- 1 Diversity and composition of cave methanotrophic communities
- 2
- 3 Kevin D. Webster,<sup>a</sup>\*# Arndt Schimmelmann,<sup>a</sup> Agnieszka Drobniak,<sup>b</sup> Maria Mastalerz,<sup>b</sup>
- 4 Laura Rosales Lagarde,<sup>c</sup> Penelope J. Boston,<sup>d</sup> Jay T. Lennon,<sup>e</sup>#
- 5
- <sup>6</sup> <sup>a</sup>Department of Earth and Atmospheric Sciences, Indiana University, Bloomington,
- 7 Indiana, USA.
- <sup>8</sup> <sup>b</sup>Indiana Geological and Water Survey, Indiana University, Bloomington, Indiana, USA.
- 9 <sup>c</sup>Department of Physical and Life Sciences, Nevada State College, School of Liberal Arts
- 10 and Sciences, Henderson, Nevada, USA.
- <sup>d</sup>NASA Astrobiology Institute, NASA Ames Research Center, Moffett Field, California,
  USA.
- <sup>e</sup>Department of Biology, Indiana University, Bloomington, Indiana, USA.
- 14 \*Present address: Kevin D. Webster, Planetary Science Institute, Tucson, AZ, USA.
- 15 #Address correspondence to Kevin D. Webster, webster@psi.edu, or Jay T. Lennon,
- 16 lennonj@indiana.edu.
- 17
- 18

19 **ABSTRACT** Methane oxidizing microorganisms (methanotrophs) are a major sink for the 20 greenhouse gas methane (CH<sub>4</sub>), and have been investigated in several environments. 21 Recent studies show that CH<sub>4</sub> consumption in caves is pervasive and is a result of active 22 methanotrophy. However, little is known about what controls the distribution and 23 abundance of methanotrophs in caves. We sampled 42 sediments from 21 caves in North 24 America to elucidate the factors shaping cave methanotroph communities. We 25 hypothesized that cave methanotroph communities should be related to cave-air CH<sub>4</sub> 26 concentrations and exhibit dispersal-limited biogeographical patterns due to the insular 27 nature of caves. Using 16S rRNA sequencing, we recovered methanotrophs from 88 % of 28 samples collected, including locations in caves where CH<sub>4</sub> concentrations were at or below 29 detection limits ( $\leq 0.3$  ppmv). Methanotrophs from the Methylocystaceae (Type II) were 30 the dominant methanotrophs as has been seen in other environments with low CH<sub>4</sub> 31 concentrations. Despite being insular ecosystems, we found that the composition of 32 methanotrophs did not vary with distance, both within and among caves. Instead, we found 33 evidence for a core microbiome, perhaps suggesting that high-affinity methanotrophs are 34 not dispersal limited. Additionally, we observed that the relative abundance of 35 methanotrophs was positively related the proportion of gravel in cave sediments and the 36 relative abundance of methylotrophs. Last, we found that the relative abundance of 37 methanotrophs was inversely correlated with cave-air CH<sub>4</sub> concentrations. Our results 38 suggest that methanotrophs in caves have influences on cave biogeochemistry beyond CH<sub>4</sub> 39 oxidation and that high-affinity methanotrophs may disperse easily into caves.

41	IMPORTANCE Recent observations have shown that the atmospheric greenhouse gas
42	methane (CH <sub>4</sub> ) is consumed by microorganisms (methanotrophs) in caves at rates
43	comparable to CH <sub>4</sub> oxidation in surface soils. Caves are abundant in karst landscapes that
44	comprise 14 % of Earth's land surface area, and therefore may be acting as a substantial
45	CH <sub>4</sub> sink. A detailed ecological understanding of the forces that shape methanotrophic
46	communities in caves is lacking. We sampled cave sediments to better understand the
47	community composition and structure of cave methanotrophs. Our results show that the
48	relative abundance of methanotrophs was negatively correlated with CH <sub>4</sub> concentrations in
49	cave air and that methanotrophic communities were similar to each other over distances of
50	10s of m to 100s of km.
51	
52	KEYWORDS Biogeography; Cave; Greenhouse Gas; Karst; Methane; Methanotroph;
53	

### 54 **INTRODUCTION**

55 Methane (CH<sub>4</sub>) is a greenhouse gas in Earth's atmosphere, an energy source for 56 humanity, and has been observed in the atmosphere of Mars (1-3). Thus, enhancing the 57 knowledge of the pathways that add and remove CH<sub>4</sub> to and from Earth's atmosphere has 58 implications for several fields. Two major processes are responsible for the removal of 59 CH<sub>4</sub> from Earth's atmosphere. Oxidation by hydroxyl radicals accounts for 90 % of 60 atmospheric-CH<sub>4</sub> removal and methane-consuming microorganisms (i.e., methanotrophs) 61 in surface soils account for an additional 5 % (4). Despite the influence that methanotrophs 62 have on regulating Earth's climate, gaps in our understanding of these organisms, like how 63 they respond to changing CH<sub>4</sub> concentrations or how they disperse through the

64 environment, remain (2, 3).

65	Rocks that host caves are widespread on Earth, have been observed on other
66	planets, and are emerging as a global CH <sub>4</sub> sink (7–10) (Figure 1). The relative importance
67	of deterministic and stochastic factors responsible for shaping the methanotroph
68	community in caves is unknown (7, 9, 11–14). Observations of atmospheric CH <sub>4</sub>
69	consumption in caves suggest that high-affinity or Type II methanotophs may be present in
70	cave environments. These organisms have broad phylogenetic affiliations (15-20), i.e. are
71	known from the Methylocystaceae, as well as several groups in clades affiliated with
72	upland soils (upland soil cluster methanotrophs). They tend to be the most prominent
73	members of the methanotrophic community in environments with low CH <sub>4</sub> concentrations
74	and high $O_2$ concentrations. They are less common in environments with high $CH_4$ and
75	low O <sub>2</sub> concentrations, like lakes or geological seeps, which are dominated by
76	methanotrophs from the Methylococcales (Type I methanotrophs) (21, 22).
77	The factors that control the distribution and abundance of methanotrophs in caves
78	are unknown. Previous work has shown that methanotrophs are frequently related to $CH_4$
79	concentrations, sediment texture, and other microorganisms, but it is unknown how these
80	factors influence methanotroph communities in caves. Type I and Type II methanotrophs
81	are known to exhibit niche partitioning between high and low CH4 subenvironments
82	respectively (22, 23). However, it is unknown if Type II methanotrophs show similar
83	deterministic changes in community composition with atmospheric to subatmospheric CH <sub>4</sub>
84	concentration gradients. Methanotrophs have also been shown to be more abundant in
85	clay-rich soils even though these soils may limit the diffusion of oxygen (24–26). Prior

86	work has documented associations between methanotroph and methylotroph abundances in
87	aquatic environments (18), and that methanotrophs interact with other microbial species
88	through volatile compounds in co-cultures (27).

89	Almost nothing is known about the importance of biogeography in structuring
90	high-affinity methanotroph communities. Previous observations have shown that
91	methanotrophs from Hawai'i exhibited close taxonomic affiliations with Arctic
92	methanotrophs (5). This suggests that these organisms may be able to readily disperse
93	through the environment. However, biota in caves are typically unique due to their
94	isolation from other environments (28, 29). Thus, it may be possible that dispersal
95	limitation is important in structuring high-affinity methanotrophic communities in caves.
06	We compled 21 caves in North America, analyzed the 16S rDNA gape, and
96	We sampled 21 caves in North America, analyzed the 16S rRNA gene, and
96 97	We sampled 21 caves in North America, analyzed the 16S rRNA gene, and developed the following hypotheses to examine the factors regulating the community
97	developed the following hypotheses to examine the factors regulating the community
97 98	developed the following hypotheses to examine the factors regulating the community composition of cave methanotrophs (Figure 1). First, we hypothesized that their
97 98 99	developed the following hypotheses to examine the factors regulating the community composition of cave methanotrophs (Figure 1). First, we hypothesized that that their community structure should mirror the structure of methanotrophic communities in other

- 103 hypothesized that high-affinity methanotroph communities in caves should be dispersal-
- 104 limited due to the insular nature of caves.

105

### 106 **RESULTS**

107	Methanotroph community structure. Methanotrophs were recovered in 88 % of
108	the cave-sediment samples including locations where the CH4 concentrations were at or
109	below our instrumental detection limits of $\leq 0.3$ ppmv (Supplemental Table 1). In samples
110	where methanotrophs were recovered, their relative abundances ranged from $1.1*10^{-5}$ to
111	$2.2*10^{-2}$ ( $Q_1 = 5.4*10^{-4}$ , $Q_2 = 2.3*10^{-3}$ , $Q_3 = 4.7*10^{-3}$ ). Cave methanotrophic communities
112	were dominated by the Methylocystaceae, but members of the Methylococcales were also
113	present (Figure 2). The median abundance of the Methylocystaceae at each site was 96 %
114	of the total methanotrophic community. Members of the Methylocystaceae accounted for
115	51 unique operational taxonomic units (OTUs). Unknown OTUs at the genus level
116	accounted for 94 % of the OTUs and virtually all (99.99 %) of the observed sequences
117	from the Methylocystaceae. The Methylococcales were more diverse than the
118	Methylocystaceae with 69 unique OTUs but only contributed 997 DNA sequences. 21
119	OTUs within the Methylococcales were identified from two undescribed families and
120	accounted for 45 % of the total observed Methylococcales.
121	Methanotrophs from other clades constituted much less of the observed diversity.
122	Seven OTUs were observed from the genus Methylibium (beta Proteobacteria:
123	Comamonadaceae), and one DNA sequence from a member of the genus Methylocella
124	(Beijerinckiaceae) was observed. Two OTUs were found from the family
125	Acidimethylosilex (Verrucomicrobia) in Cave 2 in Indiana. We did not recover members
126	from the candidate phylum NC10, or members of the USC methanotrophs.
127	Environmental influences on community structure. The CH <sub>4</sub> concentration in
128	cave air was correlated with the relative abundance of methanotrophs. The relative
129	abundances of both the Methylocystaceae ( $log_{10}(RA) = -0.5 \pm 0.3*[CH_4] - 2.4 \pm 0.4$ , $r^2 =$

0.24,  $p = 1*10^{-3}$ ) and Methylococcales ( $log_{10}(RA) = -0.5 \pm 0.2*[CH_4] - 3.6 \pm 0.3$ ),  $r^2 = -0.5 \pm 0.2*[CH_4] - 3.6 \pm 0.3$ 130 131 0.41,  $p = 1*10^{-4}$ ) increased with decreasing CH<sub>4</sub> concentration (Figure 2). An analysis of 132 covariance (ANCOVA) test revealed that there was no significant difference between the 133 slopes of the two clades (ANCOVA, F = 1.6, p = 0.21). 134 We investigated the influence of the amount of CH<sub>4</sub>, gravel, sand, silt, and clay on 135 methanotroph community composition using canonical correspondence analysis. As a 136 whole, the composition of the methanotrophic community was not related to 137 concentrations of CH<sub>4</sub>, or the volumetric proportions of gravel, sand, silt, and clay at the 138 sample locations. However, the community composition of the Methylococcales was 139 related to CH<sub>4</sub> concentrations in cave air (Canonical Correspondence Analysis, p = 0.001) 140 and the amount of clay at the sample location (p = 0.008) (Figure 4). OTUs within the total 141 methanotrophic community and within Methylocystaceae were not related to any of the 142 measured environmental parameters. 143 We tested if the proportions of gravel, sand, silt, and clay influenced the relative 144 abundance of methanotrophs using Spearman's rank correlation tests. Methanotrophs from 145 the Methylocystaceae, and thus of methanotrophic communities overall, were correlated 146 with the volumetric abundance of gravel in a sample (Spearman's rank correlation, 147 Methylocystaceae,  $\rho = 0.39$ , S = 10231, p = 0.02) (Figure 3). The relative abundance of 148 methanotrophs from the Methylocystaceae was not related to any other sediment size 149 fraction present in the samples. Furthermore, the relative abundance of the 150 Methylococcales was not related to any sediment size fraction (Spearman's rank

151 correlations, vs. Gravel p = 0.7; vs. Sand p = 0.12; vs. Silt p = 0.76; vs. Clay p = 0.90).

We examined potential interactions between methanotrophs and other microorganisms by testing if the relative abundance of methanotrophs was related to the abundance of methylotrophs in caves. The relative abundance of methanotrophs was positively correlated to the relative abundance of methylotrophs in the sampled cave sediments (Spearman's rank correlation,  $\rho = 0.71$ , S = 3849,  $p = 1*10^{-7}$ ) (Figure 5). The relative abundance of methylotrophs was not, however, related to CH<sub>4</sub> concentrations in cave air (Spearman's rank correlation,  $\rho = -0.25$ , S = 8914, p = 0.15).

159 **Biogeographical influences on community structure.** We tested the ability of 160 methanotrophs to disperse in the environment by measuring how the methanotroph 161 community changed with measures of geographic distance. Cave methanotroph 162 communities did not show significant relationships to measures of geographic distance. 163 The relative abundance of methanotrophs was not correlated with the distance from a cave 164 entrance (Spearman's rank correlation,  $\rho = 0.33$ , S = 2432, p = 0.08). Furthermore, 165 distance-decay analysis—a measure of how the spatial distance between two samples 166 affects their similarity—of the total methanotrophic community assemblage was not 167 statistically significant (Mantel test, r = 0.08, p = 0.14).

We also tested methanotroph dispersal by determining if a core group of methanotrophs was present in the samples. We defined a core community as taxa present in at least 60 % of samples. If no group of methanotrophs could consistently be found among caves, this may suggest that methanotrophs are dispersal limited. We observed a core methanotroph biome of 3 OTUs. All 3 OTUs were from unidentified genera in the Methylocystaceae. Together these taxa were present in 88 % of samples, and one OTU

was present in 86 % of all the samples. The core methanotroph biome accounted for 59 %of the obtained methanotroph sequences.

176

#### 177 **DISCUSSION**

178 Methanotroph community structure in caves. Methanotrophs are widespread in 179 cave environments. The dominance of the Methylocystaceae in cave methanotroph 180 communities is consistent with observations that have been made in other low CH<sub>4</sub> 181 environments. For example, high-affinity methanotrophs are observed to be dominant in 182 aerated riparian environments, Arctic soils, and soils associated with basaltic rocks (5, 23, 183 30, 31). Additionally, the relative abundance of methanotrophs observed in the sampled 184 cave sediments is similar to the relative abundance of methanotrophs in Arctic soils (0.026 185 to 0.589 %) (32), and from caves in Vietnam (9). However, we did not recover any USC 186 methanotrophs as have been observed in some soils that consume atmospheric CH<sub>4</sub> (16, 187 33). The lack of USC methanotrophs in our study contrasts with the methanotrophic 188 community structure present in a cave from Australia where USC methanotrophs 189 comprised 2 to 12 % of the community (10) and observations of USC- $\alpha$  methanotrophs in 190 several caves (19). Additionally, our results are somewhat inconsistent with observations 191 of methanotrophic communities from soils overlying caves in Australia where 192 Methylocella sp. and Crenothrix sp. comprised 12 to 15 % of the community (12). 193 The cave sediments in our study also contained relatively rare methanotrophs, but 194 the overall diversity of these taxa exceeded that of the dominant taxa. Minor contributors 195 to the methanotroph community consisted of the Methylococcales, as well as 196 methanotrophs from the Verrucomicrobia. The diversity of methanotrophs identified in our

study caves contrasts with the relatively depauperate assemblage found in a Spanish cave,
which reportedly consisted of only three species, *Methylocapsa aurea*, *Methylomicrobium album*, and *Methylococcus capsulatus* (34). Our results suggest that CH<sub>4</sub> oxidation in
caves may be carried out by a relatively diverse yet uncharacterized methanotrophic
bacteria.

Environmental influences on methanotroph communities. The deterministic relationships between cave methanotrophs and their environment may be best understood from the perspective of metabolism. The process of life is fundamentally centered on obtaining energy from the surrounding environment and methanotrophs are no exception. Their distributions in caves appear to result from the fact that methanotrophs rely on both CH<sub>4</sub> and oxygen to derive energy.

208 The relative abundance of methanotrophs was inversely related to CH<sub>4</sub> 209 concentrations in cave air. This pattern was documented for two major groups of 210 methanotrophs, the Methylocystaceae and the Methylococcales. While the relative 211 abundance of methanotrophs is sometimes known to positively track CH<sub>4</sub> concentrations 212 (35), methanotrophs have been observed to be negatively correlated with CH<sub>4</sub> 213 concentrations in some environments including glacial termini, geologic seeps, and a cave 214 (10, 22, 30, 36). One explanation for the inverse CH<sub>4</sub>-methanotoph relationship is that 215 methanotrophs are actively driving down CH<sub>4</sub> concentrations in locations where they make 216 up a large fraction of the total microbial community. Another possibility is that 217 methanotroph abundance in caves may be positively related to CH<sub>4</sub> fluxes, but our 218 observed pattern is a result of seasonality in cave-air flow. Many caves are known to 219 exhibit faster air flow when the external atmospheric temperature is lower than the

temperature of cave air. This leads to higher CH<sub>4</sub> concentrations in caves during the winter
(13, 37). Our samples were acquired during summer, so it is possible that our sites
experience higher CH<sub>4</sub> concentrations during the winter and thus greater CH<sub>4</sub> fluxes on an
annual basis which may explain the observed negative relationship.

224 The permeability of cave sediments also appeared to influence the methanotroph 225 community. The relative abundance of methanotrophs in caves was positively related to 226 the amount of gravel in the samples. This may be due to the fact that cave sediments with 227 greater portions of larger clasts, like gravel, may decrease the water-holding capacity while 228 increasing the permeability of cave sediments and the diffusion of gasses (38). As the 229 proportion of larger clasts in a matrix increases, the permeability of the sediment tends to 230 increase (38). This gives methanotrophs greater access to  $CH_4$  and  $O_2$  which are crucial for 231 their metabolism. Previous research observed that methanotrophs tend to be more 232 abundant in silt and clay fractions of soils because these fractions made up the bulk of the 233 soils themselves (24). This previous study did not characterize larger sediment sizes, nor 234 did it examine clay floccules that may serve to enhance permeability.

235 The abundance of clay particles influenced the community composition of the 236 Methylococcales (a less dominant group of methanotrophs in this study) as did the  $CH_4$ 237 concentration of cave air. These relationships suggest that members of the 238 Methylococcales may be partitioning themselves across different niches with regard to 239 CH<sub>4</sub> and O<sub>2</sub> availability. The Methylococcales are typically associated with environments 240 that have low O<sub>2</sub> and high CH<sub>4</sub> concentrations. Members of Methylococcales may 241 therefore be responding to dysoxic to hypoxic environments in cave sediments that result 242 from tightly packed clay-sized particles and/or extra-atmospheric CH<sub>4</sub> sources (11, 13, 14).

243	The community gradient across CH4 and clay particle size may represent the transition
244	from potentially unknown members in the Methylococcales that consume CH <sub>4</sub> at low (i.e.
245	atmospheric) concentrations to those more tolerant of lower O2 concentrations.
246	Organisms alter the physical and chemical characteristics of the environment and
247	thus may influence the distributions of other species in the environment. The positive
248	correlation between methylotrophs and methanotrophs is consistent with observations of
249	metabolic interactions between these functional groups of bacteria. For example,
250	Methylotrophs are known to consume methanol that is produced by methanotrophs (39).
251	Co-occurrences between members of Methylococcales and Methylophilaceae
252	(methylotrophic) have been observed in environmental samples, and so has the transfer of
253	methanol from methanotrophs to methylotrophs (40, 41). A similar process may be
254	occurring in caves. Alternatively, methylotrophs may simply be responding to the presence
255	of other organic compounds in cave air besides CH4. In addition to consuming methanol,
256	methylotrophs are known to consume methyl halides, like methyl chloride and methyl
257	bromide which are present as trace gases in the atmosphere (42). These gases would be
258	advected into cave air in the same fashion as atmospheric CH <sub>4</sub> (43).
259	Biogeography of methanotroph communities. Methanotrophic communities did
260	not show statistically significant relationships with local or regional assessments of

not show statistically significant relationships with local or regional assessments of
geographic distance. The composition of methanotrophic communities close to cave
entrances was not different compared to those in more interior locations. This pattern has
several possible explanations, but two seem to be the most probable. The methanotrophic
community in caves is dominated by Type II methanotrophs and these methanotrophs do
best under environmental conditions of low CH<sub>4</sub> concentrations and high O<sub>2</sub>

266	concentrations. The change from about 2 to $< 0.3$ ppmv CH <sub>4</sub> in cave air may not impose
267	strong selective pressure on Type II methanotrophs. Additionally, the lack of a pattern in
268	methanotrophic community composition may be a result of the seasonality present in cave
269	air (44, 45). Thus, CH <sub>4</sub> concentrations in cave air may be near the atmospheric
270	concentration at one point in the year and shift to sub-atmospheric concentrations at other
271	times. This could cause location to be a poor predictor of the methanotrophic assemblage
272	due to microbial physiological states like dormancy (46).
273	Likewise, cave methanotrophic communities were not spatially autocorrelated at
274	the regional scale. The lack of a distance-decay relationship in methanotrophic
275	communities from the studied caves may indicate that the scale at which distance-decay
276	relationships occur in these communities is larger than the regional scale of 500 km in our
277	study. This may suggest that dispersal is high among methanotrophs.
278	We found a core group of three OTUs that were present in the studied caves. This
278 279	We found a core group of three OTUs that were present in the studied caves. This suggests that high-affinity methanotrophs are able to move about the environment
279	suggests that high-affinity methanotrophs are able to move about the environment
279 280	suggests that high-affinity methanotrophs are able to move about the environment relatively easily. The fact that these methanotrophs accounted for 59 % of the observed
279 280 281	suggests that high-affinity methanotrophs are able to move about the environment relatively easily. The fact that these methanotrophs accounted for 59 % of the observed methanotrophs suggests these organisms are reproducing in these environments as well.
279 280 281 282	suggests that high-affinity methanotrophs are able to move about the environment relatively easily. The fact that these methanotrophs accounted for 59 % of the observed methanotrophs suggests these organisms are reproducing in these environments as well. Previous observations are also indicative of high-dispersal capabilities for these organisms.
<ol> <li>279</li> <li>280</li> <li>281</li> <li>282</li> <li>283</li> </ol>	suggests that high-affinity methanotrophs are able to move about the environment relatively easily. The fact that these methanotrophs accounted for 59 % of the observed methanotrophs suggests these organisms are reproducing in these environments as well. Previous observations are also indicative of high-dispersal capabilities for these organisms. For example, methanotrophs from a newly formed soil in Hawai'i showed close taxonomic
<ol> <li>279</li> <li>280</li> <li>281</li> <li>282</li> <li>283</li> <li>284</li> </ol>	suggests that high-affinity methanotrophs are able to move about the environment relatively easily. The fact that these methanotrophs accounted for 59 % of the observed methanotrophs suggests these organisms are reproducing in these environments as well. Previous observations are also indicative of high-dispersal capabilities for these organisms. For example, methanotrophs from a newly formed soil in Hawai'i showed close taxonomic affiliations with Arctic methanotophs (5). Additionally, microbes are known to travel

other habitable subenvironments, like recurring slope lineae or caves on Mars (8, 48), may
be decreased.

290

### 291 CONCLUSIONS

292 The structure of cave methanotroph communities appears to be best understood in 293 terms of how cave subenvironments provide methanotrophs with access to CH<sub>4</sub> and 294 oxygen. Cave methanotrophic communities showed relationships with cave-air CH<sub>4</sub> 295 concentrations and the abundance of gravel-sized clasts in cave sediments. The CH<sub>4</sub> 296 concentrations of cave air and the abundance of gravel both influence how much CH<sub>4</sub> and 297 O<sub>2</sub> methanotrophs have access to. CH<sub>4</sub> is typically at low (i.e., atmospheric) concentrations 298 in caves and as such, Methanotrophs from the Methylocystaceae were numerically 299 dominant. This community structure mirrors that of other atmospheric CH<sub>4</sub> oxidizing 300 communities (9, 32). Further, the community composition of the Methylococcales showed 301 responses to CH<sub>4</sub> concentrations and the abundance of clay in cave sediments that could 302 represent a transition from uncultivated members in the Methylococcales that consume 303  $CH_4$  at low (i.e. atmospheric) concentrations to those more tolerant of lower  $O_2$ 304 concentrations.

305 Our results suggest that methanotrophs in caves have secondary influences on the 306 cave-microbial community structure. The abundance of known methylotrophic organisms 307 is correlated with the abundance of methanotrophs. This suggests that molecules produced 308 along the methane oxidation pathway, like methanol, are leaked from methanotrophic cells

allowing for the growth of methylotrophs. Our findings suggest that CH<sub>4</sub> may be an

310 overlooked energy source for cave communities.

Methantrophs did not show relationships with measures of geographic distance in this study. This study focused on environments that are thought to present atmospheric to subatmospheric CH<sub>4</sub> concentrations year-round and in sediments that are forming from the breakdown of similar rocks. The similarity of the environments may be one reason for the observed lack a relationship with measures of distance. This result also suggests that the dispersal barrier for high-affinity methanotrophs from cave to cave may be low.

317

### 318 MATERIALS AND METHODS

Microbial Sampling. We sampled microbial communities from caves along the western front of the Appalachian fold and thrust belt, in intracratonic settings of the USA, and in the Sierra Madre of Mexico. We obtained 42 sediment samples from caves along transects from cave entrances to interiors. Samples were scraped from ~0.1 m<sup>2</sup> large areas of cave walls and floors using a spatula that had been sterilized with 70 volume % ethanol in water (49). Samples were stored on ice until they could be transferred to a -80 °C freezer.

Environmental Parameters. We measured multiple environmental parameters to
assess factors that potentially influence the composition of cave methanotrophic
communities. We measured CH<sub>4</sub> concentrations *in-situ* with Fourier Transform Infrared
Spectroscopy (FTIR) (Gasmet DX4030 – Milton Keynes, United Kingdom), or in the
laboratory with FTIR gas-chromatography (Varian – Agilent Technologies, Palo Alto,

331 California) using discrete air samples collected in the field. Some of the CH<sub>4</sub> 332 concentrations listed in this study have been previously published (14, 50), and new data 333 were collected according to methods reported in the same publications. In cases where 334 cave maps were available, the distance from the nearest entrance to the sampling location 335 was calculated along the length of the cave passages. Cave sediment grain-size 336 distributions were measured with a Malvern Mastersizer 3000 (Malvern Instruments Inc., 337 Westborough Massachusetts). Raw data from the Mastersizer were converted to % gravel, 338 % sand, % silt, and % clay sized particles by volume using the GRADISTAT software 339 package (51). Metadata used for statistical analyses are reported in Supplemental Table 2. 340 Molecular Techniques. Genomic DNA was extracted from cave sediment samples 341 with a MoBio PowerSoilTM extraction kit (MoBio, Carlsbad, California USA). About 10 342 ng of extracted DNA was used as a template for amplification by polymerase chain 343 reaction (PCR). The V4 region of the 16S rRNA gene was amplified using 515F and 806R 344 with barcoded primers designed to work with the Illumina MiSeq platform (52). DNA 345 amplification was performed using 50  $\mu$ L reactions from each sample. PCR reactions were 346 carried out by a 3 min denaturing step at 94 °C, followed by 30 cycles of 94 °C for 45 s, 50 347 °C for 60 s, and 72 °C for 90 s. A final 10-minute extension was carried out at 72 °C. 348 Quality of the PCR amplifications was checked by gel electrophoresis. Amplified DNA 349 was cleaned using a commercial kit (Beckman Coulter Agencourt AMPure XP, 350 Indianapolis, Indiana, USA). Cleaned DNA from the reactions was pooled to a final 351 concentration of 20 ng per sample. 352 We sequenced the cleaned PCR amplicons using Illumina MiSeq technology 353 (Illumina Reagent Kit v2, 500 reaction kit) at the Center for Bioinformatics and Genomics

354	at Indiana University. Data quality and unique sequences obtained from the PCR
355	amplifications were analyzed using mothur (53). DNA sequence data were aligned using
356	the Needleman algorithm. Sequences matching chimeras were removed using UCHIME
357	(54). Sequences that matched chloroplasts, Archaea, and other non-bacterial sequences
358	were also removed. OTUs were created by binning the data at 97 % sequence similarity.
359	OTUs were identified using a SILVA reference database (version 128).
360	Statistical Analyses. The R computing environment was used for all quantitative
361	analyses (55). We used 16S rRNA gene sequences to assess diversity and composition of
362	the methanotroph community. Sequences related to known methanotrophs and
363	methylotrophs were subset from the larger 16S rRNA dataset using the "subset_taxa"
364	function in the <i>phyloseq</i> software package to create OTU abundance tables and for only
365	the taxa of interest for a particular sample (56). Relationships between methylotrophs and
366	CH <sub>4</sub> concentrations were examined due to the similarities of their metabolic pathways with
367	methanotrophs. We defined the methanotrophic community as members of the
368	Methylocystaceae, Methylococcales, Methylocella, USC- $\alpha$ , and USC- $\gamma$ (17, 18). We
369	defined the methylotrophic community as members of the Methylophilaceae and
370	Methylobacteriaceae.
371	We assessed the relationships between methanotrophs and environmental
372	conditions through a series of statistical analyses. Raw abundance data were transformed
373	into fractional abundance data to avoid variations in the number of reads between samples.
374	Community dissimilarity between samples was assessed with Bray-Curtis dissimilarity.

375 Dissimilarity measures were only calculated if the number of OTUs in a phylogenetic

376 grouping was greater than 100 OTUs to avoid statistical problems arising from small

377	sample size. The effect of CH <sub>4</sub> concentrations on the methanotrophic community was
378	assessed using ANCOVA and Mantel tests. We tested the effect of sediment grain size on
379	cave methanotrophic communities using canonical correspondence analysis that was
380	performed in the R package vegan (57).
381	The spatial autocorrelation of methanotrophic communities was tested by
382	performing distance-decay analyses. Distance-decay is a spatial-autocorrelation pattern
383	that describes the decreasing similarity of biological communities with increases in
384	geographic distance (58). We performed distance-decay analysis in R by first removing
385	data from Cave 38 (Cueva de Villa Luz) since it was roughly 2000 km away from the
386	nearest samples. We then tested for a distance-decay pattern using a Mantel test in the
387	package vegan (57).
388	We characterized the methanotroph core microbiome, which is defined as a suite of
389	members shared in an environment (59). We calculated the core microbiome based on a
390	cutoff of 60 %, a value that is in the middle of ranges that have used when calculating core
391	microbiomes in other ecosystems including corals and humans (60, 61). We also chose this
392	value because it is larger than the majority of samples, but accounts for some
393	unpredictability in an organism's distribution.
394	Data and software availability. All code and data used in this study can be found
395	in a public GitHub repository ( <u>https://github.com/websterkgd/CaveMethanotrophs</u> ).
396	

# 397 ACKNOWLEDGEMENTS

This material is based upon work supported by the U.S. Department of Energy,
Office of Science, Office of Basic Energy Sciences under Award Number DE-SC0006978

400	to AS	. Additional support was provided by a National Speleological Society research grant		
401	awarded to KDW, the National Geographic Society Expeditions Council (Grant #EC0644-			
402	13 to PJB), the National Science Foundation Dimensions of Biodiversity, Grant/Award			
403	Number: 1442246 to JTL, and a US Army Research Office, Grant/Award Number:			
404	W911NF-14-1-0411 to JTL. We thank Brent Lehmkuhl for microbiological support.			
405	Members of the Bloomington Indiana Grotto facilitated access to caves in Indiana and			
406	Kentucky. We thank Rodolfo Gómez Cruz for his support at Cueva de Villa Luz (Cave			
407	38). ML Larsen, MM Muscarella, and KJ Locey aided in data analysis and interpretation.			
408	LA Durden assisted in editing an early draft of this manuscript. The authors declare no			
409	conflicts of interest.			
410				
411	REFERENCES			
412	1.	Ciais P, Sabine C, Bala G. 2013. Carbon and other biogeochemical cycles, p. 465–		
413		570. In Stocker, TF, Qin, D, Pattner, G-K, Tignor, M, Allen, SK, Boschung, J,		
414		Nauels, A, Xia, X, Bex, V, Midgley, PM (eds.), Climate Change 2013: The Physcial		
415		Science Basis. Contribution of the Intergovernmental Panel on Climate Change.		
416		Cambridge University Press, Cambride, United Kindom and New York, NY, USA.		

- 417 2. Schimmelmann A, Ensminger SA, Drobniak A, Mastalerz M, Etiope G, Jacobi RD,
- 418 Frankenberg C. 2018. Natural geological seepage of hydrocarbon gas in the
- 419 Appalachian Basin and Midwest USA in relation to shale tectonic fracturing and
- 420 past industrial hydrocarbon production. Sci Total Environ 644:982–993.
- 421 3. Webster CR, Mahaffy PR, Atreya SK, Moores JE, Flesch GJ, Malespin C, Mckay

422		CP, Martinez G, Smith CL, Martin-Torres J, Gomez-Elvira J, Zorzano M-P, Wong
423		MH, Trainer MG, Steele A, Acher Jr. D, Sutter B, Coll PJ, Freissinet C, Meslin P-
424		Y, Gough R V., House CH, Pavlov A, Eigenbrode JL, Glavin DP, Pearson JC,
425		Keymeulen D, Christensen LE, Schwenzer SP, Navarro-Gonzalez R, Pla-Garía J,
426		Rafkin SCR, Vicente-Retorillo Á, Kahanapää H, Viudez-Moreiras D, Smith MD,
427		Harri A-M, Genzer M, Hassler DM, Lemmon M, Crisp J, Sander SP, Zurek RW,
428		Vasavada AR. 2018. Background levels of methane in Mars' atmosphere show
429		strong seasonal variations 360:1093-1096.
430	4.	Kirschke S, Bousquet P, Ciais P, Saunois M, Canadell JG, Dlugokencky EJ,
431		Bergamaschi P, Bergmann D, Blake DR, Bruhwiler L, Cameron-Smith P, Castaldi
432		S, Chevallier F, Feng L, Fraser A, Heimann M, Hodson EL, Houweling S, Josse B,
433		Fraser PJ, Krummel PB, Lamarque J-F, Langenfelds RL, Le Quéré C, Naik V,
434		O'Doherty S, Palmer PI, Pison I, Plummer D, Poulter B, Prinn RG, Rigby M,
435		Ringeval B, Santini M, Schmidt M, Shindell DT, Simpson IJ, Spahni R, Steele LP,
436		Strode SA, Sudo K, Szopa S, van der Werf GR, Voulgarakis A, van Weele M,
437		Weiss RF, Williams JE, Zeng G. 2013. Three decades of global methane sources
438		and sinks. Nat Geosci 6:813-823.
439	5.	King GM, Nanba K. 2008. Distribution of atmospheric methane oxidation and
440		methanotrophic communities on hawaiian volcanic deposits and soils. Microbes
441		Environ 23:326–330.
442	6.	Degelmann DM, Borken W, Drake HL, Kolb S. 2010. Different atmospheric
443		methane-oxidizing communities in european beech and norway spruce soils. Appl

- 444 Environ Microbiol 76:3228–3235.
- Ford, Derek C; Williams P. 2013. Karst Hydrogeology and Geomorphology. Jon
  Wiley & Sons, London, England.
- 447 8. Cushing GE, Titus TN, Wynne JJ, Christensen PR. 2007. THEMIS observes
- 448 possible cave skylights on Mars. Geophys Res Lett 34:4–8.
- 449 9. Lennon JT, Nguyễn-Thùy D, Phạm TM, Drobniak A, Tạ PH, Phạm NĐ, Streil T,
- 450 Webster KD, Schimmelmann A. 2017. Microbial contributions to subterranean
- 451 methane sinks. Geobiology 15:254–258.
- 452 10. Waring CL, Hankin SI, Griffith DWT, Kertesz MA, Kobylski V, Wilson NL,
- 453 Coleman N V., Kettlewell G, Zlot R, Bosse M, Bell G. 2017. Seasonal total
- 454 methane depletion in limestone caves. Sci Rep 7.
- 455 11. Mattey DP, Fisher R, Atkinson TC, Latin J-P, Durrell R, Ainsworth M, Lowry D,
- 456 Fairchild IJ. 2013. Methane in underground air in Gibraltar karst. Earth Planet Sci
  457 Lett 374:71–80.
- 458 12. McDonough LK, Iverach CP, Beckmann S, Manefield M, Rau GC, Baker A, Kelly

459 BFJ. 2016. Spatial variability of cave-air carbon dioxide and methane

- 460 concentrations and isotopic compositions in a semi-arid karst environment. Environ
  461 Earth Sci 75:700.
- 462 13. Webster KD, Mirza A, Deli JM, Sauer PE, Schimmelmann A. 2016. Consumption
  463 of atmospheric methane in a limestone cave in Indiana, USA. Chem Geol 443:1–9.
- 464 14. Webster KD, Drobniak A, Etiope G, Mastalerz M, Sauer PE, Schimmelmann A.

465		2018. Subterranean karst environments as a global sink for atmospheric methane.
466		Earth Planet Sci Lett 485:9–18.
467	15.	Ricke P, Kube M, Nakagawa S, Erkel C, Reinhardt R, Liesack W. 2005. First
468		genome data from uncultured upland soil cluster alpha methanotrophs provide
469		further evidence for a close phylogenetic relationship to Methylocapsa acidiphila
470		B2 and for high-affinity methanotrophy involving particulate methane
471		monooxygenase. Appl Environ Microbiol 71:7472–7482.
472	16.	Kolb S. 2009. The quest for atmospheric methane oxidizers in forest soils. Environ
473		Microbiol Rep 1:336–346.
474	17.	Nazaries L, Murrell JC, Millard P, Baggs L, Singh BK. 2013. Methane, microbes
475		and models: fundamental understanding of the soil methane cycle for future
476		predictions. Environ Microbiol 15:2395–2417.
477	18.	Lau E, Nolan IV EJ, Dillard ZW, Dague RD, Semple A, Wentzell WL. 2015. High
478		throughput sequencing to detect differences in methanotrophic Methylococcaceae
479		and Methylocystaceae in surface peat, forest soil, and Sphagnum moss in
480		Cranesville Swamp Preserve, West Virginia, USA. Microorganisms 3:113–136.
481	19.	Pratscher J, Vollmers J, Wiegand S, Dumont MG, Kaster A-K. 2018. Unravelling
482		the identity, metabolic potential and global biogeography of the atmospheric
483		methane-oxidizing upland soil cluster α. Environ Microbiol.
484	20.	Edwards CR, Onstott TC, Miller JM, Wiggins JB, Wang W, Lee CK, Cary SC,
485		Pointing SB, Lau MCY. 2017. Draft genome sequence of uncultured upland soil

486		cluster Gammaproteobacteria gives molecular insights into high-affinity
487		methanotrophy. Genome Announc 5:e00047-17.
488	21.	Bowman JP. 2005. Methylococcales ord. nov., p. 248–270. In Brenner, D, Krieg, N,
489		Staley, J, Garrity, G, Boone, D, De Vos, P, Goodfellow, M, Rainey, F, Schleifer, K-
490		H (eds.), Bergey's Manual of Systematic Bacteriology. Springer, Boston, MA.
491	22.	Mills CT, Slater GF, Dias RF, Carr SA, Reddy CM, Schmidt R, Mandernack KW.
492		2013. The relative contribution of methanotrophs to microbial communities and
493		carbon cycling in soil overlying a coal-bed methane seep. FEMS Microbiol Ecol
494		84:474–494.
495	23.	Krause S, Meima-Franke M, Hefting MM, Bodelier PLE. 2013. Spatial patterns of
496		methanotrophic communities along a hydrological gradient in a riparian wetland.
497		FEMS Microbiol Ecol 86:59–70.
498	24.	Bender M, Conrad R. 1994. Methane oxidation activity in various soils and
499		freshwater sediments: Occurrence, characteristics, vertical profiles, and distribution
500		on grain size fractions. J Geophys Res 99:16531–16540.
501	25.	Sessitsch A, Weilharter A, Gerzabek MH, Kirchmann H, Kandeler E. 2001.
502		Microbial population structures in soil particle size fractions of a long-term fertilizer
503		field experiment. Appl Environ Microbiol 67:4215–4224.
504	26.	Hemkemeyer M, Christensen BT, Martens R, Tebbe CC. 2015. Soil particle size
505		fractions harbour distinct microbial communities and differ in potential for
506		microbial mineralisation of organic pollutants. Soil Biol Biochem 90:255–265.

507	27	Veraart A.I.	Garbeva P	van Beersum F	Ho A	Hordii	k CA	Meima-Franke M.
507	<i>_</i> .	<i>v</i> or uur <i>i</i> 13,	Ourocvu i i		110 1 1.	inorari	$\mathbf{K} \subset \mathbf{I}$	

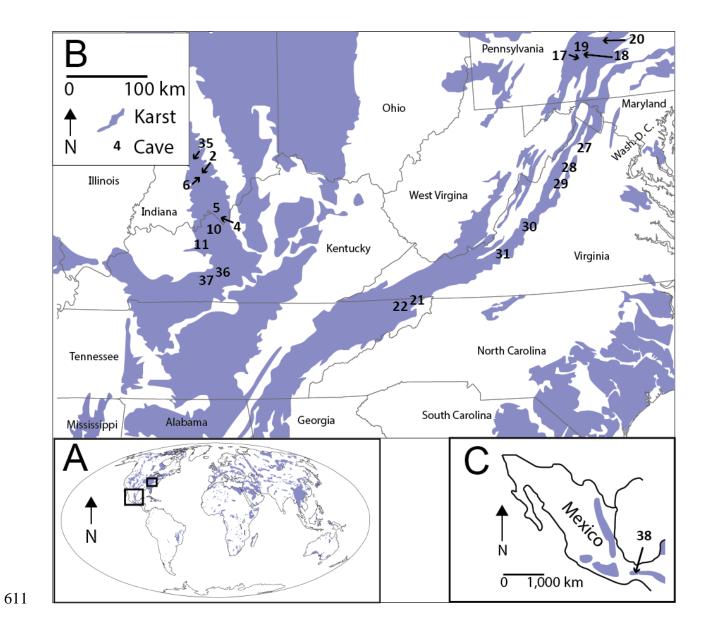
- 508 Zweers AJ, Bodelier PLE. 2018. Living apart together Bacterial volatiles influence
- 509 methanotrophic growth and activity. ISME J 12:1163–1166.
- 510 28. Gordon MS, Rosen DE. 1962. A cavernicolous form of the poeciliid fish *Poecilia*
- 511 *sphenops* from Tabasco, Mexico. Copeia 1962:360–368.
- 512 29. Howarth FG. 1993. High-stress subterranean habitats and evolutionary change in
  513 cave-inhabing arthropods. Am Nat 142:S65–S77.
- 514 30. Bárcena TG, Finster KW, Yde JC. 2011. Spatial patterns of soil development,
- 515 methane oxidation, and methanotrophic diversity along a receding glacier forefield,
- 516 Southeast Greenland. Arctic, Antarct Alp Res 43:178–188.
- 517 31. Lau MCY, Stackhouse BT, Layton AC, Chauhan A, Vishnivetskaya TA, Chourey
- 518 K, Ronholm J, Mykytczuk NCS, Bennett PC, Lamarche-Gagnon G, Burton N,
- 519 Pollard WH, Omelon CR, Medvigy DM, Hettich RL, Pfiffner SM, Whyte LG,
- 520 Onstott TC. 2015. An active atmospheric methane sink in high Arctic mineral
- 521 cryosols. ISME J 9:1880–1891.
- 522 32. Martineau C, Pan Y, Bodrossy L, Yergeau E, Whyte LG, Greer CW. 2014.
- 523 Atmospheric methane oxidizers are present and active in Canadian high Arctic soils.
- 524 FEMS Microbiol Ecol 89:257–269.
- 525 33. Lau E, Ahmad A, Steudler PA, Cavanaugh CM. 2007. Molecular characterization of
- 526 methanotrophic communities in forest soils that consume atmospheric methane.
- 527 FEMS Microbiol Ecol 60:490–500.

528	34.	Fernandez-Cortes A, Cuezva S, Alvarez-Gallego M, Garcia-Anton E, Pla C,
529		Benavente D, Jurado V, Saiz-Jimenez C, Sanchez-Moral S. 2015. Subterranean
530		atmospheres may act as daily methane sinks. Nat Commun 6:7003.
531	35.	Schütte UME, Cadieux SB, Hemmerich C, Pratt LM, White JR. 2016.
532		Unanticipated geochemical and microbial community structure under seasonal ice
533		cover in a dilute, dimictic Arctic lake. Front Microbiol 7:1–15.
534	36.	Bárcena TG, Yde JC, Finster KW. 2010. Methane flux and high-affinity
535		methanotrophic diversity along the chronosequence of a receding glacier in
536		Greenland. Ann Glaciol 51:23–31.
537	37.	Gregorič A, Zidanšek A, Vaupotič J. 2011. Dependence of radon levels in Postojna
538		Cave on outside air temperature. Nat Hazards Earth Syst Sci 11:1523–1528.
539	38.	Boadu FK. 2000. Hydraulic conductivity of soils from grain-size distribution: new
540		models. J Geotech Geoenvironmental Eng 126:739–746.
541	39.	Kalyuhznaya MG, Martens-Habbena W, Wang T, Hackett M, Stolyar SM, Stahl
542		DA, Lidstrom ME, Chistoserdova L. 2009. Methylophilaceae link methanol
543		oxidation to denitrification in freshwater lake sediment as suggested by stable
544		isotope probing and pure culture analysis. Environ Microbiol Rep 1:385–392.
545	40.	Beck DAC, Kalyuzhnaya MG, Malfatti S, Tringe SG, Glavina del Rio T, Ivanova
546		N, Lidstrom ME, Chistoserdova L. 2013. A metagenomic insight into freshwater
547		methane-utilizing communities and evidence for cooperation between the
548		Methylococcaceae and the Methylophilaceae. PeerJ 1:e23.

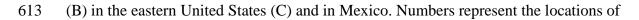
549	41.	Xia F, Zou B, Shen C, Zhu T, Gao X-H, Quan Z-X. 2015. Complete genome
550		sequence of Methylophilus sp. TWE2 isolated from methane oxidation enrichment
551		culture of tap-water. J Biotechnol 211:121–122.
552	42.	Michener JK, Vuilleumier S, Bringel F, Marx CJ. 2016. Transfer of a catabolic
553		pathway for chloromethane in Methylobacterium strains highlights different
554		limitations for growth with chloromethane or with dichloromethane. Front
555		Microbiol 7:1116.
556	43.	Barton HA, Taylor MR, Pace NR. 2004. Molecular phylogenetic analysis of a
557		bacterial community in an oligotrophic cave environment. Geomicrobiol J 21:11-
558		20.
559	44.	Kowalczk AJ, Froelich PN. 2010. Cave air ventilation and CO <sub>2</sub> outgassing by
560		radon-222 modeling: How fast do caves breathe? Earth Planet Sci Lett 289:209-
561		219.
562	45.	Gregorič A, Vaupotič J, Gabrovšek F. 2013. Reasons for large fluctuation of radon
563		and CO <sub>2</sub> levels in a dead-end passage of a karst cave (Postojna Cave, Slovenia). Nat
564		Hazards Earth Syst Sci 13:287–297.
565	46.	Jones SE, Lennon JT. 2010. Evidence for limited microbial transfer of methane in a
566		planktonic food web. Aquat Microb Ecol 58:45–53.
567	47.	Barberán A, Ladau J, Leff JW, Pollard KS, Menninger HL, Dunn RR, Fierer N.
568		2015. Continental-scale distributions of dust-associated bacteria and fungi. Proc
569		Natl Acad Sci 112:5756–5761.

570	48.	McEwen AS, Ojha L, Dundas CM, Mattson SS, Byrne S, Wray JJ, Cull SC,
571		Murchie SL, Thomas N, Gulick VC. 2011. Seasonal flows on warm Martian slopes.
572		Science (80- ) 333:740–743.
573	49.	Pašić L, Kovče B, Sket B, Herzog-Velikonja B. 2010. Diversity of microbial
574		communities colonizing the walls of a Karstic cave in Slovenia. FEMS Microbiol
575		Ecol 71:50–60.
576	50.	Webster KD, Rosales Lagarde L, Sauer PE, Schimmelmann A, Lennon JT, Boston
577		PJ. 2017. Isotopic evidence for the migration of thermogenic methane into a sulfidic
578		cave, Cueva de Villa Luz, Tabasco, Mexico. J Cave Karst Stud 79:24–34.
579	51.	Blott SJ, Pye K. 2001. Gradistat: A grain size Ddstribution and statistics package for
580		the analysis of unconcolidated sediments. Earth Surf Process Landforms 26:1237-
581		1248.
582	52.	Caporaso JG, Lauber CL, Walters WA, Berg-Lyons D, Huntley J, Fierer N, Owens
583		SM, Betley J, Fraser L, Bauer M, Gormley N, Gilbert JA, Smith G, Knight R. 2012.
584		Ultra-high-throughput microbial community analysis on the Illumina HiSeq and
585		MiSeq platforms. ISME J 6:1621–1624.
586	53.	Schloss PD, Westcott SL, Ryabin T, Hall JR, Hartmann M, Hollister EB,
587		Lesniewski RA, Oakley BB, Parks DH, Robinson CJ, Sahl JW, Stres B, Thallinger
588		GG, Van Horn DJ, Weber CF. 2009. Introducing mothur: Open-source, platform-
589		independent, community-supported software for describing and comparing
590		microbial communities. Appl Environ Microbiol 75:7537–7541.

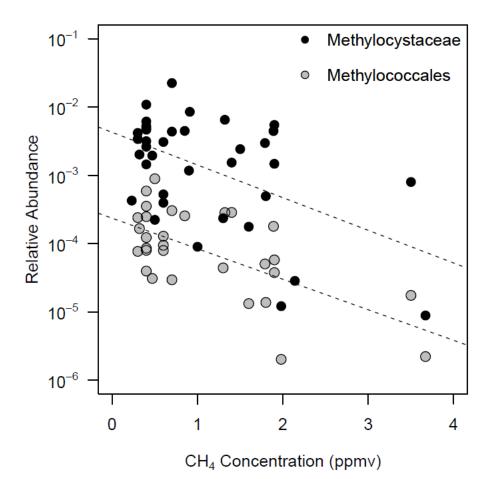
591	54.	Edgar RC, Haas BJ, Clemente JC, Quince C, Knight R. 2011. UCHIME improves
592		sensitivity and speed of chimera detection. Bioinformatics 27:2194–2200.
593	55.	R Core Team. 2017. R: A language and environment for statistical computing. R
594		Foundation for Statistical Computing, Vienna, Austria.
595	56.	McMurdie PJ, Holmes S. 2013. Phyloseq: An R package for reproducible
596		interactive analysis and graphics of microbiome census data. PLoS One 8:e61217.
597	57.	Oksanen, J.; Blanchet, F.; Kindt, R.; Legendere, P.; Minchin, P.R.; O'Hara, R.B.O.;
598		Simpson, G.L; Solymos, P.; Stevens, H.H.; Wagner H. 2018. Vegan: Community
599		ecology package. R package version v 2.4-6.
600	58.	Nekola JC, White PS. 1999. The distance decay of similarity in biogeography and
601		ecology. J Biogeogr 26:867–878.
602	59.	Shade A, Handelsman J. 2011. Beyond the Venn diagram: the hunt for a core
603		microbiome. Environ Microbiol 14:4–12.
604	60.	Björk JR, O'Hara RB, Ribes M, Coma R, Montoya JM. 2017. The dynamic core
605		microbiome: structure, dynamics and stability. bioRxiv 137885.
606	61.	Hernandez-Agreda A, Gates RD, Ainsworth TD. 2017. Defining the core
607		microbiome in corals'microbial soup. Trends Microbiol 25:125-140.
608	62.	Weary DJ, Doctor DH. 2014. Karst in the United States: a digital map compilation
609		and database. United States Geol Surv Open-File Rep 1156:1–23.
610		



612 Figure 1: (A) Locations of karst at the global scale. Inserts show the occurrence of karst

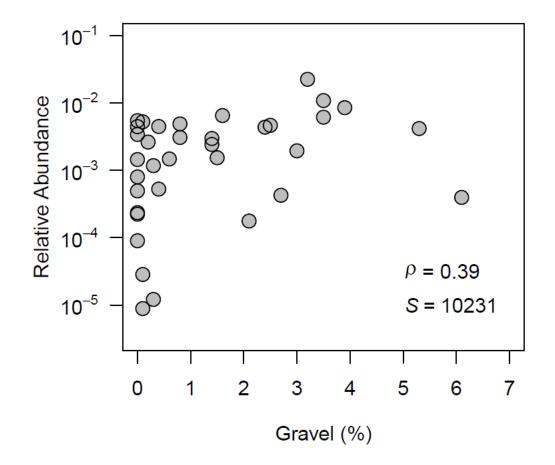


614 sampled caves in this study. Karst land cover data were obtained from (7, 62).



## 616

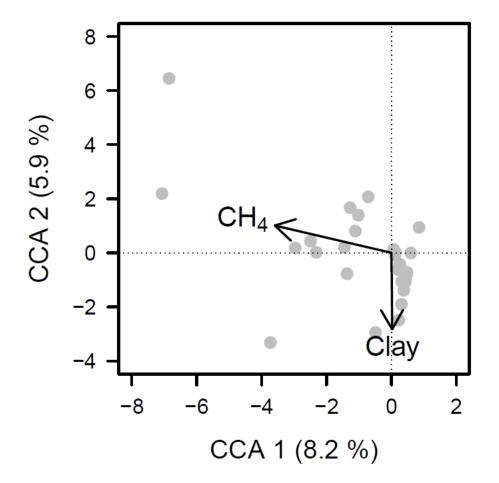
Figure 2: Relative abundances of methanotrophic community members plotted against the
CH<sub>4</sub> concentration at each sample location. Members of the Methylocystaceae, typically
associated with high-affinity methanotrophs, were more abundant than members from the
Methylococcales, typically associated with low-affinity methanotrophs.



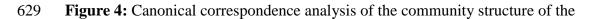
622

Figure 3. The relative abundance of methanotrophs from the Methylocystaceae plotted against the volumetric proportion of gravel in a sample (Spearman's rank correlation, p = 0.02).

626

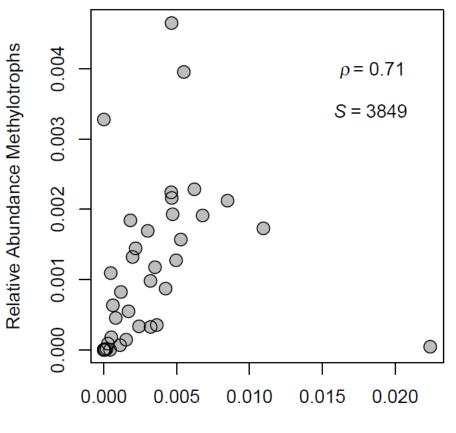


628



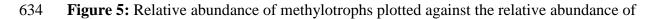
630 Methylococcales in caves. Community structure of this group was significantly related to

631 clay content and CH<sub>4</sub> concentrations.

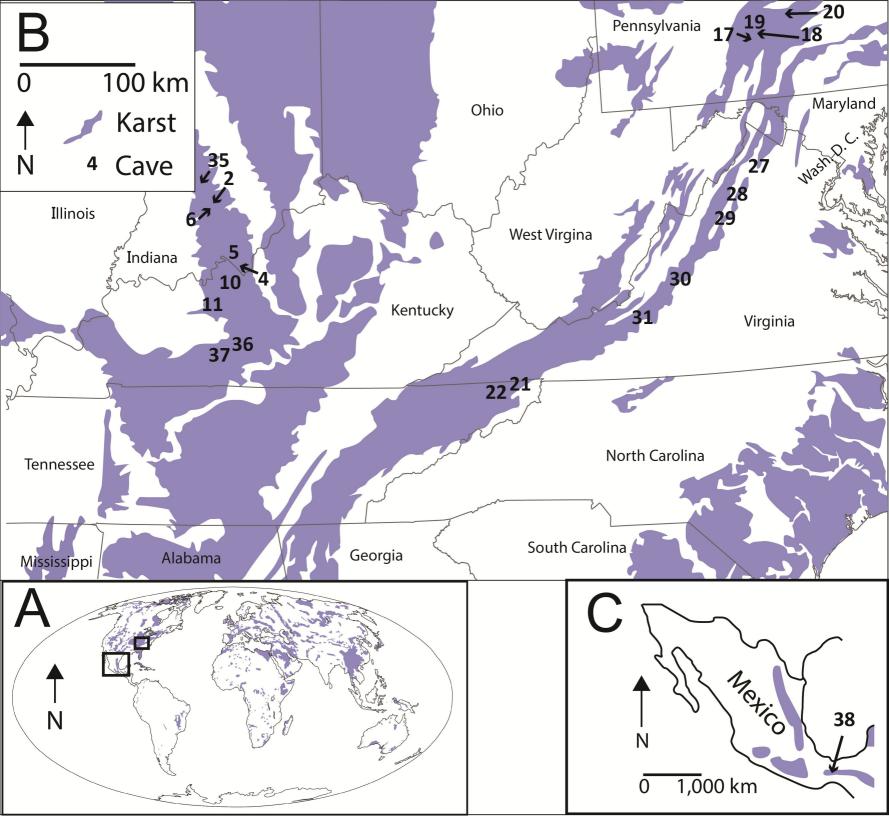


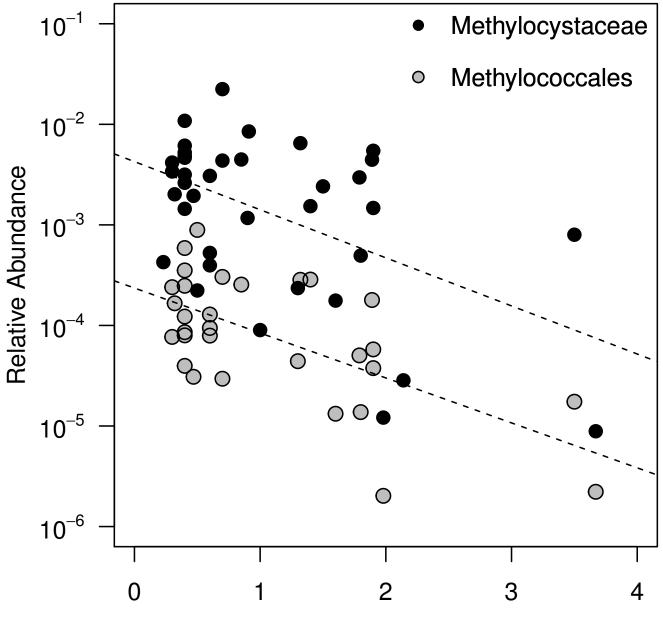
633

**Relative Abundance Methanotrophs** 

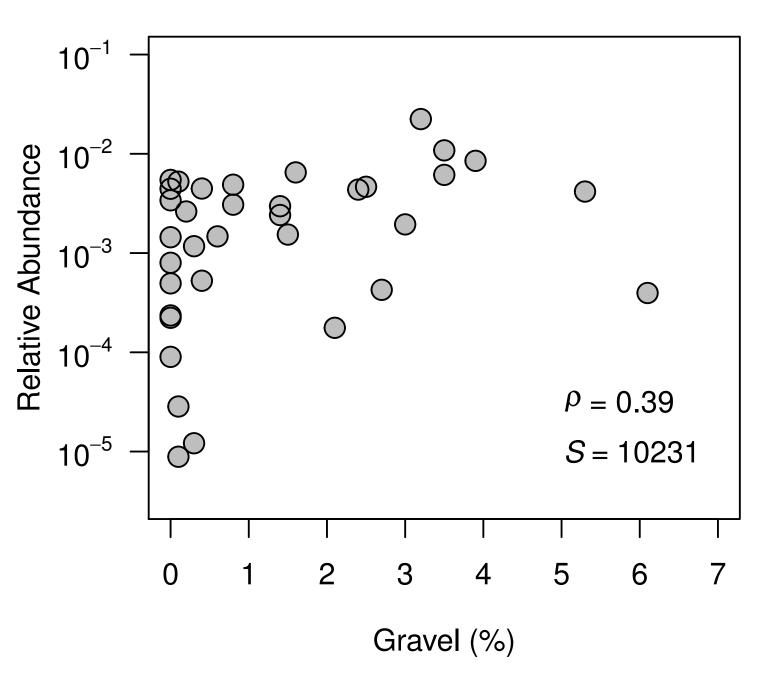


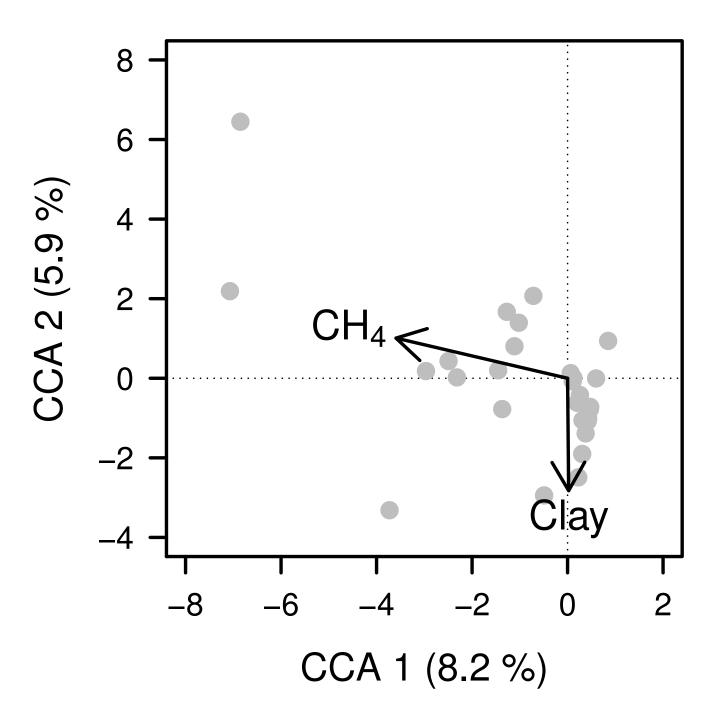
635 methanotrophs from sampled caves (Spearman's rank correlation,  $p = 1*10^{-7}$ ).



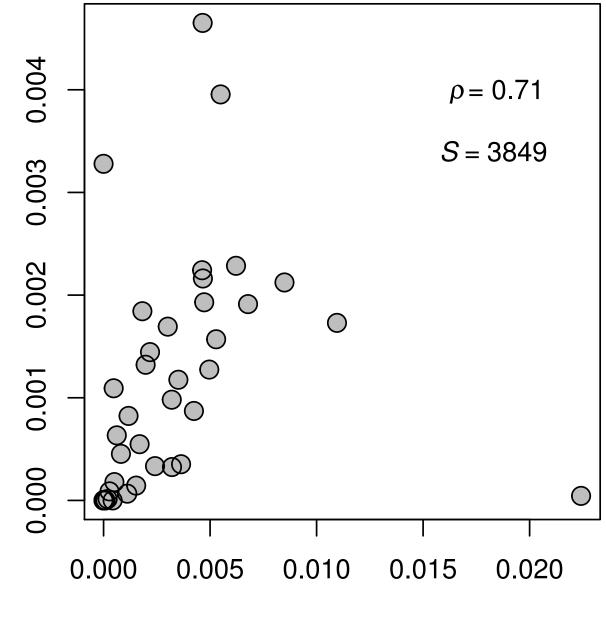


CH<sub>4</sub> Concentration (ppmv)





**Relative Abundance Methylotrophs** 



**Relative Abundance Methanotrophs**