1	Microbial communities in an ultra-oligotrophic sea are more affected by season than by distance									
2	from shore									
3										
4	Markus Haber ^{1,2} , Dalit Roth Rosenberg ¹ , Maya Lalzar ³ , Ilia Burgsdorf ¹ , Kumar Saurav ¹ , Regina									
5	Lionheart ⁴ , Yoav Lehahn ⁴ , Dikla Aharonovich ¹ , Daniel Sher ¹ , Michael D. Krom ^{1,5} , Laura									
6	Steindler ^{1*}									
7										
8	¹ Department of Marine Biology, Leon H. Charney School of Marine Sciences, University of									
9	Haifa, Israel									
10	² Department of Aquatic Microbial Ecology, Institute of Hydrobiology, Biology Centre CAS,									
11	Czech Republic									
12	³ Bioinformatics Service Unit, University of Haifa, Israel									
13	⁴ The Dr. Moses Strauss Department of Marine Geosciences, Leon H. Charney School of Marine									
14	Sciences, University of Haifa, Israel									
15	⁵ Morris Kahn Marine Research Station, Environmental Geochemistry Lab., Leon H. Charney									
16	School of Marine Sciences, University of Haifa, Israel									
17	* Corresponding author: Dr. Laura Steindler									
18	Department of Marine Biology, Leon H. Charney School of Marine Sciences, University									
19	of Haifa, 199 Aba Khoushy Ave.Mount Carmel, Haifa, 3498838, Israel									
20	Phone: +972-4-8288987 Fax: +972-4-8288267 Email: lsteindler@univ.haifa.ac.il									
21										
22	Running title: Seasonality of microbial communities in the EMS									
23	Keywords: Mediterranean Sea, SAR11, Transect, Seasonality, 16S rRNA									

24 Abstract

25 Marine microbial communities vary seasonally and spatially, but these two factors are rarely 26 addressed together. We studied temporal and spatial patterns of the microbial community 27 structure and activity along a coast to offshore transect from the Israeli coast of the Eastern 28 Mediterranean Sea (EMS) over six cruises, in three seasons of two consecutive years. The ultra-29 oligotrophic status of the South Eastern Mediterranean Sea was reflected in the microbial 30 community composition that was dominated by oligotrophic microbial groups such as SAR11 31 throughout the year, even at the most coastal station sampled. Seasons affected microbial 32 communities much more than distance from shore explaining about half of the observed 33 variability in the microbial community, compared to only about 6% that was explained by 34 station. However, the most coastal site differed significantly in community structure and activity 35 from the three further offshore stations in early winter and summer, but not in spring. Our data 36 on the microbial community composition and its seasonality from a transect into the South 37 Eastern Levantine basin support the notion that the EMS behaves similar to open gyres rather 38 than an inland sea.

39

40 Introduction

Marine microbial communities play a pivotal function in the biogeochemistry of the ocean
because of their key roles in the carbon, nitrogen and sulfur cycles (Falkowski et al. 2008). Their
composition is strongly affected by seasons with recurring microbial turnover over different
years as revealed by oceanographic time series studies conducted both at offshore stations (*e.g.*Bermuda Atlantic Time-Series Study (BATS) in the Western Atlantic Ocean and the Hawaii
Ocean Time-series (HOT) in the North Pacific subtropical gyre (Giovannoni and Vergin 2012))

47 and more coastal affected sites (e.g. the San Pedro Ocean Time Series (SPOTS) in southern 48 California (Fuhrman et al. 2006) and the western English Channel (Gilbert et al. 2012)). Spatial 49 variability of marine surface water microbial communities has been reported at various scales 50 (Ghiglione et al. 2005; Sunagawa et al. 2015). The observed differences were linked to gradients 51 in environmental conditions (Fortunato et al. 2011; Sunagawa et al. 2015; Wang et al. 2019) and 52 distinct water masses with different physico-chemical properties (Dubinsky et al. 2017; Morales 53 et al. 2018). Spatial variability due to environmental gradients is especially evident when 54 comparing nearshore and offshore microbial communities as parameters such as nutrient 55 availability, temperature, dissolved organic matter, and, especially at estuaries, salinity change 56 from the coast to the open ocean (Ghiglione et al. 2005; Fortunato et al. 2011; Quero and Luna 57 2014; Lucas et al. 2016; Wang et al. 2019). These parameters are influenced by season leading to 58 weakening or strengthening of the environmental gradients along coast to offshore transects. Few 59 studies have followed seasonal differences in the microbial community dynamics over a coastal 60 to offshore transect (Fortunato et al. 2011; Lucas et al. 2016; Wang et al. 2019), so relatively 61 little is known about how these temporal and spatial changes interact and affect the microbial 62 community composition and their functional potential. Only one of these studies (Wang et al. 63 2019) addressed community structure changes without the confounding factor of salinity, which 64 has been identified as a key factor in structuring microbial communities at estuaries (Fortunato et 65 al. 2011).

Here we followed the seasonal dynamics of microbial communities along a coast to offshore
transect in the ultra-oligotrophic Eastern Mediterranean Sea (EMS). This mostly land-enclosed
region, represents the largest body of water that is severely depleted in phosphate (Krom et al.
1991), and one of the most oligotrophic oceanic regions on Earth (Krom et al. 2010). Despite

70 being an inland sea with major anthropogenic external nutrient supply, it has been suggested that 71 it behaves similar to open ocean gyres as a result of its unusual anti-estuarine circulation (Powley et al. 2017). The nutrient cycle at offshore sites in the EMS follows predictable patterns. During 72 73 summer the water column is well stratified with a distinct deep chlorophyll maximum (DCM) 74 and very low nutrient content in the surface waters (Kress and Herut 2001; Krom et al. 2014). In 75 winter, as a result of increasingly deep-water mixing, dissolved nutrients are advected into the 76 photic zone. Due to local weather conditions (short cold and wet (often stormy) periods 77 interspersed with clear sunny ones), the phytoplankton bloom in this region starts soon after the 78 nutrients are supplied to the photic zone and increases throughout the winter to a maximum in 79 late winter (Krom et al. 2014). The dissolved phosphate in surface waters is consumed during the 80 winter phytoplankton bloom, while measurable nitrate persists (Krom et al. 1992). In the summer 81 autotrophs in offshore waters of this region tend to be phosphate and nitrogen co-limited, while 82 heterotrophic bacteria are either phosphate or phosphate and nitrogen co-limited (Thingstad et al. 83 2005; Tanaka et al. 2011; Tsiola et al. 2016). Detailed molecular characterization of the 84 microbial community in this geographic area are limited to few studies that represent only 85 snapshots of the microbial communities taken at a single time point (Feingersch et al. 2010; 86 Keuter et al. 2015; Dubinsky et al. 2017). Until now molecular studies investigating seasonal and 87 spatial differences in the microbial community structure focused on nitrogen fixing microbes 88 using clone libraries of the *nifH* (Man-Aharonovich et al. 2007; Yogev et al. 2011). However, 89 nitrogen fixation seems rare in the euphotic zone of the Eastern Mediterranean (Yogev et al. 90 2011) and thus these microbes are not expected to reflect a large part of the microbial 91 community. The goal of the present study was to examine both microbial community structure 92 and activity in time and space by amplicon sequencing of 16S rRNA genes and transcripts.

93	Surface seawater samples (10 m) were collected along a coastal-to-offshore transect in three
94	seasons (early winter, spring and summer) and for two consecutive years. The oceanographic
95	status of the system was determined from physical, chemical and remote sensing measurements,
96	and the phytoplankton community analyzed by flow cytometry and pigment analysis.
97	
98	Material and Methods
99	Sampling
100	Six one-day cruises were performed onboard the R/V Mediterranean Explorer over a two-year
101	period: two in early winter (December 1st, 2014; November 17th, 2015), two in spring (March
102	24th, 2015; March 30th, 2016) and two in summer (July 14th, 2015; July 25th, 2016). Cruises
103	followed the same coastal to offshore transect heading out from the Herzliya marina, Israel
104	(Figure 1). Four stations, labeled station 1 to 4 in coastal to offshore order, were analyzed for
105	different in situ parameters using a conductivity, temperature and depth (CTD) probe. Seawater
106	was collected from 10 m depth using twelve 8L Niskin bottles mounted on a rosette for nutrient,
107	pigment, cell count, DNA and RNA analyses. To better identify gradients along the transect,
108	CTD and nutrient data were collected for two additional stations: one located between stations 1
109	and 2, and the other between station 2 and 3. Station coordinates, bottom depths, distance
110	between stations and to the shore are summarized in Supplementary Table S1.
111	
112	CTD data and surface chlorophyll a maps

113 Satellite-based maps of surface chlorophyll concentration are derived from the Copernicus

114 Marine Environment Monitoring Service (CMEMS, http://marine.copernicus.eu/services-

115 portfolio/access-to products/). We used the level-3 Mediterranean Sea reprocessed surface

116 chlorophyll concentration

117	(OCEANCOLOUR_MED_CHL_L3_REP_OBSERVATIONS_009_073) product, which
118	consists of merged SeaWiFS, MODIS-Aqua, MERIS and VIIRS satellite data. Using multi-
119	satellite data allows continuous tracking of fine-scale chlorophyll filaments as they are advected
120	and deformed by the currents (Lehahn et al. 2017). Surface chlorophyll concentration are
121	estimated via the MedOC4 (Volpe et al. 2007) and the AD4 (D'Alimonte and Zibordi 2003)
122	algorithms for case-1 and case-2 waters, respectively. Spatial and temporal resolution is 1 km
123	and 1 day, respectively.
124	A SeaBird CTD profiler (SBE 19plus V2) was used to measure vertical profiles of temperature
125	and salinity up to a depth of 500 m. Chlorophyll a fluorescence profiles were measured with a
126	Seapoint fluorometer calibrated with bottle chlorophyll measurements and turbidity profiles with
127	a Seapoint turbidity sensor, both mounted on the CTD. Data were extracted using the Seasoft V2
128	software suite and plotted using Ocean Data View 4.7.4 (Schlitzer R. 2015, http://odv.awi.de).
129	
130	Nutrient analysis
131	Seawater for nutrient analysis was collected in 15 ml Falcon tubes pre-rinsed with sample
132	seawater. For each nutrient, duplicate non-filtered samples were frozen onboard directly after
133	collection and kept at -20°C until analysis of silicate, nitrate+nitrite and soluble reactive

134 phosphate content (within six weeks) at the service unit of the Interuniversity Institute for Marine

135 Sciences in Eilat, Israel. During the July 2016 cruise the same nutrients plus ammonia were also

- 136 measured in non-frozen samples. These samples were filtered through 0.4 µm filters, transferred
- 137 into 15 ml Falcon tubes, stored at 4°C and analysed within 24 hours at the Morris Kahn Marine

Laboratory, Sdot Yam, Israel using a SEAL AA-3 autoanalyzer. Methods are described in theSupplementary method information.

140

141 *Flow cytometry*

142 For flow cytometry sample collection, 1.5 ml triplicate seawater samples were fixed with

143 glutaraldehyde (0.125% final concentration), incubated in the dark for 10 min, stored in liquid

144 nitrogen onboard and kept at -80 °C in the lab until analysis. Then, samples were thawed at room

145 temperature and run on a BD FACSCantoTM II Flow Cytometry Analyzer Systems (BD

146 Biosciences) for counts of phytoplankton and total cell counts and an easyCyte HT Guava flow

147 cytometer (Merck Millipore) for total cell counts (see Supplementary method information for

148 details).

149

150 Pigment analysis

151 Phytoplankton community structure was identified by pigment analysis. Four to eleven L

152 seawater were filtered on Glass fiber filters (25 mm GF/F, Whatman, nominal pore size 0.7μm).

153 Filters were dried by placing their underside on a kimwipe, transferred to cryovials, stored in

154 liquid nitrogen until arrival at the laboratory, and kept at -80°C until extraction. The collected

155 cells were extracted in 1 ml 100% methanol for 2.5 h at room temperature. Extractions were

156 immediately clarified with syringe filters (Acrodisc CR, 13 mm, 0.2 µm PTFE membranes, Pall

157 Life Sciences) and transferred to UPLC vials. Samples were run on an UPLC and pigments

158 identified based on retention time and spectrum absorbance. Several known standards (DHI

159 Water and Environment Institute, (Hørsholm, Denmark)) were used to ease identification and

160 calculate pigment concentrations (see Supplementary Information for details).

161

162 DNA and RNA sample collection and extraction

163 Seawater was collected from the Niskin bottles into 10 or 20 L polycarbonate carboys. Tubes for

- 164 water transfer and the carboys were pre-rinsed three times with sample water. Samples for DNA
- 165 (5 to 11.5 L) were pre-filtered through 11 and 5 µm nylon filters (Millipore), cells collected on
- 166 0.22 µm sterivex filters (Millipore) and kept in storage buffer (40 mM EDTA, 50 mM Tris pH
- 167 8.3, 0.75 M sucrose). Samples for RNA (0.5 to 4.5 L, according to volume filtered within 15

168 minutes from Niskin bottles being on-deck) were collected without pre-filtration, directly on 0.2

169 µm filters (Supor-200 Membrane Disc Filters, 25 mm; Pall Corporation) and filters preserved in

170 RNA Save (Biological Industries). DNA and RNA samples were stored in liquid nitrogen

171 onboard and kept at -80°C in the lab until extraction. Nucleic acids were extracted at the BioRap

172 unit, Faculty of Medicine, Technion, Israel using a semi-automated protocol, which includes

173 manually performed chemical and mechanical cell lysis before the automated steps (see

- 174 Supplementary Information for details).
- 175

176 PCR amplification and sequencing of 16S rRNA genes (DNA) and transcripts (RNA) samples 177 Prior to reverse transcription, all RNA samples were tested by PCR for the presence of 178 contaminating DNA. Total RNA was reverse transcribed using the iScript cDNA synthesis kit 179 (BioRad) according to the manufacturer's instructions. A two-stage "targeted amplicon 180 sequencing" protocol (e.g. (Green et al. 2015)) was performed to PCR amplify the 16S rRNA 181 gene from cDNA and DNA (see Supplementary Information for a detailed description). The 182 primers used in the first PCR stage consisted of the 16S primer set 515F-Y and 926R (Parada et 183 al. 2016) that targets the variable V4-5 region with common sequence tags added at the 5' end as

184	described previously (e.g. (Moonsamy et al. 2013)). The first PCR stage was performed in
185	triplicates, which were pooled after validation on 1% agarose gels. Subsequently, a second PCR
186	amplification was performed to prepare libraries. These were pooled and after a quality control
187	sequenced (2x250 paired end reads) using an Illumina MiSeq sequencer. Library preparation and
188	pooling were performed at the DNA Services (DNAS) facility, Research Resources Center
189	(RRC), University of Illinois at Chicago (UIC). MiSeq sequencing was performed at the W.M.
190	Keck Center for Comparative and Functional Genomics at the University of Illinois at Urbana-
191	Champaign (UIUC).
192	
193	Sequence processing
194	After quality control of the obtained pair of fastq files for each sample, the paired reads were
195	assembled and additional QC steps taken with MOTHUR Version 1.40.2 (Schloss et al. 2009)
196	(see Supplementary Information for details). The filtered dataset was clustered into operational
197	taxonomic units (OTUs) at 97% similarity using QIIME. Representative sequences for each
198	OTU were obtained and classified using MOTHUR with the silva v128 database at 80%
199	confidence level. OTUs identified as non-prokaryotic, chloroplasts- or mitochondria-origin and
200	OTUs with <10 reads across all samples were removed. OTUs identified as SAR11 were also
201	classified against Silva v132 to get clade assignment. In total, the 56 samples had 2,820,858
202	quality sequences, which binned into 9448 OTUs. In order to avoid bias related to differences in
203	library size, all libraries were rarified to 20,000 reads per sample using the R package Vegan
204	(Oksanen et al. 2017). 9396 OTUs based on 97% similarity were retained after subsampling. Of
205	these, 8372 OTUs were found in DNA samples with 1538 to 2253 OTUs present per sample and
206	8684 OTUs in RNA samples ranging from 1604 to 2230 OTUs per sample. Read data were

- deposited in the NCBI SRA database under the project number PRJNA548664. Station 4 data are
 labeled N1200, a change made to ease reading of the article.
- 209
- 210 Sequencing controls and variance among replicates
- 211 Four negative controls from the first PCR were added randomly onto the sequencing plates to
- 212 monitor overall potential for cross contaminations in both PCR and sequencing. These negative
- 213 control samples averaged 929 reads, compared to samples averaging 50,503 reads. Accordingly,
- 214 contaminant DNA should poorly compete with sample DNA for amplification.
- 215 To assess within station variability and robustness of the observed trends, duplicates for DNA
- and RNA samples were collected and extracted for samples from station 4 for the summer 2015

cruise, the spring 2015 and 2016 cruises, as well as for station 1 for the spring 2016 cruise. We

- 218 chose these samples to have a representation from different seasons, as well as from the most
- 219 coastal and most offshore stations. Based on their Bray-Curtis dissimilarity, replicates were
- significant closer to each other than to other stations from the same cruise (paired Wilcoxon
- 221 signed-rank test, RNA: z=3.4044 and P<0.001; DNA: z=4.2812 and P<0.001).
- 222

223 Microbial community analyses

Alpha diversity parameters (Chao1 richness, Shannon H' diversity and Simpson index of

dominance) were calculated using the R package iNEXT (Chao et al. 2014; Hsieh et al. 2018).

226 Beta-diversity was estimated by pairwise calculation of the Bray-Curtis dissimilarity with the R

227 package Vegan and used for non-metric multidimensional scaling analysis. Vegan was also used

to examine significance of grouping factors (such as molecule type, season and sampling station)

229 with the ADONIS test and to examine correspondence between abiotic and biotic measurements

and variation in community composition by variation partitioning analysis and canonical

231 correspondence analysis, for which proposed explanatory variables were divided into three

232 matrices: I) Physical: distance from shore, temperature, salinity, turbidity; II) Nutrients:

233 phosphate, nitrate+nitrite, silicate; III) Biological: total fluorescence, total cell counts,

234 Prochlorococcus cell counts, Synechococcus cell counts (see Supplementary method information

235 for details of all Vegan analyses).

236 For differential OTU abundance analysis between sample groups (e.g. molecule type, season,

sampling station), OTU count data was normalized using the cumulative sum scaling method

implemented in the R package metagenomeSeq (Paulson et al. 2013) and analyzed with the R

package DEseq2 (Love et al. 2014). Significant differential OTU abundance was based on

240 Benjamini-Hochberg adjusted *P*-values.

241 Synchronous dynamics of single OTUs were analyzed by soft clustering with the R package

242 Mfuzz (Kumar and Futschik 2007) using OTUs that comprised more than 0.1% of total reads of

243 DNA and RNA samples, respectively. OTU tables were standardized using the normal SD based

244 method, fuzzifier variables were estimated (DNA: 1.19569, RNA: 1.182411) and cluster number

245 was set to 10. Only OTUs with cluster membership values of at least 0.7 were considered.

246 Similarity profile analysis (SIMPROF) (Clarke et al. 2008) based on Bray-Curtis similarity was

run in R using the clustsig package version 1.1 to test for structure in the spring 2016 samples

relating to the intrusion of coastal waters into offshore waters.

249

250 *Statistical analyses*

251 Paired Wilcoxon signed rank tests were performed to test if Bray-Curtis dissimilarity between

252 replicates from the same sample location was lower than between samples from different

253	locations from the same cruise. A Kruskal-Wallis test was used to find significant difference in							
254	alpha diversity indices between seasons. Two-tailed Mann-Whitney U tests were used to							
255	determine for each season if Bray-Curtis dissimilarity between station 1 and the other three							
256	station was higher than between the three further offshore station. Finally, Spearman correlations							
257	were performed to compare the relative abundance in RNA and flow cytometry data for							
258	Synechococcus and Prochlorococcus, to compare if distance between stations correlates with							
259	Bray-Curtis dissimilarity between samples and to compare if differences in relative abundance in							
260	RNA and DNA data are correlated. All these tests were performed in PAST version 3.14							
261	(Hammer et al. 2001). Other statistical tests were performed as described above and							
262	implemented in the used R packages.							
263								
264	Results							
265	Environmental setting							
266	The previously reported ultraoligotrophic status of the Levantine basin (Krom et al. 2005) was							
266 267	The previously reported ultraoligotrophic status of the Levantine basin (Krom et al. 2005) was reflected in our nutrient data of samples collected at 10 m (Supplementary Table S3 and S4).							
266 267 268	The previously reported ultraoligotrophic status of the Levantine basin (Krom et al. 2005) was reflected in our nutrient data of samples collected at 10 m (Supplementary Table S3 and S4). Phosphate concentrations obtained from frozen samples were near or below our detection limit							
266 267 268 269	The previously reported ultraoligotrophic status of the Levantine basin (Krom et al. 2005) was reflected in our nutrient data of samples collected at 10 m (Supplementary Table S3 and S4). Phosphate concentrations obtained from frozen samples were near or below our detection limit of 40 nM in all cruises with a maximum of 70 nM at station 1 in early winter 2014.							
266 267 268 269 270	The previously reported ultraoligotrophic status of the Levantine basin (Krom et al. 2005) was reflected in our nutrient data of samples collected at 10 m (Supplementary Table S3 and S4). Phosphate concentrations obtained from frozen samples were near or below our detection limit of 40 nM in all cruises with a maximum of 70 nM at station 1 in early winter 2014. Nitrate+nitrite concentrations for these samples ranged from < 50 nM to 240 nM at the most							
266 267 268 269 270 271	The previously reported ultraoligotrophic status of the Levantine basin (Krom et al. 2005) was reflected in our nutrient data of samples collected at 10 m (Supplementary Table S3 and S4). Phosphate concentrations obtained from frozen samples were near or below our detection limit of 40 nM in all cruises with a maximum of 70 nM at station 1 in early winter 2014. Nitrate+nitrite concentrations for these samples ranged from < 50 nM to 240 nM at the most offshore station and from <50 to 690 nM at the most coastal station. They were below our							
266 267 268 269 270 271 272	The previously reported ultraoligotrophic status of the Levantine basin (Krom et al. 2005) was reflected in our nutrient data of samples collected at 10 m (Supplementary Table S3 and S4). Phosphate concentrations obtained from frozen samples were near or below our detection limit of 40 nM in all cruises with a maximum of 70 nM at station 1 in early winter 2014. Nitrate+nitrite concentrations for these samples ranged from < 50 nM to 240 nM at the most offshore station and from <50 to 690 nM at the most coastal station. They were below our detection limit (50 nM) for frozen samples during summer cruises apart from station 1 in the July							
266 267 268 269 270 271 272 272 273	The previously reported ultraoligotrophic status of the Levantine basin (Krom et al. 2005) was reflected in our nutrient data of samples collected at 10 m (Supplementary Table S3 and S4). Phosphate concentrations obtained from frozen samples were near or below our detection limit of 40 nM in all cruises with a maximum of 70 nM at station 1 in early winter 2014. Nitrate+nitrite concentrations for these samples ranged from < 50 nM to 240 nM at the most offshore station and from <50 to 690 nM at the most coastal station. They were below our detection limit (50 nM) for frozen samples during summer cruises apart from station 1 in the July 2015 cruise (120 nM). In early winter they decreased from station 1 towards the offshore							
266 267 268 269 270 271 272 273 273	The previously reported ultraoligotrophic status of the Levantine basin (Krom et al. 2005) was reflected in our nutrient data of samples collected at 10 m (Supplementary Table S3 and S4). Phosphate concentrations obtained from frozen samples were near or below our detection limit of 40 nM in all cruises with a maximum of 70 nM at station 1 in early winter 2014. Nitrate+nitrite concentrations for these samples ranged from < 50 nM to 240 nM at the most offshore station and from <50 to 690 nM at the most coastal station. They were below our detection limit (50 nM) for frozen samples during summer cruises apart from station 1 in the July 2015 cruise (120 nM). In early winter they decreased from station 1 towards the offshore stations, whereas in spring they increased from station 1 to station 3 before decreasing again							
 266 267 268 269 270 271 272 273 274 275 	The previously reported ultraoligotrophic status of the Levantine basin (Krom et al. 2005) was reflected in our nutrient data of samples collected at 10 m (Supplementary Table S3 and S4). Phosphate concentrations obtained from frozen samples were near or below our detection limit of 40 nM in all cruises with a maximum of 70 nM at station 1 in early winter 2014. Nitrate+nitrite concentrations for these samples ranged from < 50 nM to 240 nM at the most offshore station and from <50 to 690 nM at the most coastal station. They were below our detection limit (50 nM) for frozen samples during summer cruises apart from station 1 in the July 2015 cruise (120 nM). In early winter they decreased from station 1 towards the offshore stations, whereas in spring they increased from station 1 to station 3 before decreasing again towards station 4, the most offshore station. In the summer 2016 cruise, we performed also an							

analysis on fresh (unfrozen) samples that provided a lower limit of detection (Krom et al. 2005).

277 This enabled the identification of the nutrient range during the most nutrient depleted state.

278 Concentrations of phosphate ranged from 0.2 to 2.1 nM, of nitrate from 8 to 28 nM, and of

ammonia from 1.3 to 7 nM (Supplementary Table S4) again confirming the nutrient depleted

280 status of the Levantine Basin during summer.

281 Temperature and salinity varied seasonally at the sampling depth (10 m). The lowest

temperatures were observed in spring (17.7-18.6°C) and the highest in summer (26.2-28.7°C).

283 Seasonal changes in salinity were small (<0.7 PSU from lowest to highest value) (Supplementary

284 Table S2).

285 Together CTD (Supplementary Figure 1), nutrient (Supplementary Table S3) and satellite data

286 (Figure 1) indicated that our sampling times corresponded to three distinct stages of the annual

287 cycle: the stratified water column in summer, early mixing and phytoplankton bloom in early

winter and the declining phytoplankton bloom in spring (Krom et al. 1992; Kress and Herut

289 2001). Satellite data indicated the peak of the phytoplankton bloom to be in both years in

290 February (data not shown). Based on pigment data, early winter samples were dominated by

291 haptophytes of either Phaeocystaceae or of members of the Prymnesiaceae and Isochrysidaceae

292 families (see details on pigment results in Supplementary information as well as Supplementary

Figure S2 and Supplementary Table S5). Spring samples were characterized by declining

294 phytoplankton bloom communities (Supplementary Figure S2, Supplementary Table S5) and

295 peak abundances of *Prochlorococcus* (Supplementary Figure S5B and Table S6). During

summer the water column was stratified with a mixed layer depth of 30m and characterized by

297 nutrient depleted surface waters (Supplementary Table S2) and the lowest Synechococcus

abundances at stations 2 to 4 (Supplementary Figure S5A, Supplementary Table S6).



299

Figure 1. Surface chlorophyll concentrations in the southeastern Mediterranean. Dots mark
locations of sampling stations, which are from right (coast) to left offshore: St1, St2, St3, St4.
A) Maps derived from Copernicus Marine Environment Monitoring Service (CMEMS)
merged satellite data. To reduce the area masked by clouds (white color), chlorophyll maps from
30.11.2014, 20.11.2015, 01.04.2016 and 26.07.2016 were used instead of 01.12.2014,
17.11.2015, 30.03.2016 and 25.07.2016, respectively. Furthermore, maps for 30.11.2014,

306 24.3.2015, 14.7.2015 and 26.7.2016 are composed of 5, 3, 3 and 3 consecutive images,

307 respectively. Purple dots mark the locations of sampling stations.

308 B) Surface chlorophyll concentrations along a coast-to-open-sea transect overlapping the

309 sampling stations. The transects are derived from the multi-satellite chlorophyll maps shown in

310 A, with each line corresponding to a different sampling date.

311

312 When comparing the transect sampling stations from the coast outwards, a clear difference was

313 seen between station 1, the shallowest and most coastal sampling site, and the other three more

314 offshore stations. Station 1 often differed from all other sampled stations in nutrient

315 concentrations (e.g. silicate (in spring) and nitrite+nitrate (early winter, Supplementary Tables S2

and S3)), in remotely sensed surface chlorophyll (Figure 1) and several *in situ* parameters such as

317 temperature (in early winter and spring), salinity (in summer), fluorescence (in early winter) and

turbidity (in early winter and summer, Supplementary Table S5, Supplementary Figure S1). In

319 accordance with (Berman et al. 1986), it was thus regarded as a coastal station.

320

321 Microbial community composition

322 93 of the 9396 OTUs (based on 97% similarity) formed the "core" of the microbial community

323 and were present in all samples and represented 61% of the total reads. Bacteria dominated with

324 99.68% of the total reads compared to 0.32% of reads identified as archaea. 95.3% of all reads

325 belonged to nine bacterial classes: Alphaproteobacteria (48.2% of all reads), Cyanobacteria

326 (17.6%), Gammaproteobacteria (15.4%), Flavobacteriia (5.4%), Marinimicrobia (SAR406 clade)

327 (2.4%), the Verrucomicrobia class Opitutae (2.1%), Deltaproteobacteria (1.6%), Acidimicrobia

328 (1.3%) and Betaproteobacteria (1.3%). Most Alphaproteobacteria reads belonged to the SAR11

- 329 clade (60.4% of the Alphaproteobacteria), most reads of the class Cyanobacteria were identified
- 330 as either *Prochlorococcus* or *Synechococcus* (together 85.9% of the Cyanobacteria), and the
- 331 SAR86 clade was most common within the Gammaproteobacteria (32.8% of the
- 332 Gammaproteobacteria). Relative abundances of the 13 families with >1% of total reads are
- shown in Figure 2.



Figure 2. Relative abundance of microbial taxa in DNA (left) and RNA (right) samples. Families with > 1% total reads (DNA+RNA) are shown. Families below this threshold are summarized at the phylum level, except SAR11, which are summarized as other SAR11. Phyla with < 1% of total reads are summarized at the domain level. Color bars under the graphs indicate seasons: winter- blue, spring – green, summer – orange. Sample code: Last two digits of collection year (20XX), season (Wt – early winter, Sp - spring, Sm - summer) and station. Blue vertical lines separate different cruises.

343	A significant positive correlation (Spearman correlation, r>0, P<0.05) between relative
344	abundance in DNA and RNA samples was present in 10 of these 13 families (Supplementary
345	Table S7) indicating a strong link between presence and activity. However, the composition of
346	resident (DNA) and active (RNA) microbial communities clearly differed (Figure 2) as
347	supported by non-metric multidimensional scaling analysis (NMDS) (Figure 3) and Adonis test
348	(F=22.37, R ² =0.293, P<0.001). DEseq2 analysis identified 976 OTUs (representing 74.5% of all
349	reads) with significantly different relative abundances between DNA and RNA (Benjamini-
350	Hochberg adjusted $P < 0.05$). Of these, 586 OTUs were more abundant in RNA samples and 390
351	in DNA samples. All significant OTUs from SAR11 clade I, clade II, the OM1 clade and the
352	Flavobacteriaceae were higher in the DNA than the RNA samples. Significant OTUs of the
353	Rhodobacteraceae, the SAR86 clade, the SAR116 and the Cyanobacteria subsection I family I,
354	which included Cyanobium, Synechococcus, Prochlorococcus and OTUs unclassified at the
355	genus level, were all higher in the RNA than in the DNA samples. Rhodospirillaceae and
356	Marinimicrobia had both more OTUs with higher relative abundances in RNA than in DNA
357	samples, however a few OTUs from both groups also showed the opposite trend. The remaining
358	309 significant OTUs did not sum up to $>1\%$ of total reads in any family. Given the difference
359	between DNA and RNA samples, all subsequent analyses of microbial community composition
360	(DNA) and activity (RNA) were performed separately.



361

Figure 3. NMDS plot of the RNA (empty shapes) and DNA (filled shapes) samples. Colors
 represent seasons: Spring – green, Summer – red, early Winter – blue.

364

365 Season had a larger effect on microbial structure and activity than spatial location

366 Season clearly affected both the microbial community structure (DNA) and activity (RNA)

367 (Figure 3). This seasonal effect was much larger than the effect of station location across the

368 coastal to offshore transect (DNA season $R^2=0.4980$, station $R^2=0.0584$; RNA season

 $R^2=0.5332$, station $R^2=0.0607$) with a small, significant interaction between season and station

370 (DNA: R²=0.0840; RNA: R²=0.0806) (Adonis test, details in Supplementary Table S8).

371 Season affected also the microbial diversity. For DNA samples the three alpha diversity indices

indicated lowest diversity in summer (Kruskal-Wallis test, P<0.001), when the system was most

373 oligotrophic. The same pattern was observed for RNA samples, except for one index (Simpson

diversity) that showed lowest diversity in spring (Supplementary Figure S3).

- 375 Soft clustering analysis on abundant OTUs was performed to identify OTUs that show the same
- abundance pattern across different seasons. Considering only OTUs with more than 0.1% of total

377	reads, 98 of 110 OTUs at the DNA level, and 115 of 131 OTUs at the RNA level, were
378	successfully clustered. At both the DNA and RNA level 8 of 10 clusters showed seasonal
379	preferences (Figure 4). The OTUs of these seasonal clusters contributed 63.9% and 65.0% of
380	total DNA and RNA reads, respectively. OTUs present in both clustering analyses showed
381	consistent seasonal preferences (Supplementary Table S9). No OTU was assigned to more than
382	one cluster. Cluster assignment, taxonomic identity and individual preferences of OTUs as well
383	as abundance patterns of the non-seasonal clusters are shown in Supplementary File S1.
384	DEseq2 analysis further confirmed that many OTUs differed significantly between at least two
385	seasons. 582 OTUs representing 71.3% of all reads at the DNA level and 947 OTUs representing
386	81% of all reads at the RNA level had significant differences between two seasons (Benjamini-
387	Hochberg adjusted <i>P</i> -value <0.05). Differences, in terms of relative abundance and number of
388	significantly differing OTUs, were largest when comparing summer with the other two seasons
389	(Supplementary File S2).





395	sorted by cruise and location. Sample code consists of last two digits of collection year, season
396	and station (St1 most coastal, St4 most offshore). Supplementary File S1 shows the patterns of
397	the non-seasonal clusters D4, D7, R1 and R6 and data of OTUs that did not cluster.
398	
399	DEseq2 and soft clustering results identified several taxonomic groups favoring specific seasons:
400	e.g. the Chloroflexi clade SAR202, the Marinimicrobia clade SAR406 and the
401	Deltaproteobacteria clade SAR324 were more abundant in early winter, while the Aegan-169
402	marine group of the Alphaproteobacteria family Rhodospirillaceae were more abundant in
403	summer. Lower species richness in summer, as indicated by lower Chao1 indices, was already
404	evident at the order level (Supplementary Figure S3C). The majority of orders had no OTU that
405	preferred summer over either spring or winter (Supplementary File S2).
406	At the DNA level, SAR11 was the order most affected by season, with seasonal preference
407	varying according to SAR11 clade (Supplementary Figure 4). SAR11 clade Ia was the main
408	SAR11 clade (51.4-80.3% of all SAR11 DNA reads) and the most abundant taxonomic group
409	overall (15.7-40.8% of the DNA reads). Its highest abundances were observed in summer.
410	Cluster analysis grouped its four most abundant OTUs (contributing 0.5 to 20.7% of total DNA
411	reads) with summer-associated clusters, while the other clustered OTUs were assigned to a
412	spring-associated cluster (Supplementary File 1). SAR11 clade Ia was the only SAR11 clade
413	with summer associated OTUs in the cluster analysis apart from SAR11 clade III, which had one
414	OTU but overall did not show a clear seasonal pattern. SAR11 clades Ib, II and IV had their
415	lowest relative abundances in summer. SAR11 clade II peaked in spring and early winter, while
416	clade IV was most abundant in spring (Supplementary Figure S4 B). Overall, few SAR11 OTUs

417 differed significantly between winter and spring with all but one clade III OTU preferring spring418 (Supplementary file 2).

- 419 At the RNA level, cyanobacteria subsection I was the order that was mostly affected by season.
- 420 All significant OTUs in the DEseq2 analysis (Benjamini-Hochberg adjusted *P*-value < 0.05)
- 421 (125 OTUs, 22.6% of all RNA reads) belonged to family I, with most (83) being taxonomically
- 422 assigned to *Prochlorococcus*. The vast majority of significant OTUs (109 OTUs) were
- 423 significantly more active in spring than summer and 47 were significantly more active in early
- 424 winter than summer. 43 OTUs differed significantly between spring and early winter (34 more
- 425 active in spring, 9 in early winter). The few OTUs that were significantly more active in summer
- 426 were all very rare. Overall cyanobacteria subsection I represented a larger part of the active
- 427 community in spring (35.8% average of RNA reads), followed by winter (22.4%) and summer
- 428 (12.9%) (Figure 2 B). *Prochlorococcus* made up for a larger part of the active community in
- 429 spring, whereas no clear preference between winter and spring was observed for *Synechoccocus*.
- 430 Their activity pattern correlated well with their cell abundances as determined by flow cytometry
- 431 (Spearman rs 0.75 and 0.62, both *p*<0.001 for *Prochlorococcus* and *Synechococcus*, respectively)
- 432 (Supplementary Figure S5).
- 433 Other taxonomic groups had closely related OTUs with different seasonal preferences. Details on
- their seasonal preferences are given in Supplementary Files S1 and S2, that show the cluster-
- based preferences of all abundant (>0.1 % of reads) OTUs at DNA and RNA levels and all OTUs
 with significant seasonal difference according to DEseq2 analysis (Benjamini-Hochberg adjusted *P*-value < 0.05), respectively.
- 438

439 Spatial effects: Microbial communities at coastal station 1 differ from those at offshore stations
440 in summer and early winter, but not in spring.

441 Sampling site had a weak, yet significant effect on microbial structure and activity and interacted

- 442 with season (Adonis test, see Supplementary Table S8 for details). However, no significant
- 443 correlation was found between distance between stations and Bray-Curtis dissimilarity in
- 444 microbial community structure (DNA) or activity (RNA) (Spearman correlation, P>0.05)
- 445 irrespective if the whole data set was analyzed or each season separately. Further, when testing
- 446 which environmental factors affected microbial community structure, distance between station
- 447 was not a significant factor in the DNA analysis (canonical correspondence analysis (CCA),

448 *P*>0.05, Supplementary Figure S7). In the RNA analysis distance was a significant factor in the

449 CCA, but it was no longer significant when correcting for variance explained by the biological

450 and nutrient data matrices (conditional CCA, *P*>0.05, Supplementary Figure S7).

451 As station 1 differed in environmental parameters from the three more offshore stations (see

452 environmental settings above), we tested if this difference could be detected in the microbial

453 community structure (DNA) and activity (RNA). Bray-Curtis dissimilarities were significantly

- 454 higher between station 1 and the other three stations than among the three offshore stations, in
- 455 both early winter and summer (Mann-Whitney U test, 2-tailed, winter: RNA z=-2.3219 P=0.020;
- 456 DNA: z=-2.8022 *P*=0.005; summer: RNA z=-3.1825 *P*=0.001; DNA: z=-2.8353 *P*=0.005), but
- 457 not in spring (RNA: z=-0.2967 *P*=0.767; DNA: z=-1.2858 *P*=0.199) (Figure 5). The effect was
- 458 also observed in the NMDS plot as summer and early winter samples from station 1 grouped
- 459 apart from the samples of stations 2 to 4 (Figure 3).
- 460



Figure 5. Bray-Curtis dissimilarity of A) DNA and B) RNA samples. Comparisons involving
station 1 differ significantly from comparisons among stations 2, 3 and 4 in summer and early
winter in both RNA and DNA samples

465

461

466 In summer, 22 OTUs (5.7% of all summer reads) and 27 OTUs (8.5% of all summer reads) at the 467 DNA and RNA level, respectively, differed significantly between station 1 and stations 2-4 468 (DEseq2 analysis, Benjamini-Hochberg adjusted *P*-value < 0.05). In early winter 54 OTUs (5.0%) 469 of all early winter reads) and 76 OTUs (11.8% of all early winter reads) at the DNA and RNA 470 level respectively, differed significantly between the coastal station and the offshore stations 471 (DEseq2 analysis, Benjamini-Hochberg adjusted P-value < 0.05). 5 OTUs at the DNA level and 472 2 OTUs at the RNA differed significantly between station 1 and stations 2-4 in both summer and 473 winter. Their preference was the same in both seasons suggesting a general preference for coastal 474 or offshore conditions. Supplementary File S3 provides information on the taxonomy of groups 475 enriched at station 1 or at the offshore stations 2-4. 476

477 *Offshore microbial communities are affected by intrusions of coastal water*

478 At sites close to the Israeli coast of the EMS, short-term natural pulses of nutrients in the ultra-479 oligotrophic offshore waters can occur due to intrusions of mesoscale patches of coastal water 480 that can be frequently observed by analysis of satellite-derived surface chlorophyll data (Efrati et 481 al. 2013; Dubinsky et al. 2017). These intrusions of coastal water lead to shifts in the microbial 482 communities of offshore affected sites (Dubinsky et al. 2017). Satellite chlorophyll (Figure 1) as 483 well as in situ data (Supplementary Table S5) indicated an intrusion of coastal water into 484 offshore waters taking place during the spring 2016 cruise, that affected station 4. As a result, 485 phytoplankton community structure changed, as evidenced by increased concentrations in 486 several pigments (Supplementary Table S4) and slightly higher abundances in Synechococcus 487 cell numbers at the coastal station 1 and the intrusion-affected offshore station 4 compared to the 488 offshore non-affected stations 2 and 3 (Supplementary Table S6). The affected stations grouped 489 together in hierarchical clustering of the microbial structure (DNA) and activity (RNA) 490 (Supplementary Figure S6), but DEseq2 analysis failed to identify any OTUs responsible for this 491 difference (Benjamini-Hochberg adjusted P-value > 0.05 for all OTUs). 492 493 Effects of environmental parameters on the microbial community structure and activity 494 Three data matrixes of external factors were examined for their influence on the microbial 495 structure and activity: i) physical data (consisting of distance from shore, temperature, salinity 496 and turbidity), ii) biological data: fluorescence (as proxy for chlorophyll), total cell number, 497 Prochlorococcus counts, Synechococcus counts and iii) nutrient data: concentrations of

498 phosphate, nitrate+nitrite, silicate. Together the three matrixes explained 49% of the observed

499 variability of the microbial structure (DNA samples) and 72% of the microbial activity (RNA

500	samples). Partition analysis indicated that in both analyses, the physical data matrix explained
501	most of the variability (34.2% and 54.2% for DNA and RNA, respectively), followed by the
502	biological data matrix (29.2% and 45.1% for DNA and RNA, respectively) and the nutrient data
503	matrix (11.8% and 10.3% for DNA and RNA, respectively). The physical data matrix also had
504	the largest proportion of variability not explained by any of the other two matrixes
505	(Supplementary Figure S7).
506	A CCA analysis was performed to identify the environmental factors likely responsible for the
507	observed patterns. Similar to the partition analysis, the physical matrix explained the largest
508	amount of variance, followed by the biological matrix and the nutrient matrix. Temperature and
509	salinity from the physical matrix and fluorescence from the biological matrix were significant
510	factors in the DNA and RNA analysis (both in the unconditioned and conditioned analysis,
511	P < 0.05, Supplementary Figure S7). In the nutrient matrix, only silicate was consistently
512	significant, but only in the DNA analysis ($P < 0.05$, Supplementary Figure S7).

513

514 **Discussion**

515 The ultra-oligotrophic nature of the Eastern Mediterranean Sea (EMS) is reflected in the 516 microbial community structure of this coastal-offshore transect. Typical oligotrophic groups (e.g. 517 SAR11) dominated throughout the year confirming previous snapshot studies of the EMS 518 (Feingersch et al. 2010; Dubinsky et al. 2017). The microbial communities resembled closely 519 those found at ultra-oligotrophic open ocean stations, both at the coastal station (e.g. SAR11 520 ranging from 26 ± 4 % in winter to 44 ± 6 % in summer) and even more at the offshore stations 521 (e.g. SAR11 ranging from 44 ± 8 % in winter to 47 ± 5 % in summer). These estimates are 522 similar if not higher than those found at oligotrophic ocean gyres such as the South Atlantic gyre

523	$(36 \pm 9\%)$ (Morris et al. 2012), the North Pacific gyre (ALOHA station, 44% in winter, 33% in
524	summer) (Eiler et al. 2009), the South Pacific gyre (up to 53% in the most oligotrophic station)
525	(West et al. 2016) and the Sargasso sea (North Atlantic gyre, BATS station, $33 \pm 8\%$) (Carlson et
526	al. 2009). Powley et al. (Powley et al. 2017) recently proposed that the EMS behaves similar to
527	open gyres with respect to the importance of external nutrient supply, of dissolved organic matter
528	as a key source of nutrients and ultra-oligotrophic conditions. Here we show that this is true also
529	in terms of microbial community composition. The dominance of microbial groups typically
530	considered to be oligotrophic at the coastal station (station 1) during spring and summer
531	underlines that in the EMS oligotrophic conditions are found also close to the coast. SAR11
532	relative abundances were in fact much higher (>40% of DNA reads in both spring and summer)
533	than those found at coastal stations in other parts of the Mediterranean Sea, including the North
534	Western Mediterranean and the Adriatic Sea (Alonso-Sáez et al. 2007; Quero and Luna 2014;
535	Tinta et al. 2015), highlighting the overall more oligotrophic status of the South Eastern
536	Levantine basin.
537	In our study, season affected the microbial community structure and activity much more than
538	spatial location along the coastal to offshore transect. This finding differs from other studies that
539	analyzed spatial-temporal community changes in coastal-offshore transects. In a transect that
540	sampled the Columbia river, its estuary, plume and coastal to offshore lines, Fortunato et al
541	(Fortunato et al. 2011) described a stable spatial separation between the sampling sites
542	throughout the year, while seasonal differences were observed only within site-groups. This
543	could be explained by the salinity gradient along the transect, as indicated by the high correlation
544	with salinity (Fortunato et al. 2011). However, in absence of particular salinity gradients, also
545	Wang et al. (Wang et al. 2019) found a strong influence of sample location on microbial

546 community structure when sampling off the coastal PICO station and out to the Sargasso Sea. 547 Yet, a seasonal effect could be seen there in the clustering of shelf stations, which grouped either 548 with the coastal station or with the offshore station, depending on season and by the strong 549 influence of temperature. Key differences between the study by Wang et al. (Wang et al. 2019) 550 and our study are the strength of the environmental gradient between the coastal and offshore 551 stations and the overall productivity of the coastal area. The region studied by Wang et al. (Wang 552 et al. 2019) was characterized by higher productivity and stronger shore influence on the coastal 553 stations. In the Wang et al. study primary production at the coastal station was usually above 1000 mg C m⁻³ d⁻¹, which is more than an order of magnitude higher than that of their most 554 555 offshore station (Wang et al. 2019). In the EMS, at a coastal site in Israel, Raveh et al. found that primary production would reach a maximum of about 270 mg C m⁻³ d⁻¹ and the difference 556 557 between coastal and offshore sites was only about four-fold (Raveh et al. 2015). Given that this 558 site was only 50 m from shore with a bottom depth of 5m, we expect the difference in primary 559 production between our coastal site (16 km from shore with a bottom depth of 100 m) and the 560 offshore stations to be even lower. The spatial effects on the microbial community between the 561 most coastal station 1 and the offshore stations 2 to 4 were detected in early winter and summer, 562 yet not in spring, when coastal and offshore communities were highly similar. This is in line with 563 earlier observations of the EMS, that based on chlorophyll and temperature, described a distinct 564 border between coastal and offshore waters from late spring to early winter, while this border became diffuse from winter to early spring (Berman et al. 1986). 565 566 Profound and predictable seasonal shifts in marine surface water microbial communities have

been shown in several studies (*e.g.* (Fuhrman et al. 2006, 2015; Ward et al. 2017; Galand et al.

568 2018)) and, as in our case, temperature has been frequently identified as a main environmental

569 driver (e.g. (Sunagawa et al. 2015; Lucas et al. 2016; Ward et al. 2017)). Salinity, like 570 temperature, is often an indicator for seasonal changes in hydrography and accordingly it is often 571 identified as seasonal driver of marine surface water microbial communities (e.g. (Fuhrman et al. 572 2006; Ward et al. 2017)). Despite the fact that seasonality is expected to strongly affect primary 573 productivity and nutrient availability in the EMS (Kress and Herut 2001; Krom et al. 2014; 574 Raveh et al. 2015), our measured nutrient data explained the seasonal shifts in microbial 575 communities to a lesser extent than the measured physical parameters. It is possible that the 576 effect of nutrients was not detected as several measurements, especially phosphate, were below 577 the detection limit of our analysis method. Higher sensitivity nutrient measurements using fresh 578 unfrozen samples (such as those here utilized only in the July 2016 cruise) might clarify the 579 importance of nutrients *versus* physical factors in the seasonality patterns of the microbial 580 community. Our measured parameters explained a large part of the observed variability in 581 microbial community: 49% of the DNA and 72% of the RNA data. Additional factors that were 582 not measured in this study, may be responsible for the unexplained variability. These include UV 583 radiation, which is known to penetrate deep into the clear waters of the EMS with UV doses 584 among the highest of all oceans (Tedetti and Sempéré 2006; Smyth 2011), and irradiance, a 585 factor known for structuring surface water microbial communities in the subtropical North 586 Pacific Gyre station ALOHA (Bryant et al. 2016). 587

The recurrent shifts in microbial structure and activity followed established seasonal patterns. Physical mixing of the water column in early winter seemed to be responsible for a 'resetting' of the microbial ecosystem from a low diversity state in summer, when the microbial community is nutrient depleted in the EMS community (Thingstad et al. 2005; Tanaka et al. 2011), to a high diversity state in winter, as previously proposed for the Northwestern Mediterranean Sea (Salter

592 et al. 2015). The mixing leads to an upwelling of nutrients that allow previously rare microbes to 593 grow and thrive. In addition, microbes from deeper layers are brought up and might repopulate 594 and grow in the now nutrient enriched surface waters as observed in the Western Mediterranean 595 Sea (Haro-Moreno et al. 2018). These 'upwelled' microbes might interact with surface water 596 microbes and affect the microbial community as shown in recent mesocosm experiments 597 performed in the EMS (Hazan et al. 2018). Potential candidates of 'upwelled' microbes in our 598 study include e.g. SAR202 and SAR406 groups that are typical of deeper water and were more 599 abundant in early winter than in spring or summer. In the EMS, phytoplankton is dominated by 600 nano- to micro sized organisms in winter and picophytoplankton in summer (Raveh et al. 2015). These differences were reflected in our pigment data. Seasonality in heterotrophic bacteria might 601 602 thus be due to changes in phytoplankton composition, which, likely through the different types of 603 organic carbon produced, are known to affect microbial community structure (Camarena-Gómez 604 et al. 2018).

605 Annual microbial patterns can be observed also at the level of closely related OTUs within 606 specific taxonomic groups that have different seasonal preferences, as found both in this study 607 and at the coastal PICO site (Ward et al. 2017). The seasonal preference is likely due to 608 differences in their genomic make-up resulting in different ecotypes that thrive under different 609 environmental conditions. An example is provided by the dominant SAR11 clade, where, within 610 clade 1a, the three most abundant OTUs peaked in summer, while the fourth to sixth most 611 abundant ones peaked in spring. Nevertheless, the main seasonal differences in SAR11 612 abundances were linked to SAR11 clades. They followed the general patterns found at other 613 ocean stations such as the Mola station in the Northwestern Mediterranean (Salter et al. 2015) 614 and the BATS station in the Sargasso sea (Carlson et al. 2009; Vergin et al. 2013) with some

615	differences: similarly to Mola and in contrast to BATS, in the EMS clade Ib did not replace clade
616	Ia as the dominant clade in spring, whereas in contrast to Mola we found in the EMS a
617	substantial amount of SAR11 clade IV that followed the seasonal pattern described at BATS
618	(Vergin et al. 2013).
619	Seasonal shifts in microbial communities are linked to functional differences between microbes
620	that enable better adaptation to the changing environment (Galand et al. 2018; Haro-Moreno et
621	al. 2018). These changes can have a strong impact on biogeochemical cycles. Here we provide a
622	detailed analysis of spatial and temporal changes of microbial communities in the South Eastern
623	Levatine basin. To understand how the observed microbial community shifts affect
624	biogeochemial cycles will require metagenomic and transcriptomic studies, as well as
625	ecophysiology investigations of the dominant bacterial groups.
626	

627 Acknowledgements

628 We thank all cruise participants, the crew of the R/V Mediterranean Explorer and the EcoOcean 629 foundation for their help in sampling, Dr. Rinat Bar Shalom (University of Haifa) for help in 630 cruise planning and preparation, Dr. Tanya Rivilin (Interuniversity Institute for Marine Science 631 at Eilat) for help in nutrient analysis and Dr. Stephan Green (DNA Services Facility at University 632 of Illinois at Chicago) for useful comments and suggestions on the amplicon sequencing 633 methods. This study was funded by the Israel Science Foundation grant (ISF #1243/16) to LS. 634 The seasonal cruises were supported by funding from the Leon H. Charney School of Marine 635 Sciences (Haifa University, Israel). MH was supported by an Inter-Institutional post-doctoral 636 fellowship from the Haifa University and a Helmsley Trust fellowship.

637

638 **Conflict of interest**

639 The authors declare no conflict of interest.

640

641 References

- 642 Alonso-Sáez, L., V. Balagué, E. L. Sà, and others. 2007. Seasonality in bacterial diversity in
- 643 north-west Mediterranean coastal waters: assessment through clone libraries,
- 644 fingerprinting and FISH. FEMS microbiology ecology 60: 98–112. doi:10.1111/j.1574645 6941.2006.00276.x
- 646 Berman, T., Y. Azov, A. Schneller, P. Walline, and D. Townsend. 1986. Extent, transparency
- and phytoplankton distribution of the neritic waters overlying the Israeli coastal shelf.
 Oceanologica Acta 9: 439–447.
- 649 Bryant, J. A., F. O. Aylward, J. M. Eppley, D. M. Karl, M. J. Church, and E. F. DeLong. 2016.
- 650 Wind and sunlight shape microbial diversity in surface waters of the North Pacific

651 Subtropical Gyre. The ISME Journal **10**: 1308–1322. doi:10.1038/ismej.2015.221

- 652 Camarena-Gómez, M. T., T. Lipsewers, J. Piiparinen, E. Eronen-Rasimus, D. Perez-Quemaliños,
- L. Hoikkala, C. Sobrino, and K. Spilling. 2018. Shifts in phytoplankton community
- 654 structure modify bacterial production, abundance and community composition. Aquatic
- 655 Microbial Ecology **81**: 149–170. doi:10.3354/ame01868
- 656 Carlson, C. A., R. Morris, R. Parsons, A. H. Treusch, S. J. Giovannoni, and K. Vergin. 2009.
- 657 Seasonal dynamics of SAR11 populations in the euphotic and mesopelagic zones of the 658 northwestern Sargasso Sea. The ISME Journal **3**: 283–295. doi:10.1038/ismej.2008.117
- 659 Chao, A., N. J. Gotelli, T. C. Hsieh, E. L. Sander, K. H. Ma, R. K. Colwell, and A. M. Ellison.
- 660 2014. Rarefaction and extrapolation with Hill numbers: a framework for sampling and

661	estimation	in species	diversity	v studies.	Ecological	Monograph	s 84: 45–	-67.
001						1.10110 41001		<i>. . .</i>

662 doi:10.1890/13-0133.1

- 663 Clarke, K. R., P. J. Somerfield, and R. N. Gorley. 2008. Testing of null hypotheses in
- 664 exploratory community analyses: similarity profiles and biota-environment linkage.
- Journal of Experimental Marine Biology and Ecology **366**: 56–69.
- 666 doi:10.1016/j.jembe.2008.07.009
- 667 D'Alimonte, D., and G. Zibordi. 2003. Phytoplankton determination in an optically complex

668 coastal region using a multilayer perceptron neural network. IEEE Transactions on

669 Geoscience and Remote Sensing **41**: 2861–2868. doi:10.1109/tgrs.2003.817682

670 Dubinsky, V., M. Haber, I. Burgsdorf, and others. 2017. Metagenomic analysis reveals unusually

671 high incidence of proteorhodopsin genes in the ultraoligotrophic Eastern Mediterranean

672 Sea. Environmental microbiology **19**: 1077–1090. doi:10.1111/1462-2920.13624

673 Efrati, S., Y. Lehahn, E. Rahav, and others. 2013. Intrusion of coastal waters into the pelagic

674 eastern Mediterranean: in situ and satellite-based characterization. Biogeosciences **10**:

- 675 3349–3357. doi:10.5194/bg-10-3349-2013
- Eiler, A., D. H. Hayakawa, M. J. Church, D. M. Karl, and M. S. Rappé. 2009. Dynamics of the
- 677 SAR11 bacterioplankton lineage in relation to environmental conditions in the
- 678 oligotrophic North Pacific subtropical gyre. Environmental microbiology **11**: 2291–2300.
- 679 doi:10.1111/j.1462-2920.2009.01954.x
- 680 Falkowski, P. G., T. Fenchel, and E. F. Delong. 2008. The Microbial Engines That Drive Earth's
- 681 Biogeochemical Cycles. Science **320**: 1034–1039. doi:10.1126/science.1153213

- 682 Feingersch, R., M. T. Suzuki, M. Shmoish, I. Sharon, G. Sabehi, F. Partensky, and O. Béjà.
- 683 2010. Microbial community genomics in eastern Mediterranean Sea surface waters. The
- 684 ISME Journal 4: 78–87. doi:10.1038/ismej.2009.92
- 685 Fortunato, C. S., L. Herfort, P. Zuber, A. M. Baptista, and B. C. Crump. 2011. Spatial variability
- 686 overwhelms seasonal patterns in bacterioplankton communities across a river to ocean
- 687 gradient. The Isme Journal **6**: 554.
- Fuhrman, J. A., J. A. Cram, and D. M. Needham. 2015. Marine microbial community dynamics
 and their ecological interpretation. Nature reviews. Microbiology 13: 133–146.
- 690 doi:10.1038/nrmicro3417
- 691 Fuhrman, J. A., I. Hewson, M. S. Schwalbach, J. A. Steele, M. V. Brown, and S. Naeem. 2006.
- 692 Annually reoccurring bacterial communities are predictable from ocean conditions.
- 693 Proceedings of the National Academy of Sciences of the United States of America **103**:
- 694 13104–13109. doi:10.1073/pnas.0602399103
- 695 Galand, P. E., O. Pereira, C. Hochart, J. C. Auguet, and D. Debroas. 2018. A strong link between
- 696 marine microbial community composition and function challenges the idea of functional

697 redundancy. The ISME journal. doi:10.1038/s41396-018-0158-1

- Ghiglione, J.-F., M. Larcher, and P. Lebaron. 2005. Spatial and temporal scales of variation in
 bacterioplankton community structure in the NW Mediterranean Sea. Aquatic Microbial
 Ecology 40: 229–240.
- Gilbert, J. A., J. A. Steele, J. G. Caporaso, and others. 2012. Defining seasonal marine microbial
 community dynamics. The ISME journal 6: 298–308. doi:10.1038/ismej.2011.107
- 703 Giovannoni, S. J., and K. L. Vergin. 2012. Seasonality in ocean microbial communities. Science

704 (New York, N.Y.) **335**: 671–676. doi:10.1126/science.1198078

- 705 Green, S. J., R. Venkatramanan, and A. Naqib. 2015. Deconstructing the Polymerase Chain
- 706 Reaction: Understanding and Correcting Bias Associated with Primer Degeneracies and
- 707 Primer-Template Mismatches V.M. Ugaz [ed.]. PLOS ONE **10**: e0128122.
- 708 doi:10.1371/journal.pone.0128122
- Hammer, O., D. Harper, and P. Ryan. 2001. PAST: Paleontological statistics software package
 for education. Palaeontologia Electronica 4: 9pp.
- 711 Haro-Moreno, J. M., M. Lopez-Perez, J. R. de la Torre, A. Picazo, A. Camacho, and F.
- 712 Rodriguez-Valera. 2018. Fine metagenomic profile of the Mediterranean stratified and
- 713 mixed water columns revealed by assembly and recruitment. Microbiome 6: 128.
- 714 doi:10.1186/s40168-018-0513-5
- 715 Hazan, O., J. Silverman, G. Sisma-Ventura, T. Ozer, I. Gertman, E. Shoham-Frider, N. Kress,
- and E. Rahav. 2018. Mesopelagic Prokaryotes Alter Surface Phytoplankton Production
- 717 during Simulated Deep Mixing Experiments in Eastern Mediterranean Sea Waters. Front.
- 718 Mar. Sci. 5. doi:10.3389/fmars.2018.00001
- Hsieh, T. C., K. H. Ma, and A. Chao. 2018. iNEXT: iNterpolation and EXTrapolation for species
 diversity,.
- Keuter, S., E. Rahav, B. Herut, and B. Rinkevich. 2015. Distribution patterns of bacterioplankton
 in the oligotrophic south-eastern Mediterranean Sea. FEMS Microbiol Ecol 91: fiv070.
- 723 doi:10.1093/femsec/fiv070
- 724 Kress, N., and B. Herut. 2001. Spatial and seasonal evolution of dissolved oxygen and nutrients
- in the Southern Levantine Basin (Eastern Mediterranean Sea): chemical characterization
- of the water masses and inferences on the N:P ratios. Deep Sea Research Part I:
- 727 Oceanographic Research Papers **48**: 2347–2372. doi:10.1016/s0967-0637(01)00022-x

728	Krom, M. D., S. Brenner, N. Kress, A. Neori, and L. I. Gordon. 1992. Nutrient dynamics and
729	new production in a warm-core eddy from the Eastern Mediterranean Sea. Deep Sea
730	Research Part A. Oceanographic Research Papers 39: 467–480. doi:10.1016/0198-
731	0149(92)90083-6
732	Krom, M. D., KC. Emeis, and P. V. Cappellen. 2010. Why is the Eastern Mediterranean
733	phosphorus limited? Progress in Oceanography 85: 236-244.
734	doi:10.1016/j.pocean.2010.03.003
735	Krom, M. D., N. Kress, S. Brenner, and L. I. Gordon. 1991. Phosphorus limitation of primary
736	productivity in the eastern Mediterranean Sea. Limnology and Oceanography 36: 424-
737	432. doi:10.4319/lo.1991.36.3.0424
738	Krom, M. D., E. M. S. Woodward, B. Herut, and others. 2005. Nutrient cycling in the south east
739	Levantine basin of the eastern Mediterranean: Results from a phosphorus starved system.
740	Deep Sea Research Part II: Topical Studies in Oceanography 52: 2879–2896.
741	doi:10.1016/j.dsr2.2005.08.009
742	Krom, M., N. Kress, I. Berman-Frank, and E. Rahav. 2014. Past, Present and Future Patterns in

- the Nutrient Chemistry of the Eastern Mediterranean. 49–68. doi:10.1007/978-94-0076704-1 4
- 745 Kumar, L., and M. E. Futschik. 2007. Mfuzz: A software package for soft clustering of

746 microarray data. Bioinformation **2**: 5–7. doi:10.6026/97320630002005

- 747 Lehahn, Y., I. Koren, S. Sharoni, F. d'Ovidio, A. Vardi, and E. Boss. 2017. Dispersion/dilution
- enhances phytoplankton blooms in low-nutrient waters. Nature Communications 8:
- 749 14868–14868. doi:10.1038/ncomms14868

750	Love, M. I., W. Huber, and S. Anders. 2014. Moderated estimation of fold change and dispersion								
751	for RNA-seq data with DESeq2. Genome biology 15: 550. doi:10.1186/s13059-014-								
752	0550-8								
753	Lucas, J., A. Wichels, and G. Gerdts. 2016. Spatiotemporal variation of the bacterioplankton								
754	community in the German Bight: from estuarine to offshore regions. Helgoland Marine								
755	Research 70: 16. doi:10.1186/s10152-016-0464-9								
756	Man-Aharonovich, D., N. Kress, E. B. Zeev, I. Berman-Frank, and O. Béjà. 2007. Molecular								
757	ecology of nifH genes and transcripts in the eastern Mediterranean Sea. Environ								
758	Microbiol 9: 2354–2363. doi:10.1111/j.1462-2920.2007.01353.x								
759	Moonsamy, P. V., T. Williams, P. Bonella, and others. 2013. High throughput HLA genotyping								
760	using 454 sequencing and the Fluidigm Access Array TM system for simplified amplicon								
761	library preparation: High throughput HLA 454 sequencing using the Fluidigm Access								
762	Array TM system. Tissue Antigens 81: 141–149. doi:10.1111/tan.12071								
763	Morales, S. E., M. Meyer, K. Currie, and F. Baltar. 2018. Are oceanic fronts ecotones? Seasonal								
764	changes along the subtropical front show fronts as bacterioplankton transition zones but								
765	not diversity hotspots. Environ Microbiol Rep 10: 184–189. doi:10.1111/1758-								
766	2229.12618								
767	Morris, R. M., C. D. Frazar, and C. A. Carlson. 2012. Basin-scale patterns in the abundance of								
768	SAR11 subclades, marine Actinobacteria (OM1), members of the Roseobacter clade and								

769 OCS116 in the South Atlantic. Environmental microbiology **14**: 1133–1144.

770 doi:10.1111/j.1462-2920.2011.02694.x

771 Oksanen, J., F. G. Blanchet, M. Friendly, and others. 2017. vegan: Community Ecology
772 Package,.

773	Parada, A. E., D. M. Needham, and J. A. Fuhrman. 2016. Every base matters: assessing small
774	subunit rRNA primers for marine microbiomes with mock communities, time series and
775	global field samples. Environmental microbiology 18: 1403–1414. doi:10.1111/1462-
776	2920.13023
777	Paulson, J. N., O. C. Stine, H. C. Bravo, and M. Pop. 2013. Differential abundance analysis for
778	microbial marker-gene surveys. Nature methods 10: 1200-1202. doi:10.1038/nmeth.2658
779	Powley, H. R., M. D. Krom, and P. V. Cappellen. 2017. Understanding the unique
780	biogeochemistry of the Mediterranean Sea: Insights from a coupled phosphorus and
781	nitrogen model: P and N Cycling in the Mediterranean Sea. Global Biogeochemical
782	Cycles 31 : 1010–1031. doi:10.1002/2017gb005648
783	Quero, G. M., and G. M. Luna. 2014. Diversity of rare and abundant bacteria in surface waters of
784	the Southern Adriatic Sea. Mar Genomics 17: 9–15. doi:10.1016/j.margen.2014.04.002
785	Raveh, O., N. David, G. Rilov, and E. Rahav. 2015. The Temporal Dynamics of Coastal
786	Phytoplankton and Bacterioplankton in the Eastern Mediterranean Sea. PloS one 10:
787	e0140690. doi:10.1371/journal.pone.0140690
788	Salter, I., P. E. Galand, S. K. Fagervold, P. Lebaron, I. Obernosterer, M. J. Oliver, M. T. Suzuki,
789	and C. Tricoire. 2015. Seasonal dynamics of active SAR11 ecotypes in the oligotrophic
790	Northwest Mediterranean Sea. The ISME journal 9: 347-360.
791	doi:10.1038/ismej.2014.129
792	Schloss, P. D., S. L. Westcott, T. Ryabin, and others. 2009. Introducing mothur: open-source,
793	platform-independent, community-supported software for describing and comparing
794	microbial communities. Applied and environmental microbiology 75: 7537-7541.
795	doi:10.1128/AEM.01541-09

- Smyth, T. J. 2011. Penetration of UV irradiance into the global ocean. Journal of Geophysical
 Research 116. doi:10.1029/2011jc007183
- 798 Sunagawa, S., L. P. Coelho, S. Chaffron, and others. 2015. Ocean plankton. Structure and
- function of the global ocean microbiome. Science (New York, N.Y.) 348: 1261359–
- 800 1261359. doi:10.1126/science.1261359
- 801 Tanaka, T., T. F. Thingstad, U. Christaki, and others. 2011. Lack of P-limitation of
- 802 phytoplankton and heterotrophic prokaryotes in surface waters of three anticyclonic
- 803 eddies in the stratified Mediterranean Sea. Biogeosciences 8: 525–538. doi:10.5194/bg-8-
- 804 525-2011
- 805 Tedetti, M., and R. Sempéré. 2006. Penetration of ultraviolet radiation in the marine

806 environment. A review. Photochemistry and photobiology **82**: 389–397.

- 807 doi:10.1562/2005-11-09-IR-733
- 808 Thingstad, T. F., M. D. Krom, R. F. C. Mantoura, and others. 2005. Nature of phosphorus
- 809 limitation in the ultraoligotrophic eastern Mediterranean. Science (New York, N.Y.) **309**:
- 810 1068–1071. doi:10.1126/science.1112632
- 811 Tinta, T., J. Vojvoda, P. Mozetič, I. Talaber, M. Vodopivec, F. Malfatti, and V. Turk. 2015.
- 812 Bacterial community shift is induced by dynamic environmental parameters in a
- 813 changing coastal ecosystem (northern Adriatic, northeastern Mediterranean Sea) a 2-
- 814 year time-series study: Bacterial community shift in a dynamic coastal ecosystem.
- 815 Environmental Microbiology **17**: 3581–3596. doi:10.1111/1462-2920.12519
- 816 Tsiola, A., P. Pitta, S. Fodelianakis, and others. 2016. Nutrient Limitation in Surface Waters of
- 817 the Oligotrophic Eastern Mediterranean Sea: an Enrichment Microcosm Experiment.
- 818 Microbial Ecology **71**: 575–588. doi:10.1007/s00248-015-0713-5

819	Vergin	K L	R	Beszteri	Α	Monier	and	others	2013	High-re	solution	SAR11	ecotype
017	vergin,	N . L.	, р.	Deszterr	, л.	withit	anu	ouncis.	2015.	ingn-ic	solution	SAKII	ccotype

- 820 dynamics at the Bermuda Atlantic Time-series Study site by phylogenetic placement of
- 821 pyrosequences. The ISME journal 7: 1322–1332. doi:10.1038/ismej.2013.32
- 822 Volpe, G., R. Santoleri, V. Vellucci, M. R. d'Alcalà, S. Marullo, and F. D'Ortenzio. 2007. The
- 823 colour of the Mediterranean Sea: Global versus regional bio-optical algorithms evaluation
- and implication for satellite chlorophyll estimates. Remote Sensing of Environment **107**:
- 825 625–638. doi:10.1016/j.rse.2006.10.017
- 826 Wang, Z., D. L. Juarez, J.-F. Pan, S. K. Blinebry, J. Gronniger, J. S. Clark, Z. I. Johnson, and D.
- E. Hunt. 2019. Microbial communities across nearshore to offshore coastal transects are
- 828 primarily shaped by distance and temperature. Environmental Microbiology 21: 3862–
- 829 3872. doi:10.1111/1462-2920.14734
- 830 Ward, C. S., C.-M. Yung, K. M. Davis, S. K. Blinebry, T. C. Williams, Z. I. Johnson, and D. E.
- 831 Hunt. 2017. Annual community patterns are driven by seasonal switching between
- closely related marine bacteria. The ISME journal **11**: 1412–1422.
- 833 doi:10.1038/ismej.2017.4
- 834 West, N. J., C. Lepère, C.-L. de O. Manes, P. Catala, D. J. Scanlan, and P. Lebaron. 2016.
- B35 Distinct Spatial Patterns of SAR11, SAR86, and Actinobacteria Diversity along a
- 836 Transect in the Ultra-oligotrophic South Pacific Ocean. Frontiers in microbiology 7: 234.
- 837 doi:10.3389/fmicb.2016.00234
- 838 Yogev, T., E. Rahav, E. Bar-Zeev, and others. 2011. Is dinitrogen fixation significant in the
- 839 Levantine Basin, East Mediterranean Sea?: Eastern Mediterranean Sea N2 fixation.
- 840 Environmental Microbiology **13**: 854–871. doi:10.1111/j.1462-2920.2010.02402.x
- 841