

1 **Water motion as a transformation mechanism of algal communities structure in Lake Baikal**

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3 *Lyubov S. Kravtsova, Igor B. Mizandrontsev, Svetlana S. Vorobyova, Lyudmila A. Iziboldina,*
4 *Elena V. Mincheva, Tatyana G. Potyomkina, Tatyana I. Triboy, Igor V. Khanaev, Dmitry Yu.*
5 *Sherbakov and Andrey P. Fedotov*

6

7 Limnological Institute SB RAS, Irkutsk, Russia

8

9 **Abstract.** The diversity of algal communities of phytoplankton and meio- and macrophytes was
10 investigated in Lake Baikal. Fragments of *Spirogyra* thallomes were recorded in the phytoplankton
11 community of Southern Baikal, which had never been recorded before in its composition. It was
12 also established that the structure of benthic algal communities changed in comparison with that in
13 2000 due to intense development of filamentous algae, particularly *Spirogyra*. Its lowest biomass
14 was recorded in the surf zone and wave breaking, whereas the highest biomass was registered in the
15 area of weakened effect of waves on the bottom. The cover percent of the bottom with filamentous
16 algae in different areas of the coastal zone varied from 0 to 100%. Hydraulic characteristics of
17 *Spirogyra* were the same as those of planktonic diatoms. The circulation currents and wave effect on
18 the bottom favoured transfer and distribution of *Spirogyra* from the location of its intense
19 development into the coastal area of Lake Baikal.

20 **Additional keywords:** orbital velocity, sinking velocity

21 **Introduction**

22 Alongside the effects of global warming and anthropogenic impact, algal blooms have been
23 recorded in the coastal areas of seas and lakes (Nozaki *et al.* 2003; Smith *et al.* 2006; Hiraoka *et al.*
24 2011), which cause deterioration of not only water quality but also social environments in
25 recreational areas. A similar phenomenon has been shown to exist in the ecosystem of Lake Baikal,
26 a UNESCO World Heritage site. At present, large clumps of filamentous algae thrown to the shore
27 have been found on the Baikal beaches in some areas of the lake (Timoshkin *et al.* 2016). Fragments
28 of benthic filamentous algae *Spirogyra* have been found in the phytoplankton community
29 (Kobanova *et al.* 2016; Bondarenko and Logacheva 2016), which have not been found in this
30 community before (Kozhov 1931; Popovskaya 1977). It is necessary to determine the cause(s) of the
31 emergence of *Spirogyra* cells in the phytoplankton.

32 For the past decade, changes have been recorded not only in the plankton structure (Izmest'eva *et al.*
33 2016) but also in the benthic communities of Lake Baikal (Kravtsova *et al.* 2014; Timoshkin *et al.*

34 2016). In the areas of the coastal zone confined mainly to settlements, we observe overgrowing of
35 the bottom with filamentous algae among which there are members of the genus *Spirogyra*,
36 something that is atypical of algal communities of Lake Baikal. Earlier, singular *Spirogyra* filaments
37 have been recorded only in the well-heated bays and shallow areas (sors) of Lake Baikal and its
38 tributaries (Kozhova and Iziboldina 1994; Iziboldina 2007). We hypothesise that the distribution of
39 the members of the genus *Spirogyra* in the coastal zone of the lake is attributed to circulation
40 currents of Baikal waters, which transfer these algae from places where they develop in great
41 numbers. It is known that the hydrodynamic regime together with such environmental factors as
42 temperature, light and chemical composition of water affect the biota structure of sea and freshwater
43 environments (Peters *et al.* 2006; Wolcott 2007; Wang *et al.* 2012; Liu *et al.* 2015). Moreover, the
44 water motion directly or indirectly influences hydrobionts. Specifically, characteristics of water
45 motion, including velocity of currents, rough water, dynamic pressure and turbulence, directly affect
46 mobility and transfer of aquatic organisms (Luchar *et al.* 2010; Durham *et al.* 2013; Cross *et al.*
47 2014). Wind waves and ripples affecting higher aquatic plants and benthic communities may also
48 function as limiting and optimising factors in their growth (Raspopov *et al.* 1990). An indirect effect
49 of water motion on the diversity and spatial distribution of hydrobionts occurs due to the changes of
50 sedimentation dynamics, transport of particulates and detritus in the coastal area of water bodies
51 (Snelgrove *et al.* 1988; Airoidi 1998).

52 It is interesting to know how hydrodynamics affect the flora of Lake Baikal, a unique freshwater
53 environment with a depth of over 1,630 m combining the features of sea and lake ecosystems. The
54 objective characteristic of water motion in Lake Baikal has been obtained from long-term in-situ
55 measurements, instrumental investigation and mathematical models. General patterns of formation
56 of currents, wave activity, surging, turbulence and upwelling were determined previously (Pomytkin
57 1960; Ainbund 1973; Afanasyev and Verbolov 1977; Fialkov 1983; Zhdanov *et al.* 2009; Shimaraev
58 *et al.* 2012). There are few works on the indirect effect of hydrodynamics on the biota of Lake
59 Baikal, which are devoted only to the survey of the effect of water masses on plankton (Likhoshway
60 *et al.* 1996; Jewson *et al.* 2010) and wave activity on distribution of benthic organisms in the coastal
61 area of the lake (Karabanov and Kulishenko 1990).

62 The aim of this study is to assess the role of *Spirogyra* in the structure of current algal communities
63 of Lake Baikal and contribution of hydrodynamic processes in its dissemination in the coastal zone
64 of the lake.

65 **Material and methods**

66 *Field studies*

67 We studied the algal flora of Lake Baikal in August 2016 on board the research vessel “Titov”.
68 We analysed phytoplankton of Southern Baikal to study the possible transfer of *Spirogyra* fragments
69 (Fig. 1a). Phytoplankton samples (1.5 L of water) were collected at six stations (I–VI) with a water
70 sampler at depths of 0, 5 and 15 m. Stations I–V were located at a distance of approximately 50–100
71 m off the shore. Station VI was 7 km away from the shore in the direction from Cape Listvennichny
72 to the settlement of Tankhoy. Additionally, we collected samples from depths of 25 m and 50 m. All
73 quantitative phytoplankton samples (22) were fixed in Utermel solution.

74 To estimate the recent diversity of phytobenthos, scuba divers collected 117 quantitative samples of
75 meio- and macrophytes from depths of 0–10 m in Southern, Central and Northern Baikal at 29
76 stations located at 11 sites differing in wind-wave characteristics and bottom geomorphology (Fig.
77 1b). Moreover, to assess the bottom cover percent with filamentous algae, the scuba divers mapped
78 meio- and macrophytes at 15 transects (Tr) using frames (area of 1 m²) divided into 100 equal
79 quadrats.

80 The scuba divers also collected 18 quantitative samples of meio- and macrophytes from two
81 transects directed perpendicular to the shoreline to characterise the structure of benthic algal
82 communities in Southern Baikal. One transect was located in the background region (outside the
83 impact zone of the settlement of Listvyanka) 5 km to the north of Cape Listvennichny opposite
84 Emelyanovka Valley. Another transect was located in the impact zone in Listvyanka opposite
85 Krestovka Valley (Fig. 1a). The scuba divers collected three samples in each vegetation belt at
86 depths of 0–10 m using a frame with an area of 0.16 m². The scuba divers put stones covered with
87 algae into sacks of strong fabric and lifted them onboard the vessel. The stones were then put into a
88 large cuvette with water. The algae were cut with a scalpel from the surface or brushed off. The
89 water was poured into the sieve of a mill-gauze No. 23. The algae were put into flasks and fixed in
90 4% formalin.

91 *Laboratory analysis.* Phytoplankton samples were settled in a 15–20 mL volume for about 14 days.
92 Algae were counted in 0.1 mL. Individual volumes of cells were taken into account to determine
93 algal biomass (mg m⁻³) (Makarova and Pichkily 1970). Picoplankton and cysts of chrysophytes were
94 not considered in total phytoplankton biomass.

95 To assess the transport in currents of both *Spirogyra* and diatoms, we calculated the sinking velocity
96 during settling in a laminar flow from Stokes formula:

$$97 \quad W = 2r^2g(\rho - \rho_0)/(9\mu), \quad (1)$$

98 where W is the sinking velocity of a spherical particle in the water, cm s^{-1} ; r is its radius, cm ; g is the
99 gravity acceleration (normal, 980.655), cm s^{-2} ; ρ is the particle density, g cm^{-3} ; ρ_0 is the water
100 density (1 g cm^{-3}); μ is the water dynamic viscosity, $\text{g cm}^{-1} \text{ s}^{-1}$.

101 To reduce non-spherical algal cells to a conditionally spherical shape, we calculated the equivalent
102 radius according to cell volume:

$$103 \quad r_e = \left(\frac{3V}{4\pi} \right)^{1/3} \quad (2)$$

104 The geometric form coefficient ζ was used to estimate non-spherical particles:

$$105 \quad \zeta = S/S_0, \quad (3)$$

106 where S is the particle surface area; S_0 is the sphere area of the same effective radius.

107 Dynamic coefficient of the particle shape $\Gamma(\zeta)$ in the linear area of environmental resistance was
108 calculated from Velikanov *et al.* (2013):

$$109 \quad \Gamma(\zeta) = 1 + 0.348(\zeta - 1), \quad (4)$$

110 where ζ is the geometric form coefficient.

111 The density of living *Spirogyra* was determined in the following way. First, we estimated the true
112 specific weight of dry mass (density of dry substance) of *Spirogyra* filaments. We dried them on
113 filter paper and then subjected to a solid tablet to pressure in a compression mold. The density of the
114 dry substance was calculated from the volume of this tablet and its weight. The humidity was
115 estimated from the difference between the weight of *Spirogyra* filaments dried on filter paper (until
116 a wet spot disappeared) and the weight of these filaments (quantity 20 g) dried at $103 \text{ }^\circ\text{C}$.

117 The temperature of the cell content was set as equal to the water temperature for estimating the
118 density of living *Spirogyra*. The water density at this temperature was determined from the table
119 data. The density of living *Spirogyra* was calculated from the following formula:

$$120 \quad \rho_s = v\rho_{dry} + v_w\rho_w \quad (5)$$

121 where v and v_w are the volumes of the dry substance and water, respectively; ρ_{dry} is the density of the
122 dry substance; ρ_w is the water density at the given temperature.

123 Meio- and macrophytes were sorted according to taxon level under an MBC 10 microscope at 2×8
124 magnification. Species were identified from the temporary algal preparations under an Amplival
125 microscope at 12×10 and 12×40 magnifications. Cell sizes (diameter, width and length in μm) were
126 measured with an ocular-micrometer.

127 Before weighing meio- and macrophytes on a torsion balance VT-500 with an accuracy of 0.1, we
128 dried them on filter paper until a wet spot disappeared. The data obtained were converted to 1 m² of
129 the bottom (g m⁻²).

130 The diversity of phytoplankton as well as of phytobenthos was characterised using Shannon's
131 species diversity index (Odum 1971):

$$132 \quad H = - \sum \frac{n_i}{N} \times \ln \left(\frac{n_i}{N} \right), \quad (6)$$

133 where n_i is the biomass of i -species; N is the total biomass of species in a certain habitat.

134 Algal communities (of both phytoplankton and phytobenthos) were identified from the modified
135 density index (Brotskaya and Zenkevich 1939):

$$136 \quad I = \sqrt{P \times B} \quad (7)$$

137 where P is the frequency of occurrence (the ratio of a number of samples in which a species has
138 been found to total the number of samples, %); B is the percentage of a species in the total biomass,
139 %. Species with maximal density index (I) were considered dominant, with $I > 10\%$ being
140 subdominant and species with $I < 10\%$ being considered minor.

141 We compared the structure of algal communities with that of previous years, referring to the data by
142 S. Vorobyova on phytoplankton in August 1992 (Southern Baikal – 35 quantitative samples) and L.
143 Izhboldina on phytobenthos for August in 1966–1988 (Southern, Central and Northern Baikal – 37
144 quantitative samples out of 298).

145 The dependence of the lake bottom overgrowing with filamentous algae on the hydrodynamic
146 environment was estimated by principal component analysis (PCA) using the following parameters
147 as variables: x_1 – bottom cover percent with filamentous algae, %; x_2 – composition of bottom
148 sediments; x_3 – depth, m; x_4 – width of the coastal zone, m; x_5 – wave height (h , m); x_6 – wave
149 length (λ , m); x_7 – periodicity of wave activity (τ , s); x_8 – slope ratio of the coastal zone; x_9 – bottom
150 current velocity (U_{\max} and U , m·s⁻¹); x_{10} – shear velocity (V_{sh} , m·s⁻¹) of sediment movement (0.5 mm
151 in diameter); x_{11} – coefficient of sediment mobility (K_m); x_{12} – x_{18} – hydrodynamic pressure (P , g m⁻²)
152 on vertical surface at certain depths (0.5 m; 1.5 m; 2 m; 3 m; 4 m; 5 m). In our calculations, we used
153 the highest values of wave activity (λ , τ) at the prevailing wind velocity (5–10 m·s⁻¹) during the
154 August navigation (Galazy 1993).

155 During wave activity, the water moves around circular orbits. The sizes of these circulations
156 decrease towards the shore (with the decrease in depth), and their orbits acquire the shape of flat
157 ellipses. Moreover, bottom velocities of water motion increase. According to the interaction between

158 the water flow and the bottom, we distinguish three zones during wave activity: I – offshore zone of
159 undeformed waves; II – transformation of wave (deformation zone and breaker zone); and III – surf
160 zone (Fig. 2a).

161 The maximal velocity of the water motion (in the zone of undeformed waves) was calculated from
162 the following equation (Petrov 1985):

$$U_{\max} = \frac{\pi h}{\tau \times sh(2\pi H/\lambda)}, \quad (8)$$

163 where h is the wave height, m; τ is the wave period, m; H is the depth, m; sh is the hyperbolic sine; λ
164 is the wave length, m.

165 Current velocities were calculated from the following formula (Petrov 1985) in case of a significant
166 effect of the bottom on the orbital wave component (in the zone of wave breaking):

$$U = \alpha \sqrt{1.28gh}, \quad (9)$$

168 where h is the wave height, m; g is the gravity acceleration, $\text{m}\cdot\text{s}^{-2}$; α is values varying from 0.7 to
169 1.8 with the depth decrease; in the surf zone it reaches 2 and then reduces to 0 at the end point of the
170 splash.

171 The shear velocity (the initial velocity of sediment movement) was estimated from the following
172 formula (Longinov 1963):

$$V_{sh} = 2.19K(sl)\sqrt{2rg}, \quad (10)$$

174 where V_{sh} is the shear velocity, $\text{m}\cdot\text{s}^{-1}$; $K(sl)$ is the non-dimensional coefficient depending on the
175 bottom slope (sl); r is the particle radius; $g = 9.80665 \text{ m}\cdot\text{s}^{-2}$:

$$K(sl) = 1.0074 - 0.6381sl + 0.6296sl^2 \quad (11)$$

176 The mobility coefficient of bottom sediments was calculated from the ratio of the maximal
177 horizontal component of the orbital bottom velocity and the initial velocity of sediment movement
178 (Karabanov and Kulishenko 1990):

$$K_m = U_{\max}/V_{sh}, \quad (12)$$

180 Water flow pressure on the vertical surface during wave activity was estimated from the following
181 formula (Longinov 1963):

$$P = 27h^2/H, \quad (13)$$

182 where P is the hydraulic pressure, g m^{-2} ; h is the wave height, m; H is the depth, m; numerical
183 coefficient, $\text{g m}^{-2} \text{ m}^{-1}$.

184 **Results**

188 During this study, the water temperature was 15–17 °C at a depth of up to 15 m in the coastal zone
189 and in the adjacent areas of the open pelagic zone. In 2016, 50 taxa (at a lower level than genus) of
190 planktonic algae were registered in the flora, of which 14 were diatom taxa, 3 were dinophyte taxa,
191 2 were cryptophyte taxa, 7 were chrysophyte taxa, 8 were blue-green taxa, 14 were green taxa and 2
192 were flagellate taxa. The diversity from the Shannon index varied from 2.6 to 2.8 in the coastal zone
193 at the depths of 0 m, 5 m and 15 m. Algal communities represented by 38–44 taxa, including
194 *Spirogyra* sp. with dominance of *Asterionella formosa* Hass., were found at these depths (Fig. 3a).
195 Previously, in Southern Baikal in summer (1992), the Shannon index varied from 1.8 to 2.4 at
196 different depths. At this time, the community comprised dominant species of *Rhodomonas pusilla*
197 (Bachm.) Javorn. and *Gymnodinium coeruleum* Ant.; members of the genus *Spirogyra* were not
198 recorded (Fig. 3b).

199 At the depths of 25 m and 50 m in the open pelagic area, *Spirogyra* was also recorded in the
200 phytoplankton community in 2016. *Fragilaria radians* Kütz., *Dinobryon cylindricum* Imhof., *R.*
201 *pusilla* and *G. coeruleum* dominated. In 1992, the community was dominated by the latter two
202 species distributed at these depths; *Spirogyra* was not recorded.

203 It should be noted that thallome fragments of *Spirogyra* consisting mainly of one to three cells were
204 found in phytoplankton at all studied depths whose biomass varied from 8.5 to 35.2 mg m⁻³ at 0 m, 5
205 m, 15 m, 25 m and 50 m (the abundance was 650 to 5,440 cells L⁻¹).

206 Geometric and hydraulic characteristics of both singular cells and thallome fragments of *Spirogyra*
207 (Table 1) were assessed to ascertain possibilities of their transfer with coastal currents. The same
208 characteristics were presented for comparison of diatoms: namely *A. formosa* as a dominant of the
209 phytoplankton community and *Aulacoseira baicalensis* (Meyer) Simonsen as a typical member of
210 the Baikal algal community. It was noted that the dynamic coefficient in the linear area of resistance
211 for a *Spirogyra* cell and thallome fragments of several cells is comparable with that of diatoms.

212 The density of dry *Spirogyra* calculated from the volume of the tablet and its weight was 1.36 g cm⁻³
213 ³, and humidity of the living alga was 90%.

214 The density ρ_s of the living *Spirogyra* was 1.036 g cm⁻³. The sinking velocity at 10 °C of one
215 *Spirogyra* cell with the volume of 28,600 μm^3 and equivalent radius of 19 μm was $22 \times 10^{-3} \text{ mm} \cdot \text{s}^{-1}$.

216 The density of a thin-walled diatom *A. baicalensis* calculated by us from the data provided by
217 Jewson *et al.* (2010) was 1.27 g cm⁻³. Therefore, its cell, with a volume of 14,300 μm^3 and
218 equivalent radius of 15 μm , sinks at a velocity of $90 \times 10^{-3} \text{ mm} \cdot \text{s}^{-1}$. Another diatom, *A. formosa*, sinks
219 at $15 \times 10^{-3} \text{ mm} \cdot \text{s}^{-1}$ having a cell volume of 800 μm^3 and equivalent radius of 5.76 μm . It is likely that
220 this diatom is of the same density as that of *A. baicalensis*. As *A. formosa* is able to form star

221 colonies consisting of several cells, e.g. five, its sinking velocity in this case is $44 \times 10^{-3} \text{ mm} \cdot \text{s}^{-1}$. The
222 density of living skeletonless algae at this temperature is close to water density varying within small
223 values. Even in such large cells as *Spirogyra*, the sinking velocity is lower than in a diatom of *A.*
224 *baicalensis* with a silicon exoskeleton.

225 In 2016, the benthic flora in the studied regions of Lake Baikal was more diverse in comparison
226 with that of the period before 2000. Its composition consisted of 65 taxa, 56 of them being meio-
227 and macrophytes (Table 2), 8 higher aquatic plants (*Batrachium* sp., *Elodea canadensis* Michx.,
228 *Fontinalis antipyretica* Hedw., *Myriophyllum* sp., *M. spicatum* L., *Potamogeton crispus* L. and *P.*
229 *perfoliatus* L.) and a lichen *Collema ramenskii* Elenk.

230 Among meio- and macrophytes, we detected the filamentous algae *Mougeotia* sp., *Oedogonium*
231 *flavascens*?, *Oedogonium* sp., *Spirogyra calospora*?, *Spirogyra* sp. and *Ulothrix zonata* (Web. et
232 Mohr.) Kütz. Of special interest were algae of the genus *Spirogyra* whose habitat in the previous
233 century was confined only to the coastal-sor zone (Fig. 1c), whereas at present the habitat has
234 widened significantly (Fig. 1b). Before 2000, *Spirogyra* inhabited only certain areas of the lake in
235 the form of singular filaments (Fig. 1c). Moreover, it was recorded in the grab samples collected at a
236 depth of 40–80 m (Dagarskaya Bay, Angara-Kichera Shoal and Cape Ukhan). In 2016, the
237 occurrence of *Spirogyra* was 75% along the open coastal areas in Lake Baikal. Besides *Spirogyra*,
238 filamentous algae of the genus *Oedogonium* were also widespread in the lake with an occurrence of
239 40% (Table 2). Before 2000, this alga was detected in Angara-Kichera Sor, Anga Bay, Barguzin
240 Bay (settlement of Makarovo) and opposite the town of Baikalsk. In 2016, *Oedogonium* was also
241 recorded outside the habitats mentioned above, i.e. in Listvennichny Bay, the settlement of
242 Angasolka, Cape Kotelnikovskiy, Zavorotnaya Bay and Senogda Bay.

243 The bottom cover percent with filamentous algae, predominantly with *Spirogyra*, varied between 0
244 and 100% in different regions of the coastal area at a depth of over 2 m. Velocities varied at
245 maximal wave lengths of 14 m and 21 m and periodicity of 3 s and 5.1 s, respectively, and at wave
246 heights of 1–1.2 m: U from 0.03 to 6.2 m s^{-1} and V_{sh} from 0.07 to 0.15 m s^{-1} . The orbital velocity in
247 the zone of wave profile transformation (at a depth of 3.5–3.7 m) was on average $0.32 \pm 0.03 \text{ m s}^{-1}$,
248 $1.12 \pm 0.06 \text{ m s}^{-1}$ in the zone of wave breaking (depths of 1.5–2.5 m) and $3.76 \pm 0.06 \text{ m s}^{-1}$ in the surf
249 zone at the beginning of surge (depth of up to 1.5 m). The boundaries of zones (Fig. 2a) were
250 mobile, and their location depended on gale force at Lake Baikal. At the same wave parameters, K_m
251 in these zones was higher than 1, attesting to the mobility of bottom sediments.

252 PCA analysis showed that the main percentage of variability (80%) of the whole database was
253 provided by the first two principal components. According to the first principal component, with

254 56% of total variability of the database, environmental factors such as depth, height, length and
255 periodicity of waves as well as width of the coastal zone affected the cover percent; the load of
256 variables was positive (Fig. 4a). According to the second principal component, with 24% of total
257 variability of the database, the distribution of filamentous algae was dependent on shear velocity,
258 composition of bottom sediments, depth and slope angle of the bottom in the coastal zone (Fig. 4b).
259 It is clear that the bottom cover percent with filamentous algae was independent of the water flow
260 pressure on the vertical surface, as filament strands stretched along the bottom horizontally or
261 fluctuated synchronously with the reciprocating motion of water flow. The load of variables (x_{12} – x_{18}
262 at different wave heights) on the first and second components were negative. The water flow
263 pressure on the vertical surface at a wave height of 1–1.2 m was 11–72 g m⁻² at a depth of 10–1.5 m
264 and was insufficient for detachment of filaments from the substrate (underwater video filming).
265 In space of the two first principal components, the point set is divided into two non-overlapping
266 subsets *I* and *II* (Fig. 5). Subset *I* covers the points of sampling at stations where filamentous algae
267 are often recorded. Moreover, at some stations (Cape Listvennichny, Krestovka Valley, near the
268 outlet of the Angara River, Baikalsk and Ushakovka River), the bottom cover percent with
269 filamentous algae at a depth of over 2 m reached 40–100%. At these stations, the vertical zoning of
270 the spatial distribution of meio- and macrophytes was disturbed because of mass development of
271 filamentous algae. The same subset covers the stations (settlement of Kultuk, settlement of
272 Maximikha and Ushakovka River) where we found free-lying clumps formed by filamentous algae
273 on the sandy bottom. In addition, the first subset covered the sites where the bottom cover percent
274 was up to 15%. However, the vertical zoning of algal distribution was not disturbed.
275 Subset *II* comprises stations at which historically formed zoning in the spatial distribution of meio-
276 and macrophytes remains; there were no filamentous algae or their cover percent was 1–3% (Fig. 5).
277 The diversity and specific structure of meio- and macrophyte communities of the coastal zone were
278 studied in the relatively uniform hydrodynamic environment at site 3 (Fig. 1a, Table 2) but with
279 different recreation load within its boundaries. In 2016, in the background region opposite the
280 Emelyanovka Valley, the Shannon index was not high (1.0). We revealed an algal community
281 dominated by an endemic species *Draparnaldioides baicalensis* C. Meyer et Skabitsch. (Fig. 6a). Its
282 composition was represented by 22 taxa with a total biomass of 112±58 g m⁻². The filamentous alga
283 *U. zonata* is characteristic of stony substrates in Lake Baikal. Singular filaments of three *Spirogyra*
284 morphotypes were recorded among rare species (P<10%). Their biomass was low beyond the limits
285 of balance sensitivity. The percentage of all filamentous algae was lower than 0.1% of the total
286 biomass. In 1987, the Shannon index was 1.7. At this time, the community consisted of 15 taxa with

287 the dominance of *D. baicalensis* (Fig. 6b). Members of the genus *Spirogyra* were absent among
288 subdominant and minor species in the algal community. The total biomass of the community made
289 up $90 \pm 27 \text{ g m}^{-2}$, and only *U. zonata* was recorded in this community with a biomass of 2%.

290 In 2016, the Shannon index was 1.3 in the zone of anthropogenic impact opposite the settlement of
291 Listvyanka (Krestovka Valley). The community with the dominance of *Spirogyra*, uncharacteristic
292 of the open coastal zone of Lake Baikal, formed for the first time here (for the period of
293 observations since the beginning of the previous year) (Fig. 6a). This community was represented
294 by 25 taxa with a total biomass of $109 \pm 66 \text{ g m}^{-2}$. Among minor species of filamentous algae in the
295 community we registered *U. zonata* as well as members of the genus *Oedogonium* (two taxa),
296 uncharacteristic of the open coastal zone in Lake Baikal. The contribution of filamentous algae in
297 the total biomass of the community was 58%. Earlier (in 1987), the Shannon index (2.5) was twice
298 as high. The community dominated by *Dermatochrysis reticulata* C. Meyer (Fig. 6b) included 23
299 taxa of meio- and macrophytes with a total biomass of $66 \pm 24 \text{ g m}^{-2}$. The percentage of filamentous
300 algae was 13% of the total biomass of the community, among which there were only members of the
301 genus *Ulothrix* (*U. zonata*, *U. tenerrima* Kütz. and *U. tenuissima* Kütz.); *Spirogyra* and
302 *Oedogonium* were not recorded.

303 The distribution of *Spirogyra* biomass depends on the hydrodynamic environment. Its lowest
304 content was recorded in the surf and wave breaking zones, whereas the highest values were
305 registered in the zone of weak effect of wave activity on the bottom (Fig. 2b). Moreover, the bottom
306 cover percent with filamentous algae increased in the bay with the depth increase, i.e. with the
307 decrease of wave effect on the bottom (Fig. 2c).

308 According to the visual observations of I. Khanaev during the year, the length of *Spirogyra*
309 thallomes varied in the coastal zone of Listvennichny Bay. In January, it was 7–10 cm and up to 20
310 cm in May. In June–August, filament strands could reach more than 150 cm and 10 cm in
311 September–December. The longest filament strands usually form at depths below 8 m (up to 15 m),
312 i.e. beyond the zone of wave effect on the bottom during summer storms ($h=1 \text{ m}$). During our
313 investigation, the length of filament strands reached 50–70 cm in some areas of the lake. Cells
314 forming the *Spirogyra* thallomes were represented by two size groups. Some cells ($n=38$
315 measurements) were on average $30 \pm 1 \text{ }\mu\text{m}$ wide (14–40 μm) and $152 \pm 14 \text{ }\mu\text{m}$ long (29–345 μm).
316 Others ($n=23$) were an average of $45 \pm 1 \text{ }\mu\text{m}$ wide (41–68 μm) and $208 \pm 20 \text{ }\mu\text{m}$ long (81–378 μm).

317 The mass development of *Spirogyra* affects the structural organisation of both benthic and
318 planktonic algal communities.

319 **Discussion**

320 The average directional transfer of Baikal waters forms mainly under dynamic influence of the
321 atmosphere (wind regime and pressure gradient above the water area). The system of currents at
322 Lake Baikal (Fig. 1b) comprises alongshore circulation of cyclone type covering the entire lake and
323 secondary circulations in the southern, central and northern basins of Lake Baikal (Afanasyev and
324 Verbolov 1977). In the storm period, orbital velocities with a regime probability of 0.1% near the
325 bottom can stir up the sand around the entire area of the coastal zone (Fialkov 1983). The system of
326 currents and turbulent diffusion distribute terrigenous suspended sediments brought with river
327 waters around the water area of the lake. According to a mathematical model (Mizandrontsev and
328 Sudakov 1981), suspended sediments (diameter of 0.005 mm, density of 2.65 g cm^{-3} and sinking
329 velocity of $17 \times 10^{-3} \text{ mm} \cdot \text{s}^{-1}$) are transported along the western coast of Northern Baikal at current
330 velocities of some centimeters per second for a distance of hundreds of kilometers from the mouths
331 of large tributaries located in the northern part of the lake. Offshore secondary circulations and
332 horizontal turbulent diffusion promote the removal of suspended particles in the open areas of the
333 lake (Mizandrontsev and Sudakov 1981). The diatoms *A. formosa* and *A. baicalensis* as well as fine
334 mineral sediments can be transferred by coastal currents at significant distances. This relates to
335 planktonic algae without exoskeletons with densities close to the density of the lake water and to
336 fragments of *Spirogyra* thallomes. Moreover, the density of living skeletonless planktonic algae is
337 close to 1 g cm^{-3} and can be lower than the water density (at this temperature) due to the presence of
338 gas vacuoles and fat inclusions (Henderson-Sellers 1987; Smith 1982). The transfer mechanism of
339 filaments and their fragments within the water column, considering the geometric and hydraulic
340 characteristics (Table 1), is similar to that of planktonic diatoms. The deposition rate of *A.*
341 *baicalensis* ($10,000 \mu\text{m}^3$) in the laminar water flow is $39 \times 10^{-3} \text{ mm s}^{-1}$ (Votintsev 1961) and that of
342 sea phytoplankton, particularly dinoflagellates (taking into account the equivalent radius of cells and
343 their non-spherical shape), is from $3 \times 10^{-3} \text{ mm s}^{-1}$ to $45 \times 10^{-3} \text{ mm s}^{-1}$ (Kamykowski *et al.* 1992). The
344 deposition rate of *Spirogyra* ($22 \times 10^{-3} \text{ mm s}^{-1}$) is lower than that of Baikal diatoms. As deposition
345 rates are very low, it helps both filament fragments and diatoms remain within the water column for
346 a long time. Thallome fragments of filamentous algae removed by currents from the coastal area can
347 move along the perimeter in each basin of the lake and around the entire lake (Fig. 1b). For example,
348 in the southern basin of Lake Baikal, algal fragments will be transferred from Cape Listvennichny to
349 Cape Tolsty (Fig. 1a) for 11 days at the wind with regime probability of 50% and at the velocity of
350 drift current of 9 cm s^{-1} in the middle water layer.

351 The growth and development of benthic algae in aquatic ecosystems are also closely connected with
352 hydrodynamics, in particular with wave activity (Reiter 1986; Raspopov *et al.* 1990; Nozaki *et al.*

2003; Engelen *et al.* 2005). In Lake Baikal, as in other water bodies, vegetation of dominant algae of the vegetation belts (*U. zonata* at a depth of 0–1.5 m, *T. cylindrica* (Wahl.) Ag. var. *bullosa* C. Meyer and *Didymosphenia geminata* (Lingb.) M. Schmidt. at a depth of 1.5–2.5 m and species of the genus *Draparnaldioides* at a depth from 3 m to 10–12 m) is determined by the hydrodynamic environment. A large amount of floating fragments of algae and higher plants as well as their mass clumps on the shore of Lake Baikal after summer and autumn storms has been recorded since the beginning of the first half of the previous century (Kozhov 1931; Votintsev 1961). The detachment of benthic algae from the substrate can occur under the influence of orbital and reciprocating water motions caused by wave activity (Karabanov and Kulishenko 1990). In late autumn and early winter (November–December), when the waves reach their maximal height, the algae abovementioned stop their vegetation (Izhboldina 2007). It is clear that intense development of seasonal algae occurs in summer with durable calm and low wave activity ($h=0.5$ m). At this time, algal mats of filamentous algae can be found in the coastal zone at a depth of 3 m and deeper (Kravtsova *et al.* 2014). Judging from in-situ data, in summer, *Spirogyra* forms the longest filament strands at depths of over 8–10 m. In Lake Baikal during rare summer storms with a wave height of about 1 m, bottom currents ($U_{max}=0.03\text{--}0.10$ m s⁻¹) are unable to detach filaments from the substrate, whereas at lower depths their detachment is quite probable. The filamentous algae *Spirogyra* and *Mougeotia* develop at current velocities of 0.12 m s⁻¹ and 0.29 m s⁻¹ with greater biomass in the first case and slightly lower biomass in the second case (Peterson and Stevenson 1992). In the majority of the studied regions of Lake Baikal, the orbital velocity (0.36 m s⁻¹) emerging during the storm ($h=1.0$ m) at depths of 3.3–4 m is not critical for *Spirogyra* development in comparison with the areas where the velocity at these depths can reach 0.68 m s⁻¹ (Fig. 5). Clumps of filaments freely lying on the sandy bottom after a 3-ball storm found during the field studies confirmed the wave effect on algae. In the Ushakovka River at a depth of 6 m ($U_{max}=0.18$ m s⁻¹), the filamentous clumps were likely formed from algae detached at 1.5–2.1 m, where current velocities could reach 1.02–3.80 m s⁻¹. Depending on the slope angle and width of the coastal zone, algae detached from the substrate after the storm are either transferred by near-bottom currents or washed ashore or thrown to the shore. Freely lying *Spirogyra* clumps at the bottom of the coastal zone of Lake Baikal have been also found by other researchers (Timoshkin *et al.* 2016). At lower depths (0.6–1.5 m), current velocities are significantly higher and can be 3.8–6.2 m s⁻¹ (at $h=1.0$ m). At such velocities, coarse pebbles and boulders (30 cm in diameter) are transported and the length of algal filaments shorten, e.g. in *U. zonata* from 10–15 cm to 5–6 mm (Karabanov and Kulishenko 1990). Therefore, the main limiting factor in the development of filamentous algae in

386 the zones of wave breaking and surge is mobility of bottom sediments. *Spirogyra* biomass here is
387 several times lower than at a depth of 3 m (Fig. 2*b,c*) because of the abrasive effect of sandy
388 particles (with eroding velocities minimal for binder soil) and gravel-pebble material coming into
389 motion. In the coastal zone of seas and large lakes, orbital velocities (0.5 m s^{-1}) are able to resuspend
390 fine sand particles (Rasmussen and Rowan 1997). The movement of debris (1–20 cm in diameter) in
391 water bodies is known to emerge when near-bottom current velocity exceeds critical values
392 ($V_{sh}=0.5\text{--}1.7 \text{ m}\cdot\text{s}^{-1}$) and depends on its size and bottom slope (Volkov and Ionin 1962). In Lake
393 Baikal, as in other water bodies, bottom sediments are free from algae as a result of transfer,
394 mobility, friction and turbidity of debris of different sizes. Filaments detached from the substrate
395 enter the water column. This is supported by findings of *Spirogyra* in the planktonic samples at all
396 studied depths in Southern Baikal (Fig. 3*a*).

397 Factors affecting spatial heterogeneity of benthic algae are numerous and one of them is
398 anthropogenic impact. The intense development of filamentous algae in the coastal zone of Lake
399 Baikal is caused by nutrient flux in the areas of human activity (Kravtsova *et al.* 2014; Timoshkin *et al.*
400 *al.* 2016). In particular, in July–August 2011 high concentrations of nutrients P up to 0.420 mg L^{-1}
401 and nitrate as N up to 0.20 mg L^{-1} were recorded in bottom waters of Listvennichny Bay, whereas
402 the background values for P and N in open parts of the lake are 0.007 mg L^{-1} and 0.08 mg L^{-1} ,
403 respectively (Khodzher *et al.* 2017). In addition, in 2015 the content of nutrients (mg L^{-1}) was higher
404 in comparison with that in the open lake: $\text{NH}_4^+=0.560 \pm 0.47$; $\text{NO}_2^-=0.055 \pm 0.01$; $\text{NO}_3^-=0.690 \pm$
405 0.07 ; $\text{P}_{\text{mineral}}=0.025 \pm 0.03$; $\text{P}_{\text{organic}}=0.024 \pm 0.02$; $\text{P}_{\text{total}}=0.049 \pm 0.05$ (Kulakova *et al.* 2017). Higher
406 concentrations of nutrients in the depression topography are caused by secondary pollution produced
407 by algae degrading in the coastal zone opposite the settlement of Listvyanka.

408 Despite *Spirogyra* having been registered earlier in the coastal-sor zone (Fig. 1*c*), such quantities of
409 this alga (Fig. 2*a*; Kobanova *et al.* 2016) have not been recorded in either benthic (Izhboldina 2007)
410 or planktonic algal communities (Kozhov 1931; Popovskaya 1977). This means that the specific
411 character of the recent structure of Baikal algal communities in comparison with that in the previous
412 century is the presence of *Spirogyra* in their composition (Table 2, Fig. 3*a*, Fig. 6*a*). These algae
413 have been found ubiquitously in the benthic communities of the studied habitats, except station 17
414 opposite Irinda Bay, an area characterised by elevated velocities of bottom currents. In addition,
415 secondary circulations existing in Northern Baikal do not cover this area (Fig. 1*b*). Moreover, of
416 interest are earlier findings of singular *Spirogyra* filaments outside the coastal zone at depths of 40–
417 80 m (Fig. 1*c*). Their emergence in the deep zone as well as the emergence of the thermophilic

418 diatom *A. formosa* at depths of 300–600 m may be attributed to the run-off of warm waters along the
419 slope at the boundary of the thermal bar (Likhoshway *et al.* 1996).

420 **Conclusions**

421 Hydrodynamics plays an important role in the formation of the recent structure of algal communities
422 of Lake Baikal. Under conditions of global warming, anthropogenic impact and flux of nutrients
423 into the coastal zone, we observe the bottom overgrowing with filamentous algae *Spirogyra* and
424 *Oedogonium*. During storms, filaments detach from the substrate and are washed ashore, forming
425 aggregates on the beaches or entering the water column. For the hydraulic characteristics of
426 filamentous algae, *Spirogyra* in particular are comparable with those of planktonic diatoms, and the
427 existing system of currents in Lake Baikal causes their transfer from the regions of mass
428 development and distribution outside the zones of local anthropogenic impact. Therefore, in the
429 open parts of the coastal zone of Lake Baikal remote from settlements, we find only singular
430 filaments or thallome fragments of *Spirogyra*. Bays and sors, i.e. traditional habitats of this alga,
431 also serve as a source of replenishment with algal fragments of circular currents. Only at present has
432 *Spirogyra* been registered in the plankton of the open parts of Lake Baikal despite benthic diatoms
433 being constantly recorded in the composition of plankton. In the case of rising nutrient load on the
434 coastal zone, the role of filamentous algae in the Baikal ecosystem will increase and hydrodynamics
435 will promote its dissemination. The investigation of hydrodynamic processes in the interconnection
436 with a biotic component of aquatic ecosystems plays an important role in the understanding of
437 mechanisms of their function. The algal bloom in the inland waters has become a critically
438 important issue for its impacts on natural and social environments. Long-term monitoring must
439 therefore consider the human factor controlling these blooms and their impact on water supply in
440 Lake Baikal and other large lakes threatened by accelerating eutrophication.

441 **Acknowledgements**

442 This work was supported by budget projects of the Federal Agency of Scientific Organizations Nos.
443 0345-2016-0004 and 0345-2016-0006. Verification of meio- and macrophyte taxa with the use of
444 molecular research techniques was carried out with the assistance of a grant Nos.17-44-388071 r_a.
445 The authors thank divers V. Chernykh and Yu. Yushuk for their help in sampling.

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576 macroturbulent exchange and dissipation velocity of turbulent energy in the pelagic area of
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578 Table 1. Geometric and hydraulic characteristics of *Spirogyra* and planktonic diatoms.

Object under study	Volume V , μm^3	Equivalent radius r_e , μm	Actual surface area S , μm^2	Sphere area S_0 , μm^2	Geometric coefficient, ζ	Dynamic coefficient, Γ
<i>Spirogyra</i>						
one cell	28,600	19.0	6,729	4,523	1.49	1.17
thallome	57,200	23.9	13,148	7,172	1.88	1.30
fragment of 2 cells						
thallome	85,800	27.4	20,188	9,407	2.15	1.40
fragment of 3 cells						
thallome	286,000	41.0	67,290	20,992	3.21	1.77
fragment of 10 cells						
thallome	572,000	51.5	134,580	33,323	4.03	2.05
fragment of 20 cells						
thallome	1430,000	69.9	336,450	61,382	5.48	2.56
fragment of 50 cells						
Diatoms						
one cell of <i>Aulacoseira baicalensis</i>	14,300	15.0	3,908	2,827	1.38	1.13
one cell of <i>Asterionella formosa</i>	800	5.8	907	423	2.14	1.40
five cells of <i>Asterionella formosa</i>	4,000	9.8	4,535	1,207	3.76	1.96

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Table 2. Composition of meio- and macrophytes in the coastal zone of Lake Baikal. Taxa uncharacteristic of the coastal zone in bold; total taxon number in brackets

No. of site (Fialkov, 1983)	Length of coastal zone, km	Width of coastal zone, m	Coastal slope angle, °	Depth where coastal zone bend into slope, m	Composition of meio- and macrophytes	
					before 2000	in 2016
1	40	75	3.8	5	<i>Chaetocradiella pumila</i> , <i>Cladophora compacta</i> , <i>C. kursanovii</i> , <i>Didymosphenia geminata</i> , <i>Draparnaldioides baicalensis</i> , <i>D. pilosa</i> , <i>Nostoc verrucosum</i> , <i>Tetraspora cylindrica</i> var. <i>bullosa</i> , <i>Ulothrix zonata</i> (9)	<i>Calothrix</i> sp., <i>Chaetocradiella pumila</i> , <i>Chaetomorpha curta</i> , <i>Cladophora floccosa</i> , <i>C. glomerata</i> , <i>C. kursanovii</i> , <i>Dermatochrysis reticulata</i> , <i>Dermatochrysis</i> sp., <i>Didymosphenia geminata</i> , <i>Draparnaldioides baicalensis</i> , <i>Oedogonium</i> sp. , <i>Oscillatoria amoena</i> , <i>Rivularia borealis</i> , <i>Spirogyra calospora?</i> , <i>Spirogyra</i> sp. , <i>Tolypothrix distorta</i> , <i>Ulothrix zonata</i> (17)
3	45	111	5.1	10	<i>Calothrix</i> sp., <i>Chaetocradiella pumila</i> , <i>Cladophora compacta</i> , <i>C. floccosa</i> , <i>C. kursanovii</i> , <i>Didymosphenia geminata</i> , <i>Draparnaldioides baicalensis</i> , <i>Nostoc verrucosum</i> , <i>Oscillatoria amoena</i> , <i>Schizothrix</i> sp., <i>Tetraspora cylindrica</i> var. <i>bullosa</i> , <i>Tetrasporopsis reticulata</i> , <i>Tolypothrix distorta</i> , <i>Ulothrix tenerrima</i> , <i>Ulothrix zonata</i> , <i>Ulothrix zonata</i> (16)	<i>Calothrix parietina</i> , <i>Calothrix</i> sp., <i>Chaetocradiella pumila</i> , <i>Chaetophora elegans</i> , <i>Cladophora compacta</i> , <i>C. floccosa</i> , <i>C. floccosa</i> f. <i>floccosa</i> , <i>C. floccosa</i> var. <i>irregularis</i> , <i>C. glomerata</i> , <i>C. kursanovii</i> , <i>Dermatochrysis</i> sp., <i>Didymosphenia geminata</i> , <i>Draparnaldioides arnoldii</i> , <i>D. baicalensis</i> , <i>D. pilosa</i> , <i>Nitella</i> sp., <i>Nostoc verrucosum</i> , <i>Oedogonium flavescens</i> , <i>Oedogonium</i> sp. , <i>Oscillatoria amoena</i> , <i>Phaeoplaca baicalensis</i> , <i>Schizothrix</i> sp., <i>Spirogyra</i> sp. , <i>Tetraspora cylindrica</i> var. <i>bullosa</i> , <i>Tolypothrix distorta</i> , <i>T. distorta</i> f. <i>penicillata</i> , <i>T. distorta</i> f. <i>distorta</i> , <i>Ulothrix zonata</i> (28)
8	55	200	3.4	12	<i>Chaetocradiella pumila</i> , <i>Cladophora compacta</i> , <i>C. floccosa</i> f. <i>floccosa</i> , <i>Didymosphenia geminata</i> , <i>Draparnaldioides arenaria</i> , <i>D. arnoldii</i> , <i>D. baicalensis</i> , <i>D. pumila</i> , <i>Nostoc verrucosum</i> , <i>Tetraspora cylindrica</i> var. <i>bullosa</i> , <i>Tolypothrix distorta</i> f. <i>penicillata</i> , <i>Ulothrix zonata</i> (12)	<i>Chaetocradiella microscopica</i> , <i>Chaetomorpha baicalensis</i> , <i>C. moniliformis</i> , <i>C. solitaria</i> , <i>Cladophora compacta</i> , <i>C. floccosa</i> , <i>C. floccosa</i> f. <i>floccosa</i> , <i>C. floccosa</i> var. <i>irregularis</i> , <i>C. glomerata</i> , <i>C. meyeri</i> , <i>C. pulvinata</i> , <i>Dermatochrysis</i> sp., <i>Didymosphenia geminata</i> , <i>Draparnaldioides arnoldii</i> , <i>D. baicalensis</i> , <i>Mougeotia</i> sp.? ,

10	54	241	3.3	14	<i>Didymosphenia geminata</i> , <i>Draparnaldioides arenaria</i> , <i>D. arnoldii</i> , <i>D. baicalensis</i> , <i>D. pumila</i> , <i>D. vilosa</i> , <i>Tetraspora cylindrica</i> var. <i>bullosa</i> , <i>Ulothrix zonata</i> (8)	<i>Nostoc verrucosum</i> , <i>Oedogonium</i> sp. , <i>Oscillatoria amoena</i> , <i>O. tenuis</i> , <i>Schizothrix</i> sp., <i>Spirogyra</i> sp. , <i>Tolypothrix distorta</i> , <i>Tolypothrix distorta</i> f. <i>penicillata</i> , <i>T. distorta</i> f. <i>distorta</i> , <i>Ulothrix zonata</i> (26) <i>Calothrix</i> sp., <i>Chaetocradiella pumila</i> , <i>Chaetomorpha moniliformis</i> , <i>Cladophora compacta</i> , <i>C. floccosa</i> , <i>C. floccosa</i> f. <i>floccosa</i> , <i>C. floccosa</i> var. <i>irregularis</i> , <i>C. glomerata</i> , <i>Dermatochrysis</i> sp., <i>Didymosphenia geminata</i> , <i>Draparnaldioides arenaria</i> , <i>D. arnoldii</i> , <i>D. baicalensis</i> , <i>D. pilosa</i> , <i>Gemmiphora compacta</i> , <i>Microcoleus subtorulosus</i> , <i>Microcystis muscicola</i> , <i>Nostoc verrucosum</i> , <i>Oedogonium</i> sp. , <i>Oscillatoria amoena</i> , <i>O. tenuis</i> , <i>Schizothrix</i> sp., <i>Spirogyra fluviatilis</i> , <i>Spirogyra</i> sp. , <i>Tolypothrix distorta</i> , <i>T. distorta</i> f. <i>penicillata</i> , <i>T. distorta</i> f. <i>distorta</i> , <i>Ulothrix zonata</i> (28)
12	52	154	4.1	11	<i>Cladophora floccosa</i> f. <i>floccosa</i> , <i>Draparnaldioides baicalensis</i> , <i>Tetraspora cylindrica</i> var. <i>bullosa</i> , <i>Tolypothrix distorta</i> f. <i>penicillata</i> , <i>Ulothrix zonata</i> (5)	<i>Chaetomorpha moniliformis</i> , <i>Cladophora floccosa</i> , <i>C. kursanovii</i> , <i>Didymosphenia geminata</i> , <i>Oedogonium</i> sp. , <i>Spirogyra</i> sp. , <i>Tolypothrix distorta</i> (7)
14	64	828	1.0	14	<i>Cladophora floccosa</i> f. <i>floccosa</i> , <i>Cladophora kursanovii</i> , <i>Didymosphenia geminata</i> , <i>Draparnaldioides pumila</i> , <i>Microcystis muscicola</i> , <i>Nostoc verrucosum</i> , <i>Tetraspora cylindrica</i> var. <i>bullosa</i> (7)	<i>Cladophora floccosa</i> var. <i>irregularis</i> , <i>C. glomerata</i> , <i>Draparnaldioides pumila</i> , <i>Microcoleus subtorulosus</i> , <i>Nitella</i> sp., <i>Oedogonium</i> sp.1 , <i>Oscillatoria amoena</i> , <i>O. tenuis</i> , <i>Spirogyra</i> sp. , <i>Ulothrix zonata</i> (10)
17	69	580	1.4	14	<i>Cladophora floccosa</i> f. <i>floccosa</i> , <i>C. floccosa</i> var. <i>irregularis</i> , <i>Didymosphenia geminata</i> , <i>Draparnaldioides pumila</i> , <i>Nostoc verrucosum</i> , <i>Tetraspora cylindrica</i> var. <i>bullosa</i> (6)	<i>Calothrix</i> sp., <i>Chaetomorpha moniliformis</i> , <i>Cladophora floccosa</i> , <i>C. floccosa</i> f. <i>floccosa</i> , <i>C. floccosa</i> var. <i>irregularis</i> , <i>C. glomerata</i> , <i>C. kursanovii</i> , <i>Dermatochrysis</i> sp., <i>Didymosphenia geminata</i> , <i>Draparnaldioides arnoldii</i> , <i>D. arnoldii</i> f. <i>compacta</i> , <i>D. baicalensis</i> , <i>D. pilosa</i> , <i>Nostoc verrucosum</i> , <i>Rivularia borealis</i> , <i>Schizothrix</i> sp., <i>Tolypothrix distorta</i> f. <i>penicillata</i> , <i>T. distorta</i> f. <i>distorta</i> , <i>Ulothrix zonata</i> (19)
21	56	375	2.7	18	<i>Cladophora floccosa</i> f. <i>floccosa</i> , <i>C. glomerata</i> , <i>Draparnaldioides arenaria</i> , <i>D.</i>	<i>Calothrix</i> sp., <i>Chaetomorpha moniliformis</i> , <i>Cladophora compacta</i> , <i>C. floccosa</i> , <i>C.</i>

					<i>arnoldii</i> , <i>D. baicalensis</i> (5)	<i>floccosa</i> f. <i>floccosa</i> , <i>C. floccosa</i> var. <i>irregularis</i> , <i>C. fracta</i> , <i>C. glomerata</i> , <i>C. kursanovii</i> , <i>Dermatochrysis</i> sp., <i>Didymosphenia geminata</i> , <i>Nostoc verrucosum</i> , <i>Oscillatoria amoena</i> , <i>Schizothrix</i> sp., <i>Spirogyra</i> sp. , <i>Tolypothrix distorta</i> f. <i>distorta</i> , <i>Ulothrix zonata</i> (17)
22	83	2450	0.5	22	<i>Chara</i> sp., <i>Cladophora floccosa</i> f. <i>floccosa</i> , <i>C. floccosa</i> var. <i>irregularis</i> , <i>C. fracta</i> , <i>Nitella</i> sp., <i>Nostoc pruniforme</i> , <i>N. verrucosum</i> (7)	<i>Chaetomorpha moniliformis</i> , <i>Cladophora floccosa</i> f. <i>floccosa</i> , <i>C. fracta</i> , <i>C. glomerata</i> , <i>Nostoc verrucosum</i> , <i>Oedogonium</i> sp., <i>Oedogonium</i> sp.1 , <i>Spirogyra</i> sp. , <i>Tolypothrix distorta</i> (9)
25	56	1400	0.5	12	<i>Cladophora fracta</i> , <i>C. glomerata</i> , <i>Dermatochrysis reticulata</i> , <i>Gloeothrichia pisum</i> , <i>Nostoc verrucosum</i> (5)	<i>Cladophora kursanovii</i> , <i>Didymosphenia geminata</i> , <i>Spirogyra</i> sp. , <i>Tolypothrix distorta</i> f. <i>penicillata</i> , <i>Ulothrix zonata</i> (5)
30	47	574	1.1	11	<i>Didymosphenia geminata</i> , <i>D. arenaria</i> , <i>D. pilosa</i> , <i>Tetraspora cylindrica</i> var. <i>bullosa</i> , <i>Ulothrix zonata</i> (5)	<i>Cladophora floccosa</i> , <i>C. floccosa</i> f. <i>floccosa</i> , <i>C. floccosa</i> var. <i>irregularis</i> , <i>C. glomerata</i> , <i>C. kursanovii</i> , <i>Didymosphenia geminata</i> , <i>Draparnaldioides pilosa</i> , <i>D. pumila</i> , <i>Microcystis muscicola</i> , <i>Mougeotia</i> sp.?, <i>Nitella</i> sp., <i>Nostoc verrucosum</i> , <i>Oedogonium</i> sp. , <i>Schizothrix</i> sp., <i>Spirogyra</i> sp. , <i>Tolypothrix distorta</i> , <i>T. distorta</i> f. <i>penicillata</i> , <i>T. distorta</i> f. <i>distorta</i> , <i>Ulothrix zonata</i> (19)

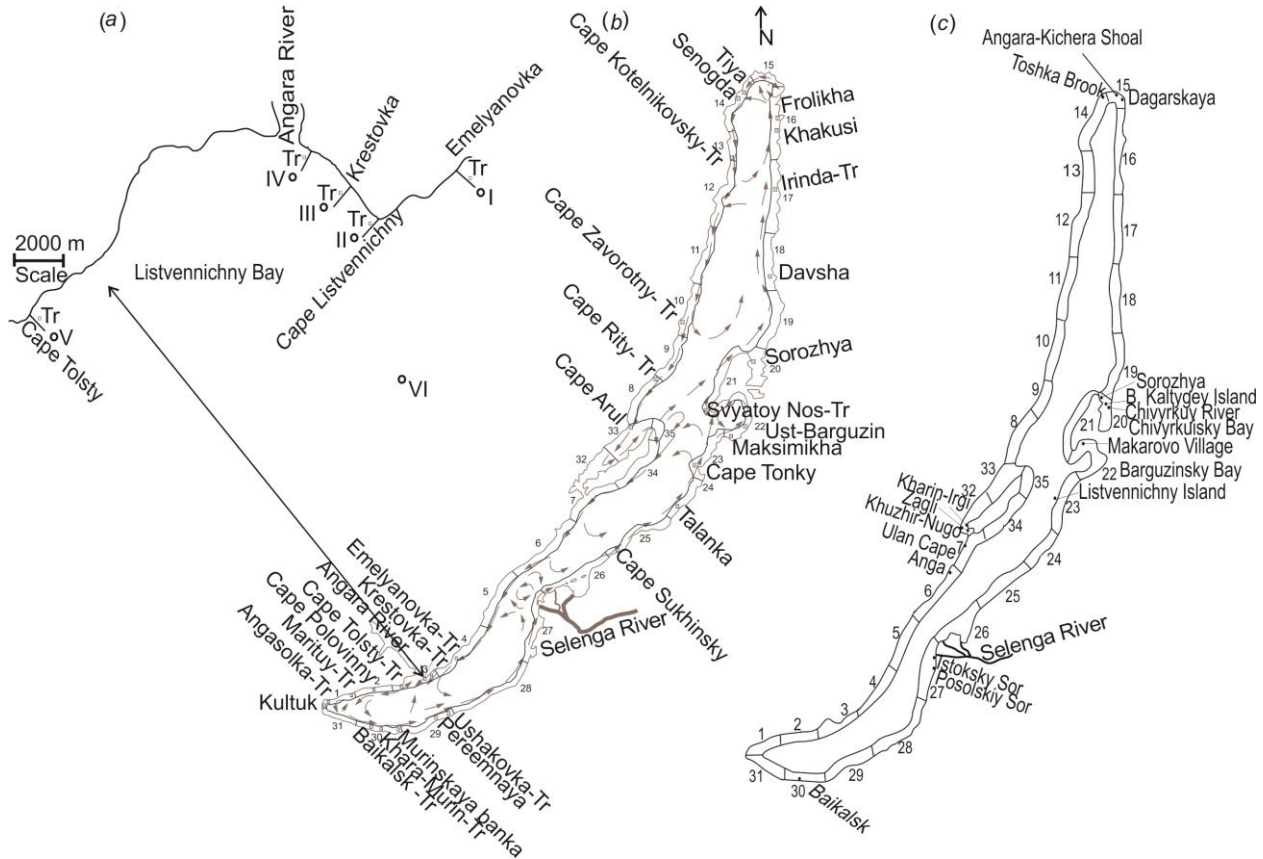


Fig. 1. (a,b) Map-scheme of algal sampling in different areas of Lake Baikal and (c) areas of findings (before 2000) of singular *Spirogyra* filaments. (a) I-VI – stations of phytoplankton sampling in Listvennichny Bay; (b) – geographic names on the scheme corresponding to points of quantitative algal samples at stations; Tr – transects at which scuba divers measured bottom cover percent with filamentous algae (%) and simultaneously collected samples; 1-35 – stations differing in wind-wave characteristics and bottom geomorphology (according to Fialkov 1983); arrows show circulation of water masses (according to Afanasyev and Verbolov 1977); (c) – points show places of *Spirogyra* findings (archive data of L.Izhboldina).

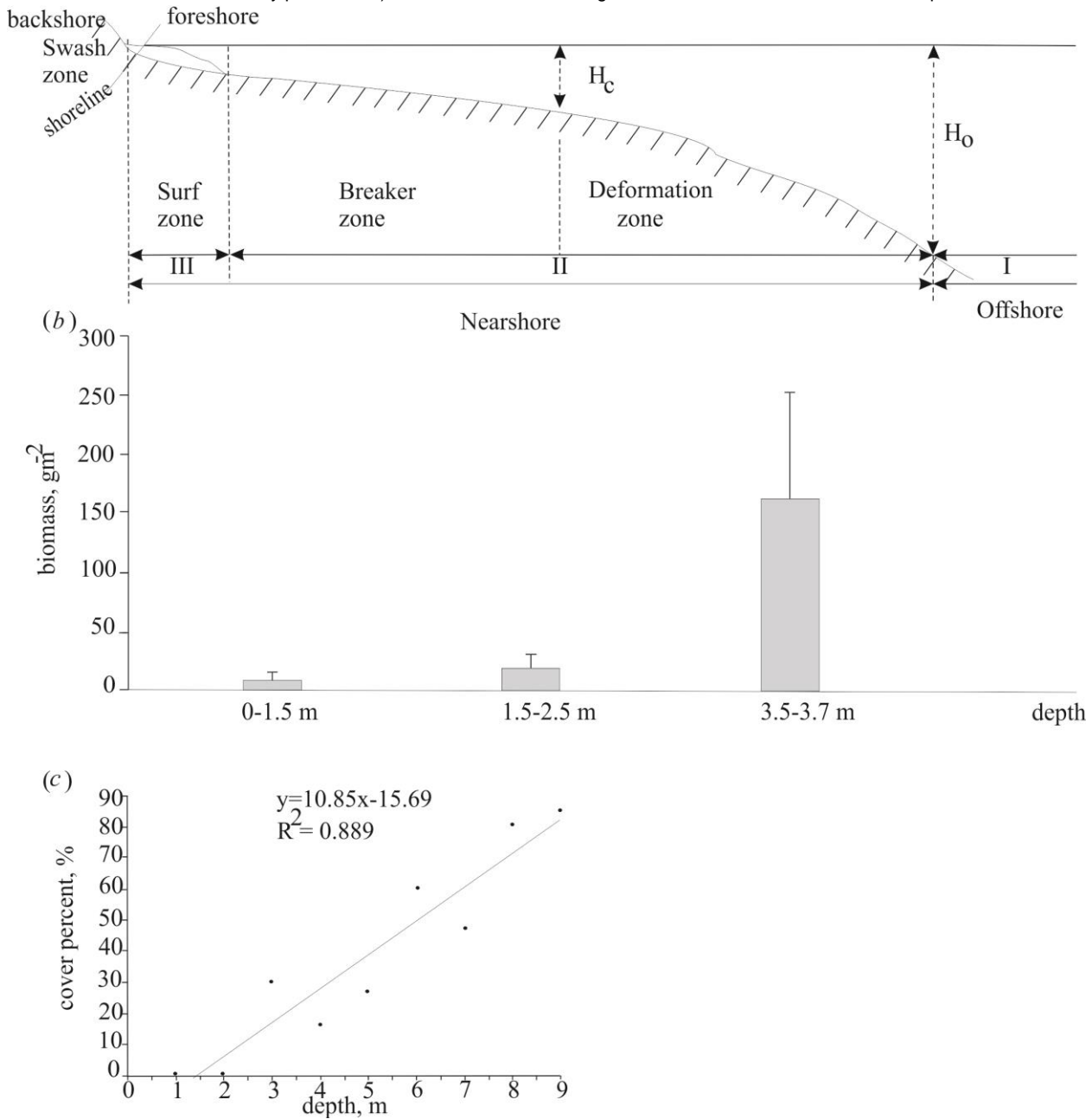


Fig. 2. (a) Zones of wave effect on the bottom, (b) biomass of *Spirogyra*, and (c) correlation between cover percent with filamentous algae and depth in the coastal zone of Lake Baikal (Krestovka Valley, August 2016). H_c - a depth with wave effect on the bottom, and H_0 - a depth without wave effect on the bottom.

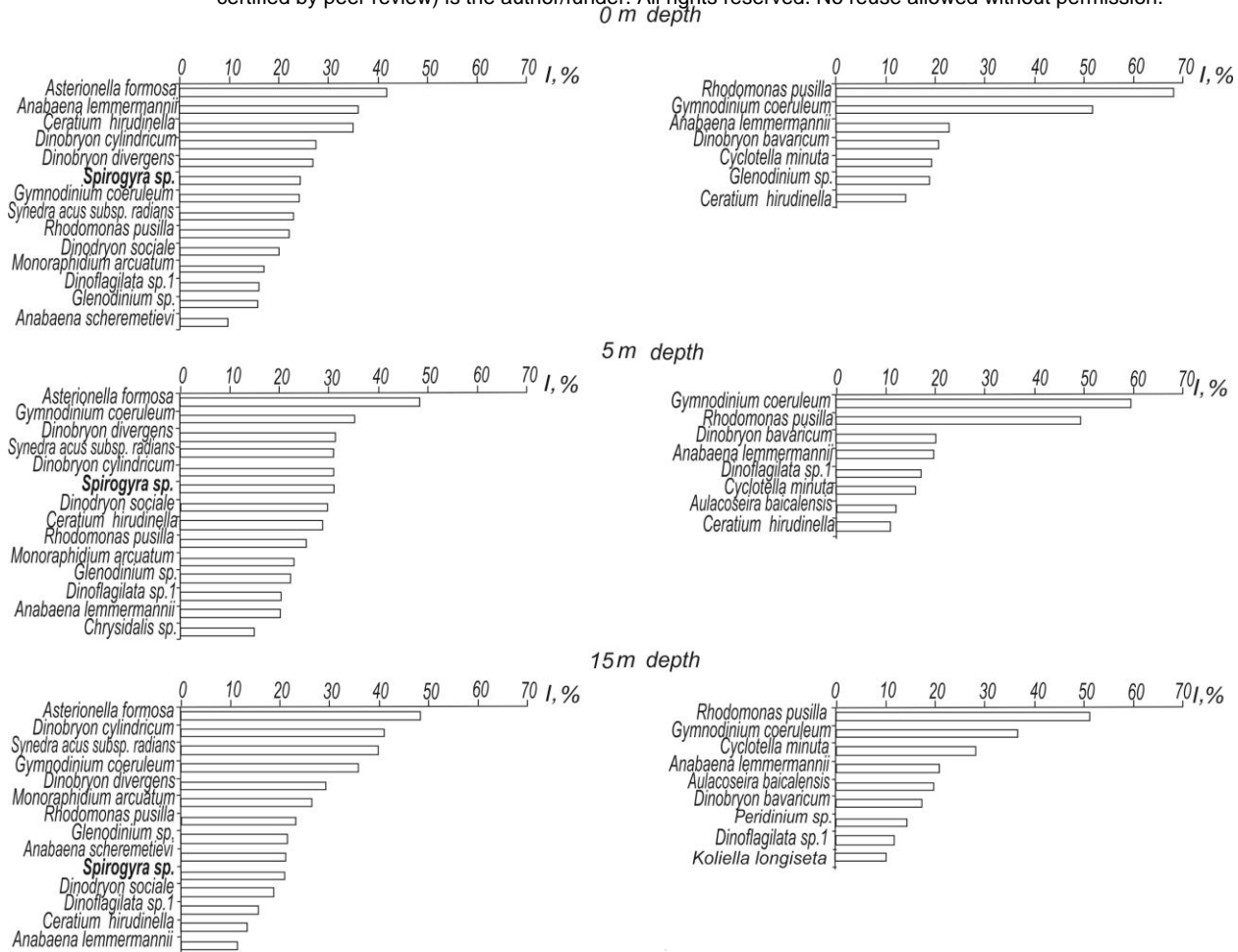


Fig. 3. Structure of phytoplankton community in Southern Baikal during different years: (a) – in 2016; (b) – in 1992. Along X-axis – density index; along Y-axis – species ranking in the order of decrease of density.

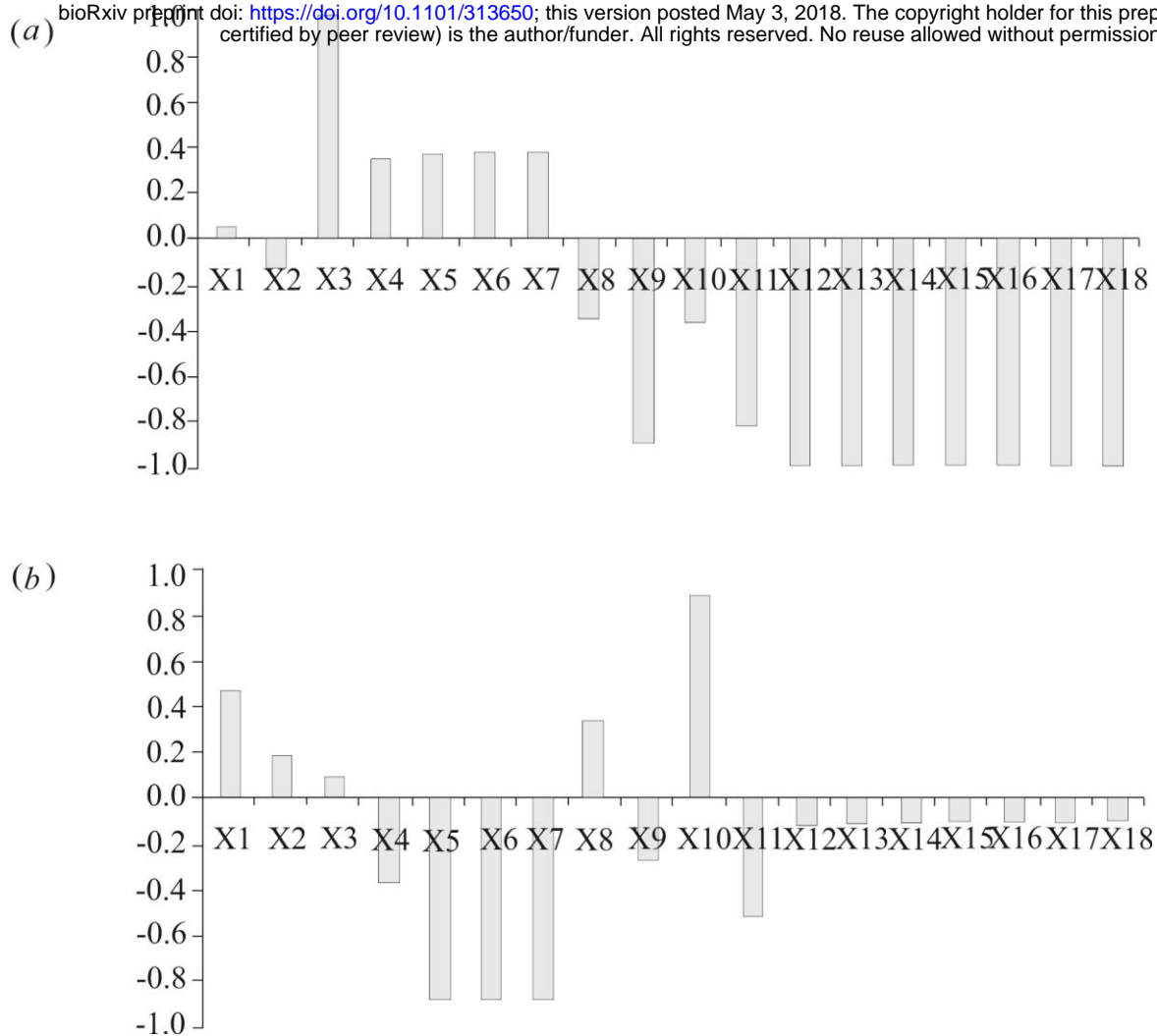


Fig. 4. Loads of variables on the first (a) and second (b) principal components. Variables along X-axis: x_1 – bottom cover percent with filamentous algae, %; x_2 – composition of bottom sediments; x_3 – depth, m; x_4 – width of the coastal zone, m; x_5 – wave height, m; x_6 – wave length, m; x_7 – periodicity of wave activity, c; x_8 – slope ratio of the coastal zone; x_9 – bottom current velocity, $\text{m}\cdot\text{s}^{-1}$; x_{10} – shear velocity of sediment movement (0.5 mm in diameter), $\text{m}\cdot\text{s}^{-1}$; x_{11} – coefficient of sediment mobility; x_{12} - x_{18} – hydrodynamic pressure ($\text{g}\cdot\text{m}^{-2}$) on vertical surface at certain depths (0.5 m; 1.5 m; 2 m; 3 m; 4 m; 5 m). Along Y-axis contribution of each variable x to total variability of characteristics.

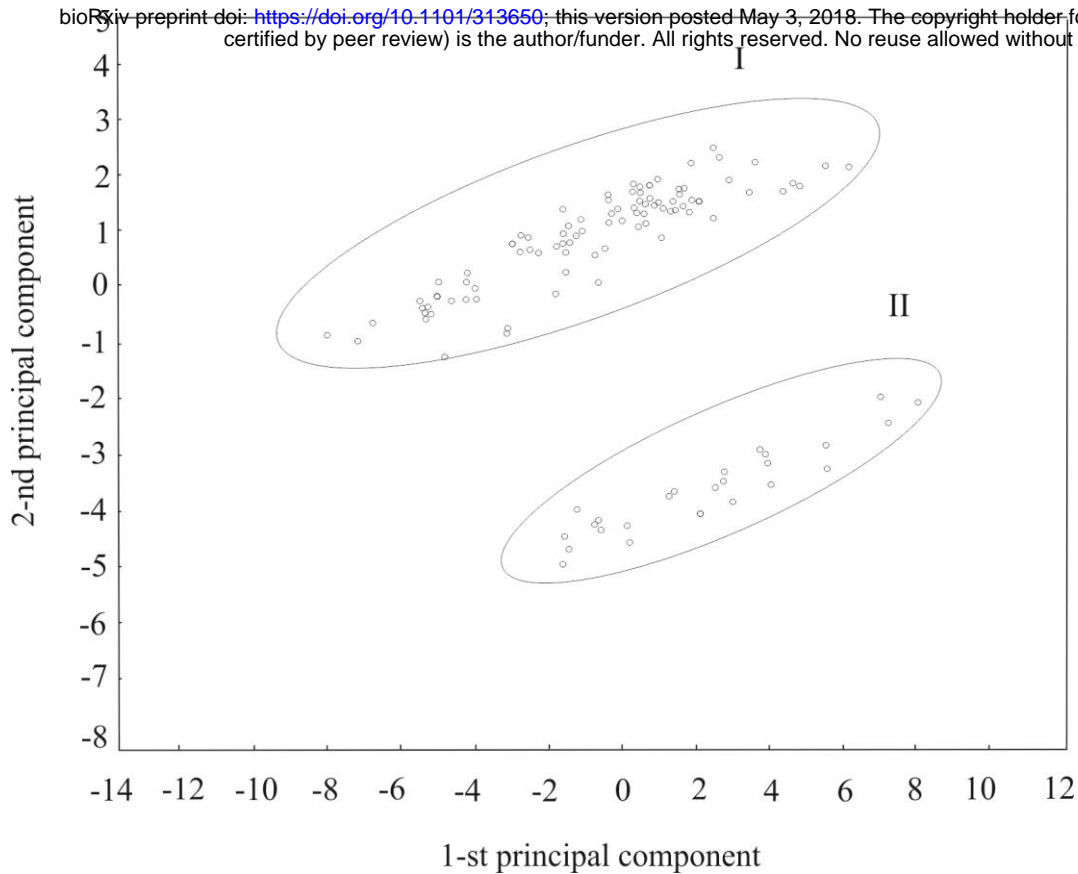


Fig. 5. Location of sampling points in the space of two principal components. Subset *I* includes points where the bottom cover percent with filamentous algae was from 15 to 100%: Angasolka, Marituy, Cape Tolsty, near Angara River, Krestovka, Cape Listvennichny, Emelyanovka, Ushakovka, Khara-Murin and town of Baikalsk. Subset *II* covers points where there were no filaments or their cover percent was 1-3%: Cape Ryty, Cape Zavorotny, Cape Kotelnikovskiy, Irinda Bay and Peninsula Svyatoy Nos.

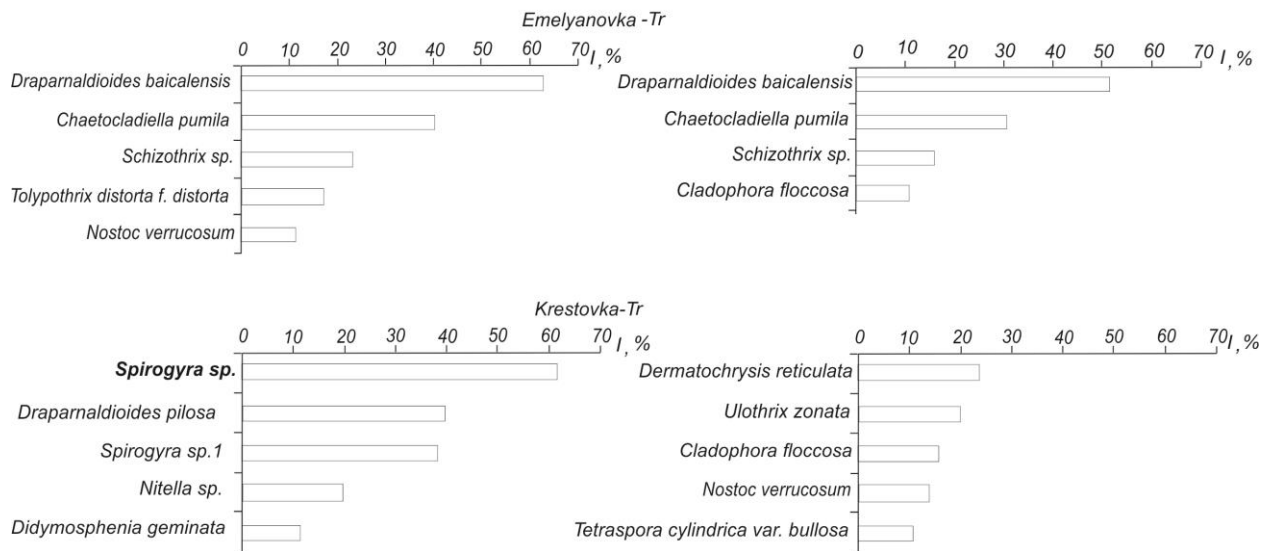


Fig. 6. Structure of meio- and macrophytes communities in the coastal zone of Lake Baikal in the area of local anthropogenic impact (Krestovka Valley, Listvennichny Bay) and background area (Emelyanovka Valley) during different years (a) – in 2016; (b) – in 1987. Along X-axis – density index; along Y-axis – species ranking in the order of decrease of density index.