelihood framework. The direct impact of a given fossil is

limited to the particular branch to which it has been assigned, although it may indirectly

influence other branches. The inclusion of a fossil conditions the estimated geographic-

transition probability matrix for that branch by imposing a spatiotemporal constraint on the

simulation process. Only the simulations resulting in a geographic range including the area of

fossil occurrence contribute to the geographic-range transition probability matrix for the

branch concerned; simulations not meeting this constraint are discarded[112]. For SFC, we used

the command '*fossil*' in DEC. We consider this to be a 'soft' constraint, because other areas

different from that in which the fossil was found could be included in the ancestral states. In

some cases, in which today's diversity is not representative of past diversity (*e.g.* due to

extreme levels of extinction), the SFC model may still overlook known fossil information. We

therefore also implemented an HFC model in which the estimation of ancestral areas was

fixed to the location of fossils. This was achieved with existing functions in the C++ version

of Lagrange, using the command '*fixnode*'. By fixing nodes to the distribution area of fossils,

we assume fossil occurrences reflect the distribution of the ancestors, *i.e.* that the fossil record

is complete. This is a strong assumption, but it makes it possible to recover all fossil ranges in

the ancestral estimations. The real scenario probably lies somewhere between the SFC and

HFC inferences.

We then compared the timing and number of range extinction and dispersal events

inferred with the three different biogeographic analyses. In DEC, range-subdivision

(inheritance) scenarios (vicariance, duplication and peripatric isolation) occur at cladogenetic

events, whereas extinction (range contraction) and dispersal (range expansion) are modeled as

stochastic processes occurring along the branches of the tree[113]. As the probability of any

extinction/dispersal event is constant along the entire length of the branch, we estimate the

periods at which range extinction and dispersal occurred by dividing the phylogeny into

intervals of 25 million years and calculating the number of branches for which

extinction/dispersal was inferred crossing a particular time interval (the same branch could

cross two continuous intervals).

Acknowledgements

The authors are very grateful to Arne Mooers, Joaquin Hortal, Juan Arroyo, and anonymous

reviewers for comments and suggestions that greatly improved the study. Previous versions of

the manuscript benefited from the comments of Gary Mittlebach, Emmanuelle Jousselin and

Jonathan Rolland. Financial support was provided by a Marie-Curie FP7–COFUND

(AgreenSkills fellowship–26719) grant to A.S.M. and a Marie Curie FP7-IOF (project 627684

BIOMME) grant to F.L.C. This work benefited from an "Investissements d'Avenir" grant

managed by the "Agence Nationale de la Recherche" (CEBA, ref. ANR-10-LABX-25-01).

Author contributions

Both authors designed the study, analyzed the data and wrote the manuscript.

Competing interests

The authors have no competing financial interests to declare.

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Reference list
1. Hillebrand, H. On the generality of the latitudinal diversity gradient. *Am. Nat.* **163,**

192–211 (2004).

2. Fuhrman, J. A. *et al.* A latitudinal diversity gradient in planktonic marine bacteria.

Proc. Natl. Acad. Sci. **105,** 7774–7778 (2008).

3. Novotny, V. *et al.* Why Are There So Many Species of Herbivorous Insects in Tropical

Rainforests? *Science* **313,** 1115–1118 (2006).

4. Jenkins, C. N., Pimm, S. L. & Joppa, L. N. Global patterns of terrestrial vertebrate

diversity and conservation. *Proc. Natl. Acad. Sci.* **110,** E2602–E2610 (2013).

5. Kreft, H. & Jetz, W. Global patterns and determinants of vascular plant diversity. *Proc.*

Natl. Acad. Sci. **104,** 5925–5930 (2007).

6. Willig, M. R., Kaufman, D. M. & Stevens, R. D. Latitudinal gradients of biodiversity:

pattern, process, scale, and synthesis. *Annu. Rev. Ecol. Evol. Syst.* **34,** 273–309 (2003).

7. Mittelbach, G. G. *et al.* Evolution and the latitudinal diversity gradient: speciation,

extinction and biogeography. *Ecol. Lett.* **10,** 315–331 (2007).

8. Crame, J. A. Taxonomic diversity gradients through geological time. *Divers. Distrib.* **7,**

175–189 (2001).

9. Alroy, J. *et al.* Phanerozoic trends in the global diversity of marine invertebrates.

Science **321,** 97–100 (2008).

10. Mannion, P. D., Upchurch, P., Benson, R. B., Goswami, A. & Trends Ecol, E. The

latitudinal biodiversity gradient through deep time. *Trends Ecol. Evol.* **29,** 42–50

(2014).

11. Mannion, P. D. *et al.* A temperate palaeodiversity peak in Mesozoic dinosaurs and

evidence for Late Cretaceous geographical partitioning. *Glob. Ecol. Biogeogr.* **21,** 898–

908 (2012).

12.   Rose, P. J., Fox, D. L., Marcot, J. & Badgley, C. Flat latitudinal gradient in Paleocene

mammal richness suggests decoupling of climate and biodiversity. *Geology* **39,** 163–

166 (2011).

13.   Marcot, J. D., Fox, D. L. & Niebuhr, S. R. Late Cenozoic onset of the latitudinal

diversity gradient of North American mammals. *Proc. Natl. Acad. Sci.* **113,** 7189–7194

(2016).

14.   Brocklehurst, N., Day, M. O., Rubidge, B. S. & Fröbisch, J. Olson's Extinction and the

latitudinal biodiversity gradient of tetrapods in the Permian. *Proc. R. Soc. B Biol. Sci.*

284, 20170231 (2017).

15.   Archibald, S. B., Greenwood, D. R. & Mathewes, R. W. Seasonality, montane beta

diversity, and Eocene insects: Testing Janzen's dispersal hypothesis in an equable

world. *Palaeogeogr. Palaeoclimatol. Palaeoecol.* **371,** 1–8 (2013).

16.   Archibald, S. B., Bossert, W. H., Greenwood, D. R. & Farrell, B. D. Seasonality, the

latitudinal gradient of diversity, and Eocene insects. *Paleobiology* **36,** 374–398 (2010).

17.   Labandeira, C. C. & Currano, E. D. The fossil record of plant-insect dynamics. *Annu.*

Rev. Earth Planet. Sci. **41,** 287–311 (2013).

18.   Powell, M. G. Latitudinal diversity gradients for brachiopod genera during late

Palaeozoic time: links between climate, biogeography and evolutionary rates. *Glob.*

Ecol. Biogeogr. **16,** 519–528 (2007).

19.   Powell, M. G., Beresford, V. P. & Colaianne, B. A. The latitudinal position of peak

marine diversity in living and fossil biotas. *J. Biogeogr.* **39,** 1687–1694 (2012).

20.   Krug, A. Z. & Patzkowsky, M. E. Geographic variation in turnover and recovery from

the Late Ordovician mass extinction. *Paleobiology* **33,** 435–454 (2007).

21.   Crame, J. A. Evolution of taxonomic diversity gradients in the marine realm: evidence

from the composition of recent bivalve faunas. *Paleobiology* **26,** 188–214 (2000).

22. Leprieur, F. *et al.* Plate tectonics drive tropical reef biodiversity dynamics. *Nat.*
Commun. **7,** 11461 (2016).

23. Mannion, P. D. *et al.* Climate constrains the evolutionary history and biodiversity of
crocodylians. *Nat. Commun.* **6,** 8438 (2015).

24. Nicholson, D. B., Holroyd, P. A., Valdes, P. & Barrett, P. M. Latitudinal diversity
gradients in mesozoic non-marine turtles. *R. Soc. Open Sci.* **3,** 160581 (2016).

25. Nicholson, D. B., Holroyd, P. A., Benson, R. B. J. & Barrett, P. M. Climate-mediated
diversification of turtles in the Cretaceous. *Nat. Commun.* **6,** 7848 (2015).

26. Coiffard, C. & Gomez, B. Influence of latitude and climate on spread, radiation and
rise to dominance of early angiosperms during the Cretaceous in the Northern
Hemisphere. *Geol. Acta* **10,** 181–188 (2012).

27. Peralta-Medina, E. & Falcon-Lang, H. J. Cretaceous forest composition and
productivity inferred from a global fossil wood database. *Geology* **40,** 219–222 (2012).

28. Shiono, T., Kusumoto, B., Yasuhara, M. & Kubota, Y. Roles of climate niche
conservatism and range dynamics in woody plant diversity patterns through the
Cenozoic. *Glob. Ecol. Biogeogr.* **27,** 865–874 (2018).

29. Jablonski, D. *et al.* Out of the tropics, but how? Fossils, bridge species, and thermal
ranges in the dynamics of the marine latitudinal diversity gradient. *Proc. Natl. Acad.*
Sci. **110,** 10487–10494 (2013).

30. Donoghue, M. J. A phylogenetic perspective on the distribution of plant diversity.
Proc. Natl. Acad. Sci. **105,** 11549–11555 (2008).

31. Wiens, J. & Donoghue, M. Historical biogeography, ecology and species richness.
Trends Ecol. Evol. **19,** 639–644 (2004).

32. Ricklefs, R. E. Global variation in the diversification rate of passerine birds. *Ecology*
87, 2468–2478 (2006).

33. Rolland, J., Condamine, F. L., Jiguet, F. & Morlon, H. Faster speciation and reduced

extinction in the tropics contribute to the mammalian latitudinal diversity gradient.

PLoS Biol. **12,** e1001775 (2014).

34. Wiens, J. J. Global patterns of diversification and species richness in amphibians. *Am.*

Nat. **70,** S86–S106 (2007).

35. Pyron, R. A. & Wiens, J. J. Large-scale phylogenetic analyses reveal the causes of high

tropical amphibian diversity. *Proc. R. Soc. London B Biol. Sci.* **280,** 20131622 (2013).

36. Cardillo, M., Orme, C. D. L. & Owens, I. P. F. Testing for latitudinal bias in

diversification rates: An example using new world birds. *Ecology* **86,** 2278–2287

(2005).

37. Weir, J. T. & Schluter, D. The latitudinal gradient in recent speciation and extinction

rates of birds and mammals. *Science* **315,** 1574–1576 (2007).

38. Condamine, F. L., Sperling, F. A. H., Wahlberg, N., Rasplus, J.-Y. & Kergoat, G. J.

What causes latitudinal gradients in species diversity? Evolutionary processes and

ecological constraints on swallowtail biodiversity. *Ecol. Lett.* **15,** 267–277 (2012).

39. Leslie, A. B. *et al.* Hemisphere-scale differences in conifer evolutionary dynamics.

Proc. Natl. Acad. Sci. **109,** 16217–16221 (2012).

40. Siqueira, A. C., Oliveira-Santos, L. G. R., Cowman, P. F. & Floeter, S. R. Evolutionary

processes underlying latitudinal differences in reef fish biodiversity. *Glob. Ecol.*

Biogeogr. **25,** 1466–1476 (2016).

41. Pyron, R. A. Temperate extinction in squamate reptiles and the roots of latitudinal

diversity gradients. *Glob. Ecol. Biogeogr.* **23,** 1126–1134 (2014).

42. Martin, P. S. & Klein, R. G. *Quaternary extinctions: a prehistoric revolution.*

(University of Arizona Press, 1989).

43. Latham, R. E. & Ricklefs, R. E. in *Species Diversity in Ecological Communities:*

Historical and Geographical Perspectives 294–314 (University Of Chicago Press,

1993).

44. Markwick, P. J. Fossil crocodilians as indicators of Late Cretaceous and Cenozoic

climates: Implications for using palaeontological data in reconstructing palaeoclimate.

Palaeogeogr. Palaeoclimatol. Palaeoecol. **137,** 205–271 (1998).

45. Roy, K. & Pandolfi, J. M. in *Climate Change and Biodiversity* (ed. Thomas, E. et al.)

160–175 (Yale University Press, 2005).

46. Hawkins, B. A., Diniz-Filho, J. A. F., Jaramillo, C. A. & Soeller, S. A. Post-Eocene

climate change, niche conservatism, and the latitudinal diversity gradient of New

World birds. *J. Biogeogr.* **33,** 770–780 (2006).

47. Dunn, R. R. *et al.* Climatic drivers of hemispheric asymmetry in global patterns of ant

species richness. *Ecol. Lett.* **12,** 324–333 (2009).

48. Eiserhardt, W. L., Borchsenius, F., Plum, C. M., Ordonez, A. & Svenning, J.-C.

Climate-driven extinctions shape the phylogenetic structure of temperate tree floras.

Ecol. Lett. **18,** 263–272 (2015).

49. Pulido-Santacruz, P. & Weir, J. T. Extinction as a driver of avian latitudinal diversity

gradients. *Evolution* **70,** 860–872 (2016).

50. Rolland, J., Condamine, F. L., Beeravolu, C. R., Jiguet, F. & Morlon, H. Dispersal is a

major driver of the latitudinal diversity gradient of Carnivora. *Glob. Ecol. Biogeogr.*

24, 1059–1071 (2015).

51. Rabosky, D. L. Extinction rates should not be estimated from molecular phylogenies.

Evolution **6,** 1816–1824 (2010).

52. Sanmartín, I. & Meseguer, A. S. Extinction in phylogenetics and biogeography: from

timetrees to patterns of biotic assemblage. *Frontiers in Genetics* **7,** 35 (2016).

53. Mayr, G. The early Eocene birds of the Messel fossil site: A 48 million-year-old bird

community adds a temporal perspective to the evolution of tropical avifaunas. *Biol.*

Rev. **92,** 1174–1188 (2016).

54.　Wilf, P., Labandeira, C. C., Johnson, K. R. & Cuneo, N. R. Richness of plant-insect

associations in Eocene Patagonia: A legacy for South American biodiversity. *Proc.*

Natl. Acad. Sci. **10,** 8944–8948 (2005).

55.　Blois, J. L. & Hadly, E. A. Mammalian response to Cenozoic climatic change. *Annu.*

Rev. Earth Planet. Sci. **37,** 181–208 (2009).

56.　Smith, R. Y. & al., et. Early Eocene plant diversity and dynamics in the Falkland flora,

Okanagan Highlands, British Columbia, Canada. *Palaeobiodivers. Palaeoenviron.* **92,**

309–328 (2012).

57.　Xing, Y., Onstein, R. E., Carter, R. J., Stadler, T. & Peter Linder, H. Fossils and a large

molecular phylogeny show that the evolution of species richness, generic diversity, and

turnover rates are disconnected. *Evolution* **68,** 2821–2832 (2014).

58.　Frederiksen, N. O. Sporomorph biostratigraphy, floral changes, and paleoclimatology,

Eocene and earliest Oligocene of the eastern Gulf Coast. *U.S. Geol. Surv.* 1–68 (1988).

59.　Greenwood, D. R. & Wing, S. L. Eocene continental climates and latitudinal

temperature gradients. *Geology* **23,** 1044–1048 (1995).

60.　Johnson, K. R. & Ellis, B. A tropical rainforest in Colorado 1.4 million years after the

Cretaceous-Tertiary boundary. *Science* **296,** 2379–2383 (2002).

61.　Morley, R. J. in *Tropical Rainforest Responses to Climatic Change* (eds. Bush, M. B.

& Flenley, J.-R.) 1–31 (Springer, 2007).

62.　Ziegler, A. *et al.* Tracing the tropics across land and sea: Permian to present. *Lethaia*

36, 227–254 (2003).

63.　Sluijs, A. *et al.* Subtropical Arctic Ocean temperatures during the Palaeocene/Eocene

thermal maximum. *Nature* **441,** 610–613 (2006).

64.   Zachos, J. C., Dickens, G. R. & Zeebe, R. E. An early Cenozoic perspective on
greenhouse warming and carbon-cycle dynamics. *Nature* **451,** 279–283 (2008).

65.   Jablonski, D., Roy, K. & Valentine, J. W. Out of the tropics: evolutionary dynamics of
the latitudinal diversity gradient. *Science* **314,** 102–106 (2006).

66.   Stebbins, G. L. *Flowering plants: evolution above the species level.* (Harvard
University Press, 1974).

67.   Meseguer, A. S. *et al.* Reconstructing deep-time paleoclimate legacies in the clusioid
Malpighiales unveils their role in the evolution and extinction of the boreotropical
flora. *Glob. Ecol. Biogeogr.* **27,** 616–628 (2018).

68.   Kindlmann, P., Schödelbauerová, I. & Dixon, A. F. G. in *Scaling biodiversity* 246–257
(Cambridge University Press, 2007). doi:10.1016/0169-5347(89)90163-8

69.   Angielczyk, K. D., Burroughs, R. W. & Feldman, C. R. Do turtles follow the rules?
Latitudinal gradients in species richness, body size, and geographic range area of the
world's turtles. *J. Exp. Zool. Part B Mol. Dev. Evol.* **324,** 270–294 (2015).

70.   Maddison, W. P., Midford, P. E. & Otto, S. P. Estimating a binary character's effect on
speciation and extinction. *Syst. Biol.* **56,** 701–710 (2007).

71.   FitzJohn, R. G. Diversitree: comparative phylogenetic analyses of diversification in R.
Methods Ecol. Evol. **3,** 1084–1092 (2012).

72.   Goldberg, E. E., Lancaster, L. T. & Ree, R. H. Phylogenetic inference of reciprocal
effects between geographic range evolution and diversification. *Syst. Biol.* **60,** 451–465
(2011).

73.   Davis, M. P., Midford, P. E. & Maddison, W. Exploring power and parameter
estimation of the BiSSE method for analyzing species diversification. *BMC Evol. Biol.*
13, 38 (2013).

74.   Silvestro, D., Schnitzler, J., Liow, L. H., Antonelli, A. & Salamin, N. Bayesian

estimation of speciation and extinction from incomplete fossil occurrence data. *Syst.*

Biol. **63,** 349–367 (2014).

75.   Ree, R. H. & Smith, S. A. Maximum likelihood inference of geographic range

evolution by dispersal, local extinction, and cladogenesis. *Syst. Biol.* **57,** 4–14 (2008).

76.   Joyce, W. G., Rabi, M., Clark, J. M. & Xu, X. A toothed turtle from the Late Jurassic

of China and the global biogeographic history of turtles. *BMC Evol. Biol.* **16,** 236

(2016).

77.   Jansson, R., Rodríguez-Castañeda, G. & Harding, L. E. What can multiple phylogenies

say about the latitudinal diversity gradient? A new look at the tropical conservatism,

out of the tropics, and diversification rate hypotheses. *Evolution* **67,** 1741–1755 (2013).

78.   Albino, A. M. & Brizuela, S. An overview of the South American fossil squamates.

Anat. Rec. **297,** 349–368 (2014).

79.   Pereira, A. G., Sterli, J., Moreira, F. R. R. & Schrago, C. G. Multilocus phylogeny and

statistical biogeography clarify the evolutionary history of major lineages of turtles.

Mol. Phylogenet. Evol. **113,** 59–66 (2017).

80.   Rodrigues, J. F. M., Olalla-Tárraga, M. Á., Iverson, J. B., Akre, T. S. B. & Diniz-Filho,

J. A. F. Time and environment explain the current richness distribution of non-marine

turtles worldwide. *Ecography* **40,** 1402–1411 (2017).

81.   Linnert, C. *et al.* Evidence for global cooling in the Late Cretaceous. *Nat. Commun.* **5,**

4194 (2014).

82.   Vellekoop, J. *et al.* Rapid short-term cooling following the Chicxulub impact at the

Cretaceous-Paleogene boundary. *Proc. Natl. Acad. Sci.* **111,** 7537–7541 (2014).

83.   Kaiho, K. *et al.* Global climate change driven by soot at the K-Pg boundary as the

cause of the mass extinction. *Sci. Rep.* **6,** 28427 (2016).

84.   Waterson, A. M. *et al.* Modelling the climatic niche of turtles: A deep-time perspective.

Proc. R. Soc. B Biol. Sci. **283,** 20161408 (2016).

85. Kerkhoff, A. J., Moriarty, P. E. & Weiser, M. D. The latitudinal species richness

gradient in New World woody angiosperms is consistent with the tropical conservatism

hypothesis. *Proc. Natl. Acad. Sci.* **111,** 8125–8130 (2014).

86. Meseguer, A. S., Lobo, J. M., Ree, R. H., Beerling, D. J. & Sanmartin, I. Integrating

fossils, phylogenies, and niche models into biogeography to reveal ancient evolutionary

history: the case of *Hypericum* (Hypericaceae). *Syst. Biol.* **64,** 215–232 (2015).

87. Ge, D. *et al.* Evolutionary History of Lagomorphs in Response to Global

Environmental Change. *PLoS One* **8,** e59668 (2013).

88. Pisano, J. *et al.* Out of Himalaya: The impact of past Asian environmental changes on

the evolutionary and biogeographical history of Dipodoidea (Rodentia). *J. Biogeogr.*

42, 856–870 (2015).

89. Condamine, F. L., Rolland, J., Höhna, S., Sperling, F. A. H. & Sanmartín, I. Testing

the Role of the Red Queen and Court Jester as Drivers of the Macroevolution of Apollo

Butterflies. *Syst. Biol.* doi.org/10.1093/sysbio/syy009 (2018).

doi:10.1093/sysbio/syy009

90. Stephens, P. R. & Wiens, J. J. Explaining species richness from continents to

communities: The time-for-speciation effect in emydid turtles. *Am. Nat.* **161,** 112–128

(2003).

91. Allen, A. P., Gillooly, J. F., Savage, V. M. & Brown, J. H. Kinetic effects of

temperature on rates of genetic divergence and speciation. *Proc. Natl. Acad. Sci.* **103,**

9130–9135 (2006).

92. Antonelli, A. & Sanmartín, I. Mass extinction, gradual cooling, or rapid radiation?

Reconstructing the spatiotemporal evolution of the ancient angiosperm genus

Hedyosmum (Chloranthaceae) using empirical and simulated approaches. *Syst. Biol.*

60, 596–615 (2011).

93.   Sun, Y. *et al.* Lethally Hot Temperatures During the Early Triassic Greenhouse.

902   *Science* **338,** 366–370 (2012).

903   94.   Crisp, M. D. & Cook, L. G. How was the Australian flora assembled over the last 65

904   million years? A molecular phylogenetic perspective. *Annu. Rev. Ecol. Evol. Syst.* **44,**

905   303–324 (2013).

906   95.   Morlon, H., Parsons, T. L. & Plotkin, J. B. Reconciling molecular phylogenies with the

907   fossil record. *Proc. Natl. Acad. Sci.* **108,** 16327–16332 (2011).

908   96.   May, M. R., Hï¿½hna, S. & Moore, B. R. A Bayesian approach for detecting the

909   impact of mass-extinction events on molecular phylogenies when rates of lineage

910   diversification may vary. *Methods Ecol. Evol.* **7,** 947–959 (2016).

911   97.   Heath, T. A., Huelsenbeck, J. P. & Stadler, T. The fossilized birth-death process for

912   coherent calibration of divergence-time estimates. *Proc. Natl. Acad. Sci.* **111,** E2957–

913   E2966 (2014).

914   98.   Spriggs, E. L. *et al.* Temperate radiations and dying embers of a tropical past: the

915   diversification of *Viburnum*. *New Phytol.* **207,** 1469–8137 (2015).

916   99.   Fritz, S. A. *et al.* Diversity in time and space: wanted dead and alive. *Trends Ecol.*

917   *Evol.* **28,** 509–516 (2013).

918   100.   Jaffe, A. L., Slater, G. J. & Alfaro, M. E. The evolution of island gigantism and body

919   size variation in tortoises and turtles. *Biol. Lett.* **7,** 558–561 (2011).

920   101.   Rodrigues, J. F. M. & Diniz-Filho, J. A. F. Ecological opportunities, habitat, and past

921   climatic fluctuations influenced the diversification of modern turtles. *Mol. Phylogenet.*

922   *Evol.* **101,** 352–358 (2016).

923   102.   Shaffer, H. B., McCartney-Melstad, E., Near, T. J., Mount, G. G. & Spinks, P. Q.

924   Phylogenomic analyses of 539 highly informative loci dates a fully resolved time tree

for the major clades of living turtles (Testudines). *Mol. Phylogenet. Evol.* **115,** 7–15
(2017).

103. Oaks, J. R. A time-calibrated species tree of crocodylia reveals a recent radiation of the
true crocodiles. *Evolution* **65,** 3285–3297 (2011).

104. FitzJohn, R. G., Maddison, W. P. & Otto, S. P. Estimating traitdependent speciation
and extinction rates from incompletely resolved phylogenies. *Syst. Biol.* **58,** 595–611
(2009).

105. Silvestro, D., Salamin, N. & Schnitzler, J. PyRate: A new program to estimate
speciation and extinction rates from incomplete fossil data. *Methods Ecol. Evol.* **5,**
1126–1131 (2014).

106. Silvestro, D., Cascales-Miñana, B., Bacon, C. D. & Antonelli, A. Revisiting the origin
and diversification of vascular plants through a comprehensive Bayesian analysis of the
fossil record. *New Phytol.* **207,** 1469–8137 (2015).

107. Gradstein, F. M., Ogg, J. G. & Smith, A. G. A Geologic Time Scale 2004. *A Geol.*
Time Scale 2004, Ed. by Felix M. Gradstein James G. Ogg Alan G. Smith (2005).

108. Gelman, A. Parameterization and Bayesian modeling. *J. Am. Stat. Assoc.* **99,** 537–545
(2004).

109. Rambaut, A. & Drummond, A. J. Tracer. Version 1.3. *Available online*
at:http://beast.bio.ed.ac.uk/ Tracer

110. Romdal, T. S., Araújo, M. B. & Rahbek, C. Life on a tropical planet: niche
conservatism and the global diversity gradient. *Glob. Ecol. Biogeogr.* **22,** 344–350
(2013).

111. Smith, S. A. Taking into account phylogenetic and divergence-time uncertainty in a
parametric biogeographical analysis of the Northern Hemisphere plant clade
Caprifolieae. *J. Biogeogr.* **36,** 2324–2337 (2009).

112. Moore, B. R., Smith, S. a., Ree, R. H. & Donoghue, M. J. Incorporating Fossil Data in

Biogeographic Inference : A Likelihood Approach. in *Proc. 52nd Annu. Syst.*

Symposium (2008).

113. Ree, R. H. & Sanmartin, I. Prospects and challenges for parametric models in historical

biogeographical inference. *J. Biogeogr.* **36,** 1211–1220 (2009).