1 **Long-distance dispersal, ice sheet dynamics, and mountaintop isolation underlie the genetic**
2 **structure of glacier ice worms**
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19 **Running title:** Population genomics of ice worms
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21 **Keywords:** RAD sequencing, Pacific Northwest, phylogeography, *Mesenchytraeus solifugus*,
22 global change biology, annelid
23

Abstract:
Disentangling the contemporary and historical factors underlying the spatial distributions of
species is a central goal of biogeography. For species with broad distributions but little capacity
to actively disperse, disconnected geographic distributions highlight the potential influence of
passive, long-distance dispersal (LDD) on their evolutionary histories. However, dispersal alone
cannot completely account for the biogeography of any species, and other factors–e.g., habitat
suitability, life history–must also be considered. North American ice worms (*Mesenchytraeus*
solifugus) are ice-obligate annelids that inhabit coastal glaciers from Oregon to Alaska. Previous
studies identified a complex biogeographic history for ice worms, with evidence for genetic
isolation, unexpectedly close relationships among geographically disjunct lineages, and
contemporary migration across large (> 1,500 km) areas of unsuitable habitat. In this study, we
analyzed genome-scale sequence data for most of the known ice worm range. We found clear

support for divergence between populations along the Pacific Coast and the inland flanks of the

Coast Mountains (mean $F_{ST} = 0.60$), likely precipitated by episodic ice sheet expansion and

contraction during the Pleistocene. We also found support for LDD of ice worms from Alaska to

Vancouver Island, perhaps mediated by migrating birds. Our results highlight the power of

genomic data for disentangling complex biogeographic patterns, including the presence of LDD.

Introduction:

For more than a century, long-distance dispersal (LDD) among presumably isolated

populations has intrigued biologists [1-4]. Historically considered rare and unpredictable, the

idea that LDD can act as a general mechanism influencing the biogeography of presumably

dispersal-limited, macroscopic organisms has gained traction in recent years, with examples

accumulating for both plants [5] and invertebrates [6-9]. Many animal vectors play an integral

role in plant and invertebrate LDD [e.g., 10, 11], however, in most cases, the resulting LDD is

limited to less than 10 km. For more extreme LDD events (e.g., greater than 100 km), the most

common animal vector is likely migratory birds, as they seasonally move by the millions over

broad spatial scales and geographic barriers, visiting similar habitats along the way [12, 13].

Through this mechanism, dispersal units (e.g., whole organisms, eggs, seeds, etc.) may be

ingested and dispersed after passing through the digestive tract [9] or by directly adhering to the

bird's exterior [12]. Thus, as long as there is an opportunity for migratory birds and dispersal

units to interact, the opportunity also exists for LDD. Physically quantifying LDD is difficult,

however, because it requires real-time sampling and searching (internal and external) of

migrating birds for hitchhiking dispersers. Moreover, because rare migratory events can affect

species distributions [14] and influence genetic differentiation among populations [15], even

thorough physical surveys of migratory birds that find no evidence for LDD cannot rule out its

presence. Therefore, alternative approaches for detecting and characterizing LDD should be

employed. Because population genomic tools are well-suited to detecting gene flow and genetic

structure among populations [e.g., 16], these tools are also well-suited to the indirect detection of

LDD, even in the absence of field observations.

Many mechanisms influence genetic relationships among taxa and a range of factors

should be considered when attempting to reconstruct biogeographic patterns. For instance, pulses

and contractions of glaciers and ice sheets have shaped the evolutionary histories of populations

and species throughout Earth's history [17-20]. These ice sheet dynamics typically affect

organisms by separating and reconnecting populations as ice cover changes. However, some

species are directly tied to ice sheets [e.g., the meltwater stonefly, 17, 21] and are therefore much

more susceptible to ice sheet influence on their evolutionary trajectories. Perhaps no species is

more directly tied to ice sheets than glacier ice worms, *Mesenchytraeus solifugus* in North

America [22] and *Sinenchytraeus glacialis* in Tibet [23]. The geographic range of *M. solifugus*

(hereafter "ice worm") follows a coastal arc from the Chugach Mountains in southeast Alaska to

the Cascade Volcanoes of Washington and Oregon [24]. Ice worms cannot tolerate temperatures

more than roughly ±7 °C from freezing and require glacier ice for survival and reproduction [25].

With such unique ecology and physiology, and a dispersal-limited life history, the evolutionary

history of ice worms since diverging from conspecifics [26, 27] should be relatively simple with

gene flow occurring during glacial periods and isolation (paired with genetic drift) driving

divergence among mountaintop-isolated populations during interglacial periods. Natural systems,

however, are often more complex than expected and indeed, the evolutionary history of ice

worms challenges general expectations of gene flow and evolutionary dynamics in ice-

dominated, mountain ecosystems.

Previous genetic studies based on one or two genetic markers identified three ice worm

lineages: a "northern" clade in southern Alaska, a "central" clade in southeast Alaska and

northern British Columbia, and a "southern" clade ranging over much of British Columbia to

southern Oregon [24, 25, 27]. Surprisingly, phylogenetic evidence supported the northern and

southern lineages as being most closely related to one another despite the central clade separating

them geographically. The most curious aspect of ice worm biogeography, however, has been the

repeated discovery of closely related ice worms on glaciers several hundred to thousands of

kilometers south of their closest genetic relatives [25, 27; P. Wimberger, unpublished data].

These disjunct northern ice worms co-occurred with, but appeared genetically distinct from, their

conspecifics (either central or southern clade ice worms) on the same glaciers. Dial *et al.* [25]

laid out three possible explanations for this pattern: wind transport, passerine-mediated dispersal,

or a more extensive previous range of the northern clade. While wind transport seems unlikely,

the potential for passerine-mediated dispersal is reasonable, particularly in light of other

examples of bird-mediated LDD [e.g., 7, 28]. The third scenario, a more extensive distribution of

the northern clade with holdover lineages inhabiting the same glacier as more recent colonizers

could indeed result in more than one distinct lineage on the same glacier. However, this pattern

may only apply to mitochondrial DNA (mtDNA) since mtDNA is maternally inherited and does

not recombine. For the nuclear genome, unless strong reproductive isolation exists between the

holdover lineages and more recent colonizers, genetic differences would be rapidly homogenized

by gene flow and recombination. Assuming no selection against migrants, a reasonable

expectation given the similarity of habitat across the ice worm range, the frequency of

contemporary versus historical mtDNA haplotypes would therefore depend upon time since

introduction, scale of migration (i.e., number of introduced haplotypes), and chance.

In this study, we leveraged a modern population genomic toolkit to add new perspective

to the age-old challenge of identifying LDD in wild populations. We also provide new insight

into how multiple factors can interact to shape the evolutionary history of species. We

hypothesized that the biogeographic history of ice worms stemmed from a confluence of factors:

extreme LDD, glacier dynamics, and mountaintop isolation. To test this hypothesis, we

generated a genome-wide single nucleotide polymorphism (SNP) data set to answer three

specific questions: (1) How do the clades previously diagnosed from a small number of markers

hold up to genome-wide scrutiny? (2) What, if any, genomic evidence exists for LDD in ice

worms? (3) How do the evolutionary relationships among ice worm populations and genetic

clusters align with glacial history in the region [e.g., 29]? Beyond a refined view of ice worm

evolution, our study confirms that LDD does occur in ice worms, providing an example of LDD

in an annelid and a rare population genomic exploration of the process. Moreover, while

considerable evidence details the existence of refugia in the Pacific Northwest (PNW) during the

Pleistocene [30], few studies have explored how ice sheet dynamics influenced the evolutionary

history of species directly tied to them [e.g., 17]. Our results reveal the profound impact that ice

sheet formation during the Pleistocene (~2.5 million – 11,700 years ago) which flowed

episodically from the crest of the Coast Mountains [29] may have had on ice worm evolution,

possibly precipitating an ongoing speciation event. Broadly, our findings highlight the power for

population genomics to capture contemporary evidence of LDD while also providing

biogeographic evidence for reconstructing the glacial history of a region.

Table 1. Sampling information and summary statistics for all ice worm populations included in this study.
Abbreviations: n = sample size, π = nucleotide diversity, Het = heterozygosity, F_{IS} = inbreeding
coefficient, AK = Alaska, BC = British Columbia, VI = Vancouver Island. π, Het, and F_{IS} were calculated
for variable sites only. Asterisks indicate populations included in both Dial *et al.* [27] and this study.

Population	Latitude, longitude	State/Prov.	Elev. (m)	n	π	Het	F_{IS}
Learnard* (LEA)	60.806, -148.721	AK	624	3	0.114	0.100	0.026
Davidson* (DAV)	59.067, -135.551	AK	986	15	0.078	0.065	0.037
Treaty* (TRE)	56.586, -130.151	BC	1376	8	0.026	0.033	-0.012
Bear (BEA)	56.096, -129.681	BC	648	6	0.039	0.046	-0.013
William Brown* (WIB)	54.611, -129.129	BC	1260	6	0.056	0.057	0.000
Jacobson* (JAC)	52.050, -126.072	BC	1249	5	0.108	0.106	0.005
White Mantle* (WHM)	50.795, -125.153	BC	1764	8	0.131	0.135	-0.005
Comox* (COM)	49.545, -125.355	BC (VI)	1881	4	0.091	0.072	0.034
Mariner (MAR)	49.460, -125.764	BC (VI)	1754	4	0.098	0.084	0.024

Methods:

Sample collection, library preparation, and SNP calling

During the summer of 2009, ice worms were collected from nine glaciers across most of

their geographic range (Figs. 1, S1; Table 1). Samples were stored in > 80% EtOH until DNA

was extracted from 59 ice worms using a Qiagen DNEasy Blood and Tissue Kit. Double-digest

restriction-site associated DNA (ddRAD) sequencing libraries were prepared following Peterson

et al. [31] with restriction enzymes EcoRI and NlaIII. During library preparation, samples were

divided into two groups and each sample was assigned a unique, variable-length barcode [32]

which was incorporated during adapter ligation. Size selection for a 350 bp ± 35 bp window was

performed with a Pippen Prep (Sage Science), and both sample groups were subsequently

amplified using PCR primers containing a group-specific barcode. The 59-sample library was

sequenced on one lane of an Illumina HiSeq4000 at the University of Illinois High-Throughput

Sequencing and Genotyping Unit with single-end, 100 bp chemistry.

Raw reads were demultiplexed, quality-filtered, and ddRAD loci were assembled *de novo*

using the *process_radtags* and *denovo_map* functions of the Stacks v1.46 pipeline [33]. We

allowed a maximum distance between stacks of 2 and a minimum read depth of 10. Next, we

applied a stringent filtering scheme to identify high-confidence SNPs that were shared among

many individuals. We only included SNPs if they were present in ≥ 5 populations, genotyped in

≥ 50% of individuals per population, and were in Hardy-Weinberg equilibrium with a minor

allele frequency of ≥ 0.025 overall. We further restricted analyses to one random SNP per locus

for all analyses except fineRADstructure (see below). All post-Stacks filtering steps were

performed in PLINK v1.07 [34] and the commands used in this study are provided on GitHub

(https://github.com/scotthotaling/ice_worm_ddRAD).

Figure 1. Ice worm populations sampled for this study. Color-coding reflects the results of a
fineRADstructure coancestry analysis. Group I (circles) and II (square) populations were generally
defined by their presence to the east or west of a key line where ice ridge(s) putatively formed during the
Pleistocene > ~25,000 years ago [29] as well as their distance to the Pacific Ocean. The deep divergence
between groups I and II is clearly evident along with more recent differentiation within each group. One
individual from the Mariner Glacier (asterisk) was admixed between the Mariner/Comox (Vancouver
Island) and Learnard (southern Alaska) clusters, indicating recent LDD.

Population genetic and phylogenetic analyses

For each population, we calculated nucleotide diversity (π), heterozygosity (Het), and the

inbreeding coefficient (F_{IS}). We also calculated a pair-wise AMOVA F_{ST} for all population

combinations [35] with the Stacks *populations* module. To test for a signature of isolation-by-

distance [IBD; 36], we estimated Euclidean distances among sites with Google Earth and tested

the correlation between geographic distance and F_{ST} with four Mantel tests performed in

GenoDive v2.0b27 [37]. The first Mantel test included all nine populations and the second

excluded both Vancouver Island populations (Mariner and Comox) to assess whether unique

histories for those populations were significantly altering results. The third and fourth Mantel

tests focused on signatures of IBD within groups "I" and "II" (see Results). We also measured

the Euclidean distance to the Pacific Ocean for each population in Google Earth Pro as an

additional spatial comparison of groups I and II. We determined if mean distances to the Pacific

Ocean differed between groups with a one-way ANOVA.

Population structure was inferred in two ways: a maximum likelihood-based method

using ADMIXTURE 1.3.0 [38] and a discriminant analysis of principal components (DAPC)

with the R package *adegenet* [39]. ADMIXTURE analyses were performed with default settings,

a range of clusters (K) from 1-12, and 25 replicates per K with the current time as the random

seed. The cross-validation (CV) error for each K was plotted to identify the best-fit K (minimized

CV across the mean of all replicates for each K). After identifying the best-fit K, we considered

the replicate that minimized CV across all 25 replicates for all K's to be the best-fit solution

overall. However, because all runs did not converge on the same result, we also inspected best-fit

solutions for other replicates of $K = 7$ (the best-fit K overall) to clarify the distribution of best-fit

solutions. For DAPC, we first used the *find.clusters* function to identify the optimal K [i.e., the K

with the lowest Bayesian Information Criterion (BIC)]. Next, to avoid over-fitting of the model,

we retained the appropriate number of principal components (PCs) according to the α-score [PCs

retained = 6, Fig. S2]. We performed a final DAPC analysis using the best-fit K and optimal

number of PCs identified in the previous two steps.

We extended our population structure analyses to infer both shared ancestry and

phylogenetic relationships in two ways: a nearest neighbor haplotype approach to infer

coancestry with fineRADstructure [40] and phylogenetic relationships inferred from singular

value decomposition estimates for quartets of tips using SVDQuartets [41] as implemented in

PAUP* v4.0a159 [42]. For fineRADstructure, we used 100,000 burn-in iterations followed by

100,000 iterations sampled every 1,000 steps for the Markov chain Monte Carlo clustering

algorithm. Next, we used 10,000 iterations of the tree-building algorithm to assess genetic

relationships among clusters. Since fineRADstructure is a haplotype-based approach, analyses

were performed using all variable sites for a given ddRAD locus (i.e., a haplotype) rather than

randomly selected single SNPs per locus. For SVDQuartets, we performed exhaustive sampling

of all possible quartets (every combination of four tips) and branch support was estimated with

100 nonparametric bootstrap replicates.

Demographic modeling

To test hypotheses of demographic history and estimate divergence times for the two

major groups identified in our population genetic and phylogenetic analyses (I and II, see

Results), we performed demographic modeling. We used fastsimcoal2 v2.603 [43] which

leverages a coalescent-based model to estimate demography from the site frequency spectrum

(SFS). We designed and tested the fit of four two-lineage models (Fig. 2A): no gene flow (M1),

unidirectional gene flow from group I into II (M2), unidirectional gene flow from group II into

group I (M3), and bidirectional gene flow (M4). All four models also included parameters for

ancestral (N_{ANC}) and current (N_I and N_{II}) effective population sizes as well as divergence time

(T_{DIV}). We maximized the number of shared SNPs between our focal groups by selecting the

four individuals with the least missing data from the same population in each group (group I =

Davidson; group II = Treaty). We also only retained loci with no missing data, which yielded

2,714 SNPs across the eight individuals. For each model, we performed 50 replicate runs, each

with 100,000 simulations and 100 cycles of a conditional maximization algorithm. We specified

the nuclear mutation rate at 3.5E-9 per site per generation, based on the estimate for *Drosophila*

melanogaster [44]. To identify the best-fit model, we calculated an Akaike Information Criterion

(AIC) score for each model and selected the best model as the one with the lowest AIC relative

to the other models. We generated 95% confidence intervals (CIs) by simulating 50 SFS

replicates from the best-fit run of the best-fit model. Next, we performed the same 50 replicate

analyses described above for each of the 50 newly simulated SFSs.

For divergence times, our final results are in generations and we refrain from converting

this estimate to years before present because the generation time for ice worms is not known.

Estimates of generation time for the family Enchytraeidae also vary widely, ranging from 21

days at 18°C [45] to a full year at 10°C [46]. Moreover, no estimates of generation time at low

temperatures (e.g., ~0°C) are available for the group. Complete details of our demographic

analyses are provided in the Supplementary Materials.

Figure 2. (A) Demographic models tested in this study which included parameters for divergence time
(T_{DIV}), effective population sizes [ancestral (N_{ANC}) and for each group (N_I, N_{II})], and migration (arrows)
when applicable. (B) The parameterized best-fit demographic model (M1). Numbers in parentheses
represent 95% confidence interval around estimates. The white box corresponds to the divergence time
(in generations) between groups I and II. Gray boxes correspond to population size parameters.

Table 2. Above the diagonal: Pair-wise AMOVA F_{ST} values for all populations included in this study. Mean F_{ST} (bottom row) refers to the average pair-wise differentiation for columnar populations versus all others. Below the diagonal (in gray): mean pair-wise shared loci for the fineRADstructure coancestry analysis (see Fig. 1).

	LEA	DAV	TRE	BEA	WIB	JAC	WHM	COM	MAR
LEA	--	0.333	0.634	0.626	0.608	0.586	0.557	0.347	0.344
DAV	74.9	--	0.567	0.557	0.541	0.527	0.515	0.290	0.295
TRE	19.8	20.0	--	0.153	0.207	0.275	0.250	0.656	0.652
BEA	19.0	19.9	127.2	--	0.165	0.250	0.222	0.650	0.645
WIB	20.1	20.5	119.8	123.0	--	0.216	0.201	0.630	0.628
JAC	21.0	20.4	99.7	99.5	104.7	--	0.172	0.604	0.600
WHM	21.6	19.1	90.7	90.9	93.0	107.8	--	0.576	0.572
COM	84.8	86.1	21.8	21.3	23.0	22.9	22.1	--	0.160
MAR	164.5	82.9	21.2	21.3	22.3	21.8	21.4	262.3	--
Mean F_{ST}	0.504	0.453	0.424	0.408	0.399	0.404	0.383	0.489	0.487

Results:

Population genetic and phylogenetic analyses

We generated 343,875,880 reads with an average of 5,828,404 sequences per individual (min. = 446,872 and max. = 40,982,490). Our total RAD data set included 360,534 unique loci. After filtering, we retained 6,019 loci and 10,392 SNPs (mean = 1.73 SNPs per locus). This final data set had genotype calls for ~65% of all SNPs. Nucleotide diversity (π) was highest in the White Mantle and Learnard populations (0.131 and 0.114, respectively) and lowest in the Treaty, Bear, and William Brown populations (0.026-0.056; Table 1). Heterozygosity followed the same pattern as π (Table 1). The inbreeding coefficient (F_{IS}) was highest in the Davidson and Comox populations (0.037 and 0.034, respectively) and lowest in Bear (-0.013) and Treaty (-0.012; Table 1). Mean differentiation (F_{ST}) for all pair-wise comparisons was 0.439. The Learnard population from southern Alaska was, on average, the most differentiated from all others (mean F_{ST} = 0.504) and White Mantle the least differentiated (mean F_{ST} = 0.383; Table 2). We detected no association between genetic and geographic distances in either of the study area-wide Mantel tests (Mantel's *r*, all populations = -0.04, *P* = 0.42; Mantel's *r*, no Vancouver Island populations = 0.15, *P* = 0.36). There was, however, a signature of IBD within group II (Mantel's *r*, group I = 0.839, *P* = 0.025) but not within group I (Mantel's *r*, group II = 0.839, *P* = 0.082).

Our DAPC analyses supported *K* = 6 as the optimal number of genetic clusters (Fig. 3A,C). ADMIXTURE results, however, supported *K* = 7 as the best-fit (Figs. 3B,C) and the SVDQuartets phylogeny largely mirrored both lines of population structure evidence (Fig. 3D). All analyses supported multiple independent genetic clusters of ice worms. Our DAPC and

ADMIXTURE results differed in two ways: (1) the best-fit DAPC result grouped the Treaty,

Bear, and William Brown populations into one cluster whereas the best-fit ADMIXTURE result

split William Brown into its own cluster. This difference accounted for the $K = 6$ versus $K = 7$

discrepancy between the approaches. (2) While both analyses identified a single individual

(MS5) from the Mariner population with genetic assignment to the Learnard cluster, DAPC

indicated full assignment of MS5 to the Learnerd cluster whereas Admixture equally assigned it

to both the Learnerd and Comox+Mariner clusters (Figs. 1, 3). Finally, because our SNP filtering

focused on overarching patterns in the data set, and likely overlooked some degree of

population-specific detail, our results are likely conservative estimates of genetic structure in the

group.

Figure 3. Population genetic structure of ice worms based upon (A) a discriminant analysis of principal
components (DAPC) for $K = 6$ and (B) ADMIXTURE results for $K = 7$. (C) Comparisons of support for
different values of K for DAPC (Bayesian information criterion, BIC; gray dashed line, left y-axis) and
Admixture (cross-validation error, CV; dark line, right y-axis). The best-fit K (white square,
ADMIXTURE; white circle, DAPC) corresponds to the K at which CV (ADMIXTURE) or BIC (DAPC)
was minimized, respectively. For ADMIXTURE, vertical black bars represent the standard deviation of
CV values for each K across 25 replicates. (D) An unrooted phylogeny of ice worms generated with

SVDQuartets. Tip colorations reflect assignments in (A). Dark circles indicate nodes with greater than
95% bootstrap support. Specimens in group I (circles) and II (squares) are denoted with different
symbols. Asterisks (A & D) highlight a single specimen, MS5, which showed evidence of shared ancestry
across geographically disjunct populations, indicating LDD.
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Our fineRADstructure results largely mirrored those from DAPC, identifying the same

six genetic clusters. As in the Admixture results, MS5 exhibited evidence for shared ancestry

between the Learnard and Comox/Mariner genetic clusters (Fig. 1). On average, MS5 exhibited a

~70/30 split of shared ancestry between Learnard and Comox/Mariner ice worms (Fig. 1).

Notably, MS5 also exhibited the highest heterozygosity of any individual in the study (and this

result did not stem from outsized coverage, Fig. S3). Our fineRADstructure results also

highlighted a primary divergence between two groups of ice worm populations (groups I and II;

Fig. 1). This split was corroborated by both SVDQuartets (Fig. 3) and F_{ST} comparisons. Mean

pairwise F_{ST} among populations within groups I and II were 0.21 and 0.30, respectively.

Between groups, however, mean pairwise F_{ST} was 0.60. Mean distance to the Pacific Ocean also

differed by 100.5 km (group I = 73.3 km, group II = 173.8 km; ANOVA, $P < 0.001$).

Demographic modeling

Our tests of demographic models for groups I and II identified a history of divergence

without gene flow (model M1) as the best fit to our data (Fig. 2B; Table S1). All other models

were rejected with ΔAIC $\geq$ 9.1 (Table S1). The second-best model (M2) included unidirectional

gene flow from group I into group II (Fig. 2A), while models M3 and M4 included gene flow

from group II into group I, resulting in ΔAIC scores $\geq$ 118.4. Groups I and II diverged

approximately ~250,000 generations ago (Fig. 2B).

Discussion:

Historical and contemporary factors, both biotic and abiotic, interact to shape the present-

day genetic structure of populations and species. Disentangling their varied contributions can be

difficult, however, particularly when evolutionary histories are muddled by unexpected events

(e.g., LDD of an organism with limited potential for active dispersal). The modern population

genomic toolkit provides historically unprecedented power to resolve biogeographical

complexity by allowing more quantitative perspectives of relatedness and greatly improved

resolution of genetic independence or similarity [47]. In this study, we used a population

genomic data set to refine understanding of the evolutionary history of the extremophile, glacier-

obligate ice worm, *M. solifugus*. Our results provide a clear genomic perspective of LDD,

showing unequivocally that migration has occurred between southern Alaska and the glaciers of

Vancouver Island ~1,900 km to the south and across the Pacific Ocean. We also provide an

independent line of biological evidence in support of the geological hypothesis that ice ridges

formed along the crest of the Coast Mountains during the Pleistocene [29].

Ice worm biogeography and LDD

The recent biogeographic history of ice worms appears to have been shaped by three

main factors: i) ice sheet dynamics, ii) mountaintop isolation from conspecifics following the

retreat of Pleistocene ice into higher elevations, and iii) LDD.

i.   Evidence for the first, overarching factor that has defined the recent evolution of ice

worms lies in our overwhelming support for the deepest divergence between two

groups (I and II) which fall largely on either side of the Coast Mountains in western

North America. During the Pleistocene (~2.5 million–11,700 years ago), western

Canada was repeatedly covered by a continental ice sheet [29]. Ice was generated in

the high peaks of the Coast Mountains and subsequently flowed west to the Pacific

Ocean and east to the interior of British Columbia from the crest of the range [29, 48,

49]. This potential western-eastern divergence in ice worms aligns with this divide,

suggesting that each group diverged in allopatry from their conspecifics. The timeline

of this divergence, however, is unclear. Our results suggest ~250,000 generations

have passed since the initial divergence but with no knowledge of generation times

for ice worms, nor related species at very low temperatures, we cannot provide a

reliable estimate of years before present. If ice worms develop rapidly, with multiple

generations per year (e.g., ≥ 3), the divergence may have occurred less than 100,000

years ago. However, if ice worms develop slowly (e.g., one generation for every five

years)—the split may have occurred over a million years ago. Either way, the split

appears more recent than estimates from mitochondrial DNA of ~1.7 million years

ago [25]. At its maximum, the Pleistocene ice sheet in the region was a ~2,000-3,000

m high convex dish with gentle interior slopes that steepened at its periphery [29, 50].

Given the sensitivity of ice worms to extreme cold [25], populations likely only

persisted on the ice sheet margins, as supported by their present-day occurrences on

the lower flanks of higher elevation, low-latitude glaciers. It is possible that, as

suggested previously [25, 27], the Boundary Ranges (the most northern subrange of

the Coast Mountains), are actually the biogeographic barrier that drove the deep

divergence in ice worms described in this study. However, without more fine-scale

population genomic sampling on both sides of the Coast Mountains (including the

Boundary Ranges), this nuance of ice worm biogeography will remain unclear (see

Fig. S4 and additional discussion in Supplementary Materials). In the same vein, the

strong genetic similarity of White Mantle to populations east of the proposed ice

ridge (Fig. 1), despite falling on its western side, indicates that either the ice ridge that

precipitated divergence among the two groups actually formed more to the west than

previously thought [29], the White Mantle population has migrated west since

divergence from conspecifics, the ice ridge itself was not a barrier driving

differentiation (as discussed above), or some combination of all three.

ii.   In western North American, the formation of the Cordilleran ice sheet was seeded by

alpine glaciers [29]. While the specific dynamics of deglaciation on valley and

drainage scales are unknown, a safe assumption is that glaciers retreated from valleys

into higher elevations, likely with ice worm populations in tow. Increasing

mountaintop isolation and subsequent genetic drift likely precipitated the more recent

differentiation within groups I and II since their initial split. Evidence for IBD within

group II supports this hypothesis. The lack of support for IBD in group I, however,

may reflect reality, the reduced power of a smaller sample size, or LDD maintaining

connections over larger spatial scales than a purely IBD model would predict.

iii.   Despite limited sampling, we were able to identify one instance of recent LDD among

geographically disparate ice worm populations. Indeed, with divided ancestry

between the Mariner/Comox and Learnard clusters, one specimen (MS5), is likely the

progeny of recent migration between the two. This indicates that LDD is both

ongoing in ice worms and perhaps not particularly rare. The most plausible

mechanism for ice worm LDD is passive dispersal of mucous-coated ice worm

cocoons sticking to the feet, beaks, or feathers of southward migrating birds [12].

Several passerines (e.g., gray-crowned rosy finches, *Leucosticte tephrocotis*) have

been observed feeding on ice worms [51, 52; S.H., personal observation] and the

presence of an ice worm clitellum [53] indicates that ice worms, like other

Mesenchytraeus species, reproduce by egg-laden cocoons [25]. The seemingly

exclusive northwest-to-southeast pattern of ice worm LDD also has temporal support

from bird migratory behavior. Late autumn ice worm reproduction [at the end of the

productive season on mountain glaciers, 54] likely occurs in concert with southward-

migrating birds stopping to feed on glaciers free of seasonal snow. In contrast,

returning spring migrants pass over the same glaciers when seasonal snowfall still

covers overwintering ice worms [25], limiting the potential for LDD in the reverse

direction.

One question remains, however, if birds are precipitating LDD in ice worms, why has it

only been observed for populations west of the Coast Mountains? This curiosity ties in to an

important question in North American biogeography: to what extent have ice sheets driven

present-day patterns of speciation and genetic differentiation among fauna of the northwest? For

ice worms, we hypothesize that populations comprising groups I and II have accumulated some

degree of reproductive isolation. This inference is supported by our demographic modeling

which rejected models involving gene flow from group II into group I. While this may be at least

partly linked to directionality in migration (e.g., bird movements), the strength in which these

models were rejected suggests that a zygotic barrier may be limiting inter-group migrants from

leaving a genomic signature of gene flow. It is also possible, and perhaps likely, that patterns of

LDD in ice worms is driven by vector migration patterns. For instance, *L. tephrocotis*, like other

songbirds [55], may preferentially follow coastlines during migration. However, until more is

known about the specific interactions of ice worms with various bird species, and by proxy, their

potential to act as LDD vectors, relating bird migrations to ice worm distributions and

demography will remain difficult. Beyond ice worms, ice sheets have been implicated as a key

driver of speciation in boreal birds [56] and phylogeographic structure of many taxa, from

nematodes to gray wolves [17, 30, 57-60], and our results clearly support these broader

implications for biodiversity accumulation and maintenance in North America.

Conclusions

In this study, we leveraged population genomic data to unravel the complex evolutionary

history of the North American ice worm, *M. solifugus.* Our results add new clarity to previous

perspectives on ice worm biogeography while also lending genomic support to the existence of

contemporary, likely passerine-mediated LDD in the group. We described two genetic grouping

(I and II) which are described with respect to the crest of the Coast Mountains, where ice ridges

formed during the Pleistocene [29]. While the phylogenetic data used in this study (i.e., the lack

of an outgroup) preclude us from diagnosing groups I and II as monophyletic, given the results

of previous studies [24, 25, 27], we predict that future efforts will diagnose them as such,

perhaps even representing two nascent species. Finally, our genomic data lend support to the

glaciological record in the region, adding a biological line of evidence to a postulated key north-

south dividing line along the crest of the Coast Mountains where ice ridges likely formed during

the Pleistocene and repeatedly propagated ice flow to the east and west [29]. This potential for

genomics to inform the geological record is intriguing and ice worms, as a rare glacier-obligate

macroinvertebrate, may be an ideal taxon for similar studies in the future.

Acknowledgements: We thank Joe Giersch for map-making assistance and the University of

Kentucky Center for Computational Sciences as well as the Washington State University Center

for Institutional Research Computing for high-performance computational resources. This

research was partially funded by NSF award #IOS-082050 to D.H.S.

Competing interests: We have no competing interests.

Author contributions: S.H., D.H.S., L.M.T, and D.W.W. conceived of and funded the study.

S.H., D.H.S., S.A.L, R.K.B., and D.W.W. collected the data. S.H. and J.L.K. analyzed the data

and wrote the manuscript with input from D.H.S., R.K.B., and D.W.W. All authors approved the

final version.

Data accessibility: Raw sequence data for this study has been submitted to GenBank under

BioProject #PRJNA479335 and code to reproduce the analyses is deposited on GitHub

(https://github.com/scotthotaling/ice_worm_ddRAD).

Funding: Field sampling for this study was supported by NSF award #IOS-0820505 to D.H.S.

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