Metabolic specializations within a bacterial community to create living rocks

Running Title: Functional guilds in living rocks

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ABSTRACT

Stromatolites are complex microbial mats that form lithified layers and ancient forms are the oldest evidence of life on earth, dating back over 3.4 billion years. Their emergence aligns with the oxygenation of the Earth's atmosphere and insight into these ancient structures would shed light on the earliest days of Earth. Modern stromatolites are relatively rare but may provide clues about the function and evolution of their ancient counterparts. Previous studies have assessed microbial diversity and overall functional potential but not at a genome-resolved level. In this study, we focus on peritidal stromatolites occurring at Cape Recife and Schoenmakerskop on the southeastern South African coastline. We identify functional gene sets in bacterial species conserved across two geographically distinct stromatolite formations and show that these bacteria may promote carbonate precipitation through the reduction of sulfur and nitrogenous compounds and produce calcium ions that are predicted to play an important role in promoting lithified mats. We propose that abundance of extracellular alkaline phosphatases, in combination with the absence of transport regulatory enzymes, may lead to the precipitation of phosphatic deposits within these stromatolites. We conclude that the cumulative effect of several conserved bacterial species drives accretion in these two stromatolite formations.

INTRODUCTION

Stromatolites are organo-sedimentary structures that date back more than 3.4 billion years, forming the oldest fossils of living organisms on Earth [1]. The emergence of Cyanobacteria in stromatolites approximately 2.3 billion years ago initiated the Great

Oxygenation Event that fundamentally altered the Earth's redox potential and resulted in

an explosion of oxygen-based and multicellular biological diversity [2]. Study of ancient

stromatolites could provide insight into how microorganisms shaped early eukaryotic

evolution but unfortunately these ancient microbial mats are not sufficiently preserved

for identification of these microbes and individual bacteria cannot be classified further

than Cyanobacteria due to morphological conservatism [1, 3]. The study of extant

stromatolite analogues will therefore help to elucidate the biological mechanisms that

led to the formation and evolution of their ancient ancestors. Modern stromatolites are

formed through a complex combination of both biotic and abiotic processes. The core

process revolves around the carbon cycle where photosynthesizing bacteria transform

inorganic carbon into bioavailable organic carbon for respiration. Bacterial respiration in

turn results in the release of inorganic carbon, which, under alkaline conditions, will bind

cations and precipitate primarily as calcium carbonate [1]. This carbonate precipitate,

along with sediment grains, can then become trapped within cyanobacterial biofilms

forming the characteristic lithified layers.

Alteration of the pH and subsequently, the solubility index (SI), through microbial cycling

of redox sensitive compounds such as phosphate, nitrogen, sulfur and other nutrients

within the biofilm may promote mineralization or dissolution of carbonate minerals. This

in turn regulates the rate of carbonate accretion and stromatolite growth. Particularly,

photosynthesis and sulfate reduction have been demonstrated to increase alkalinity

thereby promoting carbonate accretion, resulting in the gradual formation of lithified

mineral layers [1, 4]. In some stromatolite formations such as those of Shark Bay in

Australia, there is abundant genetic potential for both dissimilatory oxidation of sulfur

(which may promote dissolution under oxic conditions and precipitation under anoxic

conditions) and dissimilatory reduction of sulfate (which promotes precipitation) [5–7].

The biogenicity of stromatolites has been studied extensively in stable environments,

such as the hypersaline and marine stromatolites of Shark Bay, Australia and Exuma

Cay, Bahamas, respectively [8, 9]. Overall, Cyanobacteria, Proteobacteria and

Bacteroidetes appear to be abundant in both marine and hypersaline systems and the

Cyanobacteria are proposed to be particularly vital to these formations through the

combined effect of biofilm formation, carbon fixation, nitrogen fixation and tunneling

(endolithic) activity [6, 8–10]. It is further hypothesized that Proteobacteria and

Bacteroidetes influence the solubility index of the systems through sulfur cycling,

anoxygenic phototrophy and fermentation of organic matter [11, 12].

Peritidal tufa stromatolite systems are found along the southeastern coastline of South

Africa (SA). They are geographically isolated, occurring at coastal dune seeps

separated by stretches of coastline [13]. In these SA systems, stromatolite formations

extend from freshwater to intertidal zones and are dominated by Cyanobacteria,

Bacteroidetes and Proteobacteria [14]. A key difference between SA stromatolites and

hypersaline/marine stromatolites is chemical stability: The SA stromatolites are

impacted by a far more chemically unstable environment due to the frequent mixing of

fresh and tidal waters [15]. Thus the SA stromatolite formations face numerous

environmental pressures such as desiccation, limited inorganic phosphorus and periodic

inundation by seawater, which affect the nutrient concentrations, temperature and

chemistry of the system [15]. These formations are characterized by their proximity to

the ocean, where stromatolites in the upper formations receive freshwater from the

inflow seeps, middle formation stromatolites withstand a mix of freshwater seepage and

marine over-topping and lower formations are in closest contact with the ocean [14].

Microbial communities within these levels therefore likely experience distinct

environmental pressures, including fluctuations in salinity and dissolved oxygen [15].

While carbon predominantly enters these systems through cyanobacterial carbon

fixation, it is unclear how other members of the stromatolite-associated bacterial

consortia influence mineral stratification resulting from the cycling of essential nutrients

such as nitrogen, phosphorus and sulfur. Since the SA peritidal stromatolite systems

are in constant nutritional and chemical flux with varying influence from the fresh and

marine water sources, they present an almost ideal *in situ* testing ground for

investigating how stromatolite-associated microbial consortia interact with their

environment. Identification of conserved bacterial species across both time and space

and across varied environmental pressures would suggest that these bacteria are not

only robust but likely play important roles within the peritidal stromatolite consortia.

Using a metagenomic approach, we sought to gain insight into the foundational

metabolic processes that result in stromatolite formation. We assembled and annotated

183 putative bacterial genomes from peritidal stromatolites of two geographically

isolated sites near Port Elizabeth, South Africa. We identified several temporally and

spatially conserved bacterial species and functional gene sets, which may play a central

role in establishing and maintaining peritidal stromatolite microbial communities.

METHODS

Sample collection and DNA isolation. Samples for metagenomic analysis were

collected from Schoenmakerskop (34°02'28.2"S 25°32'18.6"E) and Cape Recife

(34°02'42.1"S 25°34'07.5"E) in January and April 2018, while samples used for

16SrRNA analysis were collected in July 2019. The samples were stored in RNAlater at

-20 °C. DNA was extracted from ~1g of sample using Zymo quick DNA Fecal/Soil

Microbe Miniprep Kit (Zymo Research, Cat No. D6010) according to the manufacturer's

instructions.

Amplicon sequence analysis. Kapa HiFi Hotstart DNA polymerase (Roche, Cat No.

KK2500) was used to generate amplicon libraries of the V4-V5 region of the 16s rRNA

ribosomal subunit gene with the primer pair E517F (5'-GTAAGGTTCYTCGCGT-3') and

E969-984 (5'-CAGCAGCCGCGGTAA-3') [16] using the following cycling parameters:

Initial denaturation at 98 °C; 5 cycles (98 °C for 45 seconds, 45 °C for 45 seconds and

72 °C for 1 minute); 18 cycles (98 °C for 45 seconds, 50 °C for 30 seconds and 72 °C

for 1 minute; final elongation step at 72 °C for 5 minutes. PCR products were purified

using the Bioline Isolate II PCR and Gel kit (Bioline, Cat. No. BIO-52060). Samples

were sequenced using the Illumina Miseq platform. Amplicon library datasets were

processed and curated using the Mothur software platform [17]. Sequences shorter than

200 nucleotides or containing ambiguous bases or homopolymeric runs greater than 7

were discarded. Sequences were classified using Naive-Bayesian classifier against the

Silva database (v132) and Vsearch software [18] was used to remove chimeras. Reads

that were 97% similar were combined into operational taxonomic units (OTUs) using the

Opticlust method [19]. Abundance data was then standardized by total, transformed by

square root and statistically analyzed using the Primer-e (V7) software package [20].

Metagenomic binning. Shotgun DNA libraries were prepared and sequenced using the

IonTorrent Ion P1.1.17 Chip technology (Central Analytical Facilities, Stellenbosch

University, South Africa). Adapters were trimmed and trailing bases (15 nts) were

iteratively removed if the average quality score for the last 30 nts was lower than 16,

which resulted in approximately 30–45 million reads per sample. Resultant

metagenomic datasets were assembled into contiguous sequences (contigs) with

SPAdes version 3.12.0 [21] using the –iontorrent and --only-assembler options with

kmer values of 21,33,55,77,99,127. Contigs were then clustered in putative genomic

bins using Autometa (Master branch, commit: a344c28) [22]. Bins were manually

curated using a tool developed within the Kwan group (table_and_cluster_col branch,

commit: 6e99c2f).

Identification of conserved taxa. Conservation of bacterial taxa was calculated using

average nucleotide identities (ANIs) of all genomic bins, which were calculated in a

pairwise manner using FastANI [23]. All genomic pairs sharing > 97% ANI were subset

and considered conserved taxa. Temporal conservation of taxa was assessed between

samples from upper formations of Schoenmakerskop and Cape Recife (Fig.1), collected

in January and April 2018 respectively. Spatial conservation between sites was

analyzed in a similar pairwise manner between upper, middle and lower regions of

Schoenmakerskop and Cape Recife.

Genome taxonomic classification. Genomes were classified using the standalone

GTDB-Tk tool (version 0.3.2) using the classify workflow and the Genome Taxonomy

Database version 89 [24]. The tool appears unable to classify genomes less than 10%

complete, as indicated in the respective tables.

Genome taxonomic clustering. Phylogeny of bacterial genomes carrying genes from

functional groups investigated here (nitrogen, sulfur, calcium and phosphate

metabolism) was inferred using JolyTree [25] with a sketch size of 5 000. JolyTree

infers phylogeny through computation of dissimilarity of kmer sketches, which is then

transformed for the estimation of substitution events of the genomes' evolution [25].

Genome annotation. Manually curated putative genomic bins were annotated using

Prokka version 1.13 [26], with GenBank compliance enabled. Protein-coding amino-acid

sequences from genomic bins were annotated against the KEGG database using

kofamscan [27] with output in mapper format. KEGG orthologs were manually collected

into functional groups and annotations were extracted from kofamscan output: Briefly, a

dictionary of genomes was made with all KO numbers supplied in a KEGG annotation

query list (i.e. functional group). Kofamscan output per genome was parsed and each

time an entry matching that within the query list was found, a count for that annotation

was increased by one, resulting in a count table of functional group KO annotations per

genome. Calcium binding proteins and calcium transporting ATPases were counted

from PROKKA annotations by searching "Leukotoxin" (Comparison of genes annotated

as leukotoxins, against the NCBI database were identified as calcium binding rather

than toxins) and "Calcium transporting ATPase" respectively.

Comparative analysis with Shark Bay. A total of 96 MAGs representative of the

microbiota associated with Shark Bay stromatolites were downloaded from MG-RAST

(https://www.mg-rast.org/linkin.cgi?project=mgp81948) [7]. These genomes were

annotated using Prokka and kofamscan and classified using GTDB-Tk as described for

the genomes of Cape Recife and Schoenmakerskop stromatolites.

Data availability. Raw 16S rRNA gene amplicon sequence files were uploaded to the

NCBI sequence read archive database in BioProject PRJNA574289. Raw reads and

binned genomes will be deposited in GenBank and respective accession numbers will

be included in the accepted version of this manuscript. Assembled metagenomes were

also uploaded to MG-RAST (Keegan et al., 2016) with the identifiers: mgm4784267,

mgm4790047, mgm4797596, mgm4797628, mgm4797629, mgm4797630,

mgm4802824, mgm4802825 and mgm4802891.

RESULTS AND DISCUSSION

Peritidal tufa stromatolites occur at several sites along the southeastern coast of South

Africa. Two stromatolite sites, Cape Recife and Schoenmakerskop, which are

approximately 3 km apart and have been well characterized with respect to their

chemical environment [14,15] were chosen for this study. Stromatolite formations at

both sites begin at a freshwater inflow and end before the subtidal zone (Fig. 1A-B) and

are exposed to different levels of tidal disturbance. Samples were collected from the

upper stromatolites at Cape Recife and Schoenmakerskop in January and April 2018 for

comparisons over time and geographic space. Additional samples were collected from

middle and lower formations for extended comparison across the two sites (Fig. 1).

Throughout this investigation, sample prefixes correspond to the site and time at which

they were collected (Table S1).

Phylogenetic distribution of microbial communities

We assessed the diversity and structure of the bacterial communities in the

stromatolites at Cape Recife and Schoenmakerskop using 16S rRNA gene amplicon

sequence analysis. All communities were dominated by Cyanobacteria, Bacteroidetes,

Alphaproteobacteria, Gammaproteobacteria and other unclassified bacteria (Fig. 2A),

which was in agreement with a previous study at Schoenmakerskop [14] and other

studies in the Dorrington laboratory (C. Damarajanan, E.W.Isemonger and R.A.

Dorrington, unpublished data). Bacterial communities associated with the upper

formations were distinct from those in middle and lower formations (R>0.6; P=0.01)

(Fig. 2B) consistent with the different environmental conditions across the system.

Binning and phylogenetic classification of putative genomes

To characterize the metabolic potential of individual bacteria, we sequenced a shotgun

metagenomic library from the 8 samples representative of the upper, middle and lower

formations at both sites well as an additional two samples from the upper formations of

each site, collected 4 months previously in January 2018. Binning of assembled

metagenomic libraries using Autometa [22], resulted in a total of 183 bacterial genome

bins from the 10 different metagenomic datasets (Table S1). These bins were manually

curated for optimal purity and completion. Fundamental characteristics of each genome

bin were calculated, and the taxonomic identity was assigned using GTDB-Tk [28]

(Table S2). The GTDB-Tk tool uses the Genome Taxonomy Database as a reference

for classification, which is based on phylogeny inferred from concatenated protein

alignments. This approach enabled the removal of polyphyletic groups and assignment

of taxonomy from evolutionary divergence. The resulting taxonomy incorporates

substantial changes in comparison to the NCBI taxonomy [28]. Equivalent NCBI

taxonomic classifications have been provided for clarity in Table S2. Using relative

coverage per sample as a proxy for abundance we found that genomes classified within

the Cyanobacteriia class were consistently dominant in all collection points, while

Alphaproteobacteria, Gammaproteobacteria and Bacteroidia were less abundant but

notable bacterial classes (Fig. 3). This distribution appears to be approximately

congruent with abundances observed in the 16S rRNA gene amplicon analyses (Fig.

2A).

Temporal and spatial conservation of bacterial species

We calculated pairwise average nucleotide identity (ANI) between all binned genomes

and defined conserved species as genomes sharing more than 97% ANI [23] (Table

S3). Species classified within the family Phormidiaceae were conserved in upper

formations in both January and April samples, as well as within the middle formations

(Table 1). Species within the genera *Acaryochloris* and *Hydrococcus,* and family

Absconditabacterales were conserved across upper formations. Seven species were

conserved across the middle formations, including species classified within genus

Rivularia and Phormidesmiaceae and Spirulinaceae families (Table 1). Species

classified within family Elainellaceae but were distinctly different species wherein

species A within family Elainellaceae was detected only in the upper pools of Cape

Recife, whilst the Elainellaceae species B was conserved across the middle formations

of both sites (Table 1). Bacterial species within genus *Microcoleus* and family

Leptolyngbyaceae were spatially conserved, and present in the January samples of

Schoenmakerskop but were not detected in metagenomes from the January sample of

Cape Recife (Table 1).

Also of interest was the presence of conserved bacterial species within the order

Absconditabacterales, which are classified under the Patescibacteria phylum.

Patescibacteria are unusually small bacteria found in groundwater and produce large

surface proteins hypothesized to help them attach or associate with other

microorganisms that perform nitrogen, sulfur and iron cycling [29]. The presence of

these conserved bacteria suggests that the inflow water seeps originate from

groundwater. There was a lack of conserved bacterial species across the lower

formations and may be due to increased variability at the subtidal interface and

potentially limited sample size. Furthermore, some bacterial genomes defined here may

represent transient bacteria that may not play any role in the stromatolite system.

Metabolic potential of binned genomes

Oxygenic photosynthesis by Cyanobacteria results in rapid fixation of carbon dioxide

and an increase alkalinity [30]. Carbonate ions bind cations such as calcium and are

precipitated under alkaline conditions, promoting the growth of stromatolite structures

[1]. Given their numerical dominance in Cape Recife and Schoenmakerskop

stromatolites, and their predicted role in other stromatolites [1], the role of carbon

cycling is likely performed by Cyanobacteria (Cyanobacteriia class). The identity of the

bacteria that cycle redox sensitive sulfur, phosphate, nitrogen and calcium, and

subsequently affect the alkalinity and solubility index enabling carbonate precipitation in

these stromatolites remain unknown. We inspected PROKKA and KEGG annotations

within stromatolite-associated bacterial genome bins to identify potential metabolic

pathways that may drive the "alkalinity engine" and promote mineral deposition and

accretion [1]. An overview of the results presented here are summarized in Fig. 4 and

Table 1.

It has been shown in experimental models that the uptake of hydrogen ions during

reduction increases the pH and results in the release of carbonate ions in stromatolites

[31]. The increased concentration of carbonate drives an increase in the saturation

index of calcite, resulting in precipitation [31]. Therefore, bacteria that reduce sulfate will

promote the growth of stromatolites, whilst oxidizers will likely drive dissolution of the

calcite precipitate. Amongst the Cape Recife and Schoenmakerskop stromatolites-

associated bacteria, the capacity for sulfate reduction was confined to only a few

genomes (Fig. S1 and S2). The complete set of genes required for assimilatory sulfate

reduction (*sat*/*met3*, *cysC*, *cysH* and *sir* genes) [32] were recovered in four genomes,

three of which were conserved *Acaryochloris* genomes (Fig. S3) and the complete set

of genes for uptake and desulfonation of alkanesulfonates (*ssuABCDE*) [33] were

detected exclusively in conserved *Hydrococus* species (Fig. S3). Alkanesulfonate

metabolism results in the release of sulfite and an aldehyde, the former of which can be

reduced by sulfite reductase (*sir* gene) [34]. Both *Acaryochloris* and *Hydrococcus*

species were detected exclusively in upper stromatolite formations and suggests

freshwater inflow may carry sulfur compounds, however, in the absence of experimental

data this remains purely speculative. Reduction of sulfate has previously been shown to

promote the precipitation of carbonates in the form of micritic crusts in Bahamian and

Australian stromatolites [7, 35] and it has been suggested that microbial cycling of sulfur

played an important role in ancient Australian stromatolites, even prior to the

emergence of Cyanobacteria [36, 37]. The cumulative removal of hydrogen by these

reduction processes would suggest conserved *Hydrococcus* and *Acaryochloris* species

may drive an alkaline pH within the system and potentially aid in calcite accretion. For a

comparative analysis of these stromatolites with other, well-studied systems, a

collection of 96 putative genomes from hypersaline Shark Bay (Australia) stromatolites

[7] was downloaded, annotated and classified as performed for the genomes from the

Cape Recife and Schoenmakerskop stromatolites (Table S4). This dataset was chosen

for comparative analysis as it is the only other set of putative genomes isolated from

stromatolite shotgun metagenomic data, however detailed functional potential analysis

of individual genomes was not carried out previously [7]. Our analysis of the functional

potential of binned genomes from hypersaline Australian stromatolites [7] showed a

greater abundance of genomes capable of both dissimilatory and assimilatory sulfate

reduction (Fig. S4) and the majority of these genomes were classified within the

Desulfobacterota phylum (Table S4). This would suggest that reduction of sulfur

compounds is important to stromatolite growth but the differing environmental pressures

(i.e. fresh vs hypersaline water) have resulted in different bacterial species performing

this conserved function. Sulphur reduction in stromatolites has been previously

attributed to strictly non-cyanobacterial, sulfate-reducing bacteria [38–40] but our

analysis would suggest that cyanobacterial *Hydrococcus* and *Acaryochloris* species

may be the drivers of sulfur reduction in the Cape Recife and Schoenmakerskop

stromatolites.

The reduction of nitrogen, nitrates and nitrites has a similar effect to sulfide, in that the

uptake of protons during the formation of ammonia, results in increased alkalinity that

could lead to calcite precipitation [7, 41–44]. The release of ammonia also aids in

calcification as the ammonia absorbs carbon dioxide, resulting in an increase in

carbonate ions [43]. Metabolism of nitrogenous compounds has likewise been proposed

to have emerged at approximately the same time as sulfur metabolism, as both

processes share similar redox states [45]. Microbial reduction of nitrogen has been

dated back 3.4 billion years [46] and ammonium availability during this time may have

sustained developing microbial life [47]. Therefore, bacteria that can fix nitrogen or

reduce nitrates/nitrites could potentially promote the growth of stromatolites and may

have added to the formation of ancient analogues. We found that several bacteria were

capable of ferredoxin-dependent assimilatory nitrate reduction (ANR) (*nirA-narB* genes)

[48], nitrogen fixation (*nifDHK* genes) and dissimilatory nitrite reduction (DNR)

(*nirBD/nrfAH* genes), all of which result in the production of ammonia [49] (Fig. S5 and

S6). The potential for assimilatory nitrate reduction was detected in several genomes,

primarily from the middle formations, including conserved *Rivularia* sp.,

Phormidesmiaceae, Phormidiaceae and Elainellaceae bacterial species (Fig. S7). The

potential for dissimilatory reduction of nitrites, via either cytoplasmic *nirBD* or

membrane-bound *nrfAH* nitrite reductases, was detected in several genomes across

both the upper and middle formations, including conserved *Hydrococcus* bacterial

species (Fig. S7). Nitrogen fixation during the night has previously been shown to be a

driver of carbonate precipitation in stromatolites [1]. Genes associated with nitrogen

fixation were identified in several bacteria associated with both upper and middle

formations, including conserved *Hydrococcus* species across the upper formations and

conserved Phormidiaceae, Spirulinaceae and Chloroflexaceae species across the

middle formations, as well as non-conserved species within the *Blastochloris* genus

(Fig. S5 - S7). These species may contribute to the growth of the stromatolites through

the formation and release of ammonia. The capacity for nitrogen fixation being retained

in Schoenmakerskop and Cape Recife is unexpected given the high seasonal

concentrations of dissolved inorganic nitrogen ranging from 95-450 μM at the two sites

[15]. Review of the genomes from Australian stromatolites revealed that 14 genomes

were capable of nitrate reduction and 7 were capable of nitrogen fixation. The genomes

were taxonomically diverse, but species within phyla Planctomycetota and

Desulfobacterota were most prominent. As with the reduction of sulfur compounds, it

would appear that the function remains conserved but that different bacteria perform

these functions across diverse conditions of these stromatolites.

Endolithic bacteria bore into carbonate substrate through dissolution facilitated by

uptake, transport and eventual release of calcium ions at the distal end of multicellular

cyanobacterial trichomes (Fig. 3) [50, 51]. Increased concentration of free calcium ions

at the surface of the stromatolite may promote calcium carbonate precipitation [31].

Certain genes are upregulated during this boring activity, including key genes encoding

calcium ATPases and calcium-binding proteins [51, 52]. We identified an incredible 1

182 gene copies encoding calcium-binding proteins and 315 calcium ATPases in the

putative genomes of bacteria from Cape Recife and Schoenmakerskop stromatolites.

The average gene copy number of calcium binding proteins and calcium ATPases was

6.5 and 1.7 respectively. The greatest abundance of gene copies was found in

conserved *Microcoleus*, Phormidiaceae and Elainellaceae genomes (Table 1, Fig. S8),

carrying 20–65 copies of the calcium binding protein gene and 1–7 copies of the

calcium ATPase gene. The potential action of these bacteria on the formation of

stromatolites is two-fold: first, the boring activity results in sand grains "welding" together

and increases the structural stability of the stromatolite [10], and secondly, the release

of concentrated calcium ions at the surface of the stromatolite could lead to rapid

mineralization on contact with either carbonate ions [1, 53] or phosphate ions [54].

Endolithic bacteria have been identified in Bahamian stromatolites (*Solentia* sp.) [10]

and may function in a similar manner to the conserved bacterial species identified here.

Among the Australian stromatolites two genomes (S_098 and S305) (Fig. S4), classified

within Elainellaceae and Phormidesmiales (Table S4), exhibited the genetic potential for

endolithic activity. These genomes carry the 22–26 gene copies of the calcium binding

protein and 3 gene copies (each) of the calcium transport ATPases. The abundances in

the Australian stromatolites were not as high as that observed in bacteria from Cape

Recife and Schoenmakerskop but the potential endolithic function appears to be

conserved within Elainellaceae species, which were conserved in the middle formations

in Cape Recife and Schoenmakerskop. However, the shared ANI of the Australian and

South African Elainellaceae species was between 75–77%, indicating that these

genomes are only distant phylogenetic relatives.

Phosphatic structures recently observed within Cape Recife stromatolites [55] are

particularly intriguing as they closely resemble ancient phosphatic stromatolites from the

Paleoproterozoic and Neoproterozoic eras in which several major oxygenation events

occurred [55]. Hydroxyapatite ($Ca_5(PO_4)_3(OH)$) is more easily precipitated than calcite

($CaCO_3$) and the release of inorganic phosphate into the biofilm by alkaline

phosphatase activity could create a nucleation point for the initiation of apatite

mineralization, the rate of which may be increased in the presence of an alkaline

environment [56]. We have identified several conserved bacterial species here that can

drive the alkalinity of the stromatolite system, as well as conserved species that

concentrate calcium ions at the stromatolite surface (Fig. 3). Apatite could readily be

precipitated if a concentrated source of phosphate ions was provided. We detected a

large number of genes encoding alkaline phosphatases among the stromatolite-

associated bacteria, with genes encoding extracellular PhoX being most abundant (190

copies across all genomes), with fewer copies of genes encoding intracellularly-

localized PhoD (118 copies) and PhoA/B (85 copies) detected [57] (Fig. S9 and S10).

Gene copies of *phoD* and *phoX* have been found to be variable in soil microbes but the

majority of bacteria carry only one copy of *phoD* and one copy of *phoX* [58, 59].

Conserved bacterial species within *Acaryochloris*, *Rivularia* and Elainellaceae groups

were particularly notable (Table 1, Fig. S11), carrying between 2–4 copies of *phoX*.

Several genomes from both sites carried genes encoding more than one type of alkaline

phosphatase, which may indicate that these bacteria can utilize different enzymes,

perhaps in response to a flux in available co-factors (i.e. calcium vs. zinc availability)

[60].

Genes encoding the phosphate-sensing regulatory enzyme (*phoR*) of the canonical

phosphate transporter (*pstABCS*) was absent in most stromatolite-associated bacteria

(Fig. S9 -11) from the Cape Recife and Schoenmakerskop sites. The loss of *phoR* could

potentially lead to toxic, uncontrolled P_i uptake [61] when external Pi levels are high.

However, a recent experiment showed that under limited phosphate conditions, the

growth of *Xanthomonas oryzae* Δ*phoR* deletion mutants was not significantly different

from wild type strains [62] corroborating earlier studies in which *E. coli* Δ*phoR* deletion

mutants had a greater biomass concentration than wild type strains under phosphate

limited conditions [63]. There may be several other genes, or potentially divergent

orthologs of *phoR* in these genomes [64, 65], which would account for the apparent

absence of *phoR*, but the widespread absence of this gene across taxonomically

diverse bacteria at Cape Recife and Schoenmakerskop, suggests that this seemingly

important regulatory protein is not advantageous to this system. It is hypothesized that

stromatolites of Cape Recife and Schoenmakerskop largely receive phosphorous from

the oceans and have consistently been shown to experience limited inorganic

phosphate availability [15, 66–68] and this consistent struggle for phosphate may have

resulted in the loss of *phoR* as its role may have become redundant. Loss of PhoR has

also resulted in increased alkaline phosphatase activity [69] and consequently, it is

possible that the combined effect of extracellular alkaline phosphatase abundance and

loss of *phoR*, could result in a relative excess of local bioavailable phosphate in the

extracellular stromatolite environment. Cyanobacteria in phosphorus-poor rivers have

been hypothesized to generate bioavailable phosphate from trapped sediments in

biofilms generating bioavailable phosphate concentrations 320-fold higher than the

surrounding water [70]. An increase in both phosphate and calcium ions could result in

the rapid precipitation of apatite, as observed in marine phosphorites, freshwater lakes

and in some soil bacteria [54, 71, 72] and could account for the phosphatic deposits

observed within the SA stromatolites [55].

Finally, conserved species classified within the genus *Rivularia* and several other non-

conserved bacterial species harbored the 11 genes required for the transport

(*phnCDEF*) and lysis (*phnGHIJKLM*) of phosphonate compounds (Fig. S9, S10 and

S12) [73]. Phosphonates are characterized by the presence of a carbon-phosphorous

bond and are biosynthesized by several organisms [74]. The C-P lyase (*phnGHIJLKM*)

breaks the C-P bond within a wide array of phosphonate substrates resulting in a

hydrocarbon and inorganic phosphate [74]. The inorganic phosphate may be used

within the bacterial cell or released aiding in accretion through increased ion

concentration [75]. However, phosphonates within an environment can prevent

precipitation of calcite by binding to crystal growth sites [76] and the degradation of

these compounds within the stromatolite formations may prevent chemical inhibition of

stromatolite growth. The potential for phosphonate degradation was also detected in 8

genomes from the Australian stromatolites (Fig. S4), including the same species

potentially capable of endolithic activity (S_098: Elainellaceae and S_305:

Phormidesmiales). The conservation of phosphonate metabolism and potential

endoliths as common features across stromatolites would indicate that these bacterial

processes are vital within the stromatolite system.

It would appear that several bacterial groups are key players in the formation of

stromatolites in Cape Recife and Schoenmakerskop, with conserved bacterial species

acting as major contributors (Table 1). We propose the redox potential and solubility

index in SA stromatolites is potentially controlled by conserved bacteria, including

conserved *Hydrococcus*, *Acaryochloris* and Phormidiaceae bacterial species in upper

formations, and Elainellaceae, Phormidiaceae, Phormdesmiaceae, *Rivularia* and

Spirulinaceae bacterial species in middle formations (Table 1), through the generation

and concentration of sulfide, ammonia and calcium ions for the rapid precipitation of

carbonate compounds and subsequent growth of stromatolite structures. We further

propose an abundance of alkaline phosphatases in combination with a suspected loss

of transport regulatory protein PhoR, may result in increased local inorganic phosphate

concentrations (Fig. 4). This free inorganic phosphate may bind with the calcium

released by boring endolithic bacteria to form the phosphatic crusts previously observed

in these stromatolites. Conservation of these functions, and in some cases taxonomic

groups, across these and Australian stromatolites would suggest these processes may

be especially essential functions within the system. The identification of conserved

bacteria performing potentially vital roles within stromatolites of Cape Recife and

Schoenmakerskop may provide insight into what cyanobacterial species may have

played a key role in the formation of ancient phosphatic stromatolites.

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CONFLICT OF INTERESTS

The authors declare no financial competing interests.

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718.

718. **Table 1.** Conservation of bacterial species (ANI > 97%) and their functional potential across upper, middle and lower stromatolite formations at Cape Recife and Schoenmakerskop.

Conserved bacteria taxonomy (ANI >97%)	Upper (Jan) CR	Upper (Jan) SCH	Upper (April) CR	Upper (April) SCH	Middle (April) CR	Middle (April) SCH	Lower (April) CR	Lower (April) SCH	Assim S reduction	Alkanesulfonate metabolism	Dissim N reduction	Assim N reduction	N fixation	Endolithic activity	Alkaline phosphatases	Phosphonate degradation
Phormidiaceae	X	X	X	X	X	X						X	X	X		
Absconditabacterales	X	X	X	X												
Acaryochloris	X	X	X	X					X						X	
Hydrococcus	X	X	X	X						X	X		X			
Leptolyngbyaceae		X	X	X												
Microcoleus		X	X	X		X								X		
Pleurocapsa	X	X		X												
Elainellaceae sp. A	X		X													
Scytonema	X		X													
Brachymonas	X		X													
Synechococcales	X		X													
Chloroflexaceae						X		X					X			
Elainellaceae sp. B					X	X		X				X		X	X	
JA-3-3Ab					X	X										
Bacteroidales					X	X										
Phormidesmiaceae					X	X		X				X				
Rivularia		X			X	X		X				X			X	X
Spirulinaceae					X	X							X			
UBA4181					X	X										
Calothrix				X		X										

Figure 1. Stromatolites were collected from four different points at two different sites

along the SA coastline. (A) Aerial view of sampling locations within the

Schoenmakerskop site and (B) sampling locations in the Cape Recife site. Samples

were collected from upper (CSU: Schoenmakerskop Inflow Jan, SU: Schoenmakerskop

Inflow April, CRU: Cape Recife Inflow Jan, RU: Cape Recife Inflow April), middle (SM1:

Schoenmakerskop Middle 1 April, SM2: Schoenmakerskop Middle 2 April, RM1: Cape

Recife Middle 1 April, RM2: Cape Recife Middle 1 April) and lower formations (SL:

Schoenmakerskop Lower April, RL: Cape Recife Lower April).

Figure 2. Distribution and abundance of bacterial taxa in stromatolite formations. (A)

Phylogenetic classification and average relative abundance (n=3) of dominant phyla in

different sample sites indicated that all stromatolite samples are dominated by

Cyanobacteria, Bacteroidetes, Alpha- and Gammaproteobacteria. (B) OTU abundance

was used to cluster stromatolite biological replicates using Bray-Curtis non-dimensional

scaling. Samples were isolated from upper (green), middle (red) and lower (blue)

stromatolites formations in Cape Recife (triangles) and Schoenmakerskop (circles).

Figure 3. Taxonomic classification of putative genome bins in stromatolites collected

from upper/inflow, middle and lower/marine formations of Schoenmakerskop and Cape

Recife. Coverage per genome has been used as a proxy for abundance and summed

for genomes within each respective taxonomic class. A key for the predicted taxonomic

classification has been provided. Abrreviations are as follows: CSU: Schoenmakerskop

Inflow Jan, SU: Schoenmakerskop Inflow April, CRU: Cape Recife Inflow Jan, RU: Cape

Recife Inflow April, SM1: Schoenmakerskop Middle 1 April, SM2: Schoenmakerskop

Middle 2 April, RM1: Cape Recife Middle 1 April, RM2: Cape Recife Middle 1 April, SL:

Schoenmakerskop Lower April, RL: Cape Recife Lower April.

Figure 4. A summarized overview of the bacterial metabolisms predicted in

stromatolites from Cape Recife and Schoenmakerskop that may influence mineral

precipitation and subsequent growth of stromatolite formations.

Figure S1. Abundance of sulfate metabolism genes in bacterial putative genomes from

bacteria associated with stromatolites in Schoenmakerskop. Gene abundance per

genome was counted from KEGG annotations detected using kofamscan [27] and are

indicated with the scale provided. Genes are identified by both their assigned KO

number and accepted abbreviation.

Figure S2. Abundance of sulfate metabolism genes in bacterial putative genomes from

bacteria associated with stromatolites in Cape Recife. Gene abundance per genome

was counted from KEGG annotations detected using kofamscan [27] and are indicated

with the scale provided. Genes are identified by both their assigned KO number and

accepted abbreviation.

Figure S3. Summary of genes encoding sulfate-reducing enzymes in stromatolite-

associated bacterial genomes. Gene abundance is indicated per the scale provided.

Genomes that do not carry any of the genes investigated here were not shown.

Phylogeny was inferred using JolyTree [24] with a sketch size of 5 000. Conserved

genomes are indicated with an asterisk within highlighted taxonomic groups. Genome

abbreviations indicate site, formation proximity to ocean and collection; CSU:

Schoenmakerskop Inflow Jan, SU: Schoenmakerskop Inflow April, CRU: Cape Recife

Inflow Jan, RU: Cape Recife Inflow April, SM1: Schoenmakerskop Middle 1,

SM2:Schoenmakerskop Middle 2, RM1: Cape Recife Middle 1, RM2: Cape Recife

Middle 1, SL: Schoenmakerskop Lower, RL: Cape Recife Lower.

Figure S4. Summarized nutrient metabolisms in genomes bins from stromatolites

sampled from Shark Bay, Australia [7]. Gene abundance per functional group per

genome was counted from KEGG annotations detected using kofamscan [27] and are

indicated with the scale provided. Genes are identified by both their assigned KO

number and accepted abbreviation.

Figure S5. Abundance of nitrogen metabolism genes in bacterial putative genomes

from bacteria associated with stromatolites in Schoenmakerskop. Gene abundance per

genome was counted from KEGG annotations detected using kofamscan [27] and are

indicated with the scale provided. Genes are identified by both their assigned KO

number and accepted abbreviation.

Figure S6. Abundance of nitrogen metabolism genes in bacterial putative genomes

from bacteria associated with stromatolites in Cape Recife. Gene abundance per

genome was counted from KEGG annotations detected using kofamscan [27] and are

indicated with the scale provided. Genes are identified by both their assigned KO

number and accepted abbreviation.

Figure S7. Summary of genes encoding nitrate/nitrite-reducing and nitrogen fixation

enzymes in stromatolite-associated bacterial genomes. Gene abundance is indicated

per the scale provided. Genomes that do not carry any of the genes investigated here

were not shown. Phylogeny was inferred using JolyTree [24] with a sketch size of 5 000.

Conserved genomes are indicated with an asterisk within highlighted taxonomic groups.

Genome abbreviations indicate site, formation proximity to ocean and collection; CSU:

Schoenmakerskop Inflow Jan, SU: Schoenmakerskop Inflow April, CRU: Cape Recife

Inflow Jan, RU: Cape Recife Inflow April, SM1: Schoenmakerskop Middle 1,

SM2:Schoenmakerskop Middle 2, RM1: Cape Recife Middle 1, RM2: Cape Recife

Middle 1, SL: Schoenmakerskop Lower, RL: Cape Recife Lower.

Figure S8. Summary of genes encoding calcium binding and transport enzymes in

stromatolite-associated bacterial genomes. Gene abundance is indicated per the scale

provided. Genomes carrying fewer than 6 genes (cumulative) in these two categories

were collapsed. Phylogeny was inferred using JolyTree [24] with a sketch size of 5 000.

Conserved genomes are indicated with an asterisk within highlighted taxonomic groups.

Genome abbreviations indicate site, formation proximity to ocean and collection; CSU:

Schoenmakerskop Inflow Jan, SU: Schoenmakerskop Inflow April, CRU: Cape Recife

Inflow Jan, RU: Cape Recife Inflow April, SM1: Schoenmakerskop Middle 1,

SM2:Schoenmakerskop Middle 2, RM1: Cape Recife Middle 1, RM2: Cape Recife

Middle 1, SL: Schoenmakerskop Lower, RL: Cape Recife Lower.

Figure S9. Abundance of phosphate metabolism genes in bacterial putative genomes

from bacteria associated with stromatolites in Schoenmakerskop. Gene abundance per

genome was counted from KEGG annotations detected using kofamscan [27] and are

indicated with the scale provided. Genes are identified by both their assigned KO

number and accepted abbreviation.

Figure S10. Abundance of phosphate metabolism genes in bacterial putative genomes

from bacteria associated with stromatolites in Cape Recife. Gene abundance per

genome was counted from KEGG annotations detected using kofamscan [27] and are

indicated with the scale provided. Genes are identified by both their assigned KO

number and accepted abbreviation.

Figure S11. Summary of genes encoding alkaline phosphatases and phosphate

transport regulatory proteins in stromatolite-associated bacterial genomes. Gene

abundance is indicated per the scale provided. Genomes that do not carry any of the

genes investigated here were not shown. Phylogeny was inferred using JolyTree [24]

with a sketch size of 5 000. Conserved genomes are indicated with an asterisk within

highlighted taxonomic groups. Genome abbreviations indicate site, formation proximity

to ocean and collection; CSU: Schoenmakerskop Inflow Jan, SU: Schoenmakerskop

Inflow April, CRU: Cape Recife Inflow Jan, RU: Cape Recife Inflow April, SM1:

Schoenmakerskop Middle 1, SM2:Schoenmakerskop Middle 2, RM1: Cape Recife

Middle 1, RM2: Cape Recife Middle 1, SL: Schoenmakerskop Lower, RL: Cape Recife

Lower.

Figure S12. Summary of genes encoding phosphonate degradation enzymes in

stromatolite-associated bacterial genomes. Gene abundance is indicated per the scale

provided. Genomes that do not carry any of the genes investigated here were not

shown. Phylogeny was inferred using JolyTree [24] with a sketch size of 5 000.

Conserved genomes are indicated with an asterisk within highlighted taxonomic groups.

Genome abbreviations indicate site, formation proximity to ocean and collection; CSU:

Schoenmakerskop Inflow Jan, SU: Schoenmakerskop Inflow April, CRU: Cape Recife

Inflow Jan, RU: Cape Recife Inflow April, SM1: Schoenmakerskop Middle 1,

SM2:Schoenmakerskop Middle 2, RM1: Cape Recife Middle 1, RM2: Cape Recife

Middle 1, SL: Schoenmakerskop Lower, RL: Cape Recife Lower.

Table S1. Summary of genomes binned per collected stromatolite sample

Table S2. Summary of characteristics of genomes binned from shotgun metagenomic

data from sampled upper, middle and lower stromatolite formations from Cape Recife

and Schoenmakerskop.

Table S3. Conserved bacterial species defined by shared ANI > 97% in stromatolite

formations from Cape Recife and Schoenmakerskop.

Table S4. Taxonomic classification of genomes from Shark Bay stromatolites

Accretion

SH ⬤   NH₃ ⬤

pH

Dissolution   SO₄   N₂/NO₃/NO₂

SO₄

phoR phoA
phoX phoD

CO₂

Photosynthesis

Precipitate

Legend:
- ⬤ SH
- ⬤ NH₃
- ⬤ Ca²⁺
- ⬤ Phosphonates
- ⬤ Hydrocarbon
- ⬤ PO₄³⁻
- ⬤ CO₃²⁻
- Ca₃(PO₄)₂
- CaCO₃

1. Alkalinity is increased in response to reduction of sulphur and nitrogenous compounds

2. Phosphate released by alkaline phosphatses and possibly phosphonate metabolism, increases free phosphate ions. Phosphonates may prevent accretion.

3. Calcium ions are released from endolithic boring activity

4. Carbon dioxide uptake from photosynthesis increases alkalinity

5. Increased concentrations of calcium, carbonate and phosphate ions leads to precipitation of calcium carbonate and calcium phosphate