# Spatiotemporal prediction of wildfire extremes with Bayesian finite sample maxima

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#### Abstract

Wildfires are becoming more frequent in parts of the globe, but predicting where and when wildfires occur remains difficult. To predict wildfire extremes across the contiguous United States, we integrate a 30 year wildfire record with meteorological and housing data in spatiotemporal Bayesian statistical models with spatially varying nonlinear effects. We compared different distributions for the number and sizes of large fires to generate a posterior predictive distribution based on finite sample maxima for extreme events (the largest fires over bounded spatiotemporal domains). A zero-inflated negative binomial model for fire counts and a lognormal model for burned areas provided the best performance. This model attains 99% interval coverage for the number of fires and 93% coverage for fire sizes over a six year withheld data set. Dryness and air temperature strongly predict extreme wildfire probabilities. Housing density has a hump-shaped relationship with fire occurrence, with more fires occurring at intermediate housing densities. Statistically, these drivers affect the chance of an extreme wildfire in two ways: by altering fire size distributions, and by altering fire frequency, which influences sampling from the tails of fire size distributions. We conclude that recent extremes should not be surprising, and that the contiguous United States may be on the verge of even larger wildfire extremes.

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# 1 Introduction

Wildfire frequency and burned area has increased over the past couple decades in the 2 United States (Dennison et al. 2014; Westerling 2016), and elsewhere (Krawchuk et al. 3 2009; Pechony and Shindell 2010). In addition to the ecological and smoke impacts associ-4 ated with increased burned area, there has been an increasing interest in extreme wildfires 5 Williams 2013) given their impact on human lives and infrastructure (Kochi et al. 2010; 6 Diaz 2012). While case studies of particular extremes provide insight into what caused 7 past events (Peterson et al. 2015; Nauslar, Abatzoglou, and Marsh 2018), predictions of 8 future extremes at a national level could inform disaster related resource allocation. The 9 term "extreme" has multiple meanings with respect to wildfires (Tedim et al. 2018), and 10 in this paper we take consider an extreme wildfire to be a fire with the largest burned area 11 over a bounded spatiotemporal domain, i.e., the block maximum within a spatial region 12 and a temporal interval (Coles et al. 2001). For example, the block maxima for widlfires 13 across the contiguous U.S. can be defined on a yearly basis (Figure 1). 14

Factors driving wildfire extremes vary in space and time (Barbero et al. 2014), but it is 15 unclear how best to account for this in a predictive model. Previous efforts have used year 16 or region-specific models, aggregating over space or time (Bermudez et al. 2009), tempo-17 rally or spatially explicit models (Mendes et al. 2010), and spatial models with year as a 18 covariate (Díaz-Avalos, Juan, and Serra-Saurina 2016). Recently, rich spatiotemporal mod-19 els have been described with linear, spatially constant covariate effects (Serra, Saez, Juan, 20 et al. 2014; Serra, Saez, Mateu, et al. 2014). However, linear, spatially constant effects 21 are suboptimal over large spatial domains with nonlinear drivers (Fosberg 1978, Goodrick 22 (2002), Preisler et al. (2004); Preisler and Westerling 2007; Balshi et al. 2009; Krawchuk et 23 al. 2009; Pechony and Shindell 2009; Vilar et al. 2010; Woolford et al. 2011; Woolford et 24 al. 2014). For example, global wildfire probability shows a hump-shaped relationship with 25 temperature and moisture (Moritz et al. 2012). Interactions among drivers also impose 26

nonlinearity, e.g., in hot and dry climates fires are fuel limited (McLaughlin and Bowers
1982), but in cold and wet climates fires are energy limited (Krawchuk and Moritz 2011).

Prediction is also complicated by uncertainty in which distribution(s) to use to assign 29 probabilities to extreme events. The generalized Pareto distribution (GPD) has frequently 30 been used (Bermudez et al. 2009; Jiang and Zhuang 2011), but the GPD requires a thresh-31 old to delineate extreme events (Davison and Smith 1990, Coles (2014)). The utility and 32 validity of a threshold for extremes in a heterogeneous region is debatable (Tedim et al. 33 2018). Recently proposed metastatistical extreme value (MEV) approaches do not require 34 such a threshold, and are based on the statistical distribution of finite sample maxima, i.e., 35 the probability distribution of the maximum value for a finite number of events (Marani 36 and Ignaccolo 2015; Zorzetto, Botter, and Marani 2016). In the MEV framework, the oc-37 currence and size of future events, and the parameters of their distributions are treated 38 as random variables which together imply a distribution for extremes. This approach has 39 roots in compound distributions (Dubey 1970; Wiitala 1999), doubly stochastic processes 40 (Cox and Isham 1980), superstatistics (Beck and Cohen 2003), and the Bayesian posterior 41 predictive distribution (Gelman et al. 2013). The link to Bayesian inference is particularly 42 useful, as it provides an easy way to propagate uncertainty forward to to predictions of 43 extremes (Coles, Pericchi, and Sisson 2003). 44

<sup>45</sup> Here, we extend the finite sample maximum approach to account for non-linear, spatially
<sup>46</sup> varying covariate effects with the goal of predicting extreme wildfire events from a sta<sup>47</sup> tistical perspective across the contiguous United States. Specifically, we aim to predict
<sup>48</sup> occurrence (where and when), and magnitude (burned area) of large wildfires at a monthly
<sup>49</sup> time scale and regional spatial scale across the contiguous United States.

# $_{50}$ Methods

## <sup>51</sup> Data description

We acquired wildfire event data for the contiguous United States from the Monitoring 52 Trends in Burn Severity (MTBS, www.mtbs.gov) program (Eidenshink et al. 2007), which 53 includes spatiotemporal information on the occurrence of wildfires in the United States 54 from 1984 to 2016. The MTBS data contain fires greater than 1000 acres ( $\approx 405$  hectares) 55 in the western U.S. and greater than 500 acres ( $\approx 202$  hectares) in the eastern U.S. For 56 consistency across the U.S., we discarded all records in the MTBS data less than 1000 57 acres, retaining 10,736 fire events (Figure 2A). Each event in the MTBS data has a discov-58 ery date, spatial point location, and final size. 59

To explain fire size and occurrence, we used a combination of meteorological variables in-60 cluding humidity, air temperature, precipitation, and wind speed. These variables were 61 selected on the basis of previous work, and also with an aim to drive a predictive model 62 with interpretable meteorological quantities. Meteorological layers were acquired from the 63 gridMET data (Abatzoglou 2013) that blends monthly high-spatial resolution (~4-km) 64 climate data from the Parameter-elevation Relationships on Independent Slopes Model 65 (Daly et al. 2008) with high-temporal resolution (hourly) data from the National Land 66 Data Assimilation System (NLDAS2) using climatologically aided interpolation. The resul-67 tant products are a suite of surface meteorological variables summarized at the daily time 68 step and at a 4-km pixel resolution. Daily total precipitation, minimum relative humidity, 69 mean wind speed, and maximum air temperature were averaged at a monthly time step for 70 each of 84 Environmental Protection Agency level 3 (L3) ecoregions for each month from 71 1984 to 2016 (Omernik 1987; Omernik and Griffith 2014). We also computed cumulative 72 monthly precipitation over the previous 12 months for each ecoregion-month combina-73 tion. We chose to segment the U.S. with level 3 ecoregions as a compromise between the 74 more numerous (computationally demanding) level 4 ecoregions, and the coarser level 2 75 ecoregions. 76

<sup>77</sup> We used publicly available housing density estimates that were generated based on the <sup>78</sup> U.S. 2000 decennial census as explanatory variables that may relate to human ignition <sup>79</sup> pressure (Radeloff et al. 2010). These are provided at decadal time steps, and spatially <sup>80</sup> at the level of census partial block groups. To generate approximate measures of hous-<sup>81</sup> ing density at monthly time intervals, we used a simple linear interpolation over time for <sup>82</sup> each block group, then aggregated spatially across block groups to compute mean housing <sup>83</sup> density for each ecoregion in each month.

# <sup>84</sup> Model development

We built two types of models: one describing the occurrence of fires within each L3 ecoregion over time (i.e., the total number of fires occurring in each ecoregion for each month from 1984 - 2016), and another describing the size of each wildfire in each ecoregion and month. For occurrence models, the response variable was a count (number of fires), and for burned area models, the response was a continuous positive quantity (size of each fire event). We used the period from 1984 to 2009 for training, witholding the period from 2010 to 2016 to evaluate predictive performance.

#### 92 Fire occurrence

We constructed four models for fire occurrence and compared their predictive performance 93 based on test-set log likelihood and posterior predictive checks for the proportion of zeros, 94 maximum count, and total count. The models differed in the distributions used in the like-95 lihood, representing counts as a Poisson, negative binomial, zero-inflated Poisson, or zero-96 inflated negative binomial random variable. The Poisson distribution is a common choice 97 for counts, and the negative binomial distribution provides an alternative that can account 98 for overdispersion. The zero-inflated versions of these distributions include a component 99 to represent extra zeros, which might be expected to work well if there are independent 100 processes that determine whether nonzero counts are possible (Lambert 1992). 101

For spatial units (ecoregions) s = 1, ..., S and time steps (months) t = 1, ..., T, each

model defines a probability mass function for  $n_{s,t}$ : the number of fires over 405 hectares in ecoregion s and time step t. For each of the four count distributions under consideration, location parameters  $\mu_{s,t}$  and (for zero-inflated models) structural zero inflation parameters  $\pi_{s,t}$  were allowed to vary in space and time. We used a log link function to ensure that  $\mu_{s,t} > 0$ , and a logit link function to ensure that  $\pi_{s,t} \in (0, 1)$ . Concatenating over spatial and temporal units, so that  $\boldsymbol{\mu} = (\mu_{s=1,t=1}, \mu_{s=2,t=1}, ..., \mu_{s=S,t=2}, ..., \mu_{s=S,t=T})$ , and similarly for  $\boldsymbol{\pi}$ , we modeled location and (when applicable) zero inflation parameters as:

$$\log(\boldsymbol{\mu}) = \alpha^{(\mu)} + \mathbf{X}\boldsymbol{\beta}^{(\mu)} + \boldsymbol{\phi}^{(\mu)} + \log(\boldsymbol{a}),$$

$$\operatorname{logit}(\boldsymbol{\pi}) = \alpha^{(\pi)} + \mathbf{X}\boldsymbol{\beta}^{(\pi)} + \boldsymbol{\phi}^{(\pi)},$$

where  $\alpha^{(\mu)}$  and  $\alpha^{(\pi)}$  are scalar intercept parameters, **X** is a known  $(S \times T) \times p$  design matrix, where p is the number of input features,  $\beta^{(\mu)}$  and  $\beta^{(\pi)}$  are column vector parameters of length p,  $\phi^{(\mu)}$  and  $\phi^{(\pi)}$  are column vector parameters of length  $S \times T$  containing spatiotemporal adjustments, and  $\boldsymbol{a}$  is a known offset vector of areas for spatial unit s = 1, 2, ..., S, repeated T times.

#### 115 Burned area

We developed five candidate models for fire size, each of which specified a different distribution for the size (burned area) of individual fire events (Reed and McKelvey 2002; Hernandez et al. 2015), including the generalized Pareto (Hosking and Wallis 1987), tapered Pareto (Schoenberg, Peng, and Woods 2003), lognormal, gamma, and Weibull distributions. We evaluated each model in terms of test set log likelihood, and posterior predictive checks for fire size extremes. We defined the response  $y_i$  as the number of hectares burned over 405 for the  $i^{th}$  fire event, which occurred in spatial unit  $s_i$  and time step  $t_i$ .

Because each burned area distribution has a different parameterization, we included
 covariate effects in a distribution-specific way. For the generalized Pareto distribution

(GPD), we assumed a positive shape parameter, leading to a Lomax distribution for ex-125 ceedances (Bermudez et al. 2009). The GPD and Lomax shape parameters are related by 126  $\kappa^{(GPD)} = 1/\kappa^{(L)}$ , and the GPD scale parameter is related to the Lomax scale and shape 127 parameters by  $\sigma^{(GPD)} = \sigma^{(L)} / \kappa^{(L)}$ . We introduced covariate dependence via the Lomax 128 scale parameter using a log link. For event i,  $\log(\sigma_i^{(L)}) = \alpha + \mathbf{X}_{(s_i,t_i)} \boldsymbol{\beta} + \phi_{s_i,t_i}$ , where  $\alpha$  is 129 an intercept parameter,  $\boldsymbol{\beta}$  is a length p vector of coefficients,  $\boldsymbol{X}_{(s_i,t_i)}$  is a row vector from 130 **X**, and  $\phi_{s_i,t_i}$  is a spatiotemporal adjustment for  $s_i$  and  $t_i$ . For the tapered Pareto model, 131 we modeled the shape parameter as  $\log(\kappa_i) = \alpha + \mathbf{X}_{(s_i,t_i)}\boldsymbol{\beta} + \phi_{s_i,t_i}$ . The lognormal model 132 included covariate dependence via the location parameter:  $\mu_i = \alpha + X_{(s_i,t_i)}\beta + \phi_{s_i,t_i}$ . The 133 gamma model used a log link for the expected value:  $\log(E(y_i)) = \alpha + \mathbf{X}_{(s_i,t_i)} \boldsymbol{\beta} + \phi_{s_i,t_i}$ . Last, 134 we modeled the Weibull scale parameter as  $\log(\sigma_i) = \alpha + \mathbf{X}_{(s_i,t_i)} \boldsymbol{\beta} + \phi_{s_i,t_i}$ . More detail on 135 the parameterization of each burned area distribution is provided in the Appendices. 136

#### 137 Accounting for nonlinear forcing

The design matrix **X** was constructed to allow for spatially varying nonlinear effects of housing density and meteorological drivers. We used B-splines to account for nonlinearity (Figure 3) and allowed the coefficients for each basis vector to vary spatially (Wood 2017). First, we constructed univariate B-splines for log housing density, wind speed, same month precipitation, previous 12 month precipitation, air temperature, and humidity, with five degrees of freedom (including an intercept) for each variable. This step generated 30 basis vectors (five for each of six variables).

To allow for spatial variation in these nonlinear effects, we added interaction effects be-145 tween each of the basis vectors and ecoregions (Brezger and Lang 2006; Kneib, Hothorn, 146 and Tutz 2009). The hierarchical nesting of ecoregion designations (Figure 2B-D) lends 147 itself to such interactions. Conceptually, coefficients in a level 3 ecoregion may be related 148 to coefficients in the level 2 ecoregion containing the level 3 region, the level 1 ecoregion 140 containing the level 2 region, and a global effect. The coefficient associated with a basis 150 vector for any level 3 ecoregion is treated as a sum of a global effect, a level 1 ecoregion 151 adjustment, a level 2 ecoregion adjustment, and a level 3 ecogregion adjustment. Thus, 152

for every univariate basis vector, we included interaction effects with ecoregion at each of the three ecoregion levels. This allows borrowing of information across space (level 3 ecoregions in a level 2 ecoregion are often adjacent), and for regions that are ecologically similar. We also included adjustments on the global intercept for each level 1, 2, and 3 ecoregion to account for spatial variation that is unrelated to climate or housing density. This specification induces sparsity in **X** that we exploit to increase the efficiency of computing  $\mu$  and  $\pi$ . In total, **X** has p = 3,472 columns, with 97% zero entries.

# <sup>160</sup> Prior specification

To avoid overfitting, we used a regularized horseshoe prior on the coefficients associated 161 with the spatially varying nonlinear effects described above (Piironen, Vehtari, and others 162 2017). This prior places high probability close to zero, while retaining heavy enough tails 163 that nonzero coefficients are not shrunk too strongly toward zero. This is consistent with 164 our prior expectation that most of the coefficients associated with the columns in X were 165 close to zero. For the zero inflated count models, we used a multivariate horseshoe to allow 166 information sharing between the zero inflated and distribution specific location parameters 167 (Peltola et al. 2014). For the remaining count models and all burned area models, this 168 was a univariate horseshoe prior. Spatiotemporal random effects were constructed using 169 a temporally autoregressive, spatially intrinsically autoregressive formulation (Besag and 170 Kooperberg 1995; Banerjee, Carlin, and Gelfand 2014). Details of these priors and the 171 resulting joint distributions are provided in the Appendices. 172

### <sup>173</sup> Posterior predictive inference for finite sample maxima

<sup>174</sup> We used the posterior predictive distribution to check each model and make inference on <sup>175</sup> extremes. The posterior predictive distribution provides a distribution for replications <sup>176</sup> of observed data  $(y^{\text{rep}})$ , and predictions of future data (Gelman et al. 2013). Concep-<sup>177</sup> tually, for a good model,  $y^{\text{rep}}$  should be similar to observed training data y, and future <sup>178</sup> predictions should be similar to future data. Distributions over both quantities can be

obtained by conditioning on y and marginalizing over model parameters  $\theta$ , e.g.,  $[y^{rep}|y] = \int [y^{rep}|\theta][\theta|y]d\theta$ .

Posterior predictive distributions facilitate model checks that compare predicted and ob-181 served test statistics (Gelman, Meng, and Stern 1996). To evaluate whether models cap-182 tured tail behavior, we compared empirical maxima  $(T(y) = \max(y))$  to the predicted 183 distribution of maxima  $T(y^{rep})$ . We also include predictive checks for the proportion of 184 zero counts, and totals for count and burned area models. Posterior predictive inference 185 for finite sample maxima is similar in spirit to the MEV approach. Both obtain a distri-186 bution over maxima by marginalizing over unknowns including the number of events, size 187 of each event, and parameters of their distributions (Marani and Ignaccolo 2015). How-188 ever, a Bayesian approach explicitly conditions on the observed data to obtain a posterior 189 distribution of parameters. 190

Seeing this connection is useful in the context of including priors and propagating uncer-191 tainty to derived parameters. For any ecoregion s and timestep t, if we define a particular 192 maximum fire size conditional on a fire having occurred as  $z_{s,t}$ , and let  $Z_{s,t}$  represent the 193 random variable of maximum fire size, then the cumulative distribution function (CDF) 194 for  $z_{s,t}$  is given by  $\Pr(Z_{s,t} \leq z_{s,t}) = F(y_{(s,t)})^{n_{s,t}}$ , where  $F(y_{(s,t)})$  is the CDF of fire size, and 195  $n_{s,t}$  is the number of wildfire events. Thus,  $\Pr(Z_{s,t} \leq z_{s,t})$  is the distribution function for 196 the finite sample maximum. The CDF for  $z_{s,t}$  can be inverted to produce a quantile func-197 tion that permits computation of prediction intervals for maximum fire sizes, conditional 198 on fires having occurred. Given a collection of posterior draws from a burned area model 199 that parameterize  $F(y_{(s,t)})$ , and a collection of posterior draws of  $n_{s,t}^{rep}$  from the posterior 200 predictive distribution of a wildfire count model, a posterior distribution for the CDF or 201 quantile function of maximum fire size can be generated which combines the two models 202 to facilitate inference on the distribution of extremes. 203

### <sup>204</sup> Parameter estimation

We used a combination of variational approximations and Hamiltonian Monte Carlo meth-205 ods to sample from the posterior distributions of count and burned area models. A varia-206 tional approximation (Kucukelbir et al. 2015) was used for count models to quickly iden-207 tify a preferred model. The best performing count model and all burned area models were 208 fit using the No-U-Turn Sampler (Hoffman and Gelman 2014). Models were fit in the Stan 209 probabilistic programming language using the rstan package (Carpenter et al. 2016; Stan 210 Development Team 2018). We ran four chains for 1000 iterations each, discarding the first 211 500 iterations as warmup. Convergence was assessed using visual inspection of trace plots, 212 with potential scale reduction statistic values  $\hat{R} \ge 1.1$  as an indicator convergence failure 213 (Brooks and Gelman 1998). 214

### <sup>215</sup> Implementation

All data processing, model fitting, and visualization were implemented with open source 216 software, primarily in the R programming language (R Core Team 2017), and wrapped in 217 a reproducible workflow via GNU Make and Docker (Stallman, McGrath, and Smith 2004; 218 Boettiger 2015). Data cleaning and transformation required the R packages assertthat 219 (Wickham 2017a), lubridate (Grolemund and Wickham 2011), Matrix (Bates and Maech-220 ler 2018), pbapply (Solymos and Zawadzki 2018), splines (R Core Team 2018), tidyverse 221 (Wickham 2017b), and zoo (Zeileis and Grothendieck 2005). Spatial data were processed 222 with raster (Hijmans 2017), rgdal (Bivand, Keitt, and Rowlingson 2018), sf (Pebesma 223 2018), and spdep (Bivand and Piras 2015). Finally, we used complot (Wilke 2017), ggre-224 pel (Slowikowski 2018), ggthemes (Arnold 2018), patchwork (Pedersen 2017), and RCol-225 orBrewer (Neuwirth 2014) for visualization. The manuscript was written in R Mark-226 down (Allaire et al. 2018). Analyses were run on an Amazon Web Services m5.2xlarge 227 EC2 instance with four physical cores and 32 GB of RAM, and the whole workflow re-228 quires  $\approx 72$  hours. All code to reproduce the analysis is available on GitHub at https: 220 //github.com/mbjoseph/wildfire-extremes (Joseph 2018). 230

# 231 **Results**

#### <sup>232</sup> Wildfire occurrence

The zero-inflated negative binomial distribution performed best on the held-out test set (Table 1), and was able to recover the proportion of zeros, count maxima, and count totals in posterior predictive checks for both the training and test data (Figure 4). All of the other count models that we considered exhibited lack of fit to at least one of these statistics in posterior predictive checks. Hereafter, we report results from the zero-inflated negative binomial model.

Minimum relative humidity and maximum air temperature had the strongest effects on 239 both the zero-inflation component and the expected value of the negative binomial com-240 ponent (Figure 5, posterior median for  $\rho$ : 0.665, 95% credible interval (CI): 0.319 - 0.861). 241 The model uncovered unique effects of meteorological variables at level 1, 2, and 3 ecore-242 gions (Figure 6). For example, a positive interaction effect between the second air temper-243 ature basis vector and the L1 Great Plains ecoregions indicates that the expected number 244 of wildfires in plains ecoregions with cold conditions is high relative to other ecoregions. 245 The Ozark/Ouachita-Appalachian forest and Ozark Highlands were also identified as 246 having region-specific temperature effects (Figure 6). Twelve month total precipitation 247 also had region specific effects in the Mississippi Alluvial and Southeast Coastal Plains 248 ecoregion, where it was associated with lower expected fire counts (Figure 6). In contrast, 249 increasing cumulative twelve month precipitation was associated with higher counts in 250 desert ecoregions (Figure 5). Housing density showed a unimodal relationship to expected 251 count (Figure 5), with lower expected counts in sparsely populated ecoregions, and higher 252 expected counts with moderately populated ecoregions. 253

Posterior 95% credible interval coverage for the number of fires over 405 hectares in the test set was 98.8%. The lowest test set interval coverage was 89.3%, in the Cross Timbers L3 ecoregion. When observed counts fell outside the 95% prediction interval, counts were larger than predicted 100% of the time. The largest difference between observed numbers and predicted 97.5% posterior quantiles (the upper limit for the 95% credible interval) occurred for the Columbia Mountains/Northern Rockies L3 ecoregion in August 2015, when 36 fires over 405 hectares occurred and at most 22 were predicted. For nearly half of the level 3 ecoregions (43 of 85), accounting for 39.7% of the land area of the contiguous U.S., the zero-inflated negative binomial model had 100% test set prediction interval coverage.

### <sup>263</sup> Wildfire burned areas

The lognormal distribution performed best on the test set (Table 2), and captured tail-264 behavior better than other burned area distributions (Figure 7). The GPD model was too 265 heavy-tailed to adequately capture the pattern in the empirical data, predicting fires far 266 larger than those observed in the training and test sets (Figure 7). The tapered Pareto 267 distribution was too light-tailed (Figure 7). The gamma and Weibull models performed 268 very poorly overall on the test set (Table 2), apparently due to a lack of congruence be-269 tween the shapes of these distributions and the actual burned area distribution. Despite 270 a poor fit to the bulk of the wildfire burned area distribution, both performed adequately 271 in the upper tails (Figure 7). Hereafter we present results for the lognormal model, which 272 had the highest test set log likelihood and captured tail behavior of the empirical fire size 273 distribution. 274

Relative humidity was the primary driver of expected burned area for a fire event (Figure 275 8A). The first basis vector for mean daily minimum relative humidity was the only coeffi-276 cient with a 95% credible interval that did not include zero (posterior median: 1.68, 95%) 277 CI: (0.8 - 2.29)). This nonlinear effect can be observed in Figure 8B as an increase in the 278 expected burned area below 20% mean daily minimum humidity. This leads to a season-279 ality gradient among ecoregions of expected fire sizes, with little or no seasonal signal in 280 typically humid ecoregions such as Marine West Coast Forests of the Pacific Northwest, 281 and seasonal oscillations in ecoregions that have periodic fluctuations between dry and hu-282 mid conditions such as the Temperate Sierras (Figure 8C). There was not strong evidence 283 that meteorological variables had spatially variable effects on expected wildfire burned 284 area. 285

Overall, 95% posterior predictive interval coverage in the test set for burned areas was 286 93%. The lowest test set coverage was 0%, for the Eastern Great Lakes Lowlands L3 ecore-287 gion, followed by 50%, for the Central California Valley L3 ecoregion, though these ecore-288 gions had just 1 and 2 wildfire events in the test set. When observed fire sizes fell outside 280 the 95% prediction interval, 24.9% of wildfires were smaller than predicted, and 75.1% of 290 wildfires were larger than predicted. The largest discrepancy between the actual size of a 291 wildfire and the predicted 97.5% posterior quantile was observed with the Wallow Fire in 292 2011 which burned 228,107 hectares, but the predicted upper limit for size was 20,756. We 293 investigate this discrepancy further in the case study below. The lognormal burned area 294 model achieved 100% interval coverage in 24 of 67 ecoregions that had wildfire events in 295 the test set, accounting for 26% of the land area of the contiguous U.S. 296

#### <sup>297</sup> Inference on extremes

By combining the output of the event count and burned area models, we derived pos-298 terior prediction intervals for the size of the largest fire in a month for each region (the 299 "burned area maximum"), integrating over uncertainty in the number of fires, as well as 300 the lognormal mean and standard deviation for burned area. We evaluated the posterior 301 distribution for the quantile function of the finite sample maximum of a lognormal distri-302 bution  $(\exp(\mu + \sigma\sqrt{2}\mathrm{erf}^{-1}(2P^{1/n} - 1)))$ , where n is the number of wildfire events,  $\mathrm{erf}^{-1}$  is 303 the inverse error function, and P is a probability) to generate prediction intervals for maxi-304 mum fire sizes by month and ecoregion, conditional on one or more fires having occurred. 305 In the holdout period from 2010 to 2016, a 99% prediction interval achieved 77.4% interval 306 coverage, with 14.8% of the burned area maxima (140 fire events) being larger than pre-307 dicted (Figure 9). As an additional check, we used the posterior distribution of the model 308 to predict the total area burned by wildfires in the test set. The model predicted the total 309 area burned over the entire contiguous United States in test period from 2010 to 2016 to 310 be 30,339,123 (95% CI: (20,496,551 - 50,446,932) and the actual value was 30,440,173. 311

<sup>312</sup> While fires over a million acres ( $\approx 404,686$  hectares) in size have happened historically in <sup>313</sup> the contiguous U.S. (Pernin 1971), no such fires were represented in in the training or test

sets. If we extrapolate, the probability of at least one fire this large in the period from 314 2010 to 2016 was estimated to be between 0.191 and 0.651 (95% CI), with a posterior 315 median of 0.348. The highest probability for such an event was 0.014 (posterior median), 316 with a 95% CI of (0, 0.237) seen for the Southwestern Tablelands ecoregion in June 2011. 317 The second highest probability was 0.004 (posterior median), with a 95% CI of (0, 0.056)318 seen for the Arizona/New Mexico Mountains ecoregion in June 2011. Aggregating spa-319 tially, we estimated monthly probabilities of a million acre wildfire. These probabilities 320 show seasonal signals corresponding to peak fire seasons, with a shift toward higher and 321 broader peaks beginning in the 21st century (Figure 10). 322

### <sup>323</sup> Error analysis case study: the 2011 Wallow Fire

To better understand how well the model could or could not anticipate notable extreme 324 events, and why, we used the largest fire in the test set as a case study. The Wallow Fire 325 was accidentally ignited on May 29, 2011 by two campers in the L3 Arizona/New Mex-326 ico Mountains ecoregion. It burned through the month of June and into early July. The 327 model underpredicted the total burned area of the Wallow Fire. Integrating over uncer-328 tainty in the predicted number of fires and expected fire size, the 99% credible interval for 329 the maximum fire size for May 2011 was (730 - 107,419) hectares, but the Wallow Fire is 330 recorded as 228,103 hectares. 331

We evaluated the contribution of each covariate to the linear predictor functions of the 332 three model components (lognormal mean for burned areas, negative binomial mean for 333 counts, and the logit probability of the zero-inflation component) to understand why these 334 predictions differed. We defined the contribution of a variable as the dot product of the 335 elements in the design matrix X corresponding to a particular driver variable (e.g., humid-336 ity), and the estimated coefficients in  $\beta$  corresponding to that variable. This provides a 337 quantitative measure of how each input variable contributes to the linear predictor for an 338 ecoregion, and incorporates the overall, level 1, level 2, and level 3 ecoregion adjustments 339 on these effects. Humidity is the primary driver of variation in the model's predictions 340 overall, and June 2011 - the month after ignition - favored more large fires, with drier, hot-341

ter conditions (Figure 11). The 99% credible interval for June 2011 was (4,258 - 428,765) hectares, which contains the true value. Had the Wallow Fire ignited two days later, the true final size would have been contained in the prediction interval. Evidently, conditions in May that drove (under)predictions of maximum burned area were not representative of the conditions over most of the Wallow Fire's duration.

Temporal mismatch aside, meteorological conditions local to the Wallow Fire differed 347 from the monthly regional means (Figure 12). In particular, wind speeds in the Wallow 348 Fire vicinity exceeded the regional monthly mean values on the date of ignition and in 340 the weeks following ignition. Over the majority of the duration of the Wallow Fire (May 350 29 to July 8), local daily conditions were drier and hotter on average than regional mean 351 monthly conditions in May, which were used to drive the statistical model. This local vari-352 ability is not represented in the regional models developed here. The failure of the model 353 to correctly predict the size of the Wallow fire suggests potential avenues for improvement, 354 discussed below. 355

# 356 Discussion

Extreme wildfires are often devastating, but perhaps they need not be surprising. By allowing the non-linear effects of weather and housing density to vary across space, this model achieves good predictive accuracy for fire extremes at a regional scale over a six year prediction window. This model predicts that extremely large wildfires, perhaps even over one million acres (404,686 hectares), have a non-negligible probability of occurrence in the contiguous United States. Such predictions can support regional wildfire management and probabilistic hazard assessment.

Driving a model with meteorological features raises challenges related to predictive uncertainty and covariate shift - a change in the underlying distribution of forcing variables, potentially outside of the historic range. Ideally, this uncertainty would be propagated forward in a predictive model, possibly through stacking of predictive distributions that are generated from multiple models of future climate dynamics (Yao et al. 2017). But, even if one had a perfect forecast, novel conditions present a challenge for predictive modeling (Quionero-Candela et al. 2009). For example, the High Plains ecoregion had its highest mean monthly precipitation, lowest 12 month running precipitation, driest, hottest, and windiest conditions in the test set period. Extrapolating beyond the range of training inputs is generally difficult, but the hierarchical spatial effect specification used here allows partial pooling among climatically similar ecoregions that can inform such predictions, unlike models fit separately to disjoint spatial regions.

Similar issues could arise when making predictions for observed but rare meteorologi-376 cal conditions. For example, mean daily minimum humidity values over 60% accounted 377 for just 3.76% of the ecoregion-months in the training data, and 0 fires occurred in such 378 months. As a consequence, there is relatively little data that can be used to inform the 379 model for such conditions, and the prior distribution which shrinks coefficients toward 380 zero may dominate the likelihood in the posterior distribution. In this case, the poste-381 rior distribution for the last basis coefficient for the partial effect of humidity is likely to 382 be close to zero. This could explain why the estimated partial effect of humidity on the 383 expected counts was less negative at the upper end of the observed humidity range, al-384 though previous work has found similarly nonlinear partial effects (Preisler et al. 2004). 385 The count model performed extremely well in this range, with 100% interval coverage for 386 the 299 ecoregion-months with mean daily minimum humidity values greater than 60%387 in the withheld test data. The model nearly always predicted zero counts with high confi-388 dence when conditions were this humid: 298 of 299 predictions made for such conditions 389 were 95% credible intervals of (0, 0). The remaining prediction had a posterior median of 390 zero, along with a 95% credible interval from 0 to 1. Monotonicity constraints could be 391 incorporated into these models via monotonic spline bases (Ramsay and others 1988), or 392 an ordered prior distribution for basis coefficients (Brezger and Steiner 2008). In this case, 393 the count model performs well under humid conditions without monotonicity constraints, 394 and there seems to be little room for performance improvements that might result from 395 such constraints. 396

<sup>397</sup> Human-caused climate change is expected to increase fire activity in the western U.S.

(Rogers et al. 2011; Westerling et al. 2011; Moritz et al. 2012; Abatzoglou and Williams 398 2016) and elsewhere (Flannigan et al. 2009), but the nonlinear effect of housing density 390 could provide additional insight into future expectations. While housing density is increas-400 ing over time in most U.S. ecoregions, some of these ecoregions are in the range of values 401 in which this increases the expected number of large fires, while others are so populated 402 that further increases would reduce the chance of a large fire. The hump-shaped effect of 403 human density on the expected number of large fires is likely driven by ignition pressure 404 and fire suppression (Balch et al. 2017). As human density increases from zero, ignition 405 pressure increases, but eventually landscapes become so urbanized, fragmented, and/or 406 fire-suppressed that wildfire risk decreases (Syphard et al. 2007; Bowman et al. 2011; 407 Bistinas et al. 2013; Knorr et al. 2013; Mcwethy et al. 2013; Syphard et al. 2017; Nagy et 408 al. 2018). At intermediate density, wildfire regimes respond to human ignition and altered 409 fuel distributions (Guyette, Muzika, and Dey 2002), but these responses depend on envi-410 ronmental context and characteristics of the human population (Marlon et al. 2008; Li et 411 al. 2009). This model indicates that the combination of moderate to high human density 412 and dry conditions would nonlinearly increase the chance of an extreme fire event. Both 413 human density and dryness are expected to increase in the future across large swaths of 414 the U.S. (Lloyd, Sorichetta, and Tatem 2017; Stavros et al. 2014, Radeloff et al. (2010)), 415 with potential implications for human mortality, health risks from smoke and particulate 416 emission, and the financial burden of wildfire management (Reid et al. 2016; Radeloff et al. 417 2018). 418

This work points to promising directions for future predictive efforts. Default choices such 419 as Poisson and GPD distributions should be checked against alternative distributions. 420 Further, the predictive skill of this model seems to suggest that ordinary events provide 421 information on extremes, which would not be the case if the generative distribution of 422 extremes was completely unique. Previous case studies have identified that extremes or 423 anomalies in climatological drivers play a role in the evolution of extreme wildfires (Peter-424 son et al. 2015), but for this work, monthly averages of climatological drivers over fairly 425 large spatial regions were used, which may smooth over anomalous or extreme conditions. 426 Enhancing the spatiotemporal resolution of predictive models could better represent cli-427

matic and social drivers and provide localized insights to inform decision-making. This
raises computational challenges, but recent advances in distributed probabilistic computing
(Tran et al. 2017), efficient construction of spatiotemporal point processes (Shirota and
Banerjee 2018), and compact representations of nonlinear spatial interactions (Lee and
Durbán 2011) may provide solutions.

The Wallow Fire case study reveals at least one limitation of increasing the spatiotem-433 poral resolution. When the model predictions are driven by covariates that are summa-434 rized in space and time (e.g. a mean across an ecoregion in a month), summary values 435 may not represent conditions that are most relevant to an event. With a discrete space-436 time segmentation, events can occur at the boundary of a spatiotemporal unit, e.g., if 437 a fire spreads into an adjacent ecoregion or ignites on the last day of the month. Large 438 wildfires can span months, and a model that only uses conditions upon ignition to pre-439 dict total burned area can fail to account for conditions that change over the course of 440 the event. Modeling ignitions as a point process in continuous space and time (Brillinger, 441 Preisler, and Benoit 2003), and explicitly modeling subsequent fire duration and spread 442 could better separate conditions that ignite fires from those that affect propagation. Such 443 an approach might be amenable to including information on fuel continuity, which is likely 444 to limit the size of extremely large fires and did not factor into the current predictions 445 (Rollins, Morgan, and Swetnam 2002; Hargrove et al. 2000). 446

To the extent that a model reflects the generative process for extreme events, the decom-447 position of contributions to the model's predictions may provide insight into attribution 448 for meteorological and anthropogenic drivers of extremes. However, a model trained to 449 represent a region-wide distribution of fire sizes will inevitably fail to capture local factors 450 that are relevant to specific events such as the Wallow Fire. If predicting the dynamics 451 of particular fire events is a goal, process-based models designed to model fire spread are 452 likely to be more appropriate than regional statistical models such as those developed 453 here. 454

This paper presents and evaluates a statistical approach to explain and predict extreme wildfires that incorporates spatially varying non-linear effects. The model reveals considerable differences among ecoregions spanning the mountain west to the great plains, deserts,
and eastern forests, and suggests a non-negligible chance of extreme wildfires larger than
those seen in over the past 30 years in the contiguous U.S. Predictive approaches such as
this can inform decision-making by placing probabilistic bounds on the number of wildfires
and their sizes, while provide deeper insights into wildfire ecology.

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# $_{780}$ Tables

Model	Holdout log likelihood
ZI Negative binomial	-3671 (70)
ZI Poisson	-4093 (77)
Negative binomial	-4298 (114)
Poisson	-4572 (139)

 Table 1. Performance of count models on the test set in descending order. Posterior

 means are provided with standard deviations in parentheses.

Model	Holdout log likelihood
Lognormal	-26341 (43)
Generalized Pareto	-26377 (45)
Tapered Pareto	-26386 (49)
Weibull	-27592 (236)
Gamma	-30675 (993)

 Table 2. Performance of burned area models on the test set in descending order. Posterior means are provided with standard deviations in parentheses.

# 781 Figures

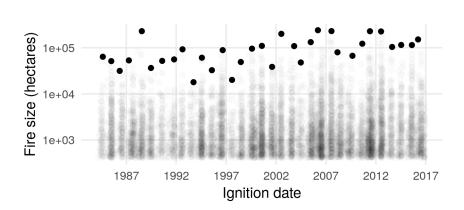


Figure 1. Sizes of wildfires over 405 hectares in the contiguous United States, from the Monitoring Trends in Burn Severity multiagency program. Each point represents a fire event, and the largest fires for each year (the block maxima) are shown as solid black points.

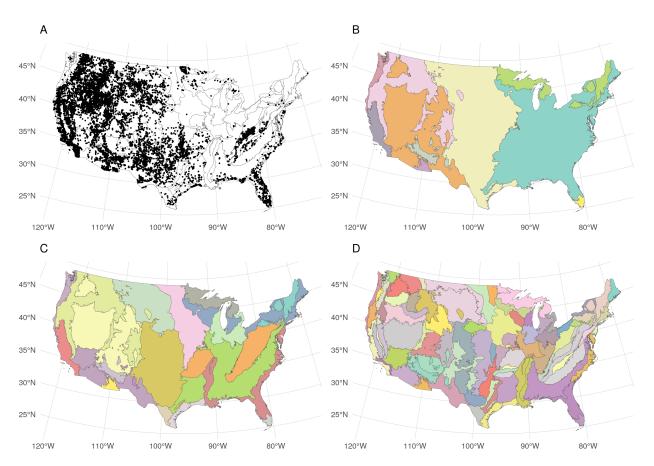
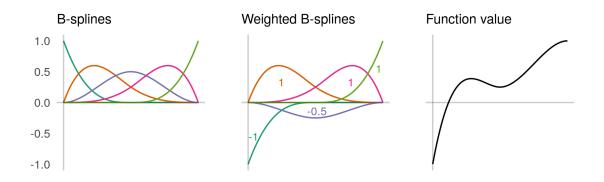
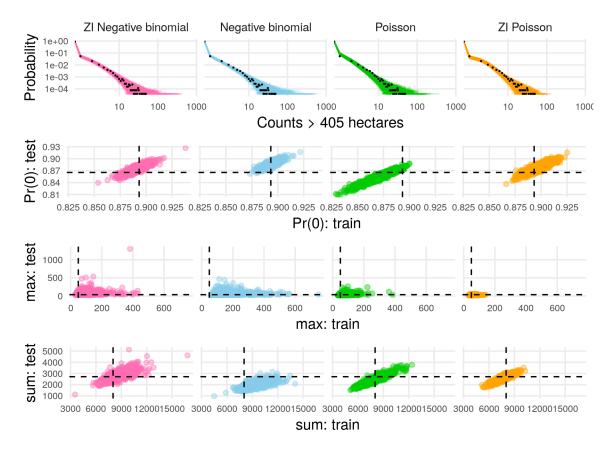


Figure 2. A. Large wildfire ignition locations are shown as points across the study region. Colors in panels B, C, and D show level 1, 2, and 3 ecoregions respectively.



**Figure 3.** Conceptual figure to illustrate the use of B-splines to construct nonlinear functions. In the left panel, five B-spline vectors are shown, which map values of an input variable (on the x-axis) to a value on the y-axis. The middle panel shows the same B-spline vectors, but weighted (multiplied) by real numbers, with the weights illustrated as annotations. These weighted B-spline vectors are summed to produce the values of a nonlinear function (right panel).



**Figure 4.** Count predictive checks. Row one shows observed count frequencies as black points and predicted frequencies as lines. Rows two, three, and four show predicted proportions of zeros, maxima, and sums (respectively) in the training and test data, with empirical values as dashed lines. Rows two through four facilitate comparison of performance on training and test sets. Ideally, model predictions cluster around the dashed lines for both the training (x-axis direction) and test (y-axis direction) sets, leading to a tight cluster of points at the intersection of the dashed lines.

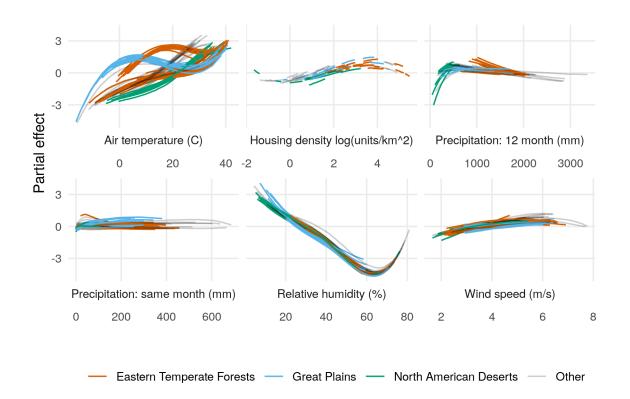


Figure 5. Partial effects on the log-transformed negative binomial mean component of the zero-inflated negative binomial model for each level 3 ecoregion, colored by level 1 ecoregion. Lines are posterior medians. Results are similar for the zero-inflation component.

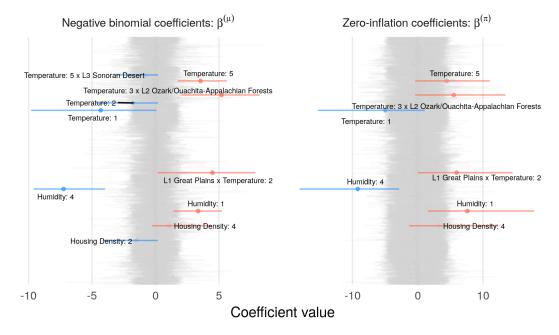
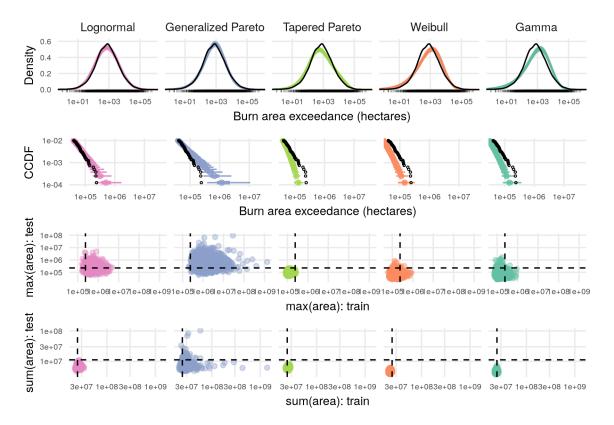


Figure 6. Caterpillar plots of zero inflated negative binomial model coefficients,  $\beta^{(\mu)}$  (left) and  $\beta^{(\pi)}$  (right). Horizontal line segments denote 95% credible intervals. Grey segments indicate coefficients with a less than 87% posterior probability of being positive or negative, and colored segments indicate coefficients that are probably positive (red) or negative (blue). B-spline vectors are indicated by colons, e.g., Humidity:1 indicates the first basis vector corresponding to humidity. Interactions between variables a and b are represented as a x b. Level 1 ecoregions are represented by L1 ecoregion name, and L2 and L3 indicate level 2 and 3 ecoregions.



**Figure 7.** Predictive checks for burned area models. The top row shows predicted density in color and empirical density for the training set in black, which reveals overall lack of fit for the gamma and Weibull models. Row two shows the complementary cumulative distribution function (CCDF) at the tails, with 95% and 50% prediction intervals shown in color and observed data as black points, which shows that the Generalized Pareto distribution predicts values that are too extreme. The third and fourth rows show checks for maximum and total burned areas in the training and test set, with observed values as dashed lines and posterior draws as colored points. These final two rows facilitate checks for summary statistics on both the training and test set, with the ideal model generating predictions (colored points) clustered close to where the dashed lines intersect.

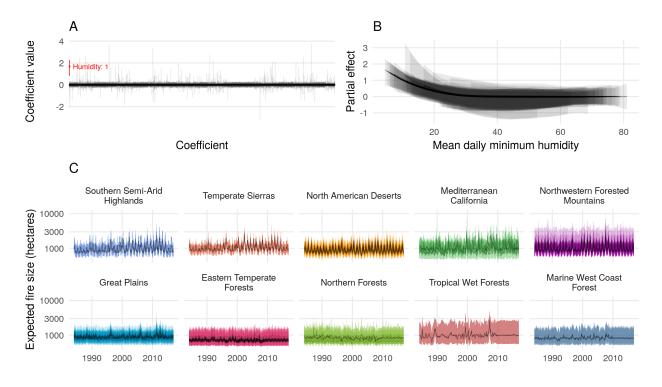
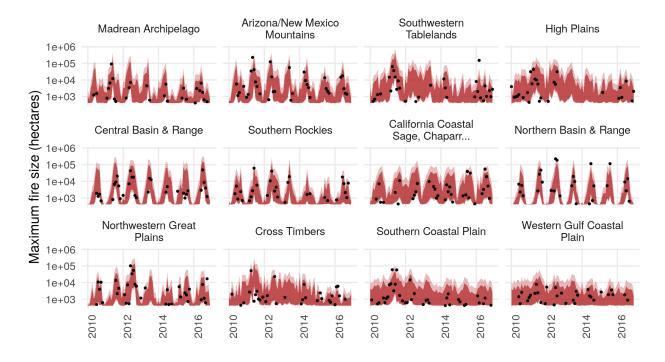


Figure 8. A. Estimated posterior medians and 95% credible intervals for each of the 3,473 coefficients associated with expected burned area. Only one coefficient - the first basis vector for humidity - had a 95% credible interval that excluded zero, shown in red. This effect is visualized in **B**. Partial effects of mean daily minimum humidity for each level 3 ecoregion, with posterior medians drawn as lines, and the 95% credible intervals as ribbons. **C**. Monthly time series of expected fire sizes for every level 3 ecoregion, faceted and colored by level 1 ecoregions sorted by mean humidity. Lines are posterior medians and ribbons are 95% credible intervals.



**Figure 9.** Posterior 99% (light red) and 95% (dark red) prediction intervals for the burned area of the largest fire event by month and level 3 ecoregion in the test set, shown for ecoregions with wildfires in more than 20 months. Empirical maxima are shown as black dots.

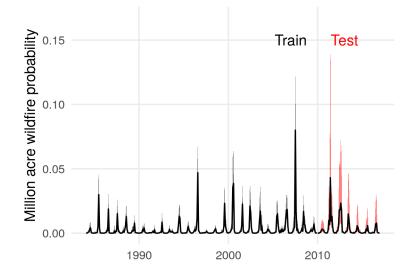


Figure 10. Estimated monthly posterior probabilities that one or more fire events exceed one million acres (404,686 hectares). The line represents the posterior median, and shaded region represents an 80% credible interval. The training period up to 2010 is shown in black, and the test period for which data were withheld during parameter estimation is shown in red.

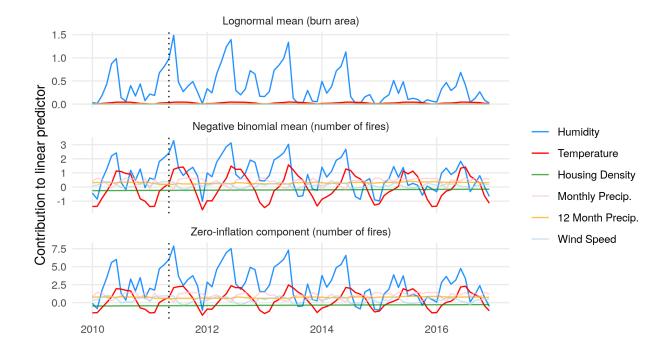
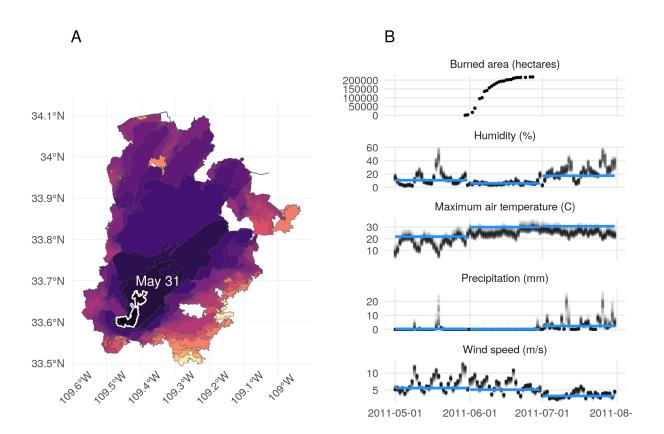


Figure 11. Posterior median contribution of each input variable to the linear predictor function of model components for the Arizona/New Mexico Mountains level 3 ecoregion from 2010-2016. A dotted vertical line marks May 2011, when the Wallow Fire ignited. Vertical positions of colored lines show contributions to the linear predictor function of each model component.



**Figure 12.** A. Progression of the Wallow Fire based on perimeter data collected from the GeoMAC database, spanning May 30, 2011 to June 27, 2011. The perimeter at the end of may is outlined in white, with brighter colors indicating later dates. B. Daily local meteorological conditions and monthly regional mean conditions for the Wallow Fire, along with the associated burned area over time. The blue line represents monthly averages of meteorological quantities computed over the entire Arizona/New Mexico Mountains ecoregion, and black points represent values extracted for "local" 4 km grid cells contained within the final burned area perimeter of the Wallow Fire.

# 782 Appendices

#### 783 **Prior specifications**

Prior distributions were chosen to regularize coefficients on the distribution specific means  $\beta^{(\mu)}$  and structural zero parameters  $\beta^{(\pi)}$ . We used a regularized horseshoe prior on these coefficients, which shrinks irrelevant coefficients towards zero, while regularizing nonzero coefficients (Piironen, Vehtari, and others 2017). For zero-inflated models, we used a multivariate version of the regularized horseshoe (Peltola et al. 2014):

$$\begin{pmatrix} \beta_j^{(\mu)} \\ \beta_j^{(\pi)} \end{pmatrix} \sim \mathcal{N} \left( \mathbf{0}, \begin{vmatrix} \tau_1^2 \tilde{\lambda}_{1,j}^2 & \rho \tau_1 \tau_2 \tilde{\lambda}_{1,j} \tilde{\lambda}_{2,j} \\ \rho \tau_1 \tau_2 \tilde{\lambda}_{1,j} \tilde{\lambda}_{2,j} & \tau_2^2 \tilde{\lambda}_{2,j}^2 \end{vmatrix} \right) \right),$$
$$\tilde{\lambda}_{m,j}^2 = \frac{c_m^2 \lambda_j^2}{c_m^2 + \tau_m^2 \lambda_j^2},$$

for each response dimension m = 1, 2 and coefficient j = 1, ..., p. Here  $\rho$  is a correlation 789 parameter,  $\tau_1$  and  $\tau_2$  are global variance hyperparameters,  $c_1$  and  $c_2$  are hyperparameters 790 that determine the amount of shrinkage on the largest coefficients, and  $\lambda_i$  is a local scale 791 parameter drawn from a half-Cauchy distribution that control the amount of shrinkage 792 applied to coefficient j (Piironen, Vehtari, and others 2017). With this prior specification, 793 information can be shared across the two response dimensions through the correlation pa-794 rameter  $\rho$ , and/or through the local scale parameters  $\lambda_i$ . For count models without struc-795 tural zeros (the Poisson and negative binomial models), this multivariate prior simplifies to 796 a univariate regularized horseshoe prior. 797

Spatiotemporal random effects were constructed using a temporally autoregressive, spatially intrinsically autoregressive formulation (Besag and Kooperberg 1995; Banerjee, Carlin, and Gelfand 2014). Temporarily suppressing the superscript that indicates whether the effects are on  $\mu$  or  $\pi$ , and denoting column t from an  $S \times T \Phi$  as  $\phi_t$  we have:

$$\boldsymbol{\phi}_{t=1} \sim \mathrm{N}(\mathbf{0}, (\tau^{(\phi)}(\mathbf{D} - \mathbf{W}))^{-1})$$

$$\boldsymbol{\phi}_t \sim \mathcal{N}(\eta \boldsymbol{\phi}_{t-1}, (\tau^{(\phi)}(\mathbf{D} - \mathbf{W}))^{-1}), \quad t = 2, ..., T$$

where  $\eta$  is a temporal dependence parameter,  $\tau^{(\phi)}$  is a precision parameter, **D** is an  $S \times S$ diagonal matrix with entries corresponding to the number of spatial neighbors for each spatial unit, and **W** is an  $S \times S$  spatial adjacency matrix with nonzero elements only when spatial unit *i* is a neighbor of spatial unit *j* ( $w_{i,j} = 1$  if *i* is a neighbor of *j*, and  $w_{i,j} = 0$ otherwise, including  $w_{i,i} = 0$  for all *i*).  $\tau^{(\phi)}$  is a precision parameter. We imposed a soft identifiability constraint that places high prior mass near  $\sum_{s=1}^{S} \phi_{t,s}^* = 0$  for all *t*.

We applied a univariate regularized horseshoe prior to all  $\beta$  coefficients in burned area models (Piironen, Vehtari, and others 2017):

$$\beta_j \sim \mathcal{N}\left(0, \tau^2 \tilde{\lambda}_j^2\right), \quad \tilde{\lambda}_j^2 = \frac{c^2 \lambda_j^2}{c^2 + \tau^2 \lambda_j^2};$$

<sup>\$10</sup> Spatiotemporal random effects were constructed in the same way as for the count models.

# **Joint distributions**

Here we provide the unnormalized posterior densities for each model. Square brackets represent a probability mass or density function. Parameterizations for model likelihoods are provided first, followed by the factorization of the joint distribution, with explicit priors.

#### <sup>815</sup> Poisson wildfire count model

<sup>816</sup> We used the following parameterization of the Poisson distribution:

$$[n|\mu] = \frac{\mu^n e^{-\mu}}{n!},$$

<sup>817</sup> where  $\mu$  is the mean and variance.

<sup>\$18</sup> The unnormalized posterior density of this model is:

$$[\boldsymbol{\beta}^{(\mu)}, \boldsymbol{\alpha}^{(\mu)}, \boldsymbol{\phi}, \boldsymbol{\sigma}^{(\phi)}, \boldsymbol{\eta}, \boldsymbol{\lambda}, \boldsymbol{c}, \boldsymbol{\tau} \mid \mathbf{N}] \propto \prod_{s=1}^{S} \prod_{t=1}^{T} [n_{s,t} | \boldsymbol{\beta}^{(\mu)}, \boldsymbol{\alpha}^{(\mu)}, \boldsymbol{\phi}_{s,t}] \times [\boldsymbol{\phi}_{1} | \boldsymbol{\sigma}^{(\phi)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_{t} | \boldsymbol{\phi}_{t-1}, \boldsymbol{\sigma}^{(\phi)}, \boldsymbol{\eta}] \times \prod_{j=1}^{p} [\beta_{j}^{(\mu)} | \lambda_{j}, \boldsymbol{c}, \boldsymbol{\tau}] [\lambda_{j}] \times [\boldsymbol{\sigma}^{(\phi)}] [\boldsymbol{\eta}] [\boldsymbol{c}] [\boldsymbol{\tau}] [\boldsymbol{\alpha}^{(\mu)}]$$

$$= \prod_{s=1}^{S} \prod_{t=1}^{T} \operatorname{Poisson}(n_{s,t} | \exp(\alpha^{(\mu)} + \mathbf{X}_{(s,t)} \boldsymbol{\beta}^{(\mu)} + \phi_{s,t})) \times \operatorname{Normal}(\boldsymbol{\phi}_{1} | \mathbf{0}, ((\sigma^{(\phi)})^{-2} (\mathbf{D} - \mathbf{W}))^{-1}) \times \prod_{t=2}^{T} \operatorname{Normal}(\boldsymbol{\phi}_{t} | \eta \boldsymbol{\phi}_{t-1}, ((\sigma^{(\phi)})^{-2} (\mathbf{D} - \mathbf{W}))^{-1}) \times \prod_{j=1}^{p} \operatorname{Normal}\left(\beta_{j}^{(\mu)} | \mathbf{0}, \frac{\tau^{2} c^{2} \lambda_{j}^{2}}{c^{2} + \tau^{2} \lambda_{j}^{2}}\right) \times \operatorname{Cauchy}^{+}(\lambda_{j} | \mathbf{0}, 1) \times (\boldsymbol{\phi}^{(\phi)} | \mathbf{0}, 1^{2}) \approx \operatorname{Dete}(|\mathbf{1}, \mathbf{1}) \approx \operatorname{Lec}(|\mathbf{0}, \mathbf{0}| + 1) = 0$$

Normal<sup>+</sup>( $\sigma^{(\phi)}|0,1^2$ ) × Beta( $\eta|1,1$ ) × Inv-Gamma( $c^2|2.5,10$ ) ×

Normal<sup>+</sup> $(\tau|0, 5^2) \times \text{Normal}(\alpha^{(\mu)}|0, 5^2).$ 

#### <sup>819</sup> Negative binomial wildfire count model

<sup>820</sup> We used the following parameterization of the negative binomial distribution:

$$[n|\mu,\delta] = \binom{n+\delta-1}{n} \left(\frac{\mu}{\mu+\delta}\right)^n \left(\frac{\delta}{\mu+\delta}\right)^{\delta},$$

- where  $\mu$  is the mean, and  $\delta$  is a dispersion parameter.
- <sup>822</sup> The unnormalized posterior density of this model is:

$$[\boldsymbol{\beta}^{(\mu)}, \boldsymbol{\alpha}^{(\mu)}, \boldsymbol{\phi}, \boldsymbol{\sigma}^{(\phi)}, \boldsymbol{\eta}, \boldsymbol{\lambda}, c, \tau, \delta \mid \mathbf{N}] \propto \prod_{s=1}^{S} \prod_{t=1}^{T} [n_{s,t} | \boldsymbol{\beta}^{(\mu)}, \boldsymbol{\alpha}^{(\mu)}, \phi_{s,t}, \delta] \times [\boldsymbol{\phi}_{1} | \boldsymbol{\sigma}^{(\phi)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_{t} | \boldsymbol{\phi}_{t-1}, \boldsymbol{\sigma}^{(\phi)}, \boldsymbol{\eta}] \times \prod_{j=1}^{p} [\beta_{j}^{(\mu)} | \lambda_{j}, c, \tau] [\lambda_{j}] \times [\boldsymbol{\sigma}^{(\phi)}] [\boldsymbol{\eta}] [c] [\tau] [\boldsymbol{\alpha}^{(\mu)}] [\delta]$$

$$= \prod_{s=1}^{S} \prod_{t=1}^{T} \text{Negative Binomial}(n_{s,t} | \exp(\alpha^{(\mu)} + \mathbf{X}_{(s,t)} \boldsymbol{\beta}^{(\mu)} + \phi_{s,t}), \delta) \times \\ \text{Normal}(\phi_1 | \mathbf{0}, ((\sigma^{(\phi)})^{-2} (\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{t=2}^{T} \text{Normal}(\phi_t | \eta \phi_{t-1}, ((\sigma^{(\phi)})^{-2} (\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{j=1}^{p} \text{Normal}\left(\beta_j^{(\mu)} | 0, \frac{\tau^2 c^2 \lambda_j^2}{c^2 + \tau^2 \lambda_j^2}\right) \times \text{Cauchy}^+(\lambda_j | 0, 1) \times \\ \text{Normal}^+(\sigma^{(\phi)} | 0, 1^2) \times \text{Beta}(\eta | 1, 1) \times \text{Inv-Gamma}(c^2 | 2.5, 10) \times \\ \text{Normal}^+(\tau | 0, 5^2) \times \text{Normal}(\alpha^{(\mu)} | 0, 5^2) \times \text{Normal}^+(\delta | 0, 5^2).$$

#### <sup>823</sup> Zero-inflated Poisson wildfire count model

<sup>824</sup> We used the following parameterization of the zero-inflated Poisson distribution:

$$[n|\mu,\pi] = I_{n=0}(1-\pi+\pi e^{-\mu}) + I_{n>0}\pi\frac{\mu^n e^{-\mu}}{n!},$$

- where  $\mu$  is the Poisson mean, and  $1 \pi$  is the probability of an extra zero.
- <sup>826</sup> The unnormalized posterior density of this model is:

$$\begin{split} [\boldsymbol{\beta}^{(\mu)}, \alpha^{(\mu)}, \boldsymbol{\beta}^{(\pi)}, \alpha^{(\pi)}, \boldsymbol{\phi}^{(\mu)}, \sigma^{(\phi,\mu)}, \eta^{(\mu)}, \boldsymbol{\phi}^{(\pi)}, \sigma^{(\phi,\pi)}, \eta^{(\pi)}, \boldsymbol{\lambda}, c, \tau, \rho \mid \mathbf{N}] \propto \\ \prod_{s=1}^{S} \prod_{t=1}^{T} [n_{s,t} | \boldsymbol{\beta}^{(\mu)}, \alpha^{(\mu)}, \boldsymbol{\beta}^{(\pi)}, \alpha^{(\pi)}, \boldsymbol{\phi}^{(\mu)}_{s,t}, \boldsymbol{\phi}^{(\pi)}_{s,t}] \times \\ [\boldsymbol{\phi}_{1}^{(\mu)} | \sigma^{(\phi,\mu)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_{t}^{(\mu)} | \boldsymbol{\phi}^{(\mu)}_{t-1}, \sigma^{(\phi,\mu)}, \eta^{(\mu)}] \times \\ [\boldsymbol{\phi}_{1}^{(\pi)} | \sigma^{(\phi,\pi)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_{t}^{(\pi)} | \boldsymbol{\phi}^{(\pi)}_{t-1}, \sigma^{(\phi,\pi)}, \eta^{(\pi)}] \times \\ \prod_{j=1}^{p} [\beta_{j}^{(\mu)}, \beta_{j}^{(\pi)} | \lambda_{j}, c, \tau, \rho] [\lambda_{j}] \times \\ [\sigma^{(\phi,\mu)}] [\sigma^{(\phi,\pi)}] [\eta^{(\mu)}] [\eta^{(\pi)}] [\alpha^{(\mu)}] [\alpha^{(\pi)}] [\rho] \prod_{m=1}^{2} [c_{m}] [\tau_{m}] \end{split}$$

#### <sup>827</sup> Zero-inflated negative binomial wildfire count model

<sup>828</sup> We used the following parameterization of the zero-inflated negative binomial distribution:

$$[n|\mu,\delta,\pi] = I_{n=0}(1-\pi+\pi\left(\frac{\delta}{\mu+\delta}\right)^{\delta}) + I_{n>0}\binom{n+\delta-1}{n}\left(\frac{\mu}{\mu+\delta}\right)^n\left(\frac{\delta}{\mu+\delta}\right)^{\delta},$$

where  $\mu$  is the negative binomial mean,  $\delta$  is the negative binomial dispersion, and , and 1 -  $\pi$  is the probability of an extra zero.

<sup>831</sup> The unnormalized posterior density of this model is:

$$\begin{split} [\boldsymbol{\beta}^{(\mu)}, \alpha^{(\mu)}, \boldsymbol{\beta}^{(\pi)}, \alpha^{(\pi)}, \boldsymbol{\phi}^{(\mu)}, \sigma^{(\phi,\mu)}, \eta^{(\mu)}, \boldsymbol{\phi}^{(\pi)}, \sigma^{(\phi,\pi)}, \eta^{(\pi)}, \boldsymbol{\lambda}, c, \tau, \rho, \delta \mid \mathbf{N}] \propto \\ & \prod_{s=1}^{S} \prod_{t=1}^{T} [n_{s,t} | \boldsymbol{\beta}^{(\mu)}, \alpha^{(\mu)}, \boldsymbol{\beta}^{(\pi)}, \alpha^{(\pi)}, \phi^{(\mu)}_{s,t}, \phi^{(\pi)}_{s,t}, \delta] \times \\ & [\boldsymbol{\phi}_{1}^{(\mu)} | \sigma^{(\phi,\mu)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_{t}^{(\mu)} | \boldsymbol{\phi}_{t-1}^{(\mu)}, \sigma^{(\phi,\mu)}, \eta^{(\mu)}] \times \\ & [\boldsymbol{\phi}_{1}^{(\pi)} | \sigma^{(\phi,\pi)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_{t}^{(\pi)} | \boldsymbol{\phi}_{t-1}^{(\pi)}, \sigma^{(\phi,\pi)}, \eta^{(\pi)}] \times \\ & \prod_{j=1}^{p} [\beta_{j}^{(\mu)}, \beta_{j}^{(\pi)} | \lambda_{j}, c, \tau, \rho] [\lambda_{j}] \times \\ & [\sigma^{(\phi,\mu)}] [\sigma^{(\phi,\pi)}] [\eta^{(\mu)}] [\eta^{(\pi)}] [\alpha^{(\mu)}] [\alpha^{(\pi)}] [\rho] [\delta] \prod_{m=1}^{2} [c_{m}] [\tau_{m}]. \end{split}$$

## <sup>832</sup> Generalized Pareto/Lomax burned area model

<sup>833</sup> We used the following parameterization of the GPD/Lomax distribution:

$$[y|\sigma,\kappa] = \frac{1}{\sigma} \left(\frac{\kappa y}{\sigma} + 1\right)^{-(\kappa+1)\kappa^{-1}},$$

- state where  $\kappa$  is a shape parameter and  $\sigma$  is a scale parameter.
- <sup>835</sup> The unnormalized posterior density of this model is:

$$[\boldsymbol{\beta}, \alpha, \boldsymbol{\phi}, \sigma^{(\phi)}, \eta, \kappa^{(L)}, \boldsymbol{\lambda}, c, \tau \mid \boldsymbol{y}] \propto \prod_{i=1}^{n_{\text{tot}}} [y_i | \boldsymbol{\beta}, \alpha, \phi_{s_i, t_i}, \kappa^{(L)}] \times [\boldsymbol{\phi}_1 | \sigma^{(\phi)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_t | \boldsymbol{\phi}_{t-1}, \sigma^{(\phi)}, \eta] \times \prod_{j=1}^{p} [\beta_j | \lambda_j, c, \tau] [\lambda_j] \times [\alpha] [c] [\tau] [\kappa^{(L)}] [\eta] [\sigma^{(\phi)}]$$

$$= \prod_{i=1}^{n_{\text{tot}}} \operatorname{Lomax}(y_i | \kappa^{(L)}, e^{\alpha + \mathbf{X}_{(s_i, t_i)} \beta + \phi_{s_i, t_i}}) \times \\ \operatorname{Normal}(\phi_1 | \mathbf{0}, ((\sigma^{(\phi)})^{-2} (\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{t=2}^{T} \operatorname{Normal}(\phi_t | \eta \phi_{t-1}, ((\sigma^{(\phi)})^{-2} (\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{j=1}^{p} \operatorname{Normal}\left(\beta_j | 0, \frac{\tau^2 c^2 \lambda_j^2}{c^2 + \tau^2 \lambda_j^2}\right) \times \operatorname{Cauchy}^+(\lambda_j | 0, 1) \times \\ \operatorname{Normal}(\alpha | 0, 5^2) \times \operatorname{Inv-Gamma}(c^2 | 2.5, 10) \times \operatorname{Normal}^+(\tau | 0, 5^2) \\ \operatorname{Normal}^+(\kappa^{(L)} | 0, 5^2) \times \operatorname{Beta}(\eta | 1, 1) \times \operatorname{Normal}^+(\sigma^{(\phi)} | 0, 1^2). \end{cases}$$

## <sup>836</sup> Tapered Pareto burned area model

<sup>837</sup> We used the following parameterization of the tapered Pareto distribution:

$$[y|\kappa,\nu] = \left(\frac{\kappa}{y} + \frac{1}{\nu}\right) \exp(-x/\nu),$$

- ss where  $\kappa$  is a shape parameter and  $\nu$  a taper parameter.
- <sup>839</sup> The unnormalized posterior density of this model is:

$$\begin{bmatrix} \boldsymbol{\beta}, \alpha, \boldsymbol{\phi}, \sigma^{(\phi)}, \eta, \nu, \boldsymbol{\lambda}, c, \tau \mid \boldsymbol{y} \end{bmatrix} \propto \\ \prod_{i=1}^{n_{\text{tot}}} [y_i | \boldsymbol{\beta}, \alpha, \phi_{s_i, t_i}, \nu] \times \\ [\boldsymbol{\phi}_1 | \sigma^{(\phi)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_t | \boldsymbol{\phi}_{t-1}, \sigma^{(\phi)}, \eta] \times \\ \prod_{j=1}^{p} [\beta_j | \lambda_j, c, \tau] [\lambda_j] \times \\ [\alpha] [c] [\tau] [\nu] [\eta] [\sigma^{(\phi)}] \end{bmatrix}$$

$$= \prod_{i=1}^{n_{\text{tot}}} \text{Tapered Pareto}(y_i | e^{\alpha + \mathbf{X}_{(s_i, t_i)} \beta + \phi_{s_i, t_i}}, \nu) \times \text{Normal}(\phi_1 | \mathbf{0}, ((\sigma^{(\phi)})^{-2} (\mathbf{D} - \mathbf{W}))^{-1}) \times \prod_{t=2}^{T} \text{Normal}(\phi_t | \eta \phi_{t-1}, ((\sigma^{(\phi)})^{-2} (\mathbf{D} - \mathbf{W}))^{-1}) \times \prod_{j=1}^{p} \text{Normal}\left(\beta_j | \mathbf{0}, \frac{\tau^2 c^2 \lambda_j^2}{c^2 + \tau^2 \lambda_j^2}\right) \times \text{Cauchy}^+(\lambda_j | \mathbf{0}, 1) \times \text{Normal}(\alpha | \mathbf{0}, 5^2) \times \text{Inv-Gamma}(c^2 | 2.5, 10) \times \text{Normal}^+(\tau | \mathbf{0}, 5^2) \times \text{Inv-Gamma}(c^2 | 2.5, 10) \times \text{Normal}^+(\tau | \mathbf{0}, 5^2) \times \text{Normal}(\phi_j | \mathbf{0}, \tau_j^2) = \lambda_j + (\phi_j | \mathbf{0}, \tau_j^2) \times \lambda_j + (\phi_j | \mathbf{0}, \tau_j^2) \times \lambda_j = \lambda_j + (\phi_j | \mathbf{0}, \tau_j^2) \times \lambda_j + \lambda$$

Cauchy<sup>+</sup>( $\nu | 0, 1$ ) × Beta( $\eta | 1, 1$ ) × Normal<sup>+</sup>( $\sigma^{(\phi)} | 0, 1^2$ ).

## <sup>840</sup> Lognormal burned area model

<sup>841</sup> We used the following parameterization of the lognormal distribution:

$$[y|\mu,\sigma] = \frac{1}{y} \frac{1}{\sigma\sqrt{2\pi}} \exp\bigg(-\frac{(\log(y)-\mu)^2}{2\sigma^2}\bigg),$$

- where  $\mu$  and  $\sigma$  are location and scale parameters, respectively.
- <sup>843</sup> The unnormalized posterior density of this model is:

$$\begin{bmatrix} \boldsymbol{\beta}, \alpha, \boldsymbol{\phi}, \sigma^{(\phi)}, \eta, \sigma, \boldsymbol{\lambda}, c, \tau \mid \boldsymbol{y} \end{bmatrix} \propto \\ \prod_{i=1}^{n_{\text{tot}}} [y_i | \boldsymbol{\beta}, \alpha, \phi_{s_i, t_i}, \sigma] \times \\ [\boldsymbol{\phi}_1 | \sigma^{(\phi)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_t | \boldsymbol{\phi}_{t-1}, \sigma^{(\phi)}, \eta] \times \\ \prod_{j=1}^{p} [\beta_j | \lambda_j, c, \tau] [\lambda_j] \times \\ [\alpha] [c] [\tau] [\sigma] [\eta] [\sigma^{(\phi)}] \end{bmatrix}$$

$$= \prod_{i=1}^{n_{\text{tot}}} \text{Lognormal}(y_i | \alpha + \mathbf{X}_{(s_i,t_i)} \boldsymbol{\beta} + \phi_{s_i,t_i}, \sigma) \times \\ \text{Normal}(\boldsymbol{\phi}_1 | \mathbf{0}, ((\sigma^{(\phi)})^{-2} (\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{t=2}^{T} \text{Normal}(\boldsymbol{\phi}_t | \eta \boldsymbol{\phi}_{t-1}, ((\sigma^{(\phi)})^{-2} (\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{j=1}^{p} \text{Normal}\left(\beta_j | 0, \frac{\tau^2 c^2 \lambda_j^2}{c^2 + \tau^2 \lambda_j^2}\right) \times \text{Cauchy}^+(\lambda_j | 0, 1) \times \\ \text{Normal}(\alpha | 0, 5^2) \times \text{Inv-Gamma}(c^2 | 2.5, 10) \times \text{Normal}^+(\tau | 0, 5^2) \times \\ \text{Normal}^+(\sigma | 0, 5^2) \times \text{Beta}(\eta | 1, 1) \times \text{Normal}^+(\sigma^{(\phi)} | 0, 1^2).$$

#### <sup>844</sup> Gamma burned area model

<sup>845</sup> We used the following parameterization of the gamma distribution:

$$[y|\kappa,\sigma] = \frac{1}{\Gamma(\kappa)\sigma^{\kappa}} y^{\kappa-1} \exp(-y/\sigma),$$

- where  $\kappa$  is a shape parameter and  $\sigma$  a scale parameter.
- <sup>847</sup> The unnormalized posterior density of this model is:

$$\begin{bmatrix} \boldsymbol{\beta}, \alpha, \boldsymbol{\phi}, \sigma^{(\phi)}, \eta, \kappa, \boldsymbol{\lambda}, c, \tau \mid \boldsymbol{y} \end{bmatrix} \propto \\ \prod_{i=1}^{n_{\text{tot}}} [y_i | \boldsymbol{\beta}, \alpha, \phi_{s_i, t_i}, \kappa] \times \\ [\boldsymbol{\phi}_1 | \sigma^{(\phi)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_t | \boldsymbol{\phi}_{t-1}, \sigma^{(\phi)}, \eta] \times \\ \prod_{j=1}^{p} [\beta_j | \lambda_j, c, \tau] [\lambda_j] \times \\ [\alpha] [c] [\tau] [\kappa] [\eta] [\sigma^{(\phi)}] \end{bmatrix}$$

$$= \prod_{i=1}^{n_{\text{tot}}} \text{Gamma}(y_i|\kappa, \kappa/\exp(\alpha + \mathbf{X}_{(s_i,t_i)}\beta + \phi_{s_i,t_i})) \times \text{Normal}(\phi_1|\mathbf{0}, ((\sigma^{(\phi)})^{-2}(\mathbf{D} - \mathbf{W}))^{-1}) \times \prod_{t=2}^{T} \text{Normal}(\phi_t|\eta\phi_{t-1}, ((\sigma^{(\phi)})^{-2}(\mathbf{D} - \mathbf{W}))^{-1}) \times \prod_{j=1}^{p} \text{Normal}\left(\beta_j|0, \frac{\tau^2 c^2 \lambda_j^2}{c^2 + \tau^2 \lambda_j^2}\right) \times \text{Cauchy}^+(\lambda_j|0, 1) \times \text{Normal}(\alpha|0, 5^2) \times \text{Inv-Gamma}(c^2|2.5, 10) \times \text{Normal}^+(\tau|0, 5^2) \times \text{Normal}^+(\kappa|0, 5^2) \times \text{Beta}(\eta|1, 1) \times \text{Normal}^+(\sigma^{(\phi)}|0, 1^2).$$

#### <sup>848</sup> Weibull burned area model

<sup>849</sup> We used the following parameterization of the Weibull distribution:

$$[y|\kappa,\sigma] = \frac{\kappa}{\sigma} \left(\frac{y}{\sigma}\right)^{\kappa-1} \exp\left(-\left(\frac{y}{\sigma}\right)^{\alpha}\right),$$

- where  $\kappa$  is a shape parameter and  $\sigma$  is a scale parameter.
- <sup>851</sup> The unnormalized posterior density of this model is:

$$[\boldsymbol{\beta}, \alpha, \boldsymbol{\phi}, \sigma^{(\phi)}, \eta, \kappa, \lambda, c, \tau \mid \boldsymbol{y}] \propto \prod_{i=1}^{n_{\text{tot}}} [y_i | \beta, \alpha, \phi_{s_i, t_i}, \kappa] \times [\boldsymbol{\phi}_1 | \sigma^{(\phi)}] \prod_{t=2}^{T} [\boldsymbol{\phi}_t | \boldsymbol{\phi}_{t-1}, \sigma^{(\phi)}, \eta] \times \prod_{j=1}^{p} [\beta_j | \lambda_j, c, \tau] [\lambda_j] \times [\alpha] [c] [\tau] [\kappa] [\eta] [\sigma^{(\phi)}]$$

$$= \prod_{i=1}^{n_{\text{tot}}} \text{Weibull}(y_i|\kappa, \exp(\alpha + \mathbf{X}_{(s_i,t_i)}\boldsymbol{\beta} + \phi_{s_i,t_i})) \times \\ \text{Normal}(\boldsymbol{\phi}_1|\mathbf{0}, ((\sigma^{(\phi)})^{-2}(\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{t=2}^{T} \text{Normal}(\boldsymbol{\phi}_t|\eta\boldsymbol{\phi}_{t-1}, ((\sigma^{(\phi)})^{-2}(\mathbf{D} - \mathbf{W}))^{-1}) \times \\ \prod_{j=1}^{p} \text{Normal}\left(\beta_j|0, \frac{\tau^2 c^2 \lambda_j^2}{c^2 + \tau^2 \lambda_j^2}\right) \times \text{Cauchy}^+(\lambda_j|0, 1) \times \\ \text{Normal}(\alpha|0, 5^2) \times \text{Inv-Gamma}(c^2|2.5, 10) \times \text{Normal}^+(\tau|0, 5^2) \times \\ \text{Normal}^+(\kappa|0, 5^2) \times \text{Beta}(\eta|1, 1) \times \text{Normal}^+(\sigma^{(\phi)}|0, 1^2).$$