Title: Extreme rainfall drives early onset cyanobacterial bloom

Running head: Extreme rainfall drives early cyanobacterial bloom

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Abstract

1. The increasing prevalence of cyanobacteria-dominated harmful algal blooms is strongly

associated with nutrient loading and changing climatic patterns. Changes to precipitation

frequency and intensity, as predicted by current climate models, are likely to affect bloom

development and composition through changes in nutrient fluxes and water column

mixing. However, few studies have directly documented the effects of extreme

precipitation events on cyanobacterial composition, biomass, and toxin production.

2. We tracked changes in a eutrophic reservoir following an extreme precipitation event,

describing an atypically early toxin-producing cyanobacterial bloom, successional

progression of the phytoplankton community, toxins, and geochemistry.

3. An increase in bioavailable phosphorus by more than 27-fold in surface waters preceded

notable increases in *Aphanizomenon flos-aquae* throughout the reservoir approximately 2

weeks post flood and ~5 weeks before blooms typically occur. Anabaenopeptin-A and

three microcystin congeners (microcystin-LR, -YR, and -RR) were detected at varying

levels across sites during the bloom period, which lasted between 3 and 5 weeks.

4. *Synthesis and applications:* These findings suggest extreme rainfall can trigger early

cyanobacterial bloom initiation, effectively elongating the bloom season period of

potential toxicity. However, effects will vary depending on factors including the timing

of rainfall and reservoir physical structure. In contrast to the effects of early season

extreme rainfall, a mid-summer runoff event appeared to help mitigate the bloom in some

areas of the reservoir by increasing flushing.

Keywords: climate change, cyanobacterial bloom, extreme rainfall, harmful algal bloom,

phosphorus, reservoir, cyanotoxins

Introduction

Harmful algal blooms (HABs) have increased in intensity, frequency, and distribution across

Canada and the world (Ho, Michalak, & Pahlevan, 2019; Orihel et al., 2012; Pick, 2016). In

Ontario alone, the number of reported HAB events between 1994 to 2009 increased by 22×, and,

by 2009, nearly half of those reported were dominated by cyanobacteria (Winter et al.,

2011). This increase in cyanobacterial growth and dominance has been strongly attributed to

increased nitrogen (N) and phosphorus (P) availability (Paerl and Huisman 2009, O'Neil et al.

2012, Sukenik et al. 2015, Paerl et al. 2016) and more recently, climatic shifts. For example,

thermal regime changes have resulted in earlier and warmer water temperatures (O'Reilly et al.,

2015; Richardson et al., 2017), increasing the competitive advantage of certain bloom-forming

cyanobacteria over other organisms (Jöhnk et al., 2008; Paerl & Huisman, 2009). Despite

predicted climatic changes in both temperature and precipitation patterns (IPCC, 2007;

McDermid, Fera, & Hogg, 2015), less attention has been given to the importance of precipitation

regime changes, which may have an even greater impact on cyanobacteria bloom dynamics due

to mixing and nutrient cycles (reviewed in Reichwaldt and Ghadouani 2012).

Extreme rainfall events, both in terms of frequency and intensity, are one of the many

climatic components predicted to change over the next quarter century (Soulis, Sarhadi, Tinel, &

Suthar, 2016). Depending on soil conditions, high intensity precipitation across a watershed may

mobilize large quantities of sediment as well as soil-bound and soluble nutrients resulting in

notable increases to in-lake nutrient bioavailability and suspended solids (Bouvy et al., 2003;

Wood et al., 2017). Increased stream flow will also increase water column mixing, alter flushing

rates and residence times, and weaken vertical stratification (Bouvy et al., 2003; Reichwaldt &

Ghadouani, 2012) - all factors known to influence cyanobacteria bloom development (Jacobsen

& Simonsen, 1993; Mitrovic, Oliver, Rees, Bowling, & Buckney, 2003; Wood et al., 2017).

However, the degree to which those factors will affect primary production in a reservoir is

directly related to event timing (e.g. early vs. late summer) and the zonal gradient created by

basin morphometry differences (Kimmel & Groeger, 1984).

Reservoirs naturally exhibit spatial heterogeneity in physical, chemical, and biological

properties (Thornton, Steel, & Rast, 1996) as the narrow, channelized flow from the catchment

(riverine zone) transitions (transitional zone) to a broader and deeper, lake-like system

(lacustrine zone). Primary production along the riverine-transitional-lacustrine gradient is

primarily mediated by light and nutrient availability, among other interacting factors (Kimmel &

Groeger, 1984). However, large precipitation events have the potential to disrupt characteristic

primary production patterns in these zones, especially if the event introduces sizable nutrient

loads, increases flushing rates, or affects other water-column properties prior to or during a

bloom period. For example, Wood et al. (2017) reported decreased cyanobacterial biomass

following an extreme rainfall event that led to water column cooling and destratification in a

shallow, eutrophic New Zealand lake. Varying responses to extreme rainfall should be

anticipated given the importance of site and event-specific factors.

Changing cyanobacterial bloom dynamics and biomass are a great concern to water

managers. Several bloom-forming cyanobacterial species such as *Aphanizomenon flos-aquae*,

Microcystis aeriginosa, and *Anabaena (Dolicospermum) flos-aquae* synthesize an array of

bioactive compounds that pose acute, chronic, and potentially fatal health risks to humans and

other organisms through dermal contact, inhalation, and/or ingestion of contaminated waters

(Chorus & Bartram, 1999; Chorus, Falconer, Salas, & Bartram, 2000; Codd, 2000). Metabolite

production as well as the composition of toxins released are complex and tied not only to the

cyanobacterial species, successional patterns, and biomass, but also to various environmental

parameters including N, P, temperature, light, pH, salinity, and micronutrients (Chorus &

Bartram, 1999), which may all be altered by intense rainfall events (Reichwaldt & Ghadouani,

2012). To date, only limited information exists regarding how toxicity may change as a result of

precipitation events.

Here, we document the effects of an extreme early summer rainfall event on

cyanobacterial bloom and toxin dynamics in a eutrophic reservoir and contrast these effects to a

second, smaller mid-summer rainfall event. Conestogo Lake (Ontario, Canada; Fig. 1), has

experienced recurring, late-summer cyanobacterial blooms dominated by *Aphanizomenon flos-*

aquae since 2004. On 23rd June 2017, the upper Conestogo watershed experienced a record

rainfall event (daily total 78 mm measured at Conestogo Dam; https://www.grandriver.ca/en/our-

watershed/Record-Rainfall-Flood-June-2017.aspx) that resulted in severe flooding and replaced

~80% of the water in Conestogo Lake during this 2-day event (Grand River Conservation

Authority (GRCA) 2018). We first evaluate how this extreme event affected the reservoir by

characterizing the physical and chemical changes across spatial zones. We then follow the

trajectory of bloom development and toxicity through the season, to better understand how

hydrological events can alter bloom risk.

Materials and methods

Sample collection

We sampled a lacustrine (43.676518, -80.718752; max depth = 17 m) and transitional

(43.708102, -80.712224; max depth = 9 m) site in Conestogo Lake once per week from 21 Jun to

17 Aug 2017. Water column profiles and samples for chemical analysis were collected at the

lacustrine site on 21 Jun (DOY 172) with biological sampling beginning on 05 Jul (DOY 186).

Sampling in the transitional zone began on 05 Jul, 2 weeks after the extreme rainfall event.

(Figure 1*a*). We measured water column profiles at 0.5 m increments using an EXO2 sonde

(YSI, Yellow Springs, OH) equipped with chlorophyll-a (μg L^{-1}), pH, temperature (°C), and

dissolved oxygen (DO, mg L^{-1}) sensors. Photosynthetically active radiation (PAR) in the water

column was also measured in 0.5 m increments using a LI-COR® Underwater Quantum Sensor

(LI-250A, LI-192, Lincoln, NE USA).

For all chemical analyses, discrete water samples were collected at 2 m and 7 m below

(lacustrine only) the water surface, and 0.5 m from the bottom using a Van Dorn bottle.

Unfiltered whole water samples at each depth were portioned into acid-washed 50 mL centrifuge

tubes for total phosphorus (TP, μg L^{-1}) and into a 125 mL dark bottle for phytoplankton

enumeration (2 m only). Additional aliquots were field-filtered using a 0.45 μm syringe filter

(Whatman) for soluble reactive phosphorus (SRP, μg L^{-1}), total dissolved phosphorus (TDP, μg

L^{-1}), ammonium (NH_4^+, mg N L^{-1}), total dissolved nitrogen (TDN, mg N L^{-1}), anions (NO_2^-,

NO_3^-, mg N L^{-1}; Cl$^-$, SO_4^-, mg L^{-1}), total dissolved iron (TDFe), mg L^{-1}), and cations (Mg^{2+}, Ca^{2+},

mg L^{-1}). All samples were transported on ice and chemical samples stored frozen until analysis

at the University of Waterloo (Waterloo, ON) or Centre for Cold Regions and Water Science

(Waterloo, ON).

Samples for total (*i.e.*, intracellular and extracellular) metabolites were pooled from four

discrete depths across the site-specific photic zone (2× Secchi depth). The method serves to

sample the photic zone as a mechanism of estimating all possible toxins in the surface and near-

surface. The estimated photic zone was split into four equal segments, sampled at each

segmented depth using a Van Dorn bottle, then pooled in a bleach-sterilized bucket. For

example, a Secchi depth measured at 1 m would be multiplied by two then divided by four to

create four discrete samplings at 0.5, 1, 1.5, and 2 m. Each pooled sample was collected in an

amber Nalgene™ polyethylene terephthalate glycol (PETG) bottle to limit adsorption (Fisher

reference: 322021-0125). All samples were stored frozen at -20 °C until analysis at the

Université de Montréal (Montréal, QC).

Chemical analyses

SRP, TDP, and TP concentrations were measured on a Cary 100 UV-Vis spectrophotometer

(Santa Clara, CA) following off-line persulfate digestion (TP samples only) as per standard

methods (Eaton, Clesceri, Greenberg, & Franson, 1998). Each N parameter was analyzed on a

separate instrument with NH_4^+ and TDN analyzed on a Westco Smartchem Analyzer (Unity

Scientific, Milford, MA) and Shimadzu TN-L analyzer (Kyoto, Japan), respectively. Anion

concentrations including NO_2^- and NO_3^- were measured on a Dionex Ion Chromatogram ICS-200

(Waltham, MA) while cations and TDFe were measured on a Perkin Elmer Optima 8000 ICP-

OES (Waltham, MA).

Phytoplankton identification and enumeration

Phytoplankton enumeration was completed by D. Findlay (Plankton-R-Us, Winnipeg, MB,

http://www.plankton-r-us.ca) as per Findlay & Kling (2003). Each sample was preserved with

4% Lugol's iodine, gravity concentrated 5-fold after 24 hrs, and finally, stored at 4 °C for ~2

months until analysis. Enumeration was completed on an inverted microscope at 125x, 400x, and

1200x with phase contrast illumination using a modified Ütermohl technique (Nauwerk, 1963)

on 10 mL aliquots of preserved sample. Only cells with viable chloroplasts were enumerated.

Cell counts were converted to mg wet weight biomass (mg m^{-3}) by approximating cell volume,

which were obtained by measurements of up to 50 cells of an individual species and applying the

geometric formula best fitted to the shape of the cell (Vollenweider, 1968). A specific gravity of

1 was assumed for cellular mass (D. Findlay, pers. communication).

Cyanobacterial metabolite analysis

Samples were prepared and screened for each of seventeen cyanobacterial compounds (Table S1)

as per Fayad *et al.* (2015) via on-line solid phase extraction ultra-high performance liquid

chromatography high resolution mass spectrometry (SPE–UHPLC–HRMS) using standards

purchased from Enzo Life Science, Abraxis, or Cyano Biotech GmbH. The limit of detection

(LOD) and limit of quantification (LOQ) were calculated for every batch of samples. Only

results that exceeded the LOQ were included in our analysis. Therefore, some samples may have

had detectable, but unquantifiable toxin concentrations.

Environmental and dam-related data

Meteorological and other reservoir-related data for this study were obtained from the GRCA's

online data portal (https://data.grandriver.ca/). Because GRCA data are provisional, all data were

passed through quality assurance and quality checking metrics to ensure sound data quality

before proceeding with analysis. Spatial data were obtained from open source portals at the

GRCA, the United States Geological Society (https://www.usgs.gov/), Statistics Canada

(https://www12.statcan.gc.ca/), and the Ontario Ministry of Natural Resources (land-use;

https://www.ontario.ca).

Data analysis and statistics

All data analysis was completed in R version 3.6.1 (R Core Team, 2018). Sonde-derived water

column profiles as well as the discrete chemical profiles were constructed using a multilevel B-

spline interpolation from the *MBA* package (Finley, Banerjee, & Hjelle, 2017) and were adjusted

to reflect the reservoir stage elevation at the time of sampling.

To compare phytoplankton community composition with environmental variables, we

used Principal Components Analysis (PCA) and redundancy analysis (RDA) including

environmental variables from the epilimnion and 0.5m above the sediment-water interface.

Species biomass data were Hellinger transformed as recommended for species linear ordinations

(Legendre & Gallagher, 2001), which normalizes for both the site and species biomass needed

for RDA. Only species that appear more than two times in the dataset were included in these

analyses. The environmental matrix included all chemical parameters at 2 m depth (epi) and

those from 0.5 m above the sediment-water interface (sed), sonde-derived measurements for the

epilimnion and hypolimnion (temperature, pH, DO) as well as physical lake parameters

(thermocline depth and metalimnion layer depth; calculated using the *rLakeAnalyzer* package

(Winslow et al., 2019)). We constructed PCA biplots with either species vectors which exceeded

a cumulative goodness-of-fit of at least 0.6 in the ordination plane or significant environmental

parameters. The higher the goodness-of-fit, the better the species is fitted on the corresponding

axis. To test for the significance of environmental factors affecting the phytoplankton

community, permutation tests (permutations = 999) on the models were completed using

anova.cca from the package *vegan.* Epilimnetic chemistry, biomass, and toxin trends were

analyzed using repeated measures ANOVA (RM-ANOVA) with a generalized least squares

model (*gls*) from the *nlme* package (Pinheiro, Bates, DebRoy, Sarkar, & R Core Team, 2018)

where time and site were fixed effects and the correlation matrix was selected based on Akaike

Information Criterion (AIC). To account for temporal autocorrelation, we nested time within

site. Finally, we tested for the correlation of biomass with toxin concentrations using Kendall's

rho.

Results

Extreme rainfall disturbance in early summer

Record rainfall (daily range 40 - 120 mm) was recorded across the Upper Conestogo watershed

on 23 Jun 2017 (DOY 174) and led to widespread flooding. This event was the largest single

rainfall in the Grand River Watershed (Luther Marsh station; ~50 km distant; records since 1950;

https://www.grandriver.ca/en/our-watershed/Record-Rainfall-Flood-June-2017.aspx).

222   Mean daily flow rates into Conestogo Lake via the Conestogo River increased from 4.56

223    $m^3 s^{-1}$ to 219 $m^3 s^{-1}$ (max 521 $m^3 s^{-1}$) in the two day period, which consequently increased the

224   total reservoir volume by 3.5 Mm^3,and is estimated to have replaced ~80% of the total reservoir

225   volume (Fig 1*b-e*) (GRCA 2018). Within several days, the reservoir storage and drawdown

226   returned to levels targeted for regular reservoir operation.

227   Reservoir volume replacement restructured the physical water column. Both the

transitional and lacustrine zones were well mixed, isothermal, and oxygenated immediately

following the event (Fig 2). Thermal stratification did not return to either location during the

sampling period. Intermittent hypolimnetic hypoxia (DO < 2 mg L^{-1}) developed in the

transitional zone while sustained hypoxia developed in the lacustrine within two weeks of the

flood event (05 July, DOY 186). Lacustrine SRP, TDP, and TP increased between 2-33× across

the sampled depths (2m: 7 to 27×, 7m: 9 to 33×, bottom [16m]: 2 to 3×; Figs. 2, 3) before and

after the extreme rainfall. Nitrogen concentrations at all measured depths were only marginally

different post-rainfall and maintained concentrations near 5.5 mg N L^{-1} for NO$_3^-$ (mean prior =

4.89 mg N L^{-1}; mean after = 5.45 mg L^{-1}) and TDN (mean prior = 5.63 mg N L^{-1}; mean after =

5.71 mg L^{-1}) (Fig. 2).

Dynamics following early summer extreme rainfall event

Mean surface water temperatures (0 - 2 m) across the reservoir ranged from 21.2 to 25.5 °C.

Nutrient concentrations (both N and P) were, on average, higher in the transitional zone than the

lacustrine zone and significantly changed through time (Table 1; RM-ANOVA). Epilimnetic

chlorophyll a, dissolved oxygen (DO), and temperature did not significantly change through time

across sites. As predicted, the transitional zone had higher concentrations of suspended solids in

the water column and lower light availability as compared to the lacustrine immediately

following the reservoir disturbance (Fig. 4).

Phytoplankton community composition. The transitional zone phytoplankton community was

composed predominately of diatoms including *Stephanodiscus niagarae* (P5530) and *Cyclotella*

psuedostelligera (P5508) and the dinoflagellate *Ceratium hirundenella* (P7644) with

approximately 5x higher total phytoplankton biomass in the transitional (total = 5742 mg m^{-3} wet

weight) than the lacustrine (total = 1113 mg m^{-3} wet weight; 2 weeks post-extreme event, Fig.

6a). In contrast to the transitional zone, the lacustrine zone contained high relative abundances of

Cryptophytes including *Cryptomonas rostratiformis* (P6565), *Cryptomonas erosa* (P6558), and

Rhodomonas minuta (P6554) and Chrysophytes (*Ochromonas globose* [P4398]), all of which

returned following the cyanobacterial bloom. Communities following the extreme rainfall event

(DOY 186) were strongly associated with PAR and epilimnetic SRP, TDP, and TP. Communities

in both zones gave way to cyanobacterial dominance (> 50% of community by biomass, Figs.S1

- S2) 20 - 24 days following the flood with apparent changes to epilimnetic temperature and pH

associated with temporal community shifts and increased presence of the filamentous nitrogen-

fixer, *Aphanizomenon flos-aquae* (Fig 5a; P1041).

Cyanobacterial dynamics. Cyanobacterial bloom duration, measured as the period during which

cyanobacterial biomass comprised > 50% of the fractional biomass, persisted between 26 and 34

days and occurred 4 to 6 weeks earlier than has been previously documented by the GRCA (S.

Cooke, pers. communication). Bloom initiation was 6 days earlier in the transitional zone as

compared to the lacustrine. Mean total cyanobacterial biomass over the sampling campaign was

1550 ± 2920 mg m^{-3} and 591 ± 1290 mg m^{-3} in the transitional and lacustrine zones, respectively

and was significantly driven by an interaction between SRP and surface temperatures (RM-

ANOVA, $F_{1,\,6} = 6.84$, $p = 0.0399$). There was no relationship between total cyanobacterial

biomass and time or site (RM-ANOVA, $F_{1,10} = 0.086$, $p = 0.77$). *Aphanizomenon flos-aquae* was

the dominant species and represented 95-97% (by biomass) of the cyanobacterial fraction across

all sites for the entirety of the sampling campaign. The highest measured *A. flos-aquae* biomass

occurred on 19 July (DOY 200, 10,900 mg m^{-3}, 4,500 cells mL^{-1}) in the transitional and near 25

Jul (DOY 206; 4,710 mg m^{-3}) in the lacustrine. *Woronichinia compacta*, *Anabaena*

(Dolichospermum) flos-aquae, and *Microcystis aeruginosa* were minor contributors (< 5 %) to

cyanobacterial biomass while *Anabaena* (*Dolichospermum) crassa*, *Planktolyngbya limnetica*,

and *Pseudoanabaena sp.* were only observed once during the campaign at low biomass (< 8 mg

m^{-3}). GRCA managers have visibly recorded cyanobacterial blooms dominated by an

Aphanizomenon sp. in Conestogo Lake from late August through September (DOY ~220 - 250)

since 2004 (S. Cooke, pers. communication).

Cyanotoxin dynamics. Four of the 17 screened cyanobacterial metabolites were detected in

Conestogo Lake. Detectable anabaenopeptin-A (AP-A), microcystin-LR (MC), -YR, and -RR

were present in 87 % of the samples ($n = 13$) but were only quantifiable in 73 % ($n = 11$; Table

2). Total microcystin concentration did not exceed 1.0 µg L^{-1} and was comparable to previously

measured concentrations during the summer months (Yakobowski, 2008). Metabolite type and

concentration varied through time (RM-ANOVA, $F_{3,44} = 9.21$, $p = 0.001$) and were different

between the two sites (RM-ANOVA, $F_{1,44} = 6.55$, $p = 0.014$). Quantifiable concentrations of all

three microcystins were present in the transitional zone, while MC-LR and MC-YR were present

in the lacustrine. MC-LR was the dominant microcystin variant across all sites with the highest

quantified values in the transitional zone at peak *A. flos-aquae* and *M. aeruginosa* biomass (Fig

6*b*, Table 2). Further, MC-YR in the lacustrine did not correspond with the MC-LR peak. AP-A

was present in both zones throughout the bloom period MC-LR concentration was significantly

correlated with the biomass of *M. aeruginosa* (Kendall rho = 0.76, $p = 0.03$). No other toxin was

correlated with the biomass of any other observed cyanobacterial species.

Dynamics following smaller mid-summer runoff event

A second, much smaller and transient increase in flow to the reservoir occurred on 24 Jul (DOY

205), associated with an ~25 mm rainfall event occurring during already wet conditions (Fig. 1).

Effects of this event were much more muted. Overall, there was a slight decrease in surface

water temperatures and decrease in the extent of oxygen supersaturation in both reservoir zones,

with minimal effects on nutrients (Fig. 2). TSS increased slightly in the lacustrine zone and

decreased in the transitional zone, suggesting the event was not of sufficient magnitude to lead to

substantive additional erosion in the upstream catchment. Light irradiation increased in both

areas (Fig. 4). In the transitional zone, this change was concurrent with a 66% decrease in total

phytoplankton biomass and ~50 % decrease in total suspended solid while the lacustrine showed

increases in both during this period (Figs. 4 & 6). Though the phytoplankton community

composition in the transitional zone did not change, the lacustrine showed a marked increase in

the cryptophyte *Katablepharis ovalis* (P6568) for the rest of the sampling period. Total

cyanobacterial toxin concentrations declined by ~75% in the transitional zone, while more muted

declines were apparent in the lacustrine. All MC variants were replaced with AP-A as the

dominant metabolite at all sites by 25 July (DOY 206).

Discussion

Globally, increases in extreme rainfall events are predicted to outpace changes in total

precipitation (Allen and Ingram 2002, IPCC 2007, Soulis et al. 2016). The impacts of extreme

rainfall on cyanobacterial blooms are inherently complex (Reichwaldt and Ghadouani 2012,

Wood et al. 2017) with the overall effect regulated by parameters including intensity, volume of

water inflow, and timing with respect to seasonality (reviewed in Reichwaldt and Ghadouani

2012). The limited studies investigating such extreme rainfall events on cyanobacterial bloom

development have identified changes to flushing rates, water column mixing, and nutrient inputs

from rainfall events as the main abiotic factors affecting cyanobacteria and phytoplankton

communities (Bouvy et al. 2003, Reichwaldt and Ghadouani 2012) – all factors that appeared to

impact the bloom dynamics in Conestogo Lake.

The record rainfall event effectively created a new high-P regime as a result of the

significant influx of P arriving from the Conestogo River, and near complete replacement of the

reservoir volume. Soluble reactive phosphorus concentrations increased ~7-27 fold in the

epilimnion following the flooding. Substantial increases in P loading, such as these, are common

following heavy rainfall, particularly if they are preceded by warm, dry periods (Jones &

Poplawski, 1998). For example, in Australian reservoirs, a 440 mm rainfall in a 3-day period

resulted in an input equivalent to 400% of the average annual in-lake P mass (Jones and

Poplawski 1998). In our study system, the Upper Conestogo watershed, which is predominantly

Tavistock Till, is particularly prone to erosion and serves as a key mechanism for particulate P

transport (Loomer & Cooke, 2003; Macrae, English, Schiff, & Stone, 2007). High flow events

can also be important to transport of soluble P species, as P is desorbed from particulate forms,

or released from soils.

Though surface water temperatures remained similar following the June extreme rainfall

event, the bottom sediments warmed from 13 to 18°C. While we cannot isolate the effects of

thermal and biogeochemical changes in bloom initiation, we note that nitrogen concentrations

changed little as a result of the flooding. This catchment has chronic high N concentrations (3 -

5 mg NO_3-N L^{-1}), associated with long-term intensive agricultural usage within the catchment. In

addition to the prevalence and eventual dominance by an N-fixing cyanobacterium, the lack of

change in N suggests that it was not likely to be the key factor leading to early bloom initiation.

Instead, it appears that the excessive supply of P initiated the development of the

observed weakly toxic early-summer cyanobacterial bloom (Figs 5 and 6). Typically,

Aphanizomenon flos-aquae dominated blooms do not develop in the Conestogo Reservoir until

late summer (S. Cooke, pers communication), ~35 to 40 days after our documented extreme

event. Interestingly, the bloom originating in the transitional zone was not only initiated a week

earlier than in the lacustrine, but it was also sustained through the typical bloom period,

highlighting the potential for elongation of the bloom season and risk associated with extreme

rainfall events. However, the intensity of the transitional bloom was directly influenced by the

mid-summer rainfall event, which flushed both biomass and toxins. The event was not of

sufficient volume to have similar effects on the deeper, larger volume lacustrine zone.

The environmental factors driving toxin production and their seasonal dynamics are quite

complex. We detected low quantities of three MC variants, MC-LR, -YR, and -RR in the early-

summer cyanobacterial bloom that dissipated following the mid-summer event. Of these, only

MC-LR was significantly correlated with the low biomass of a single taxa (*M. aeruginosa*).

However, because it difficult to assign strong relationships between the toxins and

cyanobacterial species without direct molecular studies (Vezie et al., 1998), it remains possible

that there could be the presence of a rare, toxin-producing species or strain, or multiple taxa

involved in toxin production.

MC variants are not the only cyanobacterial compounds capable of impacting human

health. Several studies using cyanobacterial extracts have reported harmful and/or toxic effects

that could not be explained solely by microcystin concentration or presence, suggesting the

possibility of other toxic compounds (Keil et al. 2002, Teneva et al. 2005, Baumann and Jüttner

2008, Smutná et al. 2014, Lenz et al. 2019). Although not often reported, improved analytical

techniques have identified numerous other bioactive compounds such as cyanopeptolins and

anabaenopeptins that are both detectable in freshwaters and often produced simultaneously with

microcystin variants (Harada et al. 1995, Welker and Von Döhren 2006, Gkelis et al. 2015,

Beversdorf et al. 2017, 2018) at similar or greater levels (Janssen 2019). For example,

Beversdorf *et al.* (2017) reported an average of 0.65 µg L^{-1} total microcystin in Lake

Koshkonong, Wisconsin, while anabaenopeptin-B and -F were measured at 6.56 µg L^{-1}

combined. Though there are no case studies of human toxicity caused by anabaenopeptins, the

compound inhibits carboxypeptidase A, and like microcystin, also inhibits protein phosphatases

with slightly overlapping inhibitory concentration ranges (Honkanen et al. 1990, Sano et al.

2001, Spoof et al. 2016). The concentrations of these toxins that could affect human health are

unknown, which has prevented the development of recreational and drinking water regulations

and/or advisories for these potentially toxic compounds. Like some microcystins, the ecological

effects of anabaenopeptins may be observed at relatively low concentrations. A recent study by

Lenz *et al.* (2019) reported induced toxicity by low concentrations (10 µg L^{-1}) of

anabaenopeptins, including AP-A, on the nematode *Caenorhabditis elegans* resulting in reduced

reproduction, reduced lifespan, and delayed hatching. Though compounds like AP-A have been

previously considered non-toxic, they may represent a new class of emerging toxins, whose

potential impacts to human health and toxicity to aquatic organisms require immediate attention

and therefore, inclusion in risk assessment for lake mitigation and monitoring programs.

While increasing bloom incidence globally has been associated with elevated nutrient

loads and with warming (Ho & Michalak, 2019; Paerl & Otten, 2016), more attention is needed

to understand the role of extreme events in altering bloom risk, and how these events may differ

in their seasonal impacts, and on spatially complex but crucial water resources, such as

reservoirs. The environmental factors driving cyanobacterial blooms are inherently complex and

dynamic, making predictions about toxicity and impact on end-user health and safety difficult to

obtain across diverse ecosystems. Extreme disturbance events, like the early summer event

described here, will only make that task more difficult since they are not only predicted to

increase but will also likely mask historical drivers of changes in bloom risk. Managers cannot

effectively manage a future ecosystem based on averages of the past. Instead, a holistic

understanding of multiple drivers of cyanobacterial bloom risk, accounting for both extreme

events, and reservoir structure, is required. Within this region, we hypothesize that early season

events, have the greatest potential to worsen bloom risk by extending the bloom season as shown

here. Later season events can increase annual nutrient loads but may also help flush existing

blooms. Here, this appears to occur during the late season runoff event with apparent effects on

toxins and cyanobacterial biomass in the transitional zone, with only minimal change apparent in

the lacustrine zone.

Author's contributions

MLL, HMB, DS, SS, SLS, and JJV conceived of the ideas and designed methodology; MLL,

DS, SS collected the data; MLL analyzed the data and lead the writing of the manuscript. All

authors contributed critically to the drafts and gave final approval for publication.

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422   *Data availability statement*

423   Data and code that support the analysis and findings in this study are available in figshare at

424   http://doi/10.6084/m9.figshare.7811963.

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Table 1. Post-flood mean (minimum - maximum) values for measured epilimnetic (2 m)

characteristics in the transitional and lacustrine zones with RM-ANOVA *p*-values. RM-ANOVA

tables are available in the supplement. Bold values are significant ($p < 0.05$).

	Chl-a mg L^{-1}	DO mg L^{-1}	Temp °C	NO$_3$-N mg L^{-1}	TDN mg L^{-1}	SRP µg L^{-1}	TDP µg L^{-1}	PAR µmol s^{-1} m^{-2}
Transitional	6.8 (1.6 - 14.0)	10.6 (7.9 - 14.5)	22.8 (21.7 - 24.0)	4.82 (3.99 - 5.65)	5.32 (4.24 - 5.86)	13.66 (1.52 - 65.2)	28.76 (11.62 - 85.77)	42.6 (6.33 - 93.9)
Lacustrine	5.4 (2.2 - 10.1)	8.6 (5.6 – 11.0)	22.3 (21.2 - 23.2)	4.37 (3.34 - 5.63)	4.74 (3.39 - 5.85)	5.89 (1.56 - 24.61)	18.53 (7.79 - 41.59)	69.1 (16.0 – 155)
Intercept	**0.001**	**< 0.001**	**< 0.001**	**< 0.001**	**< 0.001**	**0.01**	**0.001**	**< 0.001**
time	0.415	0.882	0.234	**< 0.001**	**< 0.001**	**0.014**	**0.007**	**0.003**
site	0.586	0.204	0.244	**0.001**	**0.004**	0.196	0.206	0.064

Abbreviations: chlorophyll a (Chl-a), dissolved oxygen (DO), water temperature (temp), nitrate (NO3), total dissolved nitrogen (TDN), soluble reactive phosphorus (SRP), total dissolved phosphorus (TDP), photosynthetically reactive radiation (PAR)

Table 2. Quantified cyanobacterial metabolites across sampling sites as measured by LC-HMRS

with values which exceeded the limit of detection (LOQ) presented as mean ng L^{-1} (standard

deviation ng L^{-1}), those in italics exceeded the limit of detection (LOD) but not the limit of

quantification (LOQ), and those with (-) < LOD. The metabolites MC-LA, -LY, -LW, -LF, -

HiIR, -HtyR, anatoxin, homoanatoxin, cylindrospermopsin, and cyanopeptolin A were not

detected in samples.

Date	Site	cyanobacterial secondary metabolites			
		AP-A	MC-LR	MC-RR	MC-YR
21 Jun	Lacustrine	-	-	-	-
05 Jul		-	-	-	*33.3 (16.1)*
14 Jul		*23.7 (4.9)*	-	-	*20.6 (29.1)*
18 Jul		109.3 (78.6)	406.9 (18.1)	-	-
25 Jul		189.7 (35.6)	88.2 (6.2)	*35.5 (9.2)*	*20.0 (28.2)*
01 Aug		278.8 (27.4)	16.8 (10.6)	-	59.2 (16.5)
10 Aug		619.3 (141.0)	-	-	*19.4 (1.3)*
17 Aug		344.0 (50.3)	-	-	-
05 Jul	Transitional	-	-	-	-
11 Jul		*18.6 (26.3)*	-	-	-
19 Jul		-	763.6 (154.2)	121.0 (1.6)	97.3 (36.2)
25 Jul		174.3 (34.9)	72.1 (23.3)	*27.6 (2.7)*	-
31 Jul		188.6 (108.5)	-	-	-
10 Aug		411.0 (30.6)	-	-	-
17 Aug		856.8 (37.2)	-	-	-

Abbreviations: anabaenopeptin A (AP-A), microcystin (MC-)

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Figure legends

Fig. 1 (*a*) Conestogo Lake (Mapleton Township, Ontario, Canada) and its associated watershed.

Sampling began on 21 Jun 2017 (DOY 172) in the lacustrine zone and on 05 Jul (DOY 186) in

the transitional zone. Total daily precipitation (mm) measured in upper Conestogo watershed (*b*,

gauge in Arthur, ON) was directly related to the Conestogo River mean daily flow rate (*c, n =*

48) into the reservoir (gauges in Drayton and Moorefield, ON) and the calculated reservoir

residence time (*d*). The average flow rate during the drawdown period (May – September; red

line, *c*) is typically near 0.01 m^3 s^{-1}. The extreme precipitation event on 23-24 Jun (DOY 173-

174; grey vertical bar) directly corresponds to in increased mean daily flow into the Conestogo

Reservoir, increased reservoir storage, and reduced residence time. Dark gray area around the

mean represents the standard deviations of daily measurements (flow, *n* = 48; storage, *n* = 24). A

secondary rainfall event (~25 mm) on 24 Jul (DOY 205) also impacted flow rates into the

reservoir. This figure contains information made available under GRCA's Open Data Licence

v3.0 (*a*) and v2.0 (*b - e*).

Fig. 2 Temperature (*a-b*, °C) and dissolved oxygen (*c-d*, mg L^{-1}) water column profiles collected

at 0.5 m depth increments (grey points) in the transitional (*a, c*) and lacustrine (*b, d*) zones from

21 Jun 2017 (lacustrine only; DOY 172) to 17 Aug (DOY 229). Hypoxic (red squares) zones

were detected at all sites, but anoxia (red triangles) was only detected in the lacustrine zone. TDP

(*e-f*) and TDN (*g-h*) profiles collected at 2 m and 7 m below (lacustrine only) the surface and 0.5

m from the bottom. Profile depths at each site were standardized to the dam stage elevation

(meters above sea level, m.a.s.l.) at the time of sampling. The grey regions in the x-y plane

illustrate the interaction between time and reservoir depth at each site, where the x-axis zone

indicates a later sampling date in the transitional and the y-axis grey zones illustrate differences

in site depths with reference to the dam (max elevation = 393 m.a.s.l, dam outflow = 373.5

m.a.s.l.). The first white vertical bar corresponds to the record rainfall event on (23-24 Jun 2017;

DOY 174-175) and the second to a 25 mm rainfall event on 24 Jul (DOY 205). The black

horizontal bar corresponds to the bloom period in each of the two zones.

Fig. 3 Multiplicative changes in phosphorus species and Mn (*a*), nitrogen species and DOC (*b*),

anions (*c*), cations and TDFe (*d*) prior to (21 Jun 2017, DOY 172) and after the early summer

flood event (05 Jul 2017, DOY 186) in the lacustrine zone. Samples were collected at 2 m depth

(purple), 7 m depth (blue), and 0.5 m from sediment-water interface (16m; green) in the

lacustrine zone. Values above 1× (black horizontal line) correspond to increased concentrations

while those below 1× indicate reduced concentrations. Samples at 2m for anions (*b* [NO_3^-, Cl^-,

SO_4^-) were not analyzed.

Fig. 4 TSS (*a*) and light availability (*b;* photosynthetically available radiation (PAR) at 2 m in

the transitional (green) and lacustrine (purple) zones. The horizontal bar corresponds to the

cyanobacteria bloom period in the transitional (grey) and lacustrine (black) zones while the

vertical bars are associated with the extreme rainfall event on 23 June (DOY 174) and the late

summer event on 24 Jul (DOY 205).

Fig. 5 Principal component analysis (PCA) (scaling 1) of phytoplankton community, with the

most influential phytoplankton species (*a,* Goodness of fit > 0.6) and significant (P < 0.05) and

weekly significant (P < 0.1) environmental variables (*b*). The length of the vector is proportional

to the importance of the descriptor to the sites. The day of year for each community are in boxes

and are linked through time for each reservoir zone (lacustrine [purple], transitional [green].

Phytoplankton labels: P1041; *Aphanizomenon flos-aquae*, P2235; *Ankistrodesmus spiralis*,

P6554; *Rhodomonas minuta*, P6558; *Cryptomonas erosa*, P4398; *Ochromonas globose*, P5515;

Fragilaria crotonensis, P7644 *Ceratium hirundenella*. Environmental labels: SRP_ug.L_epi;

epilimnetic soluble reactive phosphorus, SRP_ug.L_epi; sedimentary soluble reactive

phosphorus, TDP_ug.L_epi; epilimnetic total dissolved phosphorus, TP_ug.L_epi; epilimnetic

total phosphorus, PAR; photosynthetically active radiation, DO_mg.L_avg_epi; epilimnetic

dissolved oxygen, pH_avg_epi; epilimnetic pH, pH_avg_sed; sedimentary pH, temp_C_avg_epi;

epilimnetic temperature, temp_C_avg_sed; sedimentary temperature.

Fig. 6 Phytoplankton biomass (*a*) and cyanobacterial metabolite concentration trends (*b*;

exceeding the limit of quantification) for anabaenopeptin-A (AP.A) and microcystin-LR

(MC.LR), -YR (MC.YR), and -RR (MC.RR) in the transitional and lacustrine zones. The flood

event on 23-24 Jun (DOY 174-175 and secondary event on 25 July (205) are noted by vertical

bars. *Aphanizomenon flos-aquae* dominated (> 50%) phytoplankton biomass in the transitional

zone near 10 Jul 2017 (DOY 191) and in the lacustrine near 16 Jul 2017 (DOY 197). Only

cyanobacterial species that were present in more than 2 samples are plotted above. Of the seven

cyanobacterial species detected, various strains of *A. flos-aquae*, *A. flos-aquae (Dolichospermum*

flos-aquae), *M. aeruginosa*, and *W. compacta* have been documented as potential toxin

producers. Typical bloom onset occurs between DOY 220 – DOY 255 (S. Cooke, pers.

communication).

Fig. 1

630   **Fig. 2**

631 **Fig. 3**

635   **Fig. 4**

638   **Fig. 5**

642   **Fig. 6**