Intracellular amorphous calcium carbonate biomineralization in methanotrophic

gammaproteobacteria was acquired by horizontal gene transfer from cyanobacteria

Running title: ACC biomineralization in aerobic methanotrophs

Benzerara Karim[1], Millet Maxime[1], Skouri-Panet Fériel[1], Gaschignard Geoffroy[1], Mehta

Neha[1], Bezard Malo[1], Caumes Geraldine[1], Chevrier Daniel M.[2], Dezi Manuela[1], Duverger

Arnaud[1], Guigner Jean-Michel[1], Gutiérrez-Preciado Ana[3], Lefevre Christopher T.[2], López-

García Purificación[3], Menguy Nicolas[1], Monteil Caroline L.[2], Pehau-Arnaudet Gerard[4], Penard

Esthel[4], Pereiro Eva[5], Scandola Cyril[4], Travert Cynthia[1], Vantelon Delphine[6], Duprat Elodie[1],

Callebaut Isabelle[1], Moreira David[3]

[1]Sorbonne Université, Muséum National d'Histoire Naturelle, UMR CNRS 7590. Institut de

Minéralogie, de Physique des Matériaux et de Cosmochimie (IMPMC), 75005 Paris, France.

[2]Aix-Marseille Université, CEA, CNRS, BIAM, UMR7265, Institut de Biosciences et

Biotechnologies Aix-Marseille, CEA Cadarache, F-13115, Saint-Paul-lez-Durance, France

[3]Unité d'Ecologie Société et Evolution, CNRS, Université Paris-Saclay, AgroParisTech, Gif-

sur-Yvette, France

[4]Institut Pasteur, Université Paris Cité, Ultrastructural Bioimaging Unit, 75015 Paris, France

[5]ALBA Synchrotron Light Source, Cerdanyola del Vallés, Barcelona 08290, Spain

[6]Synchrotron SOLEIL, L'Orme des Merisiers, 91190 Saint-Aubin, France

For correspondence: karim.benzerara@sorbonne-universite.fr

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Abstract

Some bacteria genetically control the biomineralization of intracellular amorphous calcium carbonates (iACC) with potential implications for microbial physiology, evolution, bioremediation and biogeochemical cycling. Until now, this capacity has been documented in *Cyanobacteria*, the giant gammaproteobacterium *Achromatium* and a few magnetotactic *Pseudomonadota*. Here, we report the discovery of iACC biomineralization in members of the *Methylococcaceae*, a family of aerobic methanotrophic *Gammaproteobacteria*. A homolog of the *ccyA* gene, previously considered a diagnostic marker for iACC formation in *Cyanobacteria*, was identified in several *Methylococcaceae* genomes. Moreover, two cultivated strains, *Methylococcus geothermalis* and *Methylococcus mesophilus*, were consistently shown to form iACC granules. The *ccyA* genes in *Methylococcaceae* and *Microcystis* share stronger sequence and structural similarities than with other *Cyanobacteria*, suggesting horizontal gene transfer (HGT) from an ancestral *Microcystis*-like cyanobacterium to *Methylococcaceae*. This finding extends the known taxonomic distribution of *ccyA* and indicates that the capability to biomineralize iACC was acquired by HGT, possibly in environments such as lakes, where *Cyanobacteria* and *Methylococcaceae* commonly co-exist. The discovery of iACC in methane-oxidizing *Methylococcaceae* highlights a previously unrecognized coupling between calcium carbonate biomineralization and methane cycling in aquatic environments, suggesting that iACC formation may play an overlooked role in microbial carbon storage and local geochemical regulation.

1. Introduction

Many bacteria are capable of forming a wide range of mineral phases through a process known as biomineralization. This process may serve specific biological functions such as protection, detoxification, assistance in cell movements or storage (Cosmidis and Benzerara, 2022) and has

a significant impact on both local and global geochemical cycles, inspiring a variety of applications (Qin et al., 2020; Wan et al., 2024; Zhang et al., 2023). Among biomineralizing bacteria, several phylogenetically diverse species have been shown to form intracellular amorphous calcium carbonates (iACC). The giant gammaproteobacterium *Achromatium oxaliferum* was the first such organism discovered, as early as the late 19th century (Schewiakoff, 1893). For a long time, it remained the only known iACC-biomineralizing bacterium (Benzerara et al., 2021; Gray, 2006). However, since 2012, diverse cyanobacteria have also been found to form iACC (Benzerara et al., 2014, 2022; Couradeau et al., 2012). More recently, additional members of the *Pseudomonadota* and *Nitrospirota*, including magnetotactic bacteria capable of biomineralizing intracellular magnetite crystals (Fe(II)-Fe(III) oxides), have been shown to form iACC as well (Liu et al., 2021, 2025; Mangin et al., 2025; Monteil et al., 2020). This rapidly expanding diversity of iACC-forming bacteria suggests that many more such organisms likely remain to be discovered.

Forming iACC granules incurs an energy cost to cells, at least under certain conditions (Cam et al., 2017), suggesting that these granules likely confer adaptive benefits to the cells. However, their exact biological role remains poorly understood (Cosmidis and Benzerara, 2022). Some studies have proposed that iACC granules may buffer intracellular pH fluctuations, serve as carbon reservoirs, or act as ballast for the cells (Görgen et al., 2020). Regardless of their specific function, iACC granules involve a particular Ca homeostasis in the cells that form them, including Ca exchange between the cells and the surrounding solution with a net inward influx (Benzerara et al., 2023; Mehta et al., 2023, 2024). Studying these granules offers valuable insights: (i) Amorphous calcium carbonate (ACC) is essential to many $CaCO_3$-biomineralizing eukaryotes (Gilbert et al., 2022), and the ability to form iACC may have originated early in evolutionary history within prokaryotes before being transferred to eukaryotes (Ponce-Toledo et al., 2017). Testing this hypothesis requires further exploration of the evolutionary links

between iACC formation across different bacterial lineages. (ii) iACC-forming bacteria can

influence the local geochemical cycling of alkaline earth metals (Ca, Sr, Ba), particularly when

they bloom as large populations (Blondeau et al., 2018b; Gaetan et al., 2023). Last, (iii) some

of these bacteria are capable of sequestering toxic radioisotopes, opening new avenues for the

development of bioremediation strategies (Mehta et al., 2019, 2022a; Pamart et al., 2025).

Despite the significance of this biomineralization process, the molecular mechanisms

underlying bacterial iACC formation remain largely unexplored. In *Cyanobacteria*, iACC

forms within envelopes thought to consist of proteins and/or lipid monolayers (Blondeau et al.,

2018a). Through comparative genomics, a novel gene family named *ccyA* was identified

(Benzerara et al., 2022). This gene family lacks homologs with a known function but serves as

a diagnostic marker for iACC formation in *Cyanobacteria*. The *ccyA* gene encodes two-domain

proteins, the calcyanins, with a C-terminal domain conserved across cyanobacterial clades,

composed of several long glycine zipper (Gly-zip) motifs, and a variable N-terminal domain

that defines four distinct calcyanin subgroups, corresponding to different *Cyanobacteria* clades

(Benzerara et al., 2022). Interestingly, the N-terminal domain of one calcyanin subgroup

contains a CoBaHMA domain, which is hypothesized to play a regulatory role, possibly

interacting with anionic lipids (Gaschignard et al., 2024). Furthermore, a recent study suggests

that other genes may operate in concert with *ccyA*, such as a *cax* gene encoding a Ca^{2+}/H^+

transporter (Bruley et al., 2025).

In this study, we report the discovery of a homolog of the *ccyA* gene in *Methylococcaceae*, a

family of *Gammaproteobacteria*. Furthermore, two cultivated strains from this family were

found to form iACC, providing evidence that *ccyA* is a marker of iACC biomineralization

beyond the *Cyanobacteria* phylum. The phylogenetic distribution of *ccyA* within

Methylococcales, along with its comparison with the *ccyA* genes in *Cyanobacteria*, offer new

insights into the evolutionary origins of iACC biomineralization in *Methylococcaceae*.

2.   Materials and methods

2.1 Bacterial strains

The strain *Methylococcus geothermalis* IM1 (=JCM 33941 =KCTC 72677) was ordered from

the Riken bioresource research center in Japan (https://www.jcm.riken.jp/cgi-

[bin/jcm/jcm_number?JCM=33941](https://www.jcm.riken.jp/cgi-bin/jcm/jcm_number?JCM=33941). *M. geothermalis* is an aerobic, planktonic and non-motile,

methanotroph isolated from a hot spring soil in Korea (Awala et al., 2020). Here, it was cultured

in triplicates at 37 °C in 10 mL of modified nitrate mineral salts medium-2 (M1221;

https://www.jcm.riken.jp/cgi-bin/jcm/jcm_grmd?GRMD=1221) in a 120-mL glass serum

bottle with a 25-75 vol% of CH_4-air mix in the headspace. The genome sequence is available

under NCBI accession number CP046565.1.

The strain *Methylococcus mesophilus* 16-5^T (=KCTC 82050^T) was provided by Dr. Samuel

Imisi Awala from University of Calgary. *M. mesophilus* is a mesophilic, aerobic methanotroph

isolated from a rice field in Korea (Awala et al., 2023). It was cultured in triplicates at 21°C in

10 mL of modified nitrate mineral salts medium-2 (M1221; https://www.jcm.riken.jp/cgi-

[bin/jcm/jcm_grmd?GRMD=1221](https://www.jcm.riken.jp/cgi-bin/jcm/jcm_grmd?GRMD=1221)) in a 120-mL glass serum bottle with a 25-75 vol% of CH_4-

air mix in the headspace. The genome sequence is available under NCBI accession number

CP110921.1.

Cell growth for both strains was assessed by measuring optical density at 600 nm. The pH of

the cell suspension was measured over the growth of the cells with a precision better than 0.1.

2.2 Fourier Transform infrared spectroscopy (FTIR)

FTIR spectroscopy analyses of both strains were performed to detect bulk signals from iACC inclusions, following the methodology outlined by (Mehta et al., 2022b). Cells were pelleted from 15 mL of a culture grown to an optical density of 0.2, washed with MilliQ® water, dried and ground in an agate mortar. Attenuated total reflectance Fourier transform infrared (ATR-FTIR) spectra were acquired in the mid-IR region (4000–600 cm^{-1}) using a NICOLET 6700 FTIR spectrometer equipped with a diamond internal reflection element (IRE), a KBr beam splitter and a deuterated triglycine sulfate (DTGS) detector. Spectra were collected by averaging 64 scans at resolution of 4 cm^{-1}.

2.3 Synchrotron-based X-ray absorption near edge structure (XANES) spectroscopy

The relative quantification of Ca in iACC within *M. geothermalis* cells was achieved using XANES spectroscopy, following the approach detailed by Mehta et al. (2023). For this, 15 mL of a cell suspension were harvested by centrifugation, washed, and dried at 37 °C for 48 hours. The dried biomass (~5 mg) was then diluted in 35 mg cellulose, pressed into a 10 mm diameter pellet, and mounted onto Ca-free holders. XAS at the Ca K-edge was performed on the LUCIA beamline at the SOLEIL synchrotron (Saint-Aubin, France) (Flank et al., 2006; Vantelon et al., 2016). The fixed exit double-crystal monochromator was equipped with Si (111) crystals. Spectra were recorded in the fluorescence mode using a 60 mm^2 mono-element silicon drift diode detector (Bruker). In the pre-edge region (3940–4030 eV), spectra were recorded with a step size of 2 eV and a counting time of 1 s. In the edge region (4030–4070 eV), spectra were recorded with a step size of 0.25 eV and a counting time of 1 s. The sample holder was positioned at a 24° angle relative to the fluorescence detector to minimize self-absorption effects and was maintained at a temperature of 77 K using a liquid He-cryostat to minimize potential beam damages and localized heating. Energy calibration was performed using a calcite reference, with the first inflection point set to 4045 eV.

2.4 Scanning electron microscopy (SEM)

Cell suspensions of both *M. geothermalis* and *M. mesophilus* were filtered onto 0.22 μm Isopore

polycarbonate filters and rinsed with 10 mL of MilliQ® water. The filters were air-dried at

ambient temperature and mounted on aluminum stubs using double-sided carbon tape. Before

analysis, the filters were coated with a thin layer of carbon via evaporation. Scanning electron

microscopy (SEM) analyses were performed using a Zeiss™Ultra55 SEM with an electron

accelerating voltage of 10 kV, a working distance of 7.5 mm and a 60 μm aperture. Images

were captured using a backscattered electron detector (AsB detector). Energy-dispersive x-ray

spectrometry (EDXS) data were acquired with a Bruker QUANTAX EDS detector, enabling

hyperspectral imaging (HyperMap), and the data were analysed with the Esprit software

(Brucker).

2.5 Transmission electron microscopy

A 0.5 mL cell suspension was harvested by centrifugation, rinsed twice in 1 mL of MilliQ®

water before resuspension in 0.5 mL of MilliQ® water. A drop of 3 μL was deposited onto a

formvar-coated TEM grid and air-dried. Observations were carried out using a Jeol 2100F TEM

microscope operating at 200 kV, equipped with a Schottky emitter, a STEM device for Z-

contrast imaging in high-angular annular dark-field (HAADF) mode, and a Jeol Si(Li) X-ray

detector for elemental analyses via energy dispersive x-ray spectrometry (EDXS).

2.6 Synchrotron-based cryo-soft X-ray tomography (c-SXT)

Three-dimensional X-ray imaging of *M. geothermalis* cells was performed at the ALBA

synchrotron (Barcelona, Spain) using cryo-transmission X-ray microscopy at the MISTRAL

beamline (Sorrentino et al., 2015). The cryofixation process involved plunge-freezing at the

Institut Pasteur (Paris, France) using a Leica EM GP. A Quantifoil TEM grid (Q-R2/1-C2100)

was loaded in the chamber of the plunge freezer. Four microliters of a two-day-old *M. geothermalis* culture (OD = 0.251) and one microliter of a 100-nm gold nanoparticles suspension (BBI Solutions, 5X concentrated) were deposited on the front of the grid. After blotting for four seconds from the back, grids were plunged into a container filled with liquid ethane cooled with liquid nitrogen (-196°C), ensuring proper vitrification of the samples. Subsequently, the samples were shipped from Paris to the ALBA synchrotron in a dry shipper and transferred to the cryo-chamber of the Mistral beamline for imaging.

A tilt series of projections was collected from −70° to +70° in 1° increments (except otherwise mentioned), with an incident X-ray energy of 520 eV. Each projection was exposed for 3 seconds. A 40-nm Fresnel zone plate was used, achieving an effective pixel size of 12 nm. The projections were normalized for incoming flux and deconvolved using the measured point spread function of the optical system (Otón et al., 2016). Projections were aligned using Etomo with the 100-nm gold nanoparticle fiducials. Tomographic reconstruction, along with simultaneous iterations reconstruction technique (SIRT) deconvolution, were performed using IMOD (Otón et al., 2016). Volume segmentation, calculations and visualization of tomograms were carried out using the Microscopy Image Browser and ImageJ (Belevich et al., 2016).

2.7 Genome collection and phylogenomic analyses

The 858 high-quality genome sequences of species of the order *Methyloccoccales* available in the Genome Taxonomy Database (GTDB) version r220 (Parks et al., 2022) were downloaded. They included seven families according to the GTDB taxonomy: *Cycloclasticaceae* (43 genomes), *Methylococcaceae* (112 genomes), *Methylomonadaceae* (679 genomes), *Methylothermaceae* (13 genomes), UBA1147 (6 genomes), UBA2778 (3 genomes), and DRLZ01 (2 genomes). Eight additional gammaproteobacterial genomes were downloaded to serve as outgroup in phylogenomic analyses: *Coxiella burnetii* cb109 (GCA_000300315.1),

Legionella pneumophila Hextuple_2q (GCA_000277025.1), *Reinekea blandensis* MED297 (GCF_000153185.1), *Dichelobacter nodosus* (GCF_001263335), *Francisella tularensis* FSC147 (GCA_000018925.1), *Nitrosococcus oceani* ATCC 19707 (GCF_000012805.1), *Halorhodospira* sp. (GCA_007121755.1), and *Alkalilimnicola ehrlichii* MLHE-1 (GCF_000014785.1). Protein-coding genes were predicted with Prodigal (Hyatt et al., 2010) when annotations were missing.

GTDB-Tk (Chaumeil et al., 2022) was used to identify 120 conserved protein markers encoded in the whole collection of *Methylococcales* and outgroup genome sequences. Individual markers were aligned using MAFFT L-INS-I v7.310 (Katoh et al., 2019). Poorly aligned sites were removed using trimAl (Capella-Gutiérrez et al., 2009) with the –strictplus option and the resulting trimmed alignments were concatenated using an in-house Python script, yielding an alignment of 36,916 amino acid sites. A Maximum Likelihood tree based on this supermatrix was inferred using IQ-TREE v3.0.1 (Wong et al., 2025) with the LG+G substitution model (Le and Gascuel, 2008). Branch statistical support was calculated with 1,000 ultrafast bootstrap replicates (Hoang et al., 2018). A tree including only the *Methylococcales* species containing the *ccyA* gene was constructed using the same concatenated markers and the same method described above.

2.8 Detection of *ccyA*

The detection of *ccyA* was conducted using the in-house tool pCALF (standing for python CALcyanin Finder, https://github.com/K2SOHIGH/pcalf/tree/main), which uses four hidden Markov model (HMM) profiles describing the C-terminal glycine zipper triplication (called the $(GlyZip)_3$ motif), specific of calcyanins in *Cyanobacteria* as described in detail by Gaschignard et al. (2024).

2.9 Sequence analysis

Calcyanin protein sequences detected in *Methylococcales* were compared with those found in *Cyanobacteria*. For this purpose, multiple sequence alignments were performed using MAFFT L-INS-I v7.310 (Katoh et al., 2019), then checked and further analyzed using Hydrophobic Cluster Analysis (HCA) (Bitard-Feildel et al., 2018). Aligned sequences were rendered on the basis of ESPript 3.0 results (Gouet et al., 2003). 3D structure predictions were performed using the AlphaFold3 server (Abramson et al., 2024).

3. **Results**

3.1 Detection and phylogenetic distribution of *ccyA* in the order *Methylococcales*

The *ccyA* gene was identified as a diagnostic marker of iACC biomineralization in *Cyanobacteria*. Moreover, homologs of *ccyA* were also reported in four gammaproteobacteria of the order *Methylococcales* (Benzerara et al., 2022). To investigate this further, we systematically searched for *ccyA* homologs across the 858 genome sequences of *Methylococcales* species available in the Genome Taxonomy Database (GTDB) version r220 (Parks et al., 2022) (Table S1).

We detected *ccyA* in 39 genomes of species belonging to this order, distributed in the genera *Methylococcus* (including *M. geothermalis* and *M. mesophilus*), *Methylospira* (*Candidatus* Methylospira mobilis), and *Methylumidiphilus* (many strains affiliated with *M. alinensis*) (Table 1). According to the GTDB taxonomy, all these species were assigned to the family *Methylococcaceae*; none were associated with the species-rich *Methylomonadaceae* or with the other families of the order *Methylococcales* (Figure 1 and Figure S1). The species possessing *ccyA* occur in diverse environments, including the water column and sediment of freshwater lakes, compost, active municipal landfills, and rice field soils (Table 1 and Table S2). In all identified genomes, the protein calcyanin was consistently annotated as a hypothetical protein.

According to the GTDB metadata, most of the genomes containing *ccyA* (34 out of 39) were metagenome-assembled genomes (MAGs), and only five originated from cultured isolates (Table 1). Only two of these isolates were available: *Methylococcus geothermalis* IM1 and *Methylococcus mesophilus* 16-5, which were used in our subsequent studies.

Figure 1: Maximum likelihood phylogenomic tree of 858 *Methylobacterales* taxa available in the Genome Taxonomy Database (GTDB version 220) (see Table S1). The tree is based on the concatenation of 120 conserved proteins. Taxa containing the *ccyA* gene are highlighted in green. Taxa that do not contain the *ccyA* gene are collapsed. Taxon names correspond to the GTDB taxonomy. The green arrow indicates the most likely point for the acquisition of *ccyA* by an ancestor of several *Methylococcaceae* lineages. Ultrafast bootstrap support (UFBS) values >90% are indicated on branches (black dots: 99-100%; gray dots: 90-99%). The complete, un-collapsed tree is available in Fig. S1.

Genome ID	Environmental biome	Geographic location	GTDB genus	Type	Strain availability	Bibliography
GCF_012769535.1	Geothermal environment	South Korea	*Methylococcus*	Isolate	Available axenic culture, *M. geothermalis*	Awala SI et al (2020)
GCF_016106025.1	Compost from a depth of 5 cm	Russia: Moscow region	*Methylococcus*	Isolate	No response from the person who isolated the strain	Patent RU2020129522A
GCF_026247885.1	Rice field soil; 10-15 cm deep	South Korea	*Methylococcus*	Isolate	Available axenic culture, *M. mesophilus*	Awala SI et al (2023)
GCF_029478315.1	Mixture of strains (e.g., ATCC 33009; BKM 2116) from several laboratories followed by selection under specific conditions	missing	*Methylococcus*	Isolate	*M. capsulatus* , No response from the person who isolated the strain	Patent RU2613365C1
GCF_009498235.1	Acidic Sphagnum-dominated peat bog	Russia: Tver region	*Methylospira*	Isolate	*Candidatus* Methylospira mobilis, strain was lost (pers. Comm. S. Dedysh)	Danilova OV et al. (2016).
GCA_028688525.1	landfill. NEUS_C leachate from a leachate well at an active municipal landfill in the North Eastern United States	USA	*Methylococcus*	MAG	NA	Grégoire et al. (2023). M
GCA_003242955.1	Stratified humic lake	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Rissanen AJ, et al. (2018).
GCA_903912685.1	Lake sediment	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903831675.1	Lake sediment	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903831425.1	Lake sediment	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903929635.1	Lake sediment	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903882955.1	Lake water	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903848025.1	Lake water	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903896155.1	Lake water, Alinen Mustajärvi, 3m in depth	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903921065.1	Lake water, Alinen Mustajärvi, 4.4m in depth	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903908275.1	Lake water, Alinen Mustajärvi, 5.1m in depth	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903899245.1	Lake water, Alinen Mustajärvi, 5.8m in depth	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903834355.1	Lake water, Alinen Mustajärvi, 6.3 m in depth	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903877805.1	Lake water, Alinen Mustajärvi, 3m in depth	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903833575.1	Lake water, Alinen Mustajärvi, 4m in depth	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903835705.1	Lake water, Alinen Mustajärvi, 5.5m in depth	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903908235.1	Lake water, Alinen Mustajärvi, 6m in depth	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903911995.1	Lake water, Alinen Mustajärvi, 4m in depth	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903920655.1	Lake water, Alinen Mustajärvi, 4.5m in depth	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903823855.1	Lake water, Alinen Mustajärvi, 4.5 m deep	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903882195.1	Lake water, Alinen Mustajärvi, 5.5m in depth	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903898235.1	Lake water, Alinen Mustajärvi, 6m in depth	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903857375.1	Lake water, Alinen Mustajärvi, 3.5m in depth	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903835535.1	Lake water, Alinen Mustajärvi, 4m in depth	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903870715.1	Lake water, Alinen Mustajärvi, 4.5m in depth	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903901165.1	Lake water, Alinen Mustajärvi, 5m in depth	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903869035.1	Lake water, Alinen Mustajärvi, 5.5m in depth	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903879085.1	freshwater sediment, Alinen Mustajärvi	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903833305.1	freshwater sediment, Alinen Mustajärvi	Finland; Lake Alinen-Mustajarvi	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903878515.1	Lake water, 3.9m in depth	Sweden; Lake Kiruna	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903893045.1	Lake water, 4.2m in depth	Sweden; Lake Kiruna	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903882675.1	Lake water	Sweden; Lake Kiruna	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903881155.1	Lake water	Sweden; Lake Kiruna	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)
GCA_903906125.1	Lake water, 0.2 m in depth	Sweden; Lake Bengtsgolen	*Methyloumidiphilus*	MAG	NA	Buck M et al. (2021)

Table 1: List of genomes of *Methylococcales* containing *ccyA*. Main information about the genomes, including whether they are derived from cultivated strain or metagenomes, the genera they are affiliated with following GTDB classification, their environmental and geographic origin.

3.2 Comparison of *Cyanobacteria* and *Methylococcaceae* calcyanin sequences

The calcyanin protein consists of two domains. The C-terminal $(GlyZip)_3$ domain, characterized by a triplicated long glycine zipper, is highly conserved across cyanobacteria (Benzerara et al., 2022). In contrast, the N-terminal domain is more variable and defines four calcyanin types called W (or CoBaHMA), X, Y and Z, the latter three being restricted to specific clades (*Gloeomargarita, Fischerella* and closely related genera, and *Microcystis*, respectively). A multiple sequence alignment of the C-terminal domains revealed conservation of the signature motifs of the $(GlyZip)_3$ domain in the *Methylococcaceae* (Figure 2), which included the repeated basic Gly-X(3)-Gly pattern within each GlyZip motif, interrupted at the midpoint by a highly conserved Gly-Pro dipeptide. Additionally, several polar and aromatic amino acids were notably conserved, particularly in the third GlyZip motif. Some variability was observed among members of the *Methylococcaceae* as in the whole set of cyanobacteria sequences, which principally occurred outside the GlyZip motifs. Three-dimensional (3D) structure predictions using AlphaFold3 suggested that the GlyZip motifs form a hairpin structure of tightly packed alpha-helices, though predicted with low pLDDT confidence scores (Figure 3). Importantly, AlphaFold3 failed to accurately predict the spatial arrangement between these motifs, indicating that the $(GlyZip)_3$ domain may adopt an as-yet-unknown 3D fold.

These sequence comparisons further reveal that the *Methylococcaceae* C-terminal $(GlyZip)_3$ domains are more similar to Z-type calcyanins of *Microcystis* than to other types found in other cyanobacteria. Strong and specific sequence similarities are especially apparent in the first GlyZip motif (Figure 2). Additionally, *Microcystis* Z-type calcyanins and *Methylococcaceae* calcyanins also share features in their N-terminal domain, which lacks strongly hydrophobic amino acids and is enriched in alanine and polar residues (Fig. S2). However, despite these compositional similarities, no significant sequence alignments could be established between

298   these N-terminal sequences, which are predicted by AlphaFold3 to form coiled-coil structures,

299   again with low confidence (Fig. S2).

Figure 2: Comparison of the sequences of calcyanin GlyZip motifs of *Methylococcaceae* and cyanobacteria. Four calcyanin sequences representative of thmae different clades were selected for cyanobacteria, while for *Methylococcaceae*, six non-redundant ones were extracted from Table 1. The *Fischerella* sp. NIES-4106 calcyanin sequence was chosen as it has three GlyZip motifs, instead of two for most of the calcyanin sequences of this clade. Identities are shown in white on a black background, while similarities are highlighted with colors, according to amino acid properties and position in the alignment (with highly similar positions indicated in bold). The periodic motif involving glycine (or a small amino acid) every four residues is highlighted by numbers relative to the central conserved proline (gold star), and can be followed from one GlyZip motif to the next using vertical dotted lines. Highly conserved polar/aromatic positions are highlighted with stars. The lilac pink box highlights a region which is highly conserved (HC) between the *Microcystis* and *Methylococcaceae* sequences. The secondary structures (α-helices) as predicted by AlphaFold3 (pLDDT<70) on the *Methylococcus geothermalis* sequence are shown on top of the alignment.

Moreover, the similarity between the Z-type calcyanins of *Microcystis* and the new calcyanins

of *Methylococcaceae* is strongly supported by phylogenetic analysis. Indeed, a maximum

likelihood tree based on the GlyZip motifs conserved in the C-terminal domain of all calcyanins

showed that the *Microcystis* and *Methylococcaceae* sequences are closely related with maximal

statistical support (Figure 4). Interestingly, the relationships among the *Methylococcaceae*

sequences perfectly mirrored those based on the phylogenomic analysis of a large set of

conserved markers of the corresponding *Methylococcaceae* species (Figure 4), indicating that

the *ccyA* gene has followed the speciation pattern of these species.

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Figure 3: Ribbon representation of the 3D structure of the GlyZip motifs, as predicted by Alphafold3
using the *Methylococcus geothermalis* calcyanin sequence. The GlyZip motifs were superimposed with
each other. The periodic motifs involving glycine (or a small amino acid) every four residues are colored
on the ribbons, and highly conserved charged/polar/aromatic positions (stars in Figure 2) are depicted
in atomic details. The region highly conserved (HC) between the *Microcystis* and *Methylococcaceae*
sequences is highlighted in lilac pink. For note, the Alphafold3 per-residue pLDDT values are < 70,
probably due to the fact that no template could be found in the protein data bank to support the relevance
of this prediction (and not because of a dubious prediction, a hypothesis ruled out by in-depth analysis
of the model, in particular highlighting the common central core around the Gly-Pro sequence). Structure
uncertainties concern the more variable N- and C-terminal extremities of the GlyZip motifs, which in
particular deviate from the packing imposed by the periodic repetitions of glycine in motifs 2 and 3.
343

Figure 4: Maximum likelihood phylogenetic tree of the GlyZip domain of the calcyanin sequences. Calcyanin sequences of *Methylococcaceae* species are highlighted in green. The different types of cyanobacterial calcyanins (X, W/CoBaHMA, Y, and Z) are indicated. For the *Methylococcaceae*, a species tree based on the concatenation of 120 conserved markers is shown on the left (this tree is rooted on an outgroup composed of *Methylococcaceae* sp. GCA_902812445.1, *Methylotetracoccus* sp. GCA_025458755.1, *Methylotetracoccus* sp. GCA_004168015.1, and *Methylotetracoccus oryzae* GCF 006175985.1). Numbers on branches show ultrafast bootstrap values. The sequences used the comparison of calcyanin sequences in Figure 2 appear in bold here.

3.3 Experimental evidence of iACC formation by *Methylococcaceae* cultivated strains

Both genomes of the cultivated strains *Methylococcus geothermalis* IM1 and *Methylococcus mesophilus* 16-5 contain the *ccyA* gene, which is indicative of iACC formation in cyanobacteria. Based on this, we searched for the presence of iACC in these strains. Laboratory cultures exhibited a typical exponential growth curve, with a doubling time of 5.8±0.1 h for *M. geothermalis* and 13.1±0.8 h for *M. mesophilus* (Fig. S3).

Attenuated total reflectance Fourier transform infrared (ATR-FTIR) spectroscopy performed on cell pellets showed unambiguously the presence of ACC in both *Methylococcaceae* cultivated strains (Figure 5 and Table S3 for band assignment). Indeed, spectra showed a band at around 860 cm^{-1}, detected in both *M. geothermalis*, *M. mesophilus*, a synthetic ACC reference and iACC-forming cyanobacteria but absent in a reference calcite and non-iACC-forming cyanobacteria, and that was shown to be characteristic of ACC (Mehta et al., 2022b). Other features observed in the spectra were bands at 1641 and 1528 cm^{-1}, identified as amide I and amide II bands (Figure 5). Bands at 1444 and 1391 cm^{-1} detected in both *M. geothermalis* and *M. mesophilus*, are similar to those observed at 1474 and 1395 cm^{-1} in ACC but they are not exclusive to iACC-forming cyanobacteria and likely correspond to methyl groups in lipids and/or C-O groups in carboxylic acids.

This finding was further corroborated by x-ray absorption near-edge structure (XANES) spectroscopy at the Ca K-edge conducted on *M. geothermalis* cells. The derivative of the Ca K-edge XANES spectrum of *M. geothermalis* cell pellets closely matched that of synthetic ACC and iACC in cyanobacteria, distinguishing it from calcite and non-iACC-forming cyanobacteria (Fig. 6; Mehta et al., 2023). Specifically, the B1 feature in the derivative spectrum, which was shifted compared to calcite, indicated the presence of ACC. Linear combination fitting (LCF) of the XANES data revealed that approximately ~57% of the Ca in *M. geothermalis* was in the

381　form of ACC, with the remaining 43% of Ca likely complexed by organics and/or associated

382　with polyphosphates (Fig. 6).

Figure 5: (A) ATR-FTIR spectra of *Methylococcus geothermalis* and *M. mesophilus* in the mid-IR range compared with reference spectra. Reference spectra are given for calcite, synthetic ACC, iACC-forming cyanobacteria (Cyano(ACC+)) and non-iACC-forming cyanobacteria (Cyano(ACC-)). The dashed rectangle envelops the zoomed regions of mid-IR ACC spectra shown in the inset 1 (1800–900 cm^{-1}) in (B), and inset 2 (950–820 cm^{-1}) in (C). Numbers on the spectra correspond to the band numbers and their assignment are provided in Table S3. Band 11 at 860 cm^{-1} is indicative of ACC.

Figure 6: (A) Derivate XANES spectra at the Ca K-edge of *Methylococcus geothermalis* cell pellets compared with reference derivate spectra of calcite, synthetic ACC, iACC-forming cyanobacteria and non-iACC-forming cyanobacteria. The presence of a peak noted B1 indicated by a dashed line is indicative of ACC. (B) Linear combination fitting (LCF) of the derivate XANES spectrum of *Methylococcus geothermalis* cell pellets at the Ca K-edge. The solid line shows experimental data and the dotted line corresponds to the LCF result.

Consistently, both scanning electron microscopy (SEM) and scanning transmission electron microscopy (STEM) in the high-angle annular dark-field (HAADF) imaging mode combined with energy dispersive x-ray spectrometry (EDXS) showed that cells of both strains contained

Ca-rich inclusions with some Mg but no S and no P, interpreted as calcium carbonate inclusions (Fig. 7 and Fig. S4). Besides, cells also contained granules rich in phosphorus (P), with some Mg and K, interpreted as polyphosphate inclusions (Rivas-Lamelo et al., 2017).

Figure 7: Scanning electron microscopy analyses of cultivated *Methylococcaceae*. (A-C) *Methylococcus geothermalis*. (A) Backscattered electron image. (B) Overlay of Ca (blue) and P (yellow) EDXS maps showing iACC and polyphosphate inclusions, respectively. (C) Characteristic EDXS spectra of polyphosphate (orange) and iACC (blue) inclusions. Emission lines of Mg Ka, P Ka, K Ka and Ca Ka are indicated. (D-F) Same analyses for *M. mesophilus*.

Last, cryo-soft X-ray tomography (c-SXT) was used to image vitrified *M. geothermalis* cells in a near-native state (15 cells in total, 6 cells used for volume reconstruction). Ca-rich inclusions were observed as regions of higher intensity (Figure 8A and Figure S5). These inclusions are most evident in the tomograms due to their strong absorption contrast and small, spherical morphology (Figure 8A). The absorption contrast from Ca-rich inclusions is much higher than other structures inside the cell (Figure 8B). Clearly defined Ca-rich inclusions were segmented to determine their average size and distribution, yielding a mean diameter of 118 ± 45 nm (n = 58) (Figure 8C). Other structural features, such as intracytoplasmic membranes and phosphate inclusions (Awala et al., 2020), can be distinguished from the smaller Ca-rich inclusions in the

467 tomograms but not from each other. For whole-cell segmentation, these structures were

468 therefore grouped as a single material. Figure 8D shows the 3D volume reconstruction of the

469 two cells from Figure 8A, highlighting the outer cell membrane, intracytoplasmic

470 membranes/phosphate inclusions, and Ca-rich inclusions. At least 10 Ca-rich inclusions were

471 found per cell.

Figure 8: Cryo-soft X-ray tomography (c-SXT) of two ice-vitrified *Methylococcus geothermalis* cells. The image intensity corresponds to X-ray absorption by intracellular material at 520 eV—an energy within the "water window," where absorption by C and N atoms provides contrast of biological structures without chemical staining. At this energy, absorption by O atoms is minimal, allowing X-ray transmission through both the ice layer and hydrated cells. In contrast, the Ca $L_{2,3}$-edge at 350 eV enables detection of Ca atoms within the cells. (A) Virtual mid-cell z-slice of two *M. geothermalis* cells in reconstructed tomogram (an additional tomogram of two different cells is shown in Figure S3, along with tomogram videos for each dataset). Intensity scale (rainbow colour code) refers to absorption values. Asterisk marks position of a Ca-rich inclusion. Scale bar is 500 nm. (B) Absorption profile extracted from horizontal line in (A) (asterisk marks position of a Ca-rich inclusion). (C) Particle size distribution of Ca-rich granules measured from segmented volumes of the tomogram. (D) Volume reconstruction of cells in (A) shown in two perspectives with Ca-rich granules in pink, intracytoplasmic membranes/phosphate inclusions in beige and outer cell membrane in grey.

4. Discussion

4.1 Formation of iACC inclusions by bacteria remains an overlooked process

Here we report the discovery of iACC in members of the *Methylococcaceae*. This clade expands the growing list of bacteria known to biomineralize ACC, which already includes *Achromatium* (Benzerara et al., 2021; Gray, 2006; Schewiakoff, 1893), numerous phylogenetically diverse members of the cyanobacteria (Benzerara et al., 2022) and several magnetotactic *Pseudomonadota*, spanning both *Gammaproteobacteria* and *Alphaproteobacteria* (Liu et al., 2021; Mangin et al., 2025; Monteil et al., 2020). Collectively, these discoveries suggest that iACC biomineralization is more widespread than previously recognized but has often gone unnoticed, likely due to the analytical challenges associated with detecting ACC. Indeed, amorphous carbonates are highly labile and require dedicated techniques for reliable identification (Addadi et al., 2003).

For example, although the ultrastructures of *M. geothermalis* and *M. mesophilus* had previously been examined using electron microscopy, iACC granules were not reported (Awala et al., 2020, 2023). This omission can be explained both by the distinct objectives of those studies that did not targeted mineralogy and the absence of chemical analyses suited to the detection of iACC. In contrast, our combined use of STEM-EDXS on intact, untreated cells together with Ca K-edge XANES and/or FTIR spectroscopy enabled the unambiguous detection of iACC in both species. Moreover, iACC granules were likely lost during the ultramicrotomy preparations employed by Awala et al. (2020 & 2023) since standard protocols involving chemical fixation, post-fixation and dehydration are known to dissolve iACC in cyanobacteria (Blondeau et al., 2018a; Li et al., 2016). It is therefore possible that the numerous electron-translucent vesicles or intracellular inclusions previously interpreted as polyphosphate or glycogen granules by Awala et al. (2020, 2023) actually represent dissolved or transformed iACC granules. However, this hypothesis can only be rigorously tested using cryo-electron microscopy of vitreous sections (CEMOVIS).

An additional and pivotal line of evidence for iACC biomineralization in the *Methylococcaceae* was the detection of the *ccyA* gene, previously identified as a diagnostic marker of iACC formation in cyanobacteria (Benzerara et al., 2022). Our findings extend the diagnostic value of *ccyA* beyond cyanobacteria, although so far, no other clades represented in current genomic databases appear to carry this gene—unless their versions of *ccyA* have diverged substantially. Importantly, however, the absence of *ccyA* does not necessarily imply the absence of iACC formation. For instance, some magnetotactic *Pseudomonadota* and *Achromatium* both form iACC despite lacking *ccyA*, suggesting that multiple, distinct molecular pathways may give rise to iACC biomineralization (Mangin et al., 2025).

Last, numerous MAGs of iACC-forming *Methylococcaceae* that we have detected originate from chemically stratified boreal lakes and ponds (Table 1; Buck et al., 2021). Interestingly, MAGs from other bacteria phylogenetically close to iACC-forming magnetotactic bacteria were also found in the same stratified lakes and ponds in Finland and Sweden (Mangin et al., 2025), suggesting that iACC biomineralization might be a geochemically important process in these lakes.

4.2 Relationship between *Methylococcaceae* metabolism and iACC biomineralization

The mechanisms governing iACC biomineralization remain poorly understood. It has been proposed that bacterial primary metabolism may play an important role in driving this process (Görgen et al., 2020). In *Achromatium*, a sulfur-oxidizing bacterium, iACC formation is linked to sulfide oxidation to elemental sulfur (Yang et al., 2019). This reaction consumes protons, thereby increasing intracellular pH. The resulting shift in carbonate equilibrium favors the conversion of bicarbonate (HCO_3^-) to carbonate ions (CO_3^{2-}), which in turn promotes calcium carbonate precipitation. Conversely, when elemental sulfur is oxidized to sulfate, protons are released, decreasing intracellular pH and causing iACC dissolution. Thus, iACC function as a

dynamic pH buffer: consuming protons through their dissolution and releasing them through their precipitation.

A similar hypothesis has been proposed for cyanobacteria, where photosynthesis elevates intracellular pH and thereby increases carbonate ion concentration, favoring iACC precipitation (Li et al., 2016). While these models hold for *Achromatium* and cyanobacteria, the present study shows that a proton-consuming primary metabolism is not strictly required for iACC biomineralization. In *Methylococcaceae*, aerobic methanotrophy does not consume protons, as shown by equation (1):

$$CH_4 + 2O_2 = CO_2 + 2H_2O \quad (1)$$

Moreover, the CO_2 produced by aerobic methanotrophs should promote calcium carbonate dissolution as shown by Krause et al. (2014) and according to carbonate equilibria:

$$CaCO_3 + CO_2 + H_2O = Ca^{2+} + 2HCO_3^- \quad (2)$$

However, a portion of this CO_2 may be diverted into small organic molecules via RuBiSCO-mediated carbon fixation, as shown in *Methylococcus*, and may therefore not react with $CaCO_3$ (Oshkin et al., 2021). Overall, iACC precipitation in *Methylococcaceae* likely relies on an additional, secondary metabolic process—yet to be identified—that consumes protons and converts CO_2 into carbonate ions so that:

$$CO_2 + H_2O = H_2CO_3$$

$$H_2CO_3 + Ca^{2+} + 2OH^- = CaCO_3 + 2H_2O$$

Alternatively, if CO_2 fixation exceeds CO_2 release from methanotrophy, the net CO_2 production by the cell would become at least sometimes negative, shifting the carbonate equilibrium towards carbonate ions and favoring ACC precipitation. This hypothesis aligns with recent findings in magnetotactic iACC-forming *Alphaproteobacteria*, whose primary metabolism does not induce alkalinization but appears to support CO_2 fixation (Mangin et al., 2025).

Finally, calcium uptake is essential for storage within iACC. Ca^{2+}/H^+ protein exchangers (Cax

proteins) are plausible candidates and we have detected *cax* genes in *Methylococcaceae*

genomes (data not shown). However, these transporters are generally inferred to function in

calcium efflux rather than uptake in bacteria (Waditee et al., 2004), which contradicts the

requirements for iACC formation. Future studies are needed to elucidate the specific metabolic

and transport processes enabling iACC biomineralization in *Methylococcaceae*. Aerobic

methane-oxidizing bacteria are widespread in many aquatic systems and have a significant

environmental and ecological role since they are key in the regulation of emissions of CH_4, a

potent greenhouse gas, and provide CH_4-derived organic carbon compounds to surrounding

heterotrophic microorganisms (Khanongnuch et al., 2022). How the capability to biomineralize

iACC in some of these organisms may play an overlooked regulatory role in these important

biogeochemical processes remains to be understood.

4.3 Evolutionary origin of iACC biomineralization in *Methylococcaceae*

The *ccyA* gene has an ancient origin within cyanobacteria, explaining its broad phylogenetic

distribution across the phylum, with multiple secondary losses in modern lineages (Benzerara

et al., 2022). This evolutionary history gave rise to the diversification of several calcyanin

families, characterized by distinct N-terminal domains, termed CoBaHMA, X, Y, and Z. The

X, Y, and Z types are restricted to specific clades, i.e. *Gloeomargarita*, *Fischerella* (and closely

related genera), and *Microcystis*, respectively. In contrast, the CoBaHMA-type calcyanin

occurs in several distantly related lineages and has therefore been proposed as the most ancient

form of calcyanin.

The calcyanin sequences identified in *Methylococcaceae* have specific characteristics not found

in the four cyanobacterial types but share several similarities with the *Microcystis* (Z-type)

calcyanins. This is unexpected given the vast evolutionary distance between *Methylococcaceae*

and all the cyanobacterial clades. Phylogenetic analysis strongly supports the relationship

between the *Methylococcaceae* and cyanobacteria sequences (Figure 4). The most parsimonious explanation for this observation is that *Methylococcaceae* acquired *ccyA* through horizontal gene transfer (HGT) from a cyanobacterial ancestor closely related to *Microcystis*. The fact that the phylogeny of the *Methylococcaceae* calcyanin sequences mirrors that of the corresponding species, strongly suggests a single HGT acquisition by an ancestral *Methylococcaceae* species, followed by subsequent divergence in several *Methylococcaceae* subgroups, including many independent losses in the contemporary *Methylococcaceae* lineages lacking *ccyA*. Such independent losses have also been proposed to explain the current distribution of *ccyA* in the contemporary cyanobacteria lineages (Benzerara et al., 2022).

About 1/3 of known *Microcystis* genomes contain *ccyA*, which was also likely ancestral in this genus and lost secondarily in several lineages (Gaetan et al., 2023). Consistently with the presence of *ccyA*, abundant iACC-forming *Microcystis* strains have been observed in several lakes (Gaetan et al., 2023). One plausible ecological setting for the HGT of *ccyA* from *Microcystis* to *Methylococacceae* is exemplified by Lake Cadagno, a meromictic lake (i.e., permanently stratified with an anoxic monimolimnion). At its chemocline, methane removal is driven by aerobic methane oxidizers which are physically associated with cyanobacteria (Milucka et al., 2015). More precisely, physical associations between *Microcystis* and diverse *Methylococcales* are common in lakes and ponds, where they may be responsible for an overlooked sink for methane (Li et al., 2021). Similarly, microbial consortia comprising cyanobacteria and methanotrophic bacteria are employed in oxygen-deficient wastewater treatment plants to remove methane (Safitri et al., 2021). In these consortia, cyanobacteria produce O_2 that fuels aerobic methanotrophy, while the resulting CO_2 is taken up by the phototrophs. Such tight metabolic coupling provides an ideal setting for close and sustained interactions between cyanobacteria and *Methylococcales*, potentially facilitating HGT. One of

628 these HGT events appears to have transmitted the capability to biomineralize iACC from

629 cyanobacteria to *Methylococcaceae.*

Conclusions

This study shows that aerobic methanotrophs belonging to the *Methylococcaceae* family form iACC. They acquired the *ccyA* gene, a diagnostic marker of iACC biomineralization in cyanobacteria, through HGT from an ancestral *Microcystis*-like cyanobacterial donor. This HGT event may have taken place in a permanently stratified lake or in any other anoxic-oxic transition zone in aquatic habitats, where *Microcystis* and *Methylococcaceae* can co-exist. *Methylococcaceae* forming iACC may be abundant in some environments. Yet, how iACC are formed, in particular, how much of the carbon from methane oxidation is stored within iACC and diverted from the production of small organic carbon molecules by these bacteria remains to be quantified.

Data Availability Statement

All gene sequences were retrieved from GTDB as indicated in the Material and Method section and no new sequence was generated by this study. Microscopy and spectroscopy data are available on Zenodo reference number xxx.

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Conflict of interest statement

659 The authors declare no conflicts of interest.

Author contributions (see CRediT)

Benzerara Karim: conceptualization, formal analysis, writing – original draft, supervision,

funding acquisition. **Millet Maxime**: conceptualization, methodology, formal analysis, writing

– original draft. **Skouri-Panet Fériel**: methodology, formal analysis, writing – review and

editing. **Gaschignard Geoffroy**: methodology, formal analysis. **Mehta Neha**: methodology,

formal analysis, writing – review and editing. **Bezard Malo**: formal analysis. **Caumes

Geraldine**: conceptualization, formal analysis. **Chevrier Daniel M**.: formal analysis, writing

– original draft. **Dezi Manuela**: funding acquisition. **Duverger Arnaud**: methodology.

Guigner Jean-Michel: methodology. **Gutiérrez-Preciado Ana**: formal analysis. **Lefevre

Christopher T**.: conceptualization, funding acquisition, writing – review and editing. **López-

García Purificación**: conceptualization, funding acquisition, writing – review and editing.

Menguy Nicolas: methodology. **Monteil Caroline L.**: conceptualization, funding acquisition,

writing – review and editing. **Pehau-Arnaudet Gerard**: methodology. **Penard Esthel**:

methodology. **Pereiro Eva**: methodology. **Scandola Cyril**: methodology. **Travert Cynthia**:

methodology. **Vantelon Delphine**: methodology, writing – review and editing. **Duprat Elodie**:

conceptualization, methodology, supervision, formal analysis. **Callebaut Isabelle**:

conceptualization, methodology, formal analysis, supervision, writing – original draft. **Moreira

David: conceptualization, methodology, formal analysis, supervision, writing – original draft,

funding acquisition.

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