

Planktonic protist diversity across contrasting Subtropical and Subantarctic waters of the southwest Pacific

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1 ABSTRACT

2 Planktonic protists are an essential component of marine pelagic ecosystems where they mediate important
3 trophic and biogeochemical functions. Although these functions are largely influenced by their taxonomic
4 affiliation, the composition and spatial variability of planktonic protist communities remain poorly
5 characterized in vast areas of the ocean. Here, we investigated the diversity of these communities in
6 contrasting oceanographic conditions of the southwest Pacific sector (33-58°S) using DNA metabarcoding
7 of the 18S rRNA gene. Seawater samples collected during twelve cruises (n = 482, 0-2000 m) conducted
8 east of New Zealand were used to characterize protist communities in Subtropical (STW) and Subantarctic
9 (SAW) water masses and the Subtropical Front (STF) that separates them. Diversity decreased with latitude
10 and temperature but tended to be lowest in the STF. Sample ordination resulting from the abundance
11 of amplicon single variants (ASVs) corresponded to the different water masses. Overall, *Dinophyceae*
12 (34% of standardized total number of reads) and Chlorophyta (27%) co-dominated the euphotic zone,
13 but their relative abundance and composition at class and lower taxonomic levels varied consistently
14 between water masses. Among Chlorophyta, several picoplanktonic algae species of the *Mamiellophyceae*
15 class including *Ostreococcus lucimarinus* dominated in STW, while the *Chloropicophyceae* species
16 *Chloroparvula pacifica* was most abundant in SAW. *Bacillariophyta* (7%), *Prymnesiophyceae* (5%), and
17 *Pelagophyceae* (3%) classes were less abundant but showed analogous water mass specificity at class
18 and finer taxonomic levels. Protist community composition in the STF had mixed characteristics and
19 showed regional differences with the southern STF (50°S) having more resemblance with subantarctic
20 communities than the STF over the Chatham Rise region (42-44°S). Below the euphotic zone, Radiolaria
21 sequences dominated the dataset (52%) followed by *Dinophyceae* (27%) and other heterotrophic groups
22 like Marine Stramenopiles and ciliates (3%). Among Radiolaria, several unidentified ASVs assigned to
23 *Spumellarida* were most abundant, but showed significantly different distribution between STW and SAW
24 highlighting the need to further investigate the taxonomy and ecology of this group. This study represents
25 a significant step forward towards characterizing protistan communities composition in relation to major
26 water masses and fronts in the South Pacific providing new insights about the biogeography and ecological
27 preferences of different taxa from class to species and genotypic level.

28 **Highlights**

- 29 • Water-mass preference of different taxa emerged at class, species and genotypic level.
- 30 • *Mamiellophyceae* green algae dominated in subtropical waters.
- 31 • *Dinophyceae* and *Chloropicophyceae* green algae dominated in subantarctic waters.
- 32 • A diverse assemblage of Radiolaria dominated the mesopelagic zone.
- 33 • Small rather than large taxa dominated phytoplankton blooms in subtropical waters.

34 **Keywords**

35 Planktonic protist, taxonomic diversity, 18S rRNA metabarcoding, Biogeography, southwest Pacific,
36 Subtropical, Subantarctic, Subtropical Front

37 **Competing interests**

38 The authors declare no competing financial interests.

39 1. INTRODUCTION

40 Planktonic protists, including phototrophic, heterotrophic and mixotrophic single-celled eukaryotes, have
41 key roles in the functioning of marine ecosystems (Caron et al. 2012). Phytoplankton are responsible
42 for 50% of global primary productivity (Field et al. 1998). Most of this primary production is consumed
43 and processed by heterotrophic protists (i.e. microzooplankton) before becoming available for larger
44 zooplankton and higher trophic levels (Calbet and Saiz 2005; Calbet and Landry 2004; Zeldis and Décima
45 2020). From a biogeochemical perspective, the microbial production, consumption and remineralization
46 of organic matter is at the core of global biogeochemical cycles including the nitrogen and carbon cycles,
47 and is pivotal in regulating the ocean's capacity to sequester atmospheric CO₂ via the biological carbon
48 pump (Boyd et al. 2019; Turner 2015).

49 The trophic and biogeochemical processes driven by microbial communities are influenced by their
50 taxonomic composition, which is tightly coupled to physico-chemical conditions. With increasing evidence
51 of climate change effects on the physico-chemical status of the ocean (e.g. warming, increased stratification
52 and reduced nutrient supply, and acidification) (Henley et al. 2020; Pörtner et al. 2014; Sarmiento et
53 al. 2004) it becomes imperative to better characterize the biogeography and distributions of microbial
54 communities in relation to oceanographic provinces (Cavicchioli et al. 2019). This is required to establish
55 a conceptual framework and baseline upon which future environmental change can be evaluated.

56 The diversity and dynamic nature of microbial communities has precluded a comprehensive characteri-
57 zation of species composition and distributional patterns across at relevant temporal and spatial scales
58 (Wietz et al. 2019). Extensive application of DNA metabarcoding approaches during the last 10 years
59 have contributed significantly to this end by characterizing the diversity of marine protist communities
60 over a wide range of temporal and spatial scales with unprecedented taxonomic resolution and coverage
61 (Santoferrara et al. 2020). Despite these efforts there are still vast ocean regions like the southwest
62 (SW) Pacific Ocean that due to its large size and remoteness remain largely unexplored with regards to
63 high-throughput sequencing characterization of protist communities composition and spatial distribution
64 of major taxonomic groups. This study contributes to fill this gap by investigating protist communities in
65 relation to major water masses and oceanographic fronts characteristic of the SW Pacific waters east of

66 New Zealand (Figure 1).

67 New Zealand's continental mass interrupts the converging flows of the South Pacific subtropical gyre
68 and the northward excursions of the Antarctic Circumpolar Current (ACC). The mixing of the warm
69 and saltier subtropical water (STW) with the cold, relatively fresh subantarctic water (SAW) (Boyd et al.
70 1999) results in the genesis of oceanic fronts and semi-permanent eddies with distinctive signatures in
71 water properties extending along the eastern margin off New Zealand (Fernandez et al. 2018). To the
72 north, the East Auckland Current (EAUC) brings STW sourced partially by the Tasman Front (Sutton and
73 Bowen 2014). At about 37°S the EAUC turns south to become the East Cape Current (Stanton et al. 1997)
74 extending the STW inflow to the Chatham Rise where it separates from the coast to the east as part of
75 the Subtropical Front (STF) (Deacon 1982). The STF is characterised by strong temperature gradients
76 and a sharp salinity contrast that intensifies near the rise (Smith et al. 2013), up to 4°C and 0.7 practical
77 salinity units respectively over 1° latitude in this region (Belkin and Gordon 1996). This transitional
78 zone separating waters of subtropical origin from the subantarctic ones is known as the Subtropical Front
79 Zone (SFTZ) (Deacon 1982) and it is bounded by the northern (N-STF) and southern (S-STF) branches
80 of the STF. The SFTZ can be up to 500 km wide in the Tasman Sea region before it gets constricted
81 around the South Island of New Zealand where gradients set in motion, guided by the continental slope,
82 the geostrophic flow associated with the S-STF branch and its coastal expression, the Southland Current
83 (SC). The mean transport of the SC is about 8 Sv ($1 \text{ Sv} = 10^6 \text{ m}^3 \text{ s}^{-1}$) with 10% corresponding to STW
84 and 90% to SAW (Sutton 2003). The SC advects this mix of STW and SAW northwards off the east
85 coast and reaches south of the Chatham Rise through the Mernoo Gap and the Bounty Through. Further
86 east and along the flanks of Campbell Plateau, the flows associated with the Subantarctic Front (SAF)
87 carry the largest portion of SAW, about 50 Sv into the region south and east of the Chatham Rise (Bowen
88 et al. 2014; Stanton and Morris 2004). Access of SAW onto the plateau from the east occurs through
89 the bathymetric gaps, saddles and ridges where waters then become isolated from the neighbouring
90 circulation and significantly contribute to the development of oceanographic and climatic processes such
91 as subantarctic mode water formation (Forcén-Vázquez et al. 2021). Southeast of the Chatham Rise and
92 away from the plateau the SFTZ re-emerges as a 150 km wide band with the typical signatures of the
93 STF-N and STF-S fronts (Sutton 2001).

94 STW and SAW have contrasting biogeochemical characteristics (Boyd et al. 1999; Bradford-Grieve
95 et al. 1999; Chiswell et al. 2015; Heath 1985; Sherlock et al. 2007). North of Aotearoa New Zealand, STW
96 is oligotrophic (low macro- and micronutrients) and phytoplankton production is considered to be limited
97 by nitrogen (Zentara and Kamykowski 1981) with pervasive nitrogen-fixation by diazotrophs (Ellwood
98 et al. 2018; Law et al. 2012). The STF is a dynamic region, characterized by strong temperature and
99 salinity gradients (Sutton 2001) where high levels of vertical and lateral mixing of nitrogen-limited STW
100 and macronutrient-rich SAW (Chiswell 2001), leads to regionally elevated annual net primary production
101 (Murphy et al. 2001; Pinkerton et al. 2005). In SAW iron is the primary limiting nutrient for phytoplankton
102 growth (Banse 1996; Boyd et al. 1999) although silicate and light can become limiting at times in SAW
103 extending southeast of Aotearoa New Zealand which is considered high-nutrient, low-chlorophyll, low-
104 silicate (HNLC-LSi) region (Boyd et al. 2010; Dugdale et al. 1995). These conditions are typically
105 associated with SAW north of the Subantarctic Front (SAF), which is an area commonly referred to as the
106 Subantarctic Zone (SAZ) (Trull et al. 2001) or the Subantarctic Water Ring (Longhurst 2007). In the SAZ,
107 increasing availability of dissolved silica southwards shifts the Polar Frontal Zone extending between the
108 SAF and the Polar Front to ‘standard’ Southern Ocean HNLC conditions (Rigual-Hernández et al. 2015).

109 Several studies have characterized microbial community composition in STW and SAW east of New
110 Zealand, using microscopy (Chang and Gall 1998), pigment (Delizo et al. 2007) and flow-cytometry
111 (Hall et al. 1999). These regional studies have focused mainly on the STF zone or coastal communities
112 (Chang et al. 2003; Hall et al. 2006), while studies analyzing wider phytoplankton distribution across
113 STW and SAW have targeted specific groups such as coccolithophores (Chang and Northcote 2016;
114 Saavedra-Pellitero et al. 2014). More process-oriented studies have also provided partial information
115 on phytoplankton composition in SAW and STW east of New Zealand (Chiswell et al. 2019; Ellwood
116 et al. 2013; Peloquin et al. 2011). These studies have described the prevalence of larger cells and diatoms
117 through winter and spring in the more productive waters of the STF, compared to STW and SAW. Diatom-
118 and autotrophic flagellate-dominated communities have been reported in STW on the northern flank of the
119 Chatham Rise during spring while dinoflagellates and small flagellates are documented as dominating
120 the eukaryotic phytoplankton in SAW (Bradford-Grieve et al. 1997; Chang and Gall 1998). Diatom
121 and coccolithophore species composition of sediment trap fluxes on the northern (STW-influenced) and

122 southern (SAW-influenced) flanks of the Chatham Rise highlight the importance of these phytoplankton
123 groups in the region (Wilks et al. 2021). However, there is surprisingly only little information available on
124 the taxonomic composition of phytoplankton communities prevailing in open-ocean waters away from
125 the STFZ over the Chatham Rise region (Chang and Northcote 2016; Chiswell et al. 2019; Peloquin et al.
126 2011; Twining et al. 2014). Further east of this region (170 °W), phytoplankton community composition
127 from polar to equatorial waters have been characterized using pigment analysis (DiTullio et al. 2003)
128 whereas a more recent study applied DNA metabarcoding analysis to investigate microbial diversity
129 patterns in relation to physico-chemical gradients and oceanographic features (Raes et al. 2018). DNA
130 metabarcoding has also been recently applied to investigate protist diversity changes across the Southland
131 Current, a coastal expression of the STF that flows along the eastern margin of New Zealand's South
132 Island (Allen et al. 2020). However, the taxonomic composition of protistan communities associated with
133 open-ocean water masses in the SW Pacific and across major oceanographic fronts that separate them is
134 still lacking.

135 The aims of the present study are: 1) to characterize the diversity of protistan communities in STW and
136 SAW east of New Zealand and across the STFZ that separates these water masses, and 2) to investigate the
137 spatial distributional patterns of the main protistan taxonomic groups and species in relation to physical
138 and chemical variability of the main water masses east of Aotearoa New Zealand. Specifically, we
139 want to know how (dis-)similar are protist communities in the biogeochemically contrasting STW and
140 SAW? What are the main environmental factors responsible for these differences? Which are the main
141 taxonomic groups associated with each water mass and their environmental preferences? To do so we have
142 applied DNA metabarcoding analysis (18S rRNA) to > 450 samples collected during 12 oceanographic
143 voyages conducted over several years (2009-2017) and different seasons across STW and SAW east of
144 New Zealand. This sequence data together with core physico-chemical and biological parameters (e.g.
145 temperature, salinity, mixed-layer depth (MLD), macronutrients, total and size-fractionated chlorophyll
146 a) provides the most comprehensive dataset of protistan plankton diversity in STW and SAW in the SW
147 Pacific and contributes significantly towards building a robust baseline against which future changes in the
148 region can be evaluated.

Cruise	Start	End	Project	Water mass	min Lat	max Lat	Season	N stations	N Samples
TAN0902	30-01-09	03-02-09	BiophysMoorings	SAW // STF // STW	-46.62	-41.23	Summer	3	28
TAN0909	27-10-09	30-10-09	BiophysMoorings	SAW // STF // STW	-46.64	-41.2	Spring	3	32
TAN1006	06-05-10	08-05-10	BiophysMoorings	SAW // STF // STW	-46.64	-41.19	Autumn	3	33
TAN1103	19-02-11	21-02-11	BiophysMoorings	SAW // STF // STW	-46.61	-41.31	Summer	3	34
TAN1113	29-09-11	01-10-11	BiophysMoorings	SAW // STF // STW	-46.63	-41.22	Spring	3	34
TAN1203	17-02-12	05-03-12	SOAP	STF	-44.61	-43.48	Summer	10	10
TAN1204	19-03-12	21-03-12	BiophysMoorings	SAW // STF // STW	-46.64	-41.26	Autumn	4	32
TAN1212	19-09-12	05-10-12	Spring Bloom II	STW	-37.87	-37.51	Spring	19	105
KAH1303	08-03-13	14-03-13	Bay of Plenty	STW	-39.17	-38.76	Autumn	12	37
TAN1516	05-12-15	21-12-15	Chatham Rise	STF	-44.42	-42.72	Summer	20	39
TAN1604	14-05-16	21-05-17	Cross-shelf Exchange	STW	-36.18	-33.38	Autumn	7	42
TAN1702	18-03-17	29-03-17	Campbell Plateau	SAW	-54.26	-46.77	Autumn	13	52
TAN1802	13-02-18	13-02-18	SO-RossSea	SAW	-58.03	-58.03	Summer	1	5

Table 1. Summary of cruises from which samples were collected. Information includes the cruise identification code, start and end dates, the project or region, the water masses surveyed, latitudinal and seasonal coverage, number of stations and samples collected in each cruise.

149 2. METHODS

150 2.1. Study area and sample collection

151 Seawater samples and data were collected during 12 research cruises conducted in SW Pacific waters east
152 of Aotearoa New Zealand between 2009 and 2017 (Figure 1). The dataset covered 100 stations distributed
153 between 54.3 and 33.4 °S with seawater samples (n = 482) collected over the 0-2000 m depth range and
154 covering spring, summer and autumn periods (Table 1). The number of DNA samples from STW (n=269)
155 were 2-fold higher than those from SAW (n=120) and STF (n=94) mainly due to the large number of
156 samples from the Spring Bloom II voyage (TAN1212,(Chiswell et al. 2019)) (Table 1). The seasonality
157 coverage was similar among the three different regions (STW, SAW, STF) but was biased against winter
158 with most samples collected during spring, summer and autumn periods. (Table 1; Figure S1). Details
159 about latitudinal and seasonal coverage of each water mass and the sample density distribution of analysed
160 DNA samples, together with physico-chemical variables, are shown in Figure S2.

161 Samples were collected from 10 L Niskin bottles attached to a CTD rosette in association with

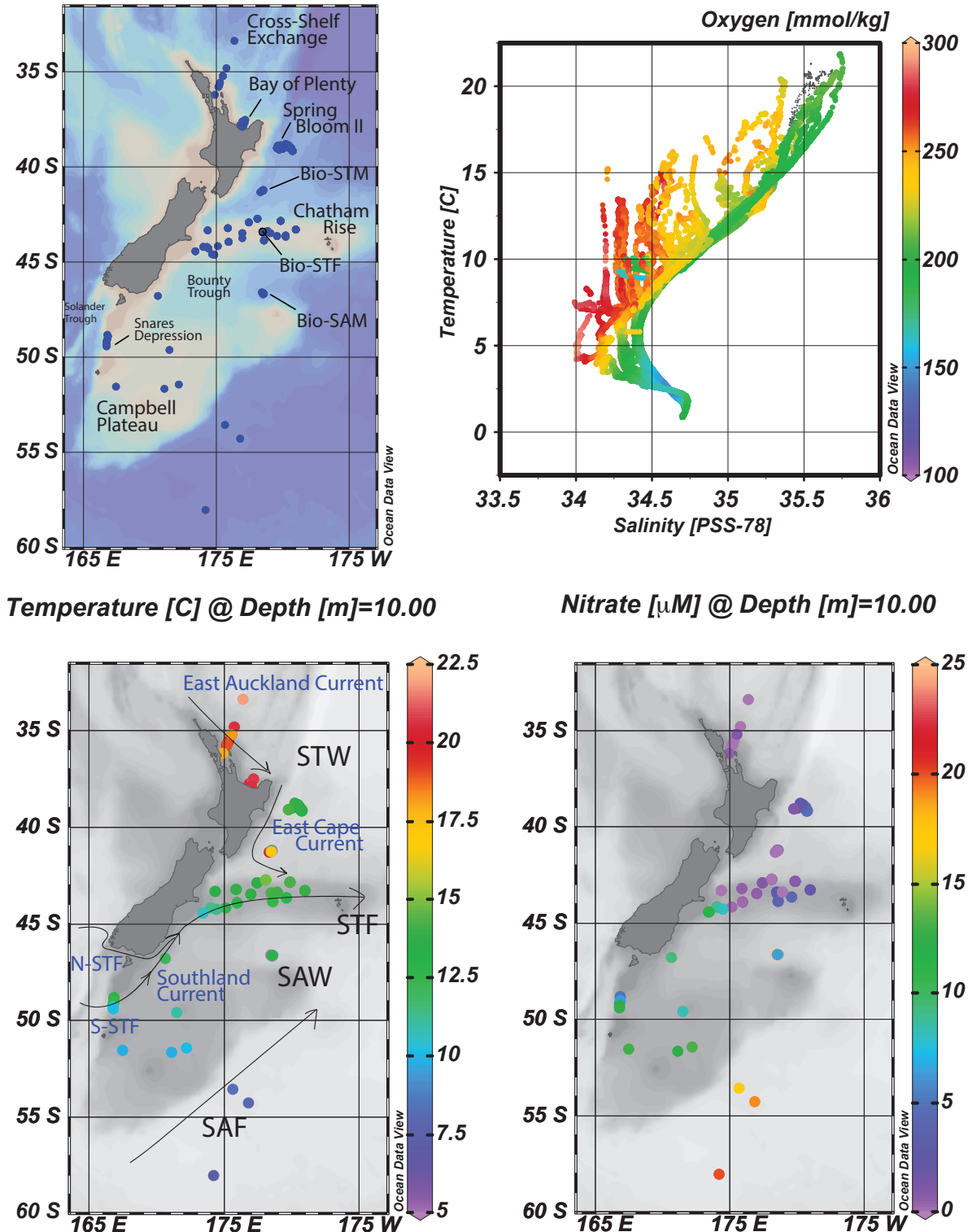


Figure 1. Study area. (A) Map of the study area with the sampling sites locations. (B) T-S diagram with oxygen concentration. Surface (C) temperature (C) and (D) nitrate concentration (D) at sampling sites in relation to major water masses and currents and fronts of the study region. North (N-STF) and South Subtropical Front (S-STF) adapted from Smith et al. 2013

162 a Seabird 9plus CTD, equipped with temperature, salinity, dissolved oxygen, fluorometer and beam
163 transmissiometer sensors. K_d PAR was estimated from chlorophyll *a* (Chl *a*) (Morel and Maritorena 2001).
164 The euphotic zone depth (Zeu) was defined as the depth where downwelling PAR irradiance was 1%
165 of incident irradiance (E_0). The MLD was defined as the shallowest depth where density exceeded the
166 5 m value by 0.03 kg/m^3 (Gardner et al. 1995). During the TAN1516 voyage samples were collected
167 with a Niskin bottle deployed manually down to 10 m depth and from the R/V *Tangaroa* Underway
168 Flow-Through System (TUFTS) system equipped with temperature, salinity, and fluorescence sensors.
169 Seawater samples for nutrients, Chl *a* and DNA were sampled from the Niskin bottles using acid-washed
170 silicone tubing and filtered through different types of filters for processing, as outlined below.

171 **2.2. Nutrients, total and size-fractionated chlorophyll a**

172 Samples for nutrients were filtered through 25 mm-diameter Whatman GF/F filters into clean 250 ml
173 polyethylene bottles and frozen at -20°C until analysis using an Astoria Pacific API 300 microsegmented
174 flow analyzer (Astoria-Pacific, Clackamas, OR, United States) according to the colorimetric methods
175 described in (Law et al. 2011).

176 For total Chl *a*, 250-400 mL seawater were filtered under low vacuum ($<200 \text{ mm Hg}$) through 25
177 mm GF/F filters. These were folded and wrapped in aluminum foil or placed in Secol envelopes and
178 stored at -80°C or in liquid nitrogen until analysis. For size-fractionated Chl *a* ($0.2\text{-}2 \mu\text{m}$, $2\text{-}20 \mu\text{m}$,
179 $>20 \mu\text{m}$) 400-500 mL were filtered sequentially through 47 mm polycarbonate filters by vacuum. Filters
180 were folded and stored in 1.5 mL cryovials at -80°C until analysis using 90% acetone extraction by
181 spectrofluorometric techniques on a Varian Cary Eclipse fluorometer following method APHA 10200 H
182 (Baird 2017)

183 **2.3. DNA samples collection and extraction**

184 Seawater samples of 1.5–5.0 L were filtered either through $0.22 \mu\text{m}$ filters (47 mm-diameter polyether-
185 sulfone, Pall-Gelman) using low vacuum or through $0.22 \mu\text{m}$ Sterivex filter units (Millipore) using a
186 peristaltic pump (Cole-Palmer). Disc filters were then folded and placed in cryovials and sterivex units
187 were filled with RNAlater and flash-frozen in liquid nitrogen prior to storing at -80°C . Disc filters were
188 cut in two halves first and then into small pieces using a sterile blade. Each half was placed in separate

189 tubes and lysed in parallel (2 h at 65 °C on a Boekel thermomixer set at 750 rpm) using the Nucleospin
190 Plant kit Midi Kit (Macherey-Nagel, Duren, Germany). The 100 µL of PL2 buffer recovered from each
191 halved filter was then pooled together and the DNA extraction procedure continued with the Mini version
192 of the Nucleospin Plant kit.

193 For Sterivex filters DNA was extracted using a Tris-buffered lysis solution containing EDTA, Triton
194 X 100 and lysozyme (pH = 8.0) and the Qiagen DNeasy Blood & Tissue. Briefly, to collect cells that
195 eventually detached from the filter surface, the RNAlater present in the filter unit was collected into a 2 mL
196 Eppendorf tube using a syringe and then centrifuged (13,000 rpm, 10 min). The pellet was resuspended
197 using 1 mL of the lysis solution and pipetted back into the original Sterivex. The cartridge was secured
198 using Parafilm, put into a 50 mL falcon tube and incubated in a shaking incubator overnight (75 rpm, 37
199 °C). 1 mL of buffer Qiagen buffer AL and 40 µL of proteinase K (20 mg/mL) was then added into the
200 Sterivex. After securing the Sterivex, as described previously, the filter unit was incubated for 2 hours (75
201 rpm, 56 °C). Followed the incubations the lysate was recovered from the cartridge and DNA extraction
202 and purification continued following manufacturer's instructions in the Qiagen DNeasy Blood & Tissue.

203 **2.4. PCR amplification, amplicon sequencing and sequence processing**

204 The V4 region of the 18S rRNA gene was amplified using the eukaryotic primers V4F_illum (CCAGCAS-
205 CYGCGGTAATTCC) and V4R_illum (ACTTTCGTTCTTGATYRATGA) with Illumina overhang adapters
206 (TCGTCGGCAGCGTCAGATGTGTATAAGAGACAG and GTCTCGTGGGCTCGGAGATGTGTATAA-
207 GAGACAG) following procedures described in (Gutierrez-Rodriguez et al. 2019). PCR reactions were
208 prepared in 50 µL using 2x KAPA HiFi HotStart ReadyMix, (0.3 nM dNTP, 2.5 mM MgCl₂), 0.5 µM of
209 each primer and of DNA template. The thermocycling profile included 95 °C/3 min, 10 cycles (98 °C/10s,
210 44 °C/20s, 72 °C/15s), 15 cycles (98 °C/10s, 62 °C/20s, 72 °C/15s) and 72 °C/7min).

211 Amplicon sequencing was conducted at the Genotoul GeT core facility (Toulouse, France) using an
212 Illumina Miseq and a 2 x 250 cycles Miseq kit version 2. 482 samples were sequenced generating a total
213 of 9166190 reads, with a median sequencing depth across all samples of 18954 reads per sample (range
214 = 6539 – 36551). Obtained sequences were processed using the DADA2 pipeline (Callahan et al. 2017)
215 following the procedure described by (Trefault et al. 2021). Taxonomic assignation was made against

216 PR2 database version 4.12 (<https://pr2-database.org/>), yielding 16861 amplicon single variants (ASVs)
217 assigned to protistan taxa. Details on the number of samples, reads and ASVs associated with each water
218 mass are shown in (Table S1).

219 **2.5. Pre-processing of ASV table and diversity analysis**

220 We standardized the ASV table to the sequencing depth in each sample by normalizing the relative
221 abundance to the mean number of sequences obtained across samples (median sequencing depth * (number
222 of reads per ASV /total number reads per sample)). The relative contribution of specific groups in
223 different water masses and regions were estimated from the sum of standardized reads across the samples
224 considered.

225 Similarity analysis was undertaken using non-multidimensional scaling (nMDS) and ANOSIM using
226 phyloseq (McMurdie and Holmes 2013) and vegan (Oksanen et al. 2019) R packages. Differences in
227 species abundance across waters masses and regions was analysed using negative binomial generalized
228 linear models coded in the DESeq2 package (Love et al. 2014). For analysis of higher taxonomic rank
229 (division, class), distribution and their relation to environmental variables, the tax_glom function in
230 phyloseq was used to agglomerate the previously standardized ASV table into the chosen taxonomic level.

231 **2.6. Data accessibility**

232 Physico-chemical, biological and geographic data including measurements of temperature, salinity, oxygen
233 and transmissivity obtained with a Seabird 9plus; inorganic nutrients; total and size-fractionated chl *a*);
234 and estimated mixed-layer and euphotic zone depth can be found PANGAEA repository archive data
235 sets (submitted to PANGAEA). Raw sequence data are available on NCBI under accession number
236 PRJNA756172 (). The quality filtered ASV table together with the taxonomy and metadata table used to
237 build the phyloseq object are provided as csv tables. R scripts for data processing and figures can be found
238 in https://github.com/agutierrez2001/Catalyst_Biogeography.

239 3. RESULTS

240 3.1. Physical, chemical and biological variability

241 STW was identified as those waters with surface salinity >35 psu (range = 35.1 - 35.6) (Figure 2) and
242 included samples collected during the Cross-shelf Exchange II (TAN1604), the Bay of Plenty (KAH1303),
243 the Spring Bloom II (TAN1212) voyages, and several cruises that visited the northern mooring site
244 (Bio-STM) of the Biophysical Moorings long-term monitoring program (Nodder et al. 2016) (Figure 1,
245 Table 1 and Figure S1). The Subtropical Front (STF) had salinity values ranging between 34.4 and 35.0 psu
246 (Figure 2) and included samples collected on the Chatham Rise during TAN1516 (Fisheries Oceanography
247 IV) and several Biophysical Mooring voyages as well as those collected in the STF upstream of the
248 Chatham Rise, as it flows northwestward between the South Island and the Campbell Plateau (TAN1702)
249 (Figure 1, Table 1, Figure S1). Typically SAW had salinity values <34.4 (Figure 2) and included samples
250 collected at the southern Biophysical Mooring site located in the Bounty Trough (Bio-SAM), and at
251 several stations on Campbell Plateau, and in the SAF south of the plateau (TAN1702, TAN1802) (Figure 1,
252 Table 1, Figure S1).

253 Sea-surface temperature was on average lowest in SAW (10.7 ± 2.4 °C, mean \pm standard deviation,
254 sd), intermediate in the STF (13.1 ± 1.7 °C) and highest in STW (16.1 ± 3.2 °C) (Figure 2). Temperature
255 showed greater overlap among water masses and regions than surface salinity (Figure 2). STW sampled
256 during the Spring Bloom II voyage, for instance, showed surface temperature consistently lower than
257 15 °C (12.5 - 14.5 °C) (Figure 2). Nitrate concentrations were lowest in STW (2.98 ± 1.96 $\mu\text{mol/L}$),
258 intermediate and more variable in STF waters (4.28 ± 4.17 $\mu\text{mol/L}$) and highest in SAW (12.17 ± 4.02
259 $\mu\text{mol/L}$), consistent with HNLC conditions of these southern waters (Figure 1D).

260 Chl *a* concentration in the surface mixed-layer was on average higher in the STF (0.65 ± 0.65 $\mu\text{g/L}$)
261 compared to STW (0.38 ± 0.31 $\mu\text{g/L}$) and SAW (0.37 ± 0.23 $\mu\text{g/L}$) (ANOVA, $F(2,220) = 14.2$, $p < 0.0001$).
262 Most of the DNA samples included in this study were taken from oligotrophic and mesotrophic waters
263 (surface Chl *a* <0.5 $\mu\text{g/L}$) with only a few collected from waters with Chl *a* concentrations >1.0 $\mu\text{g/L}$. The
264 smallest size-fraction (0.2 – 2.0- μm Chl *a*) dominated the phytoplankton communities across all water
265 masses but more so in STW (Chl *a* <2.0 $\mu\text{g/L}$ \approx 75 % of total Chl *a*) compared to SAW and STF (40-50%,

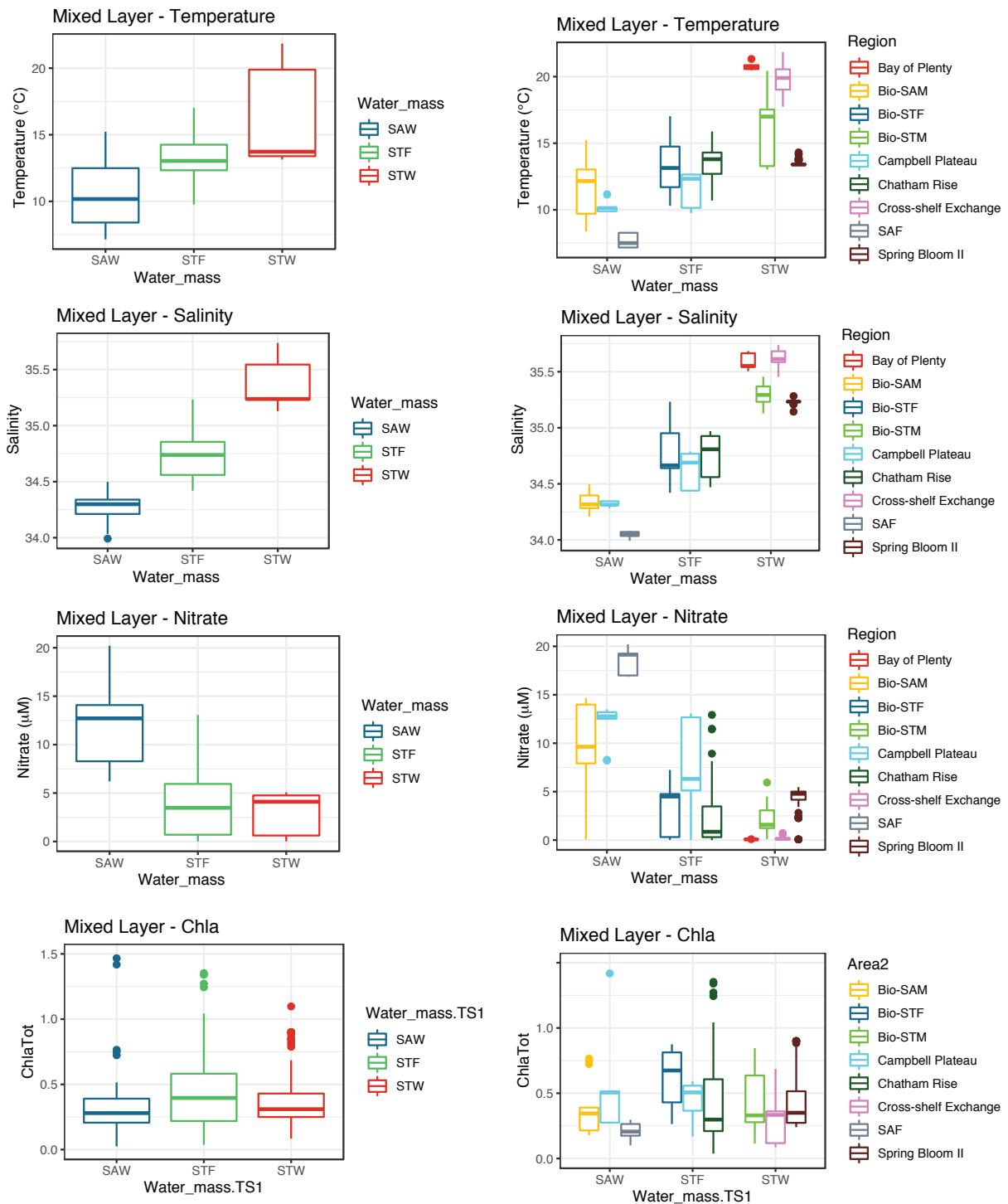


Figure 2. Surface mixed-layer physico-chemical variability. Box-plot representation of surface mixed-layer temperature and salinity, nitrate and chlorophyll a concentration in each water mass. Box-plots show the median, the first and third quartiles (lower and upper hinges) and the values within (line) and outside (dots) the $\pm 1.5 * IQR$ (IQR, interquartile range).

266 Figure S3). The contribution of $>20\text{-}\mu\text{m}$ size-fraction to surface mixed layer Chl *a* was higher in SAW
267 and STF, and although it remained on average relatively low ($<15\%$), it occasionally reached $> 50\%$ levels
268 (Figure S3).

269 A closer look revealed regional differences within each water mass (Figure 2). In SAW for instance,
270 SAF surface waters were colder and fresher than those on the Campbell Plateau (TAN1702) and in the
271 Bounty Trough (Bio-SAM site) (Figure 2). Surface nitrate concentrations were lower in the Bounty
272 Trough compared to Campbell Plateau and SAF, consistent with the southwards strengthening of HNLC
273 conditions. *Chla_{ML}* concentration was higher on Campbell Plateau (0.62 ± 0.48 , mean \pm sd) compared
274 to surface waters in the Bounty Trough (0.33 ± 0.20) and the SAF (0.21 ± 0.07) (Figure 2), although
275 differences were only significant between the Campbell Plateau and SAF to the south (one-way ANOVA,
276 $F(2,33) = 4.494$, $p = 0.019$).

277 Within STW, surface temperature and salinity were highest in northernmost waters sampled during
278 the Cross-shelf Exchange II voyage (TAN1604) and lowest in STW waters sampled during the Spring
279 Bloom II voyage (TAN1212) conducted at the beginning of austral spring (September-October) (Figure 2).
280 Temperature and salinity at the Biophysical Mooring subtropical site (Bio-STM) were intermediate on
281 average and had a greater range that reflected the wider temporal variability covered by multiple visits
282 conducted at different times of the year over several years (Table 1). Nitrate concentrations showed
283 the opposite trend with highest values associated with the colder and deep-mixed waters, and lower
284 values reflecting warmer and stratified waters of the Bay of Plenty and Cross-shelf Exchange II voyages
285 (Figure 2).

286 Regional differences were also observed between the S-STF flowing north of the Campbell Plateau
287 through the Snares Depression (Figure 1), which transported colder and fresher waters, and the STF further
288 north flowing eastwards over the Chatham Rise (STF, Chatham Rise) (Figure 2). Nitrate concentrations
289 tended to be higher in STF waters adjacent to Campbell Plateau than on the Chatham Rise reflecting the
290 HNLC nature of the plateau. Relatively high nitrate concentrations ($>10 \mu\text{mol/L}$) were also measured
291 during summer (TAN1516) on the Chatham Rise at some stations located in the southwestern flank of the
292 STF with colder (10.7 and $11.3 \text{ }^\circ\text{C}$) and fresher (34.47 and 34.56) characteristics of surface mixed-layer
293 waters indicating a SAW influence (Figure 1).

294 **3.2. Alpha-diversity – Species richness**

295 Species richness and the Shannon diversity index estimated in the euphotic zone were on average lower
296 in the STF compared to SAW and STW (Figure 3A). Highest diversity in STW was observed in the
297 northernmost waters surveyed during the Cross-shelf Exchange II voyage and at the Bio-STM site, which
298 included samples collected during various cruises that spanned multiple years and seasons. Protist species
299 richness during the Spring Bloom II voyage was substantially lower than these other STW locations
300 (Figure 3B). In SAW diversity was lower on Campbell Plateau compared to open ocean regions adjacent to
301 flows of the SAF (Figure 3B) located to the north and south of the plateau, respectively (Figure 1). Within
302 the STF, diversity was also lower in upstream waters of the S-STF located further south (46-49 °S) than
303 in the STFZ in the Chatham Rise region (43-45 °S) (Figure 1, Figure 3). Here, higher diversity values
304 were associated with the Bio-STF site visited multiple times in different seasons during the Biophysical
305 Mooring program (2009-12)(Table 1) than during the summer TAN1516 voyage, which covered the
306 Chatham Rise region more extensively (Figure 1).

307 Species diversity in the euphotic zone tended to decrease with latitude (model I linear regression,
308 $F(1,236) = 25.6$, $R^2 = 0.10$, $p\text{-value} < 0.0001$), although differences in mean diversity values were observed
309 among water masses and regions with similar latitudes (Figure S4). Similar relationships were observed
310 between diversity and temperature (model I linear regression, $F(1,208) = 59.1$, $R^2 = 0.20$, $p\text{-value} < 0.0001$
311) with regional differences modulating this trend (Figure S5). Within STW for instance, species richness in
312 the euphotic zone was higher in oligotrophic waters and decreased with increasing Chl *a*, being generally
313 higher in warmer waters (Figure S6). Samples from the STF presented lower species richness in the
314 euphotic zone compared to SAW and STW across the entire range of Chl *a* and nitrate concentrations (
315 Figure S6).

316 Diversity patterns in the aphotic zone were investigated on the Biophysical Moorings dataset only
317 (Bio-STM, -STF and -SAM, $n = 113$). Diversity in the aphotic layer of the STW and SAW sites was lower
318 than in the sunlit layers (Figure 3C). In the STF, however, this was higher in the aphotic compared to the
319 euphotic layer, which resulted in similar estimates of species richness in the aphotic zone of the different
320 water masses unlike the diversity pattern observed for the euphotic zone (Figure 3).

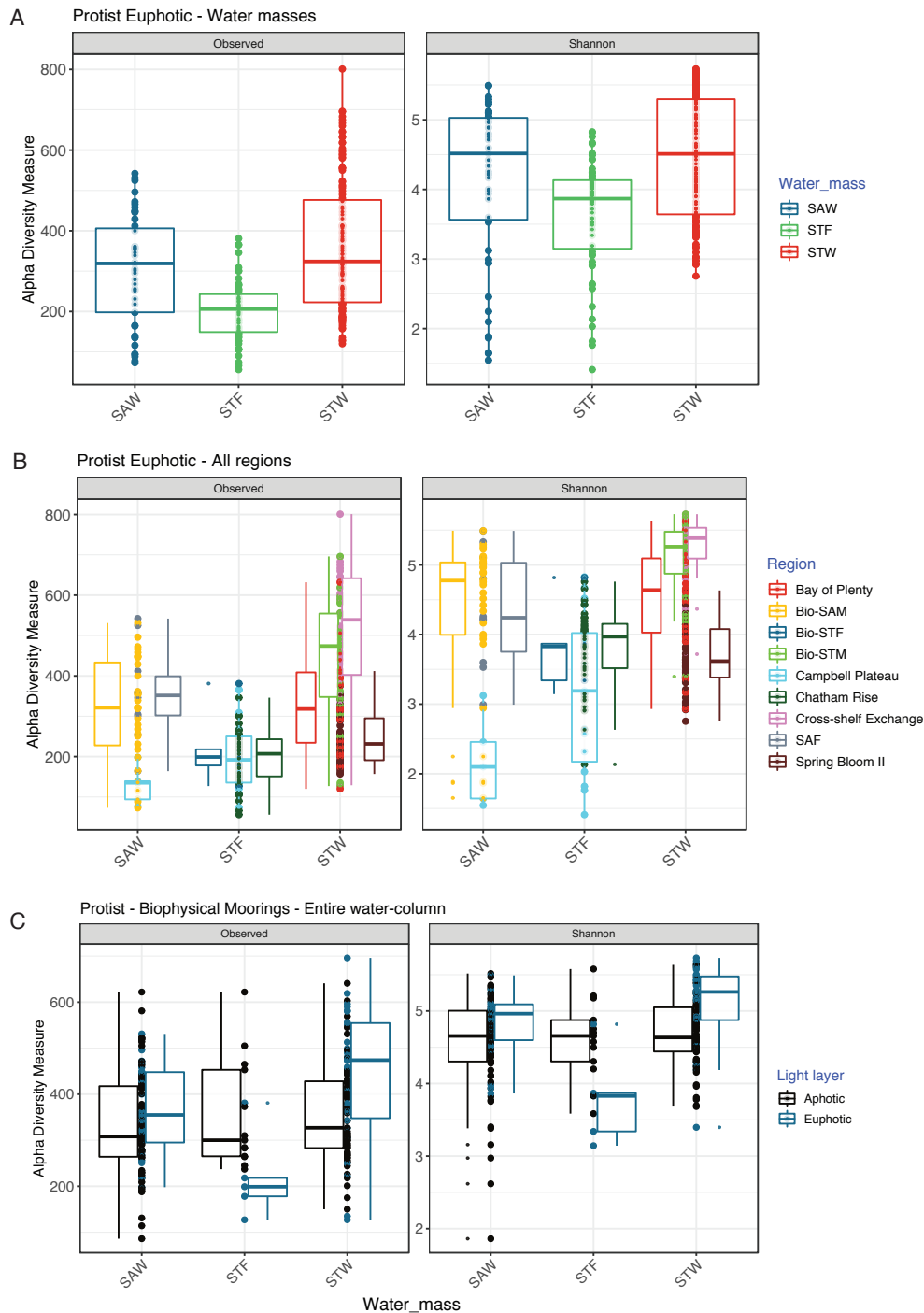


Figure 3. Species richness and diversity index estimated for (A) the euphotic zone of each water masses (SAW, STF, and STW); (B) the euphotic zone of each regions and (C) the aphotic zone of the Biophysical Mooring programme sites SAF (Subantarctic Front) and Campbell Plateau correspond to the same voyage TAN1702 (April 2017).

321 **3.3. Beta-diversity of protistan communities**

322 To explore the similarities between protistan communities in different geographic samples we performed
323 principal component analysis on ASV abundance. The first analysis with all samples yielded two main
324 clusters corresponding to samples from the euphotic and aphotic zone (Figure 4A). A second analysis
325 focused on the euphotic zone cluster samples (n = 239) according to different water masses although
326 certain overlaps also occurred, particularly between the STF and SAW samples (Figure 4B). Different
327 regions tended to cluster separately as well, with STW samples from the Spring Bloom II voyage forming
328 a separate cluster from that of the Cross-shelf Exchange and Bay of Plenty cruises (Figure 4B).

329 To investigate the influence of different oceanographic drivers on the community composition we
330 performed PERMANOVA analysis with continuous variables of physical (T, Sal), chemical (NO_3^- con-
331 centration - euphotic zone median NO_3^-) and biological (Chl *a* concentration - water-column median)
332 parameters. The analysis was conducted on a subset of samples (n=182) for which these measurements
333 were available (STW=34, STF=53, SAW=95). All variables yielded statistical significance ($p < 0.001$) with
334 salinity explaining most of the variability (F.Model = 34.4, $R^2 = 0.13$, $P < 0.001$) followed by temperature
335 ($R^2 = 0.08$) and nitrate ($R^2 = 0.06$) with chl *a* concentration explaining a minor fraction of this variability
336 ($R^2 = 0.02$) (Figure S7) (Table S2). Overall this set of variables left 69% of the variability unexplained. The
337 second PERMANOVA analysis included categorical variables (e.g. Water mass and region) showed that
338 while Water mass explained 16% of the variability (F. Model = 28.0, $P < 0.001$) – similar to that explained
339 by salinity – the region explained an additional 27% of the variability (F.Model = 15.4, $P < 0.001$) and up
340 to 51% of the variability together with physico-chemical and biological continuous variables included in
341 the first PERMANOVA analysis (Table 2).

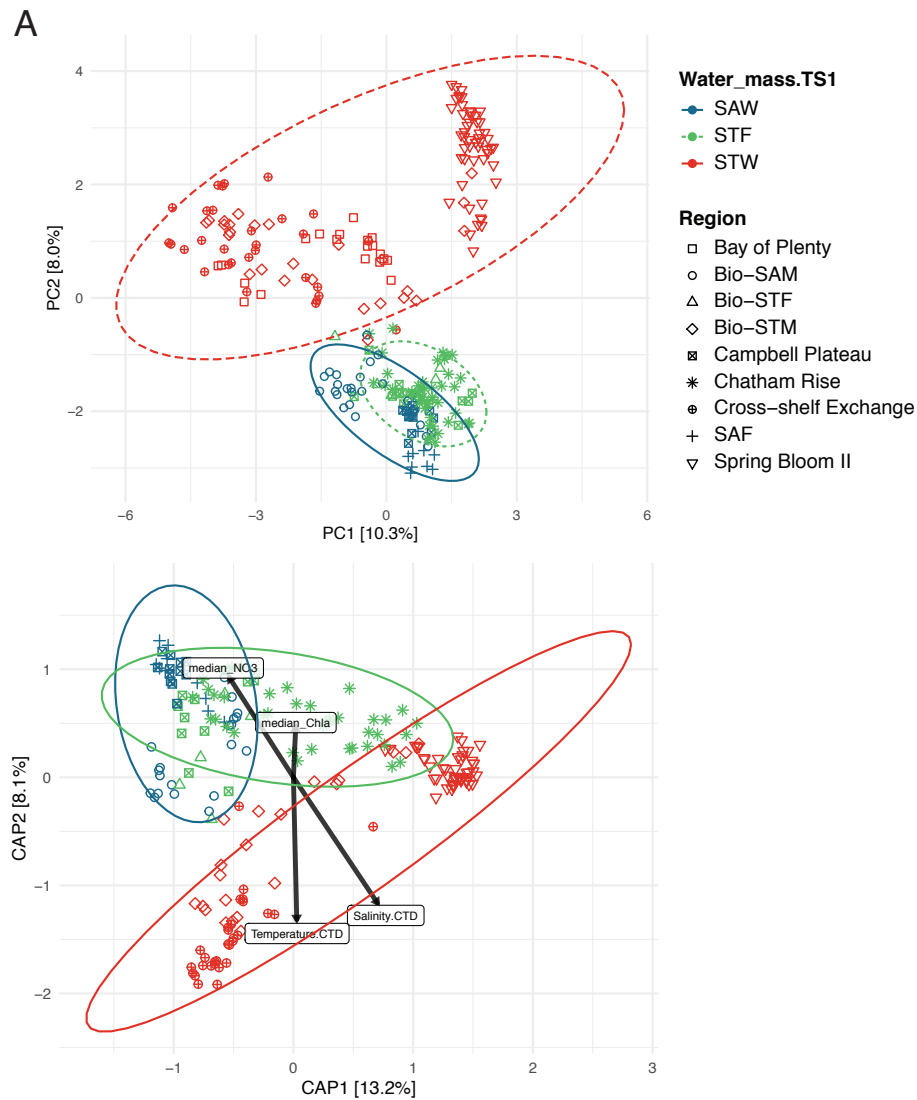


Figure 4. Figure 4. (A) Principal component analysis based on ASV composition of euphotic samples only color coded by water masses and shapes for regions/voyages (n=240). (B) Biplot of redundancy analysis (RDA) computed at species (ASV) level in the euphotic zone for which T, Sal, Nitrate and Chl *a* measurements were available. Arrows indicate the sign and strength of the correlation between community composition and environmental variables that were significant in PERMANOVA analysis (n=197) samples from Chatham Rise TAN1516 lack CTD and MLD data

Variable	Df	SumsOfSqs	MeanSqs	F.Model	R ²	Pr(>F)
Water-mass	2	9.321	4.6606	27.6669	0.16054	0.001
Region	6	15.558	2.593	15.3932	0.26796	0.001
Temperature	1	1.459	1.4588	8.6597	0.02512	0.001
Salinity	1	1.05	1.0503	6.2351	0.01809	0.001
median NO ₃	1	1.076	1.0763	6.3893	0.01854	0.001
median Chl <i>a</i>	1	1.128	1.1283	6.6977	0.01943	0.001
Residuals	169	28.469	0.1685		0.49032	0.001
Total	181	58.062			1	

Table 2. Summary of PERMANOVA analysis including the Water-mass and Region as categorical variables in addition to the continuous environmental variables. Temperature and salinity represent the surface values. Nitrate (NO₃⁻) and chlorophyll *a* (Chl *a*) median concentration calculated for samples within the euphotic zone. Analysis was conducted with the Adonis function of the vegan R package

342 **3.4. Division and class level taxonomic composition**

343 Dinoflagellate reads (syndiniales excluded) dominated the sequencing datasets from samples taken in the
344 euphotic zone (34% of total - sequencing depth-normalized – reads), with Chlorophyta accounting for
345 27%. Ochrophyta, constituted mainly by *Bacillariophyta* (7%) and *Pelagophyceae*(3%) and *haptophytes*
346 belonging to *Prymnesiophyceae* class(5%) were the other most important phytoplankton divisions, while
347 Stramenopiles_X (mainly through Marine Stramenopiles, MASTs, 4%) Radiolaria (4%), and Ciliophora
348 (3%) contributed most among the heterotrophic protist. Although such groups were consistently dominant,
349 their relative contributions and particularly, their composition at class (Figure 5 and Figure 6) and finer
350 taxonomic resolution (see subsection 3.5.) varied between water masses.

351 For STW samples for instance, where dinoflagellates and Chlorophyta co-dominated the overall
352 dataset (30% of sequences each), followed by Ochrophyta (10%) with Haptophyta, Stramenopiles_X and
353 Radiolaria accounting for a lower percentage of metabarcodes (7% each). *Mamiellophyceae* accounted

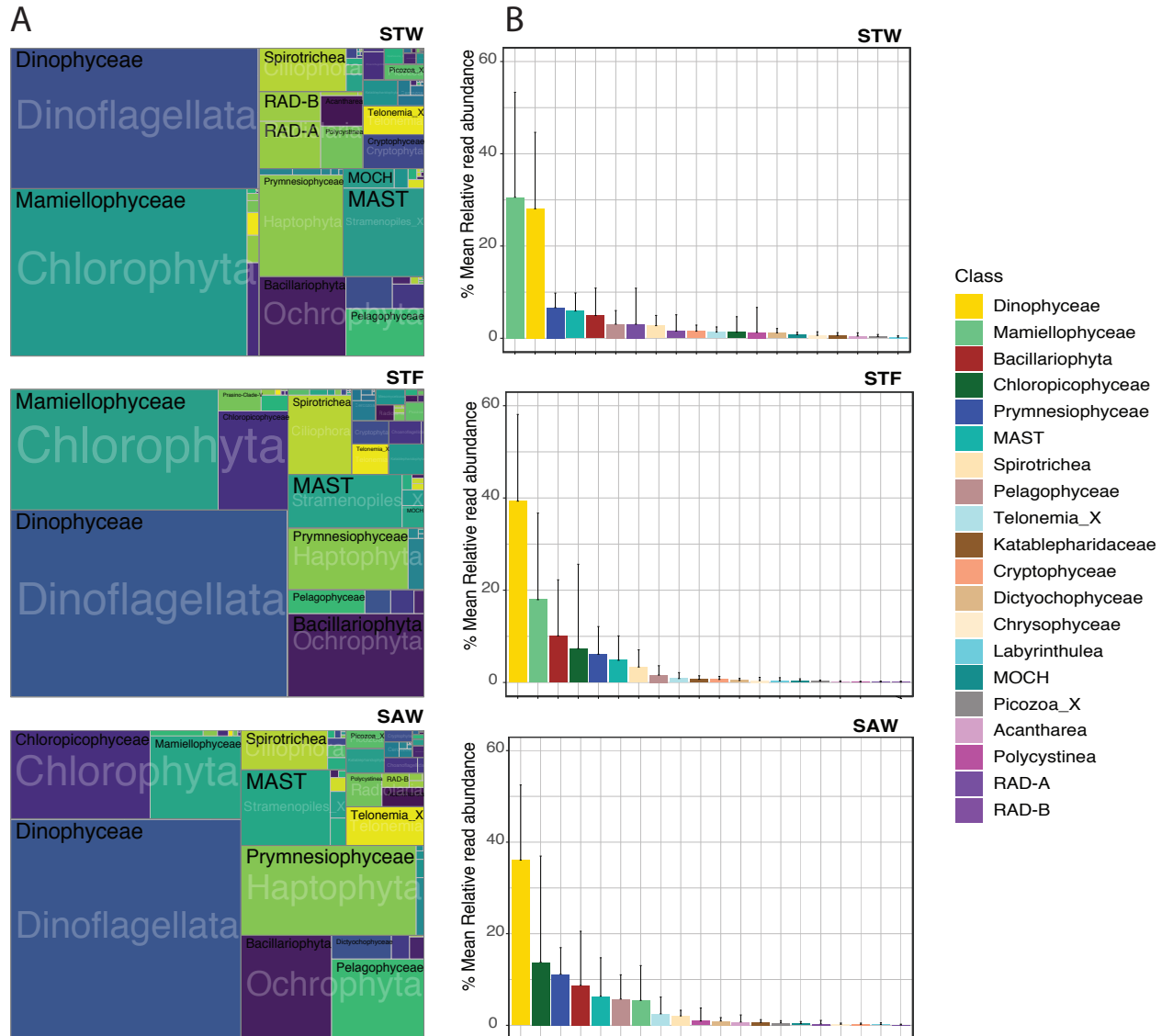


Figure 5. Protist community composition at division/class level (syndiniales excluded) in the euphotic zone of STW, STF and SAW. A) The area of each taxonomic group in the treemap represents the read abundances affiliated to each group standardized to the median sequencing depth across samples [$\text{median sum otus} * (\text{otu reads} / \text{sum (otu reads)})$]. B) Barplots represent the mean relative read abundance of most abundant classes across different water masses (error bars are the standard deviation of the mean).

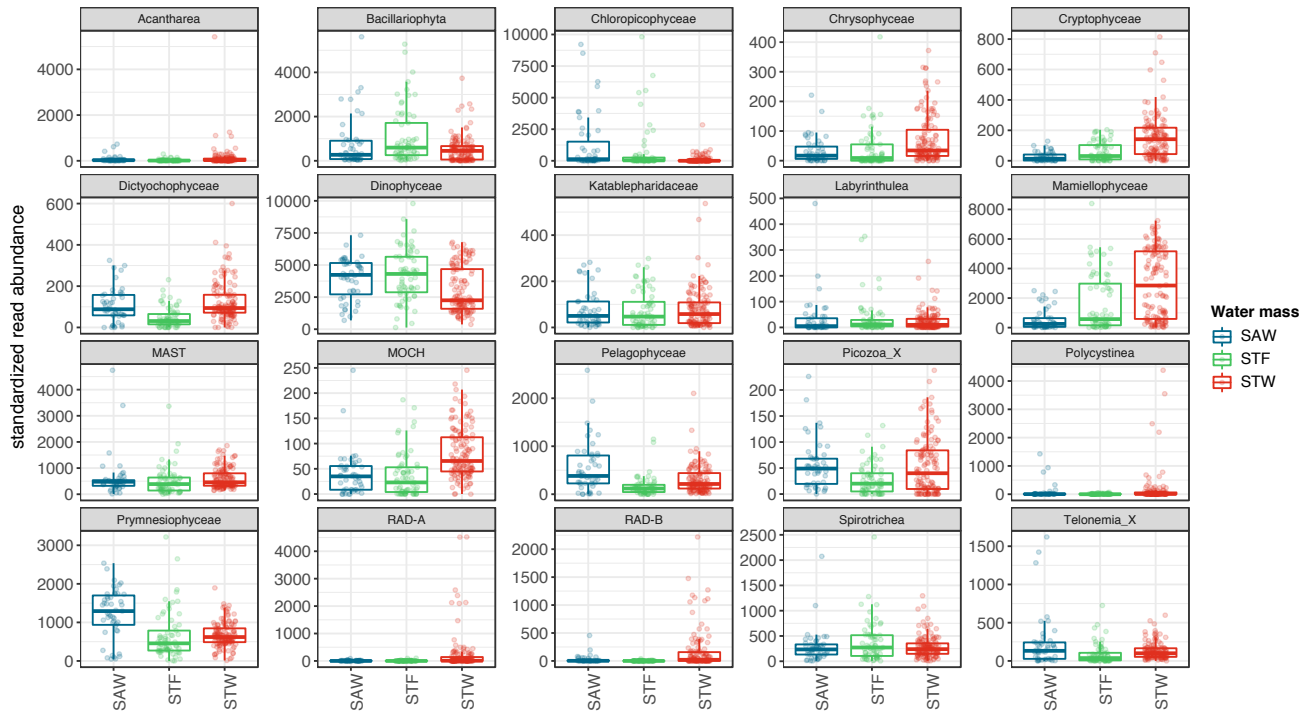


Figure 6. Box-plots showing standardized read abundance in subtropical, subantarctic and subtropical front of the twentieth most abundant protist classes in the euphotic zone. Box-plots show the median, the first and third quartiles (lower and upper hinges) and the values within the $\pm 1.5 * IQR$ (IQR, interquartile range) (line). Points represent values of single samples.

354 for the vast majority of reads affiliated to Chlorophyta (>95%) while *Chloropicophyceae* represented
 355 only a minor fraction of sequences affiliated to this division (Figure 5). *Prymnesiophyceae* was the
 356 most abundant class of Haptophyta (6%). *Bacillariophyta* (5%) and *Pelagophyceae* (3% each) classes
 357 dominated over *Dictyochophyceae* (1%) and *Chrysophyceae* (1%) among Ochrophyta, while *Cryptophyceae*
 358 (*Cryptophyta*, 2%) accounted for a relatively larger percentage. The class Among Heterotrophic groups
 359 MASTs (*Stramenopile_X*, 5%), RAD-A (*Radiolaria*, 3%) and ciliates of the class *Spirotrichea* contributed
 360 most to the STW metabarcoding data (Figure 5 and Figure 6).

361 The SAW sequencing dataset was clearly dominated by dinoflagellates (37% total reads), followed
 362 by Chlorophyta (18%), Ochrophyta (15%) and Haptophyta (12%) with a more even share of the total
 363 number of reads compared to STW (Figure 5). At class level, *Chloropicophyceae* (14%), was clearly the
 364 most abundant group of green algae followed by *Prymnesiophyceae* (12%), *Bacillariophyta* (8%) and
 365 *Pelagophyceae* (6%) phytoplankton classes (Figure 5). The heterotrophic component in the SAW metabar-

366 coding data was dominated by MASTs (6%) followed by ciliates (3%) while the relative contribution
367 of Radiolaria in the euphotic zone was minor (<0.5%) and mainly attributed to Polycystinea rather than
368 RAD-A &-B classes (Figure 5 and Figure 6).

369 The STF protist metabarcoding data presented intermediate characteristics between STW and SAW
370 (Figure 5). As in SAW, dinoflagellates (39%) dominated the data although the contribution of Chlorophyta
371 (26%) was on average higher and closer to levels found in the STW metabarcoding dataset. Accordingly,
372 *Mamiellophyceae* was the second most abundant class of green algae (17%), although the contributions
373 of *Chloropicophyceae* (10%), and *Pyramimonadophyceae* (1.5%) were also substantial. The division
374 Ochrophyta (12%) accounted for a similar fraction of phytoplankton reads as in STW and SAW, but in STF
375 the phytoplankton community was clearly dominated by *Bacillariophyta* (10%) with minor contributions
376 from *Pelagophyceae* (2%) and *Dictyochophyceae* (0.5%) classes. The heterotrophic component was
377 dominated by MASTs (5%) and ciliates (4%) while the contribution of Radiolaria (<0.2%) was on average
378 lower than in STW and SAW datasets (Figure 5 and Figure 6).

379 To investigate protist community composition below the euphotic zone we used the Bio-physical
380 moorings dataset that covered systematically the entire water column at the Bio-STM and Bio-SAM
381 (0-3100 and 0-2800 m, respectively) sites and the Bio-STF site on the Chatham Rise crest (0-350 m)(n
382 = 113 samples). Overall, Radiolaria (Polycystinea) dominated the aphotic metabarcoding data (52% of
383 protist reads) followed by dinoflagellates (27%). Heterotrophic groups such as MASTs (3%) and ciliates
384 (3%) contributing substantially less (Figure 7 and Figure S8). Phytoplankton groups such as diatoms,
385 green algae, and prymnesiophytes contributed at similar levels (3% each), likely reflecting mixing layers
386 extending below the euphotic zone. The relative contribution of Radiolaria was similar in SAW (57%) and
387 STW (50%) samples but much lower in the STF samples (35%) where together with dinoflagellates (36%)
388 they co-dominated protistan reads below the euphotic zone (Figure 7 and Figure S8). Among Radiolaria,
389 Polycystinea class, mainly through the Spumellarida order, was most abundant (25-35% of total reads)
390 although Acantharea (10%), RAD-B (8%) and to a lesser extent RAD-A (1%) classes also contributed to
391 the dominance of this group (Figure 7 and Figure S8).

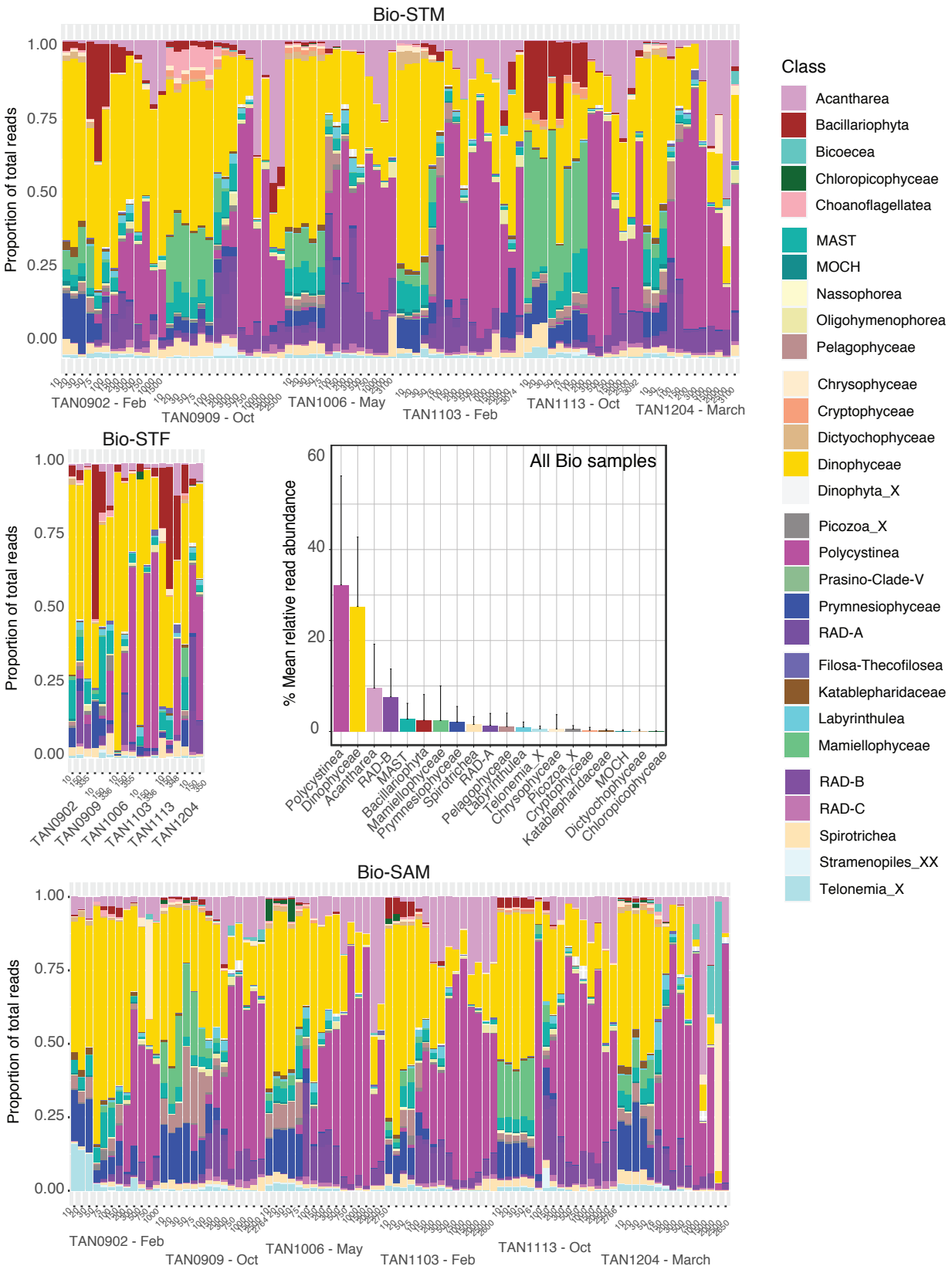


Figure 7. Relative read abundance of main protistan classes in samples collected throughout the water column (0-2000 m) during multiple voyages to the Biophysical Mooring program sites in STW (Bio-STM), STF (Bio-STF) and SAW (Bio-SAM) and mean contribution for the whole sampling program (n = 113)(error bars as in Figure 4. 24/20

392 Vertically, the dominance of Radiolaria became more prominent in the mesopelagic and bathypelagic
393 samples where they often represented >75% of total reads (Figure 7). Polycystinea class tended to
394 dominate across all sample depths, with increasing abundance at depths >300 m. Sequences affiliated
395 with Acantharea and RAD-B showed similar abundance but followed opposite vertical distributional
396 trends, with RAD-B being more abundant in shallower samples (100-500 m) while Acantharea abundance
397 increased with depth and peaked in bathypelagic samples (>1000 m) (Figure 7).

398 **3.5. Genus and species community composition**

399 Species composition also varied among the metabarcodes from the physically and biogeochemically
400 distinct water masses (Figure 8, Figure 9). In STW, the green algae (Chlorophyta) was dominated by the
401 species *Ostreococcus lucimarinus* followed by *Bathycoccus prasinus*. These *Mamiellophyceae* species
402 together with *Micromonas commoda* and other *Micromonas* species (*M. bravo* I, II, and *M. pusilla*)
403 accounted for the majority of the sequences affiliated to green algae in the euphotic zone of STW (Figure 8,
404 Figure S9 and Figure S10). Reads affiliated to several dinoflagellate species such as *Gymnodinium*
405 sp., *Heterocapsa rotundata* and *Gyrodinium spp.* were among the most abundant in STW dataset.
406 *Gymnodinium* sp., and *H. rotundata* were more abundant at the Bio-STM whereas *Gyrodinium helveticum*
407 was prevalent in STW of the EAUC surveyed during the Cross-shelf Exchange voyage (Figure S10).
408 Diatom ASVs identified as Polar-centric Mediophyceae_X sp. and *Minidiscus trioculatus* were the most
409 abundant diatoms reads in STW metabarcodes, particularly in the Spring Bloom II voyage. Among
410 ASVs assigned to pelagophytes, an unidentified Pelagophyceae_XXX_sp. (ASV_0058) and *Pelagomonas*
411 *calceolata* ASVs were the most abundant ones (Figure 8, Figure S10 and Figure S9). Among the class
412 *Prymnesiophyceae*, *Phaeocystis globosa* (ASV_0065) was the most abundant ASV (Figure 8) with several
413 ASVs belonging to *Gephyrocapsa oceanica* and *Chrysochromulina* spp. mainly contributing to the overall
414 dominance of the latter genus within the class (Figure S9, Figure S10).

415 Among the STF metabarcodes, the Mamiellophyceae *O. lucimarinus* was also the most abundant
416 species overall, although the relative contribution of *Chloroparvula pacifica* (asv_0014, *Chloropico-*
417 *phyceae*) increased substantially compared to STW (Figure 8, Figure S9). It is worth noting the high
418 abundance of sequences affiliated to *Chloropicon sierburthii* in addition to *Chloroparvula pacifica*, which

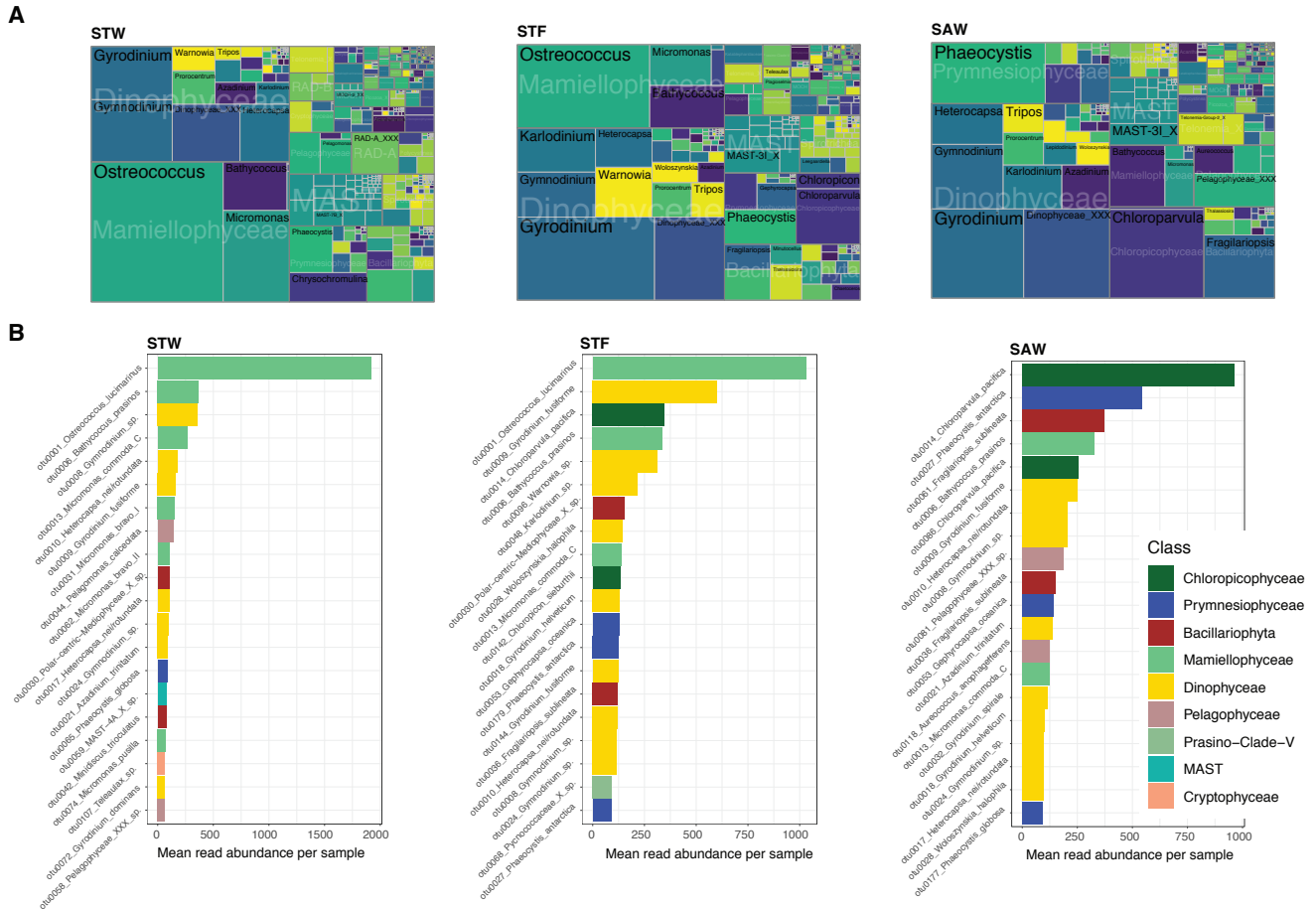


Figure 8. Water mass genus and species abundance (A) Treemaps showing the community composition at class/genus level in the euphotic zone of the STW, STF and SAW. The area of taxonomic group is proportional to the read abundances affiliated to each group standardized to the median sequencing depth across samples [$\text{median sum otu} * (\text{otu reads} / \text{sum (otu reads)})$]. (B) Mean standardized read abundance of most abundant ASVs and assigned species, color coded for their class affiliation, in the euphotic zone of the different water masses.

419 contributed to the overall increase in the relative abundance of *Chloropicophyceae* in the STF dataset
420 (Figure 8, Figure S9). Reads assigned to the heterotrophic dinoflagellate *Gyrodinium fusiforme* was the
421 second most abundant in STF samples, with other ASVs affiliated with *Warnowia* sp. and *Karlodinium*
422 sp. appearing among the most abundant dinoflagellate species (Figure 8, Figure 9 and Figure S9). The
423 higher relative abundance of Bacillariophyta observed in STF was mainly driven by ASVs affiliated to
424 unidentified Polar-centric Mediophyceae species and *Fragilariopsis sublineata* (Figure 8 and Figure S9)
425 which tended to be more abundant in the STFZ of the Chatham Rise region (TAN1516) and the S-STF
426 flowing north of C. Plateau (TAN1702), respectively (Figure S10). The identification of *F. sublineata*
427 should be taken with caution the V4 region of the two sequences included in PR2 is 100% similar to the
428 annotated sequence of *Fragilariopsis Kerguelensis* present in PR2 making impossible to unambiguously
429 assign that ASVs to one of this two species. To reflect this ambiguity the ASVs assigned to *F. sublineata*
430 in PR2 are referred as *F. sublineata/kerguelensis* throughout the text (see discussion). Among *Prymnesio-*
431 *phyceae*, *Phaeocystis* spp. was the dominant genus but most reads in this case belonged to *P. antarctica*
432 instead of *P. globosa* (Figure 8, Figure S9). ASVs assigned to the prymnesiophyte *G. oceanica* also
433 increased substantially in STF compared to STW datasets. *Pelagophyceae* in the STF metabarcodes was
434 dominated by *Aureococcus* and *Pelagococcus* spp. (Figure S9) by although the class relative contribution
435 was relatively low (Figure 5). The relative abundance of *Cryptophyceae* and *Dictyophyceae* remained
436 minor overall (<2%) (Figure 5), but both groups showed changes in their specific composition across
437 the water masses metabarcode datasets. Among *Cryptophyceae*, *Plagioselmis prolunga* and *Teleaulax*
438 *gracilis* increased substantially from STW to STF samples while the relative contribution of sequences
439 assigned to *Teleaulax* sp. decreased. Changes within *Dictyophyceae* were less substantial but showed an
440 increase in the relative abundance of *Dictyocha speculum* and *Pseudochattonella farcimen* from STW to
441 STF metabarcodes (Figure S9).

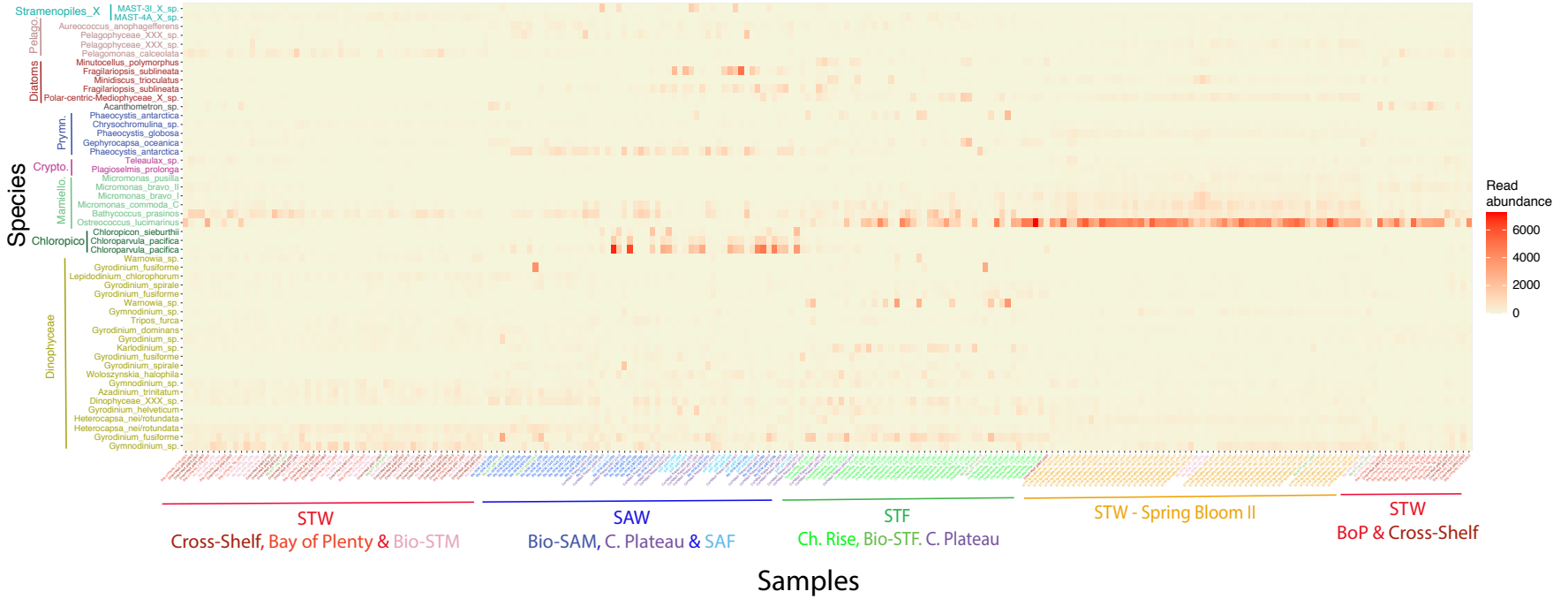


Figure 9. Heatmap showing the standardized read abundance of the 50 most abundant species (Y-axis) across samples collected in the euphotic zone (X-axis). Samples were clustered using nMDS and Jaccard distance and sample labels color coded according to the water mass and region they were collected from. Species were organized and color coded by class affiliation.

442 Among SAW samples, *Chloropicophyceae* ASV_0014 assigned to *Chloroparvula pacifica* was the
443 most abundant taxa, with other ASVs of this species (e.g., asv_0086) contributing also to the dominance
444 of this class over Mamiellophyceae in SAW (Figure 8, Figure 9, Figure S9). Among Mamiellophyceae,
445 reads from *B. prasinos* became the most abundant followed by *M. commoda* while the contribution of *O.*
446 *lucimarinus* in SAW samples was minor (Figure 8 and Figure S9). The increase in the relative abundance
447 of Prymnesiophyceae reads observed in SAW samples was driven mainly by *Phaeocystis* spp., with
448 *P. antarctica* emerging as the second most abundant species in SAW metabarcodes (Figure 8). Other
449 *Phaeocystis* species (*P. globosa*, *P. cordata* and *Phaeocystis* spp.) contributed also to the dominance of this
450 genus within prymnesiophytes (Figure S9 and Figure S10). Among Bacillariophyta, ASVs assigned to *F.*
451 *sublineata/keruelensis* (ASVs 0061 and 0036) were the dominant in SAW samples (Figure 8 and Figure 9)
452 while *Thalassiosira* sp., Polar-centric Mediophyceae sp. and other diatom genus contributed substantially
453 less (Figure S9 and Figure S10). Increasing abundance of pelagophyte reads in SAW (Figure 5) was
454 mainly due to *Pelagophyceae*_XXX_sp (ASV_0081), the same ASV that dominated in STW samples,
455 and to *Aureococcus anophagefferens* which were among the most abundant taxa in SAW metabarcodes
456 (Figure 8 and Figure S9).

457 To investigate how the relative abundance of the species identified varied between the water masses,
458 we ran a differential gene expression analysis based on the negative binomial distribution (DESeq) using
459 water mass as a categorical variable. For the euphotic zone, this yielded 70 and 35 ASVs out of 3984
460 ASVs that were significantly more or less abundant (p-value < 0.01) in STW compared to SAW datasets,
461 respectively (Figure S11A). Species that showed greatest differences (>10 log₂-fold changes) in their
462 relative abundance were not necessarily among the most abundant species in each water mass. Among the
463 species associated with STW, we found the larger changes in relative abundance in the diatoms ASVs
464 assigned to Polar-centric Mediophyceae, *Minutocellus polymorphus*, and *Minidiscus trioculatus*; the
465 prasinophytes *O. lucimarinus* and several *Micromonas* species; the prymnesiophytes *Phaeocystis globosa*
466 and *Chrysochromulina* sp., and dinoflagellates *Warnovia* sp., *H. rotundata* and number of other species
467 (Figure S11). Among species with preferences for SAW we found several diatom species including *F.*
468 *sublineata* and *Cylindrotheca closterium*, the prasinophyte *B. prasinos*, the pelagophytes *Pelagococcus* sp.
469 and *A. anophagefferens*, and several dinoflagellates including the heterotrophic species *G. fusiforme* and

470 *G. spirale* and mixoplankton *Karlodinium veneficum* (Figure S11).

471 Different ASVs affiliated to the same species often showed distinct preference for STW and SAW.
472 Most abundant *Chloroparvula pacifica* ASVs (e.g. asv_0086 and asv_0014, Figure 8) were associated
473 with SAW, while less abundant ASVs (e.g. asv_0532 and asv_0336) were more associated with STW (
474 Figure S11A). Similarly, most abundant ASVs of *P. antarctica* (asv_0011, asv_0027) showed preference
475 for SAW (10-30-log₂ fold negative change) while much less abundant asv_0218 showed greater affinity
476 for STW. This intraspecific variability was observed within unidentified species (e.g. Dinophyceae_XXX_
477 sp. and Pelagophyceae_XXX_sp.) which included ASVs with opposite affinities for STW and SAW (
478 Figure S11A).

479 DESeq analysis between the STF and STW also depicted specific differences as observed between
480 STW and SAW, with some additional diatoms ASVs assigned to *Thalassiosira* sp. and *Pseudo-nitzschia*
481 *delicatissima* enriched in the STF dataset (Figure S11B). In most cases, the distinctive species pattern
482 observed between the STF and either STW or SAW metabarcodes coincided with those identified from
483 the comparison between STW and SAW described above (Figure S11B). For instance, ASVs of Polar-
484 centric Mediophyceae species found in higher abundance in STW compared to SAW, were also associated
485 preferentially with STF compared to SAW while *F. sublineata*, and *C. closterium* diatoms ASVs, that were
486 enriched in SAW compared to STW, were also preferentially associated with the STF rather than STW
487 (Figure S11B). Only few ASVs, such as those assigned to the prasinophyte *Chloropicon sieburthii* and
488 dinoflagellate *Gonyaulax* sp., were distinctively associated with STF waters (Figure S11C).

489 The most abundant reads found below the euphotic zone belong to the polycystinean order Spumellarida
490 Group I family (Figure S8). Several ASVs contributed to the dominance of Spumellarida Group I family,
491 but different ASVs dominated in each water mass dataset (Figure S12 and (Figure S13). ASV0023 and
492 ASV0039 affiliated with the Group I family of Spumellarida dominated in SAW, while ASV0064 and
493 ASV0085 were most abundant in STF (Figure S13). Conversely, most abundant radiolarian species in
494 STW were unidentified species assigned to the acantharean order Chaucanthida (Figure S13). Several
495 less abundant Radiolaria ASVs showed significantly different abundances across the aphotic depth layers
496 of different water masses suggesting differences in their ecological preferences (Figure S14). Reads
497 assigned to known photosynthetic species also showed differential abundance below the euphotic zone

498 in the different water masses datasets. Most notably the higher abundance of reads affiliated to the
499 *Mamiellophyceae* *O. lucimarinus* and the *Prymnesiophyceae* *P. globose* in STW compared to SAW yielded
500 significant differences in their abundance below the euphotic zone as well. A pattern observed also for
501 less prominent phytoplankton species like the diatom *Asterionellopsis glacialis* or *Thalassiosira rotula*
502 that despite being less abundant in the euphotic zone yielded significant differences in their abundance
503 below the euphotic zone.

504 **4. DISCUSSION**

505 The taxonomic composition of protistan communities in SAW and STW in the SW Pacific and across
506 the main frontal zones of the region (STF, SAF) has not been extensively characterized. Previous
507 studies have typically focused on phytoplankton communities and their variability across SA and ST
508 waters flanking the STFZ over the Chatham Rise (Chang and Gall 1998; Delizo et al. 2007; Hall et al.
509 1999) but broader biogeographic studies in the SW Pacific are scarce (DiTullio et al. 2003). By using DNA
510 metabarcoding, we provide a comprehensive taxonomic characterization of the protistan communities
511 associated with STW and SAW at southern temperate latitudes. Although samples included in this study
512 were collected in different seasons, the seasonal coverage was similar for STW and SAW (Table 1 and
513 Figure S1), allowing a meaningful comparison among the microbial communities associated with these
514 water masses throughout an average year.

515 **4.1. Protistan community structure in STW and SAW of the southwest Pacific**

516 Species richness decreased latitudinally and with temperature (Figure S4 and Figure S5) as expected from
517 global diversity patterns observed and modelled for marine bacterial and phytoplankton communities
518 (Barton et al. 2010; Fuhrman et al. 2008; Ibarbalz et al. 2019) as well as from previous DNA-based reports
519 in the SW Pacific region (Raes et al. 2018). Consistent with this trend, species richness was higher in
520 warmer STW than in SAW while lowest diversity was associated with the STF (Figure 3).

521 The relative low diversity observed in the STF could be related to the increased phytoplankton biomass
522 and productivity typically associated with the STF across the annual cycle (Murphy et al. 2001; Pinkerton
523 et al. 2005) and the dominance of fewer ‘bloom-forming’ species in this highly productive zone (Chang

524 and Gall 1998). The fact that species richness within STW was lowest during the more productive spring
525 bloom conditions (TAN1212) is consistent with the view that more productive waters such as found in the
526 STF, during the spring bloom, and locally on the Campbell Plateau (Gutiérrez-Rodríguez et al. 2020) tend
527 to decrease protistan diversity. However, the lower diversity associated with STF, relative to STW and SAW,
528 was systematically observed across the different levels of nitrate and Chl *a* concentrations encompassed in
529 this study (Figure S6) suggesting that other factors may contribute to this pattern. Interestingly, the lower
530 diversity in the STF relative to STW and SAW was buffered below the euphotic zone (Figure 3C). Similarly,
531 the temperature-driven latitudinal pattern described globally for epipelagic plankton also disappeared
532 below the euphotic zone (Ibarbalz et al. 2019). The decoupling between the diversity patterns in the sunlit
533 and dark ocean suggested by these results are somewhat contrary to the connectivity between the epi-
534 and bathypelagic zones as inferred by the high correspondence of bacterial communities and processes
535 between these realms (Mestre et al. 2018; Ruiz-Gonzalez et al. 2020). The reasons for these differences
536 are unclear and highlight the need of further studying ecological processes that shape microbial diversity
537 throughout the entire water column.

538 Despite the regional and seasonal variability encompassed within both STW and SAW (Table 1,
539 Figure 2) we observed systematic differences in the taxonomic composition associated with these water
540 masses (Figure 4). Such water-mass specificity has been previously observed for the prokaryotic (Agogué
541 et al. 2011; Galand et al. 2010; Seymour et al. 2012; Techtmann et al. 2015) and eukaryotic components
542 (Hamilton et al. 2008; Raes et al. 2018) of microbial communities across different oceans. Among
543 environmental drivers, salinity rather than temperature or nitrate concentration was the physico-chemical
544 variable that explained best the compositional (dis-)similarities among euphotic samples (Figure 4,
545 Figure S7). Bray-Curtis dissimilarity indices of surface bacterial communities across the Southland
546 Current was also strongly correlated with salinity (Baltar et al. 2016). These results support the view
547 that STW and SAW east of New Zealand are better conceptualized as bioregions or provinces rather than
548 habitats *sensu* (*sensu* Martiny et al. 2006), where (phyto)plankton communities reflect oceanographic
549 processes and history in addition to contemporary physico-chemical conditions.

550 Samples from the STF itself were also distinguished from those in SAW and STW based on their
551 taxonomic composition, although they showed a greater overlap (Figure 4) that reflected the active mixing

552 and transition nature of such frontal zones. This overlap was particularly evident between samples
553 collected at the Bio-STF and Bio-SAM sites located on the Chatham Rise and its subantarctic flank,
554 respectively, and between the S-STF and SAW over the Campbell Plateau, which suggests a stronger
555 physical connectivity and ecological affinity of the STF with SAW than STW. Similarly, the horizontal
556 mixing and phytoplankton community size structure in the STF zone has been reported to be more tightly
557 coupled across SAW-influenced than STW-influenced water types (Safi et al., submitted). Nevertheless,
558 the distinct protistan communities observed in STW and SAW, and to a lesser degree in the STF, highlights
559 the role of oceanographic features such as the STF as boundaries that influence the diversity of oceanic
560 microbial communities in large oceanic provinces (Baltar et al. 2016; Raes et al. 2018).

561 **4.2. Taxonomic composition of phytoplankton community**

562 Our results showed the overall dominance of dinoflagellates and Chlorophyta across all water masses,
563 followed by Bacillariophyta, *Prymnesiophyceae* and *Pelagophyceae* (Figure 5). Yet consistent differences
564 in the relative contribution of these taxonomic groups between water masses emerged at class and species
565 taxonomic classification levels (Figure 6, Figure 8, Figure S9). Analysis of intra-specific diversity revealed
566 differences in the distribution of ASVs of the same species suggesting the presence of different ecotypes in
567 some cases (e.g. *Chloroparvula pacifica*, *P. antarctica*) and current taxonomic gaps within certain groups
568 that remain to be characterized (e.g. *Pelagophyceae*_XXX; *Dinophyceae*_XXX) (Figure S11). Below we
569 discuss the distributional patterns of major taxonomic groups, highlighting different taxonomic ranks to
570 shed some light on their ecological preferences.

571 **4.3.3. Chlorophyta (Green algae)**

572 The relative contribution of the two main green algae classes, *Mamiellophyceae* and *Chloropicophyceae*,
573 showed opposite distribution patterns (Figure 5). *Mamiellophyceae* was the most abundant class in STW
574 and constituted the bulk of green algae that dominated these waters samples, while its relative abundance
575 decreased across the STF to reach lowest levels in SAW (Figure 6). Picoplanktonic algae *O. lucimarinus*
576 was clearly the most abundant species in STW and STF (Figure 8, Figure 9) in agreement with a previous
577 metabarcoding analysis conducted across the Southland Current (Allen et al. 2020) where this species
578 abundance peaked in neritic STW inshore of the main current core and decreased towards SAW end of the

579 coast-offshore sampling transect (Allen et al. 2020). *O. lucimarinus* was also among the most abundant
580 species of *Mamiellophyceae* in a 18S rRNA metabarcoding survey conducted on coastal waters globally
581 (Tragin and Vaultot 2019).

582 The dominance of picoplanktonic *Mamiellophyceae* in STW is consistent with the greater contribution
583 of $<2 \mu\text{m}$ Chl *a* (80%,) observed in this water mass (Figure S4). It is worth noting that the highest
584 abundance of this group was observed during the onset of the Spring bloom samples (TAN1212, Figure S10)
585 when *Mamiellophyceae* accounted for 40-75% of 18S rRNA reads, while diatom contributions remained
586 around 10% over the 3-weeks of sampling (Figure S15). Among the multiple surveys of the Bio-STM
587 site, *Mamiellophyceae* contribution tended to be highest during early spring coinciding with the onset of
588 the spring bloom (Figure 7). *Mamiellophyceae* and *O. lucimarinus* were the most abundant phytoplankton
589 class and species in the STF samples along the Chatham Rise (TAN1516, Figure S15 and Fig. S10)
590 particularly on the STW influenced northern flank of the rise (Figure S16). The abundance and presence
591 of some prasinophyte classes, including *Mamiellophyceae*, have often been assessed from their diagnostic
592 pigment prasinoxanthin. Quantitative application of pigment-based approaches showed that prasinophytes
593 dominated the community in the STFZ and its subtropical flank across the Chatham Rise (170 °E) (Delizo
594 et al. 2007) and further east (170°W) (DiTullio et al. 2003), in agreement with our DNA based results.
595 In the Indian sector of the SO, a latitudinal study also found the highest contribution of prasinophytes
596 associated with the STF (Iida and Odate 2014). Broader application of pigment approaches have revealed
597 that prasinophytes can contribute substantially to vernal blooms at temperate latitudes (Bustillos-Guzman
598 et al. 1995; Gutiérrez-Rodríguez et al. 2011; Latasa et al. 2010; Nunes et al. 2018). High abundance of
599 several species of prasinophytes including *Ostreococcus* spp. and *Micromonas* spp. have been recently
600 reported during the onset of the North Atlantic spring bloom from 16S rRNA amplicon sequencing analysis
601 (Bolaños et al. 2020) and at more temperate latitudes of the Eastern North Atlantic using 18S rRNA
602 metabarcoding (Joglar et al. 2021). The deep mixing layers (>100 m) during the New Zealand STW
603 spring bloom II voyage (TAN1212) (Chiswell et al. 2019), where prasinophytes dominated, supports the
604 idea that this picophytoplankton group thrives under high-nutrient, high-mixing conditions playing an
605 important role in the development of spring blooms, characteristic of temperate latitudes. Overall, our
606 results highlight the wide ecological breadth of *Mamiellophyceae* and certain species like *O. lucimarinus*

607 which tend to dominate across a wide range of physical, chemical and trophic conditions encountered
608 within STW.

609 *Chloropicophyceae* (previously defined as Prasinophytes clade VII, Lopes dos Santos et al. 2017a)
610 showed an opposite trend to *Mamiellophyceae*, with the highest relative abundance associated with SAW
611 samples (Figure 5 and Figure 6). Culture representatives of *Chloropicophyceae* and 18S rRNA sequences
612 have been obtained from tropical and subtropical latitudes of the north and south Pacific (Lopes dos Santos
613 et al. 2017a,b; Tragin and Vaultot 2018) but to our knowledge this is the first report of their presence
614 and high abundance in subantarctic waters. The majority of *Chloropicophyceae* reads were assigned
615 to a reference sequence corresponding to *Chloroparvula pacifica* and included several ASVs one of
616 which (ASV0014) was the most abundant protist ASV found in SAW (Figure 8). *Chloropicophyceae*
617 has been suggested as the dominant group of green algae in meso- and oligotrophic oceanic waters in
618 contrast with the preference of *Mamiellophyceae* for more nutrient-rich coastal environments (Lopes
619 dos Santos et al. 2017b; Shi et al. 2009; Tragin and Vaultot 2018). In this study, *Chloropicophyceae* were
620 most abundant in SAW which are considered HNLC, suggesting that the preference of this group for
621 meso-/oligotrophic conditions reported for typically macronutrient limited waters could also encompassed
622 iron limited HNLC conditions. Furthermore, field experiments have suggested that phytoplankton growth
623 in HNLC regions in the subarctic Pacific and the Southern Ocean is co-limited by B vitamins and iron
624 micronutrients (Bertrand et al. 2007; Koch et al. 2011; Panzeca et al. 2006). Genome analysis of one
625 species of *Chloropicophyceae* (*Chloropicon primus*) indicates that this group might be able to synthesize
626 thiamine, in contrast to *Mamiellophyceae*, which depends on exogenous vitamin B1 or related precursors
627 supplied by B1-synthesizing marine bacteria or other algae (Lemieux et al. 2019; Paerl et al. 2015). The
628 potentially significant ecological role of B1 and B-vitamins, in general, in regulating and shaping the
629 taxonomic composition of phytoplankton communities, is rarely considered and is still not well understood
630 (Sañudo-Wilhelmy et al. 2014) but could contribute to explain the contrasting distribution patterns of both
631 classes.

632 Although *Chloropicophyceae* abundance occasionally peaked in the Bounty Trough (Bio-SAM sam-
633 ples) the high relative contribution of this group in SAW was mainly due to the high abundance systemati-
634 cally observed on Campbell Plateau and the S-STF flowing north of the plateau (Figure 1, Figure S15,

635 Figure S16). In the S-STF, *Chloropicon sieburthii* made a substantial contribution in addition to the
636 more dominant *Chloroparvula pacifica* (Figure S10). Whether this regional preference was linked to the
637 bathymetric and hydrographic characteristics of the plateau (Forcén-Vázquez et al. 2021; Neil et al. 2004),
638 the natural iron fertilization hypothesized for the region (Banse and English 1997; Gutiérrez-Rodríguez
639 et al. 2020) or a combination of these and other aspects cannot be concluded from our study. Moreover,
640 an ASV belonging to this genus was also found to contribute substantially to protistan communities in
641 coastal waters of the California Current Ecosystem (Gutierrez-Rodriguez et al. 2019), highlighting the
642 need of further studies to better understand the ecological drivers beyond coastal-oceanic trophic gradients
643 responsible for the water mass preferences of such phytoplankton groups and species.

644 4.3.4. *Dinophyceae (Dinoflagellates)*

645 Dinoflagellates relative abundance tended to be higher in SAW and the STF compared to STW metabar-
646 codes (Figure 5, Figure 6), consistent with previous microscopy-based observations (Chang and Gall
647 1998). ASVs affiliated to *Gyrodinium* genus and particularly *G. fusiforme* were identified as the most
648 abundant species in agreement with a previous study in the Southland Current where DNA barcodes of
649 *Karlodinium* and *Gymnodinium*, *Gyrodinium helveticum* and *G. spirale* were also retrieved and among
650 the most abundant species (Allen et al. 2020). Many *Gyrodinium* species prey on bacteria and algae
651 (Hansen 1992; Jang et al. 2019; Jeong et al. 2008). They can constitute an important component of
652 microzooplankton biomass in coastal and oceanic environments (Jeong et al. 2010; Sherr and Sherr 2007)
653 including high latitude waters (Archer et al. 1996; Olson and Strom 2002; Strom et al. 2001) where they
654 have shown the capability of cropping down iron-stimulated diatom blooms (Saito et al. 2006). While
655 species of *Gyrodinium* were prevalent across all water masses in our study, their abundance was higher in
656 more productive STF waters (Figure 8, Figure S9), where higher Chl *a* concentrations was accompanied
657 by increased abundance of diatoms and larger phytoplankton cells, confirming their pivotal importance in
658 pelagic foodwebs as the link between primary producers and metazoan zooplankton (Zeldis and Décima
659 2020).

660 **4.3.5. Bacillariophyta (Diatoms)**

661 Diatoms tended to be more abundant in the STF metabarcodes compared to STW and SAW although
662 relatively high contributions (>30%) were at times attained in all water masses (Figure 6, Figure 7). Most
663 abundant ASVs in STF waters were identified as polar-centric Mediophyceae_X sp., *Thalassiosira* sp.
664 and *Fragilariopsis sublineata* (Figure 8, Figure S9), in the later case mainly due to their abundance in
665 the southern STF flowing next to C. Plateau (Figure S10). Several diatom species including *F. sublin-*
666 *eata/keruelensis*, *F. cylindrus*, *Chaetocerus peruvianus*, and *Cylindrotheca closterium* were significantly
667 more abundant in the STF compared to STW, but not compared to SAW (Figure S11) supporting the
668 greater resemblance of diatoms assemblage between STF and SAW.

669 In SAW, the most abundant diatom ASVs were identified as closely related to *F. sublineata/keruelensis*
670 (Figure 8) consistent with the preference of *Fragilariopsis* species for SAW inferred from microscopy
671 analysis (Chang and Gall 1998). *F. sublineata* has been reported to dominate in sea ice algal biomass and for
672 being well adapted to low light conditions (McMinn et al. 2010); however it is seldom reported among the
673 dominant species in Southern Ocean surface waters, where other species like *F. curta* and *F. keruelensis*
674 tend to dominate (Mohan et al. 2011; Olguín and Alder 2011; Quéguiner et al. 1997) supporting the
675 identification of ASVs assigned to *F. sublineata/F. keruelensis* in this study to *F. keruelensis*. However,
676 the taxonomic assignment of the most abundant *F. sublineata/F. keruelensis* ASVs (e.g., asv_0036, asv_
677 0061) were only closely related to referenced sequences in PR2 (the bootstrap value at species level
678 assignment of ASVs <50%) and their sequence showed 5 to 7 mismatches with the annotated sequences
679 (Figure S17) which highlights the intraspecific diversity of the *Fragilariopsis* spp. The low silicate
680 characteristics of SAW east of New Zealand (Dugdale et al. 1995) is likely a key factor responsible for the
681 southward increase of heavily silicified diatoms like *Fragilariopsis* spp. which tended to be lowest showed
682 in the Bio-SAM, intermediate on C. Plateau and highest in southern most waters of the SAF (Figure S10)
683 in a way consistent with their tendency to dominate south of the SAF (Assmy et al. 2013; Pinkernell and
684 Beszteri 2014). Furthermore, the shift in the relative abundance of the dominant ASVs assigned to *F.*
685 *sublineata/keruelensis* (ASV0036 and ASV0061) observed between subantarctic waters north (Bio-SAM,
686 C. Plateau) and south of the SAF (SAF) (Figure S11, Figure S17) suggests potential differences in their
687 silicate requirements for phytoplankton growth.

688 In STW, in addition to unidentified Polar-centric Mediophyceae_X sp., other small species such as
689 *Minidiscus trioculatus* and *Minutocellus polymorphus* were identified among the most abundant diatoms
690 (Figure 8, Figure S9), consistent with the dominance of these small diatom taxa in neritic-modified STW
691 of the upstream Southland Current (Allen et al. 2020). While most common genera reported in STW
692 (and STF) by microscopy analysis (e.g. *Thalassiosira* spp., *Chaetoceros* spp., *Guinardia* spp.) were
693 also detected by DNA metabarcoding, the small nano-sized species revealed as numerically dominant
694 by DNA approaches can be overlooked by microscopy (Chang and Gall 1998). Diatoms are generally
695 conceptualized as the microplankton group associated with new production and high export potential
696 (Legendre and Lefevre 1995; Uitz et al. 2006; Vidussi et al. 2001). However, there are increasing evidence
697 supporting the importance of small nano- and even pico-sized diatoms in both coastal and oceanic systems
698 (Arsenieff et al. 2020; Buck et al. 2008; Hernández-Ruiz et al. 2018; Lomas et al. 2009). Our results
699 showing the dominance of *M. trioculatus* and *M. polymorphus* in STW particularly during the more
700 productive conditions of the Spring Bloom II and the STF over the Chatham Rise (Figure 9, Figure S10)
701 further support the important role played by small diatoms in driving open-ocean phytoplankton production
702 (Leblanc et al. 2018).

703 **4.3.6. Pelagophyceae**

704 Pelagophyceae showed the opposite trend compared to diatoms and have their lowest abundance associated
705 with the STF zone (Figure 6). Their abundance and relative contribution increased towards SAW (Figure 5,
706 Figure 6 in agreement with pigment-based dominance of *Pelagophyceae* in the SA waters east of C.
707 Plateau (DiTullio et al. 2003). The relative abundance of this class and *Pelagomonas* species also
708 decreased following natural or experimental iron addition experiments in HNLC waters of the SO (Irion
709 et al. 2020; Thiele et al. 2014). These observations are consistent with the physiological advantage in iron
710 uptake of pelagophytes over other small eukaryotic phytoplankton groups (Timmermans et al. 2005) and
711 indicates a competitive advantage for pelagophytes under oligotrophic conditions. Vertically, the relative
712 contribution of this class increased with depth (Figure 7, Figure S16) in agreement with their preference for
713 deeper layers (Cabello et al. 2016; Gall et al. 2008; Latasa et al. 2017) and their physiological adaptation to
714 low light and high nutrient environments (Dimier et al. 2009; Dupont et al. 2015). This vertical segregation

715 was evident in both STW and SAW samples despite the different specific composition observed in
716 each water mass (Figure S9) with *Pelagomonas calceolata* (ASV0044) and unidentified pelagophyte
717 (Pelagophyceae_XXX.sp, ASV0081) being the most abundant species in STW and SAW, respectively
718 (Figure 8)(Figure S9). *P. calceolata* is a widespread species (Andersen et al. 1996; Moon-van der Staay
719 et al. 2001). Whether the ubiquity of this species is bound to high genetic diversity or physiological
720 versatility is not clear. In our study, several ASVs were assigned to *P. calceolata* and while the most
721 abundant one showed preference for STW, other less abundant ASVs were significantly more abundant
722 in SAW (Figure S11). Similarly, we found different water mass preferences among ASVs assigned to
723 unidentified pelagophytes, with some preferentially associated with STW or SAW but interestingly none
724 with STF (Figure S11). While these observations suggest that different ASVs may reflect ecologically
725 relevant different units (Rodríguez et al. 2005) they also highlight the importance of culture isolations and
726 species characterization to better determine the diversity of pelagophyte assemblages.

727 **4.3.7. Prymnesiophyceae**

728 Prymnesiophyceae were prevalent across all water mass metabarcodes (Figure 5) but tended to be more
729 abundant in SAW (Figure 6). Overall, their relative contribution to eukaryotic phytoplankton assemblages
730 was lower than depicted by pigment-based analyses of open ocean microbial communities (Andersen
731 et al. 1996). The prevalence of 19'hexanoyloxyfucoxanthin pigment marker in oceanic waters and the
732 application of quantitative methods (e.g. CHEMTAX) have shown that *Prymnesiophyceae* represents
733 between 20-50 % of the phytoplankton community in oceanic waters (Andersen et al. 1996; DiTullio
734 et al. 2003; Latasa et al. 2005; Liu et al. 2009; Swan et al. 2016). Such dominance has been also depicted
735 by improved genomic approaches that revealed the extremely high genetic and functional diversity of
736 non-calcifying prymnesiophytes (Cuvelier et al. 2010; Liu et al. 2009). In our study, non-calcifying
737 species, mainly assigned to *Phaeocystis* spp. and *Chrysochromulina* spp., dominated the group (Figure S9)
738 in agreement with DNA-based studies in the SW Pacific region (Sow et al. 2020; Wolf et al. 2014). The
739 abundance and relative contribution of *Phaeocystis* spp. was lowest in STW, intermediate in STF and
740 peaked in SAW while *Chrysochromulina* spp. followed the opposite trend with higher contributions
741 associated with STW (Figure S9). The dominance of *Phaeocystis* spp. in SAW was mainly driven by *P.*

742 *antarctica* (Figure 8), corresponding with the prominence of this group in the Southern Ocean (Verity
743 et al. 2007) and observed decreasing abundance observed from SAW towards STW of the SW Pacific
744 region during austral autumn-Winter (Sow et al. 2020). These results contrast with the similar spatial
745 distribution *P. antarctica* spp. metabarcodes observed between contrasting conditions on and off the
746 Kerguelen Plateau (Irion et al. 2020). Strains assigned to *P. globosa* and *P. cordata* were also detected in
747 all water masses although they tended to be more prevalent and abundant in STW compared to SAW in the
748 New Zealand region (Sow et al. 2020). Coccolithophores are an important component of phytoplankton
749 communities in the Southern Ocean region extending from the STF to the Polar Front known as the
750 Great Calcite Belt (Balch et al. 2016; Chang and Northcote 2016). *Gephyrocapsa oceanica* was the most
751 prevalent and abundant coccolithophore species found in our study. ASVs assigned of this species were
752 found across all water masses but tended to be most abundant in the STF (Figure S9; Figure S10) in
753 agreement with previous microscopy-based studies in this region of the SW Pacific (Rigual-Hernández
754 et al. 2020; Saavedra-Pellitero et al. 2014). *Emiliana huxleyii*, which generally dominate coccolithophore
755 assemblages in this region (Chang and Northcote 2016; Saavedra-Pellitero et al. 2014), and in the Southern
756 Ocean (Balch et al. 2016; Holligan et al. 2010) showed very low abundances across the different water
757 masses surveyed in this study (data not shown).

758 **4.3.8. Cryptophyceae**

759 In our datasets, the contribution of *Cryptophyceae* metabarcodes was relatively low on average (<3%)
760 but showed increasing abundance from SAW to STW where they represented up to 10% of protistan
761 reads in the euphotic zone (Figure 6). The genus and species composition of this group also differ
762 substantially between STW and SAW in our dataset (Figure S9). Similar water-mass preference was
763 depicted by quantitative pigment analysis in the same STFZ region over the Chatham Rise, where
764 cryptophytes contribution in STW (47-63% chl *a*) was higher than in SAW (6% chl *a*) in one of the two
765 consecutive springs surveyed (Delizo et al. 2007). Cryptophytes are an ubiquitous phytoplankton group
766 with widespread distribution from coastal to open oceanic systems and from tropical to polar latitudes
767 (Buma et al. 1992; Nunes et al. 2019; Piwosz et al. 2013). They have been reported to form blooms
768 in coastal embayments (Jeong et al. 2013; Johnson et al. 2013) and coastal Antarctic waters favoured

769 by low salinity conditions (Moline et al. 2004; Nunes et al. 2019; Schofield et al. 2017). The higher
770 contributions we observed in STW relative to SAW, however, argues against the direct influence of salinity
771 on cryptophytes at least in open-ocean waters. Cryptophytes have been also observed to respond positively
772 to iron fertilization in HNLC waters of the North Pacific (Sato et al. 2009; Suzuki et al. 2009) suggesting
773 that their lower abundance in SAW in our study could be related to iron-limited conditions characteristic of
774 the subantarctic region. The contribution of cryptophytes in STW was highest during the open-ocean spring
775 bloom (Spring Bloom II-TAN1212 voyage) and in shelf-slope stations of the EAUC current (Cross-shelf
776 Exchange-TAN1604 voyage) consistent with their preference for more nutrient-rich conditions (Carreto
777 et al. 2016; Fuller et al. 2006; Latasa et al. 2010). Significant contributions by cryptophytes has also been
778 observed in open ocean waters of the NW Mediterranean at the termination of the spring bloom (Vidussi
779 et al. 2000) where they even dominated the surface mixed layer community at highly stratified stations.
780 Interestingly, the higher contribution of cryptophytes in our study occurred towards the end of the spring
781 bloom (TAN1212)(Figure S16), coincident with strong surface stratification and biomass accumulation
782 (Chiswell et al. 2019), and supporting the importance that stratification may have on this group compared
783 to salinity. *G. cryophile* and other species of *Teleaulax* have been reported as mixotrophic (Schneider et al.
784 2020), which could favor their increase in later stages of the spring bloom when the coincident decrease of
785 nutrients and increase of potential preys tend to favor mixotrophy (Mitra et al. 2014).

786 4.3.9. Heterotrophic and mixotrophic protists below the euphotic zone

787 We used the samples collected during six voyages of the Biophysical Moorings time-series between 2009-
788 2012 to investigate the protistan community composition below the euphotic zone (Table 1). Metabarcoding
789 datasets here were clearly dominated by Radiolaria (Figure 7), a holoplanktonic amoeboid group with
790 widespread distribution in modern oceans (Biard et al. 2016; Suzuki and Not 2015). Radiolaria are mainly
791 heterotrophic protists with many mixotrophic species in the photic zone bearing endosymbiotic microalgae
792 that can contribute substantially to primary production in oligotrophic oceans (Caron et al. 1995; Decelle
793 et al. 2015). While Radiolaria are found throughout the entire water column, their contribution to plankton
794 biomass (Biard et al. 2016; Boltovskoy and Correa 2016) and metabarcodes (Ollison et al. 2021) tends to
795 be greater in the mesopelagic ocean, in a way consistent with our metabarcoding results. In this study,

796 significant contributions of Radiolaria were mainly constrained to the aphotic zone (Figure 7). This depth-
797 related pattern, contrasts with previously reported abundance of Radiolaria, and symbiotic Collodaria, in
798 the sunlit ocean (Vargas et al. 2015). We found substantial contributions of photosymbiotic Collodaria at
799 times, particularly in SAW (Figure S9), but the vertical distribution of Radiolaria below the euphotic zone
800 suggested they were mainly composed of heterotrophic species. The high copy number of 18S rDNA in
801 Radiolaria may contribute to their high relative abundance metabarcoding datasets ((Gutierrez-Rodriguez
802 et al. 2019; Vargas et al. 2015)); however, it is unlikely to be responsible for their dominance, particularly
803 in relation to dinoflagellates and ciliates, which are also known to have high copy numbers (Gong et al.
804 2013; Piredda et al. 2017). Moreover, the positive relationship between cell length and 18S rDNA copy
805 number (Biard et al. 2017; Zhu et al. 2005) and the higher C and N density (mass: volume) of Radiolaria
806 compared to other protists (Mansour et al. 2021) suggest that gene-based relative abundance of these groups
807 was likely reflected in their relative contribution to the community biomass.

808 Utilising their sticky pseudopodia and large size, Radiolaria dwelling below the euphotic zone can
809 effectively intercept sinking particles and act as flux-feeders influencing the quality and quantity of
810 vertical fluxes (Ohman et al. 2012; Stukel et al. 2019). The presence of mineral skeletons, made of
811 silica (Polycystinean groups) or strontium sulfate (e.g. Acantharea), provides substantial mineral ballast
812 (Takahashi 1983) conferring them a key role in vertical export that is supported by their common presence
813 in sediment traps (Bernstein et al. 1987; Gutierrez-Rodriguez et al. 2019; Michaels et al. 1995; Preston
814 et al. 2020) and their enrichment in suspended and sinking particles (Duret et al. 2020). Despite their
815 abundance and their important role in biogeochemical processes (Biard et al. 2016; Guidi et al. 2016)
816 little is known about the vertical distribution of Radiolaria particularly in the meso- and bathypelagic
817 ocean (>500 m) (Biard and Ohman n.d.; Boltovskoy 2017; Llopis Monferrer et al. 2021; Ollison et al.
818 2021). In our study, we found an opposite distribution between Acantharea and RAD-B, which showed
819 preference for the upper (<500 m) and deeper (>500 m) mesopelagic samples, respectively (Figure 7).
820 Among Acantharea, most sequences were assigned to the order *Chauchantida* (Figure S9, Figure S14),
821 which has been found in sinking particles collected in the Southern Ocean (Duret et al. 2020) and the
822 California Current (Gutierrez-Rodriguez et al. 2019; Preston et al. 2020). In a recent study conducted in
823 the Southern Ocean, RAD B was reported to be enriched in small (<10 μ m) suspended particles relative

824 to sinking particles (Duret et al. 2020) while RAD-A were consistently found in sinking particles reaching
825 abyssal depths in the California Current Ecosystem (Preston et al. 2020). The higher contribution of
826 RAD-B relative to RAD-A we found in our study is consistent with the tendency of RAD-B to remain
827 suspended compared to RAD-A with greater sinking potential. Among the polycystinean Radiolaria,
828 Spumellarida was the most important order, in a way consistent with observed abundance of this order
829 in sinking particulate organic matter collected in sediment traps deployed in mesopelagic and abyssal
830 depths in the California Current Ecosystem (Gutierrez-Rodriguez et al. 2019; Preston et al. 2020). Several
831 ASVs assigned to Spumellarida Group I were among the most abundant species in our study (Figure S13).
832 Interestingly, some of these ASVs showed preference for SAW while others were more abundant in STW,
833 highlighting the need to improve our taxonomic knowledge of this group.

834 Relative contributions of ciliates were also higher below the euphotic zone, mainly driven by ASVs
835 affiliated with Spirotrichea class (Grattepanche et al. 2016). Most abundant ASVs assigned to the order
836 Strombidiida (Oligotrichia, Spirotrichea) and *Leegaardiella* sp. (Choreotrichia, Spirotrichea), which have
837 been reported below the euphotic zone at meso- and bathy-pelagic depths (Duret et al. 2020; Grattepanche
838 et al. 2016). In addition to Spirotrichea, class Oligohymenophorea and Nassophorea contributed substan-
839 tially in both STW and SAW, mainly sustained by ASVs assigned to OLIGO5 (*Oligophymenophorea*)
840 and *Discotrichidae* (*Nassophorea*), both previously reported in mesopelagic waters (Duret et al. 2020).
841 The dominance of these groups was consistent across different water masses off eastern New Zealand
842 although some species like *Leegaardiella* (Oligotrichia) and *Strombidium_k_sp* (Choreotrichia) showed
843 preference for STW and SAW, respectively (Figure S11). Ciliates below the euphotic zone feed on
844 bacteria and small protists associated with particulate organic matter (Caron et al. 2012). Several species
845 also have the potential to engage in photoautotrophy and phagotrophy (Leles et al. 2017). By doing so,
846 ciliates play an important role within planktonic food webs contributing to trophic transfer and nutrient
847 recycling in the dark ocean. The high taxonomic diversity and abundance of heterotrophic protists in this
848 and previous studies (Duret et al. 2020; Grattepanche et al. 2016; Ollison et al. 2021; Zoccarato et al.
849 2016) highlights their importance in planktonic systems below the euphotic zone and emphasises how
850 little we know about their ecological role in the food web functioning. Further studies focusing on the
851 taxonomic and functional diversity of heterotrophic protists have the potential to shed light on the trophic

852 and biogeochemical processes that transform organic matter in the dark ocean, and hence significantly
853 improve our understanding of the biological carbon pump and natural deep-ocean carbon sequestration.

854 **5. CONCLUSIONS**

855 The spatial diversity patterns observed are in agreement with global trends of decreasing diversity at higher
856 latitudes and colder waters. However, deviations from this general pattern were also observed regionally.
857 Species richness and diversity of protist communities in the STF were systematically lower compared to
858 adjacent STW and SAW waters in the northern and southern regions of the STFZ surveyed, highlighting
859 the importance of oceanographic features in determining regional diversity. Dinoflagellates and green
860 algae co-dominated the protist community in the euphotic zone but water-mass specificity emerged at lower
861 taxonomic levels within these and other major taxonomic groups and the community composition varied
862 consistently between water masses. Within green algae for instance, *Mamiellophyceae* class dominated
863 in STW driven by several species showing different regional abundance, while *Chloropicophyceae* class
864 became dominant in SAW where several ASVs assigned to *Chloroparvula pacifica* appeared among
865 the most abundant taxa. Interestingly, other less abundant ASVs identified as *Chloroparvula pacifica*
866 showed statistically significant preference for STW. Analogous intra-specific variability was observed
867 within species belonging to other phytoplankton classes with widespread distribution (e.g. *P. antarctica*,
868 prymnesiophytes; *P. calceolata*, pelagophytes) suggesting the genotypic diversity may be linked to
869 ecological traits that influence distribution patterns. Although chl *a* levels comprised in this study were
870 relatively low, small rather than large taxa dominated the phytoplankton proliferations associated with
871 spring bloom conditions and the STFZ suggesting that picoplankton can also be important for primary
872 and export production either directly or through zooplankton grazing. The mesopelagic zone was clearly
873 dominated by radiolarian sequences, supporting the importance of this group for the functioning of the
874 dark ocean. Taxonomic assignment revealed a diverse assemblage of Radiolaria and a taxon-specific water
875 mass and vertical distribution patterns. However, further research is needed about the ecology of these
876 organisms to link this compositional variability to their function in the system.

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1403 *Oceanography* 149, pp. 16–26.

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1416 **8. AUTHOR CONTRIBUTION STATEMENT**

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1418 curation, Visualization, Writing - Original Draft & Funding acquisition
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1420 Original Draft & Funding acquisition.
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- 1427 • Matt Pinkerton: Writing - Review, Editing & Funding acquisition.
- 1428 • Scott D Nodder: Conceptualization, Investigation, Data curation, Writing - Original Draft & Funding
1429 acquisition.

1430 **List of Tables**

1431 **Table. 1** Summary of cruises from which samples were collected. Information includes
1432 the cruise identification code, start and end dates, the project or region, the water
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1434 collected in each cruise.

1435 **Table. 2** Summary of PERMANOVA analysis including the Water-mass and Region as
1436 categorical variables in addition to the continuous environmental variables. Tem-
1437 perature and salinity represent the surface values. Nitrate (NO_3^-) and chlorophyll
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1439 Analysis was conducted with the Adonis function of the vegan R package

1440 **Table. S1** Supplementary Table S1. Table with the number of samples, the median number
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1443 **Table. S2** Supplementary Table S2. Summary of PERMANOVA analysis results conducted
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1567 of these ASVs abundance (standardized read abundance) in each euphotic sample
1568 in relation to latitude and color coded by region.

1569

Planktonic protist diversity across contrasting Subtropical and Subantarctic waters of the southwest Pacific

1570

1571

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1575

Supplementary Material

1576 **9. SUPPLEMENTARY MATERIAL**

Water mass	N samples	N reads (median)	N ASVs
All	482	15913	16861
SAW	120	17158	7377
STF	91	16431	4261
STW	271	15630	12407

Table S1. Supplementary Table S1. Table with the number of samples, the median number of reads per sample and the number of ASVs in the entire dataset and within each water mass.

Variable	Df	SumsOfSqs	MeanSqs	F.Model	R2	Pr(>F)
Temperature	1	4.929	4.9288	21.639	0.08489	0.001
Salinity	1	7.831	7.8308	34.38	0.13487	0.001
median NO ₃	1	3.362	3.3617	14.759	0.0579	0.001
median Chla	1	1.625	1.6246	7.133	0.02798	0.001
Residuals	177	40.316	0.2278		0.69436	
Total	181	58.062			1	

Table S2. Supplementary Table S2. Summary of PERMANOVA analysis results conducted to calculate the significance of environmental variables ability to explain ASV composition on a subset of 188 samples collected from STW (n=34), STF (n=53), SAW (n=95) for which measurements of presented variables were available. Analysis was conducted with the Adonis function of the vegan R package.

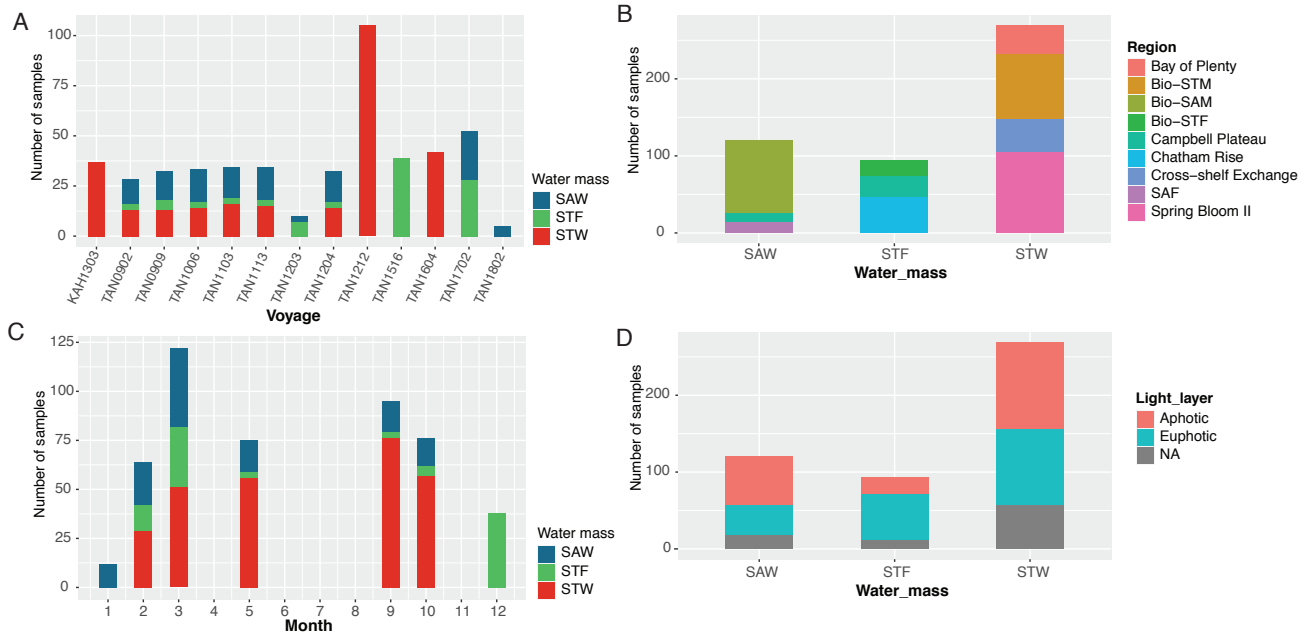


Figure S1. Supplementary Figure S1. Number of DNA samples from each water mass surveyed in different (A) cruises, (B) regions, (C) months of the year, and (D) photic zone.

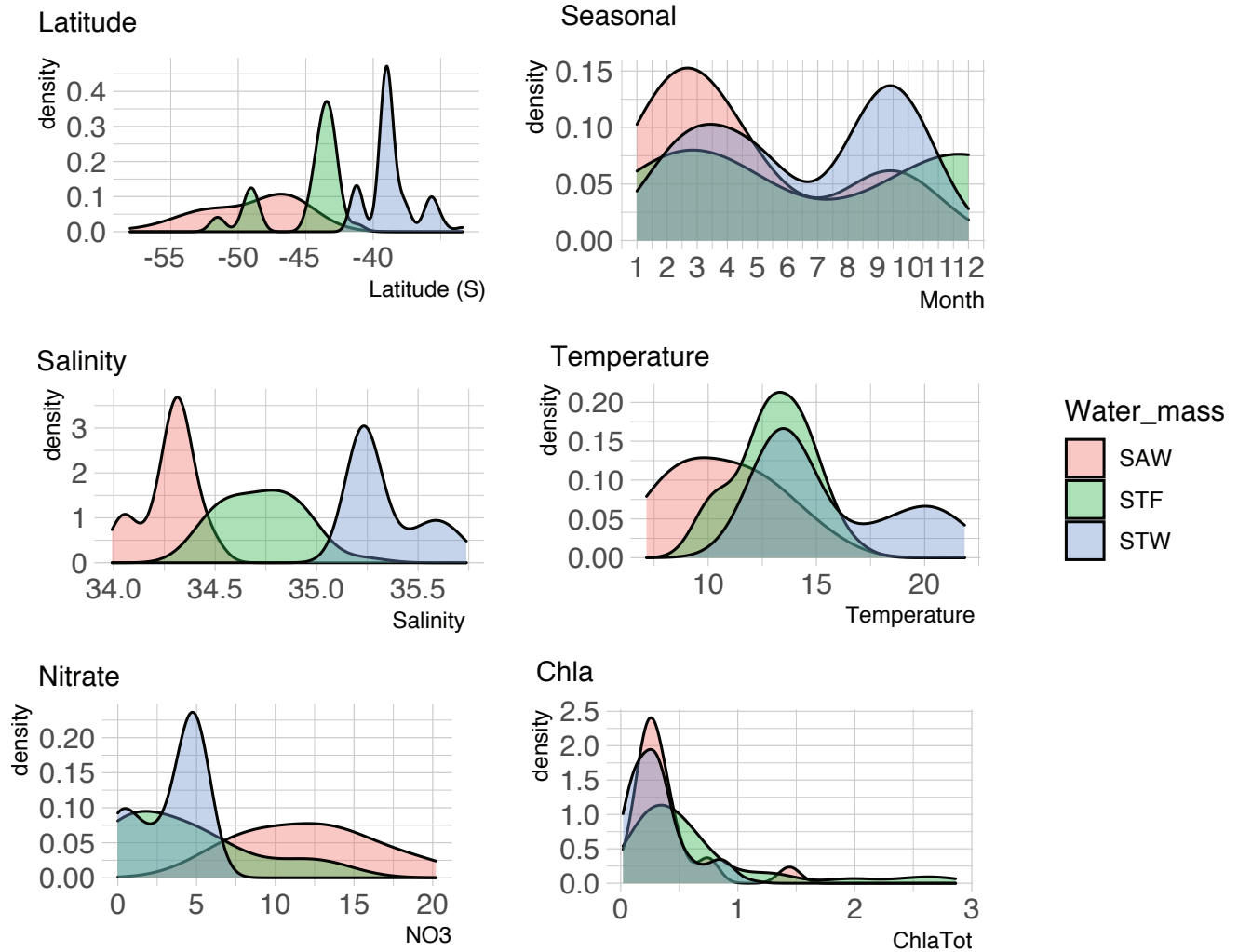


Figure S2. Supplementary Figure S2. Density distribution of DNA samples across latitude, season, and mixed-layer temperature, salinity nitrate, and chlorophyll *a* concentrations in each water mass surveyed.

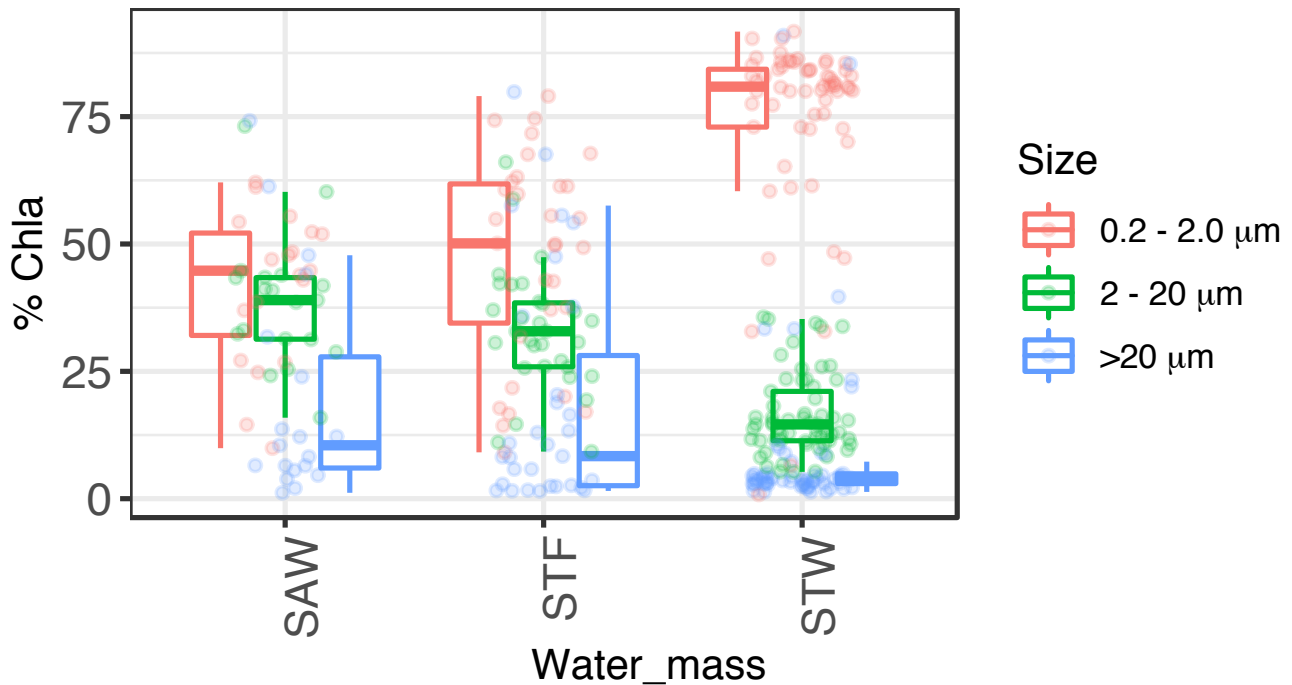


Figure S3. Supplementary Figure S3. Phytoplankton community size-structure. Box-plots show the chlorophyll a concentration associated to pico-, nano-, and micro-phytoplankton nominal size fractions (0.2-2 μm, 2-20 μm, >20 μm) in the surface mixed-layer of each water masses. Each dot correspond to a single sample. Box-plots show the median, first and third quartiles and the values within the $\pm 1.5 * IQR$ (IQR, interquartile range). The figure includes data from TAN1516/Chatham Rise, TAN1702/Campbell Plateau, TAN1212/Spring Bloom, TAN1203/SOAP; n = 102.

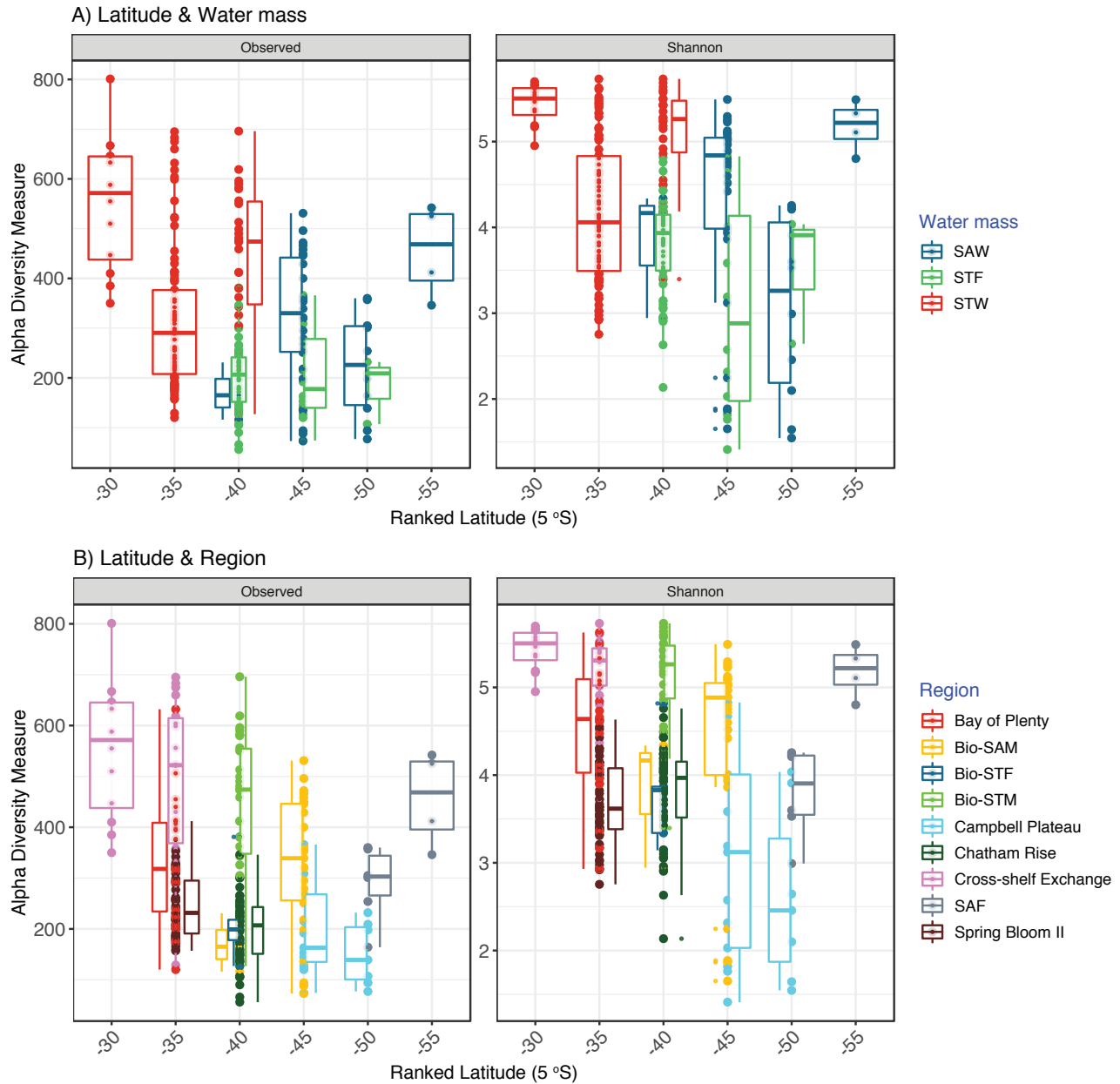


Figure S4. Supplementary Figure S4. Protist diversity richness and Shannon diversity index binned in 5° latitude bins and color coded by (A) water mass and (B) regions.

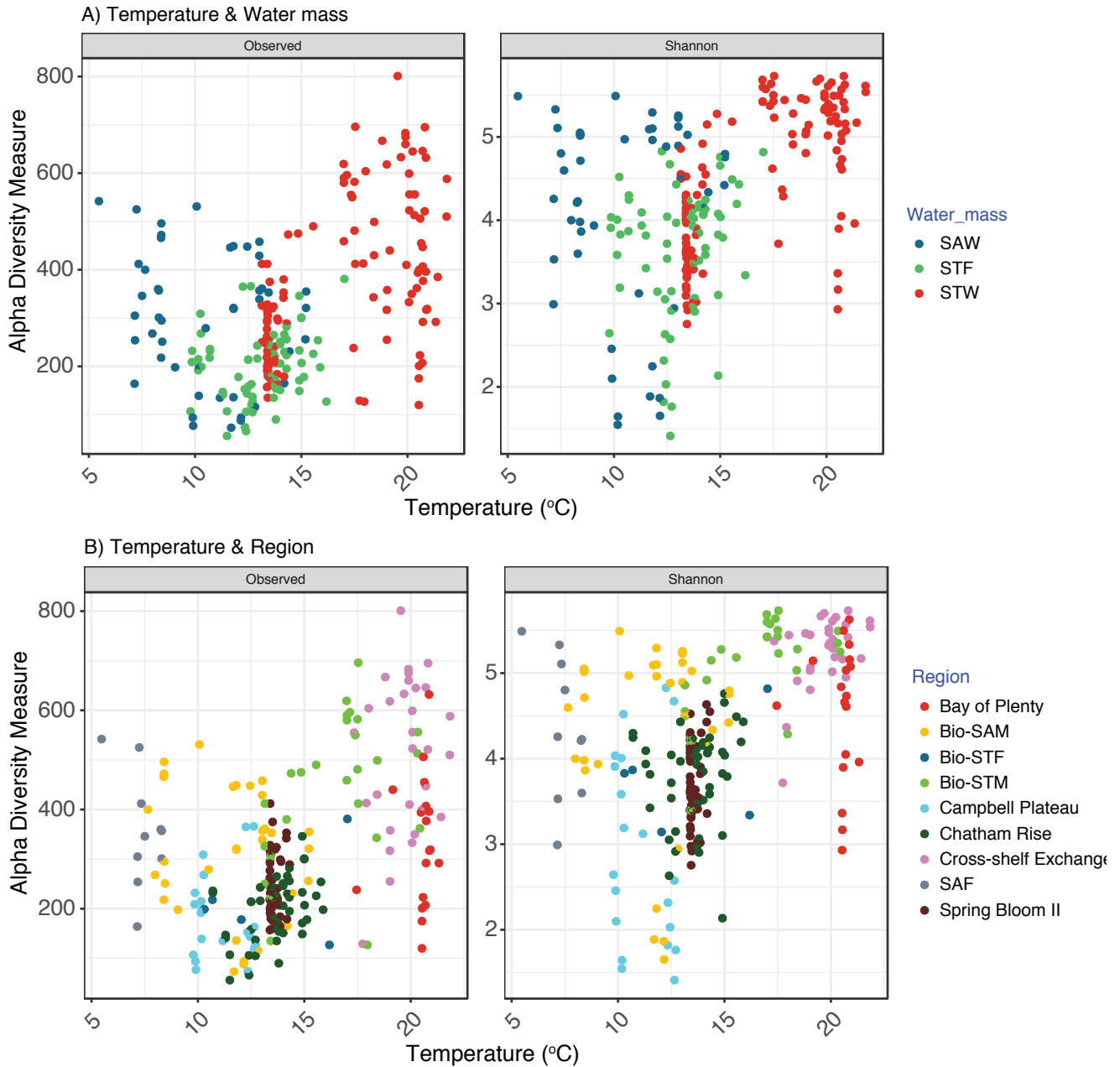


Figure S5. Supplementary Figure S5. Diversity and temperature. Protist diversity richness and Shannon diversity index binned in 5° latitude bins and color coded by (A) water mass and (B) regions.

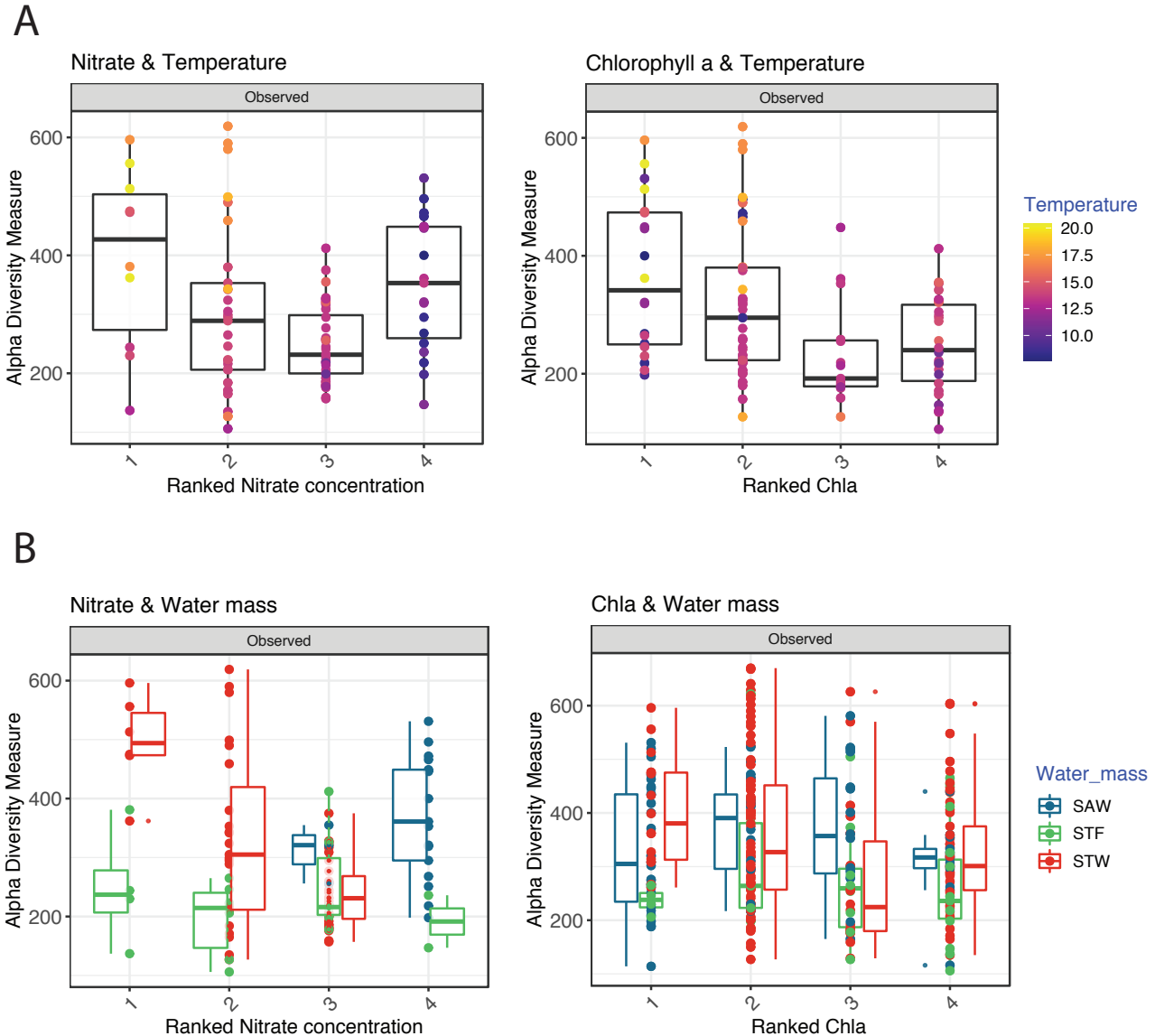


Figure S6. Supplementary Figure S6. Diversity and trophic conditions. Mean protist diversity richness in relation to nitrate and chlorophyll a concentrations in the euphotic zone and binned in four ranks and color coded by (A) temperature and (B) water mass. Function ntile was used to break individual measurements into buckets of the same size.

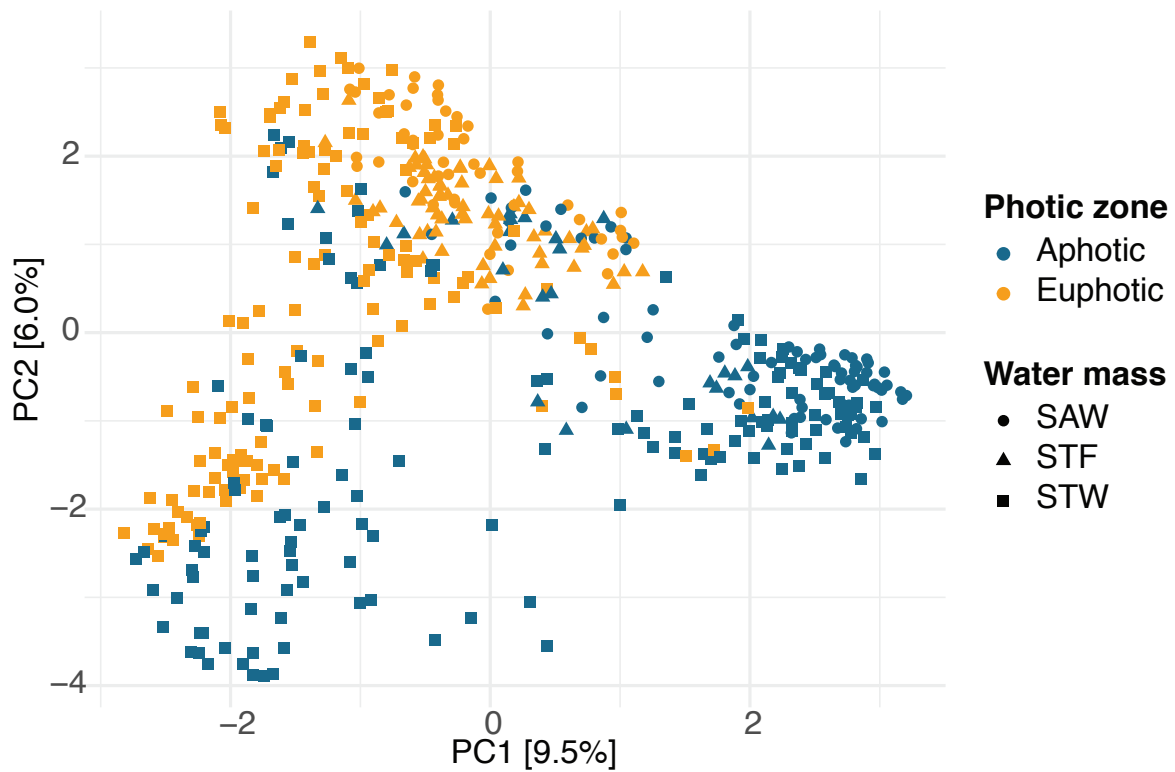
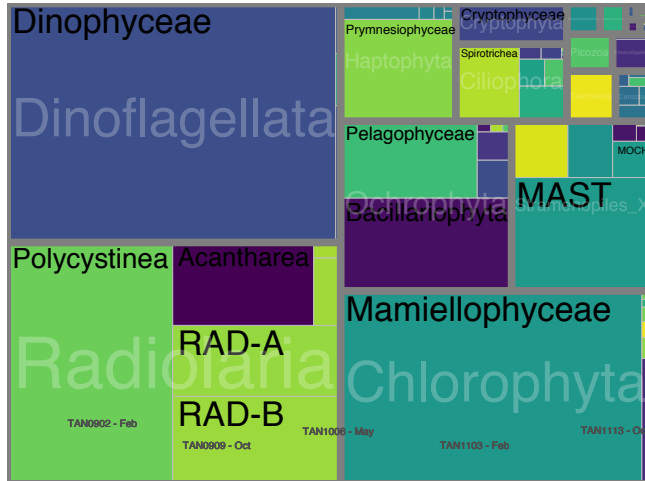
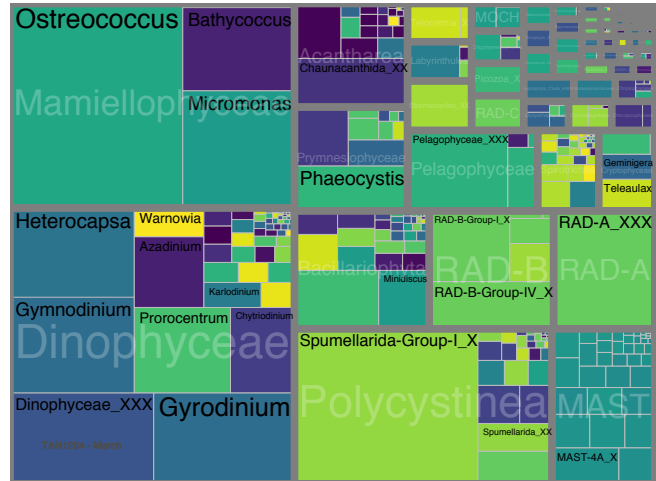


Figure S7. Supplementary Figure S7. Principal component analysis based on ASV composition of all samples coded by symbol color and shape, respectively, for the light layer and water masses where the samples were collected from.

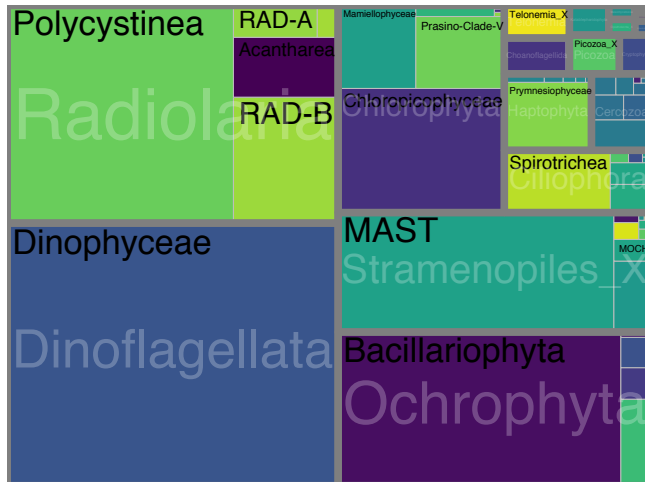
STW - aphotic - Class



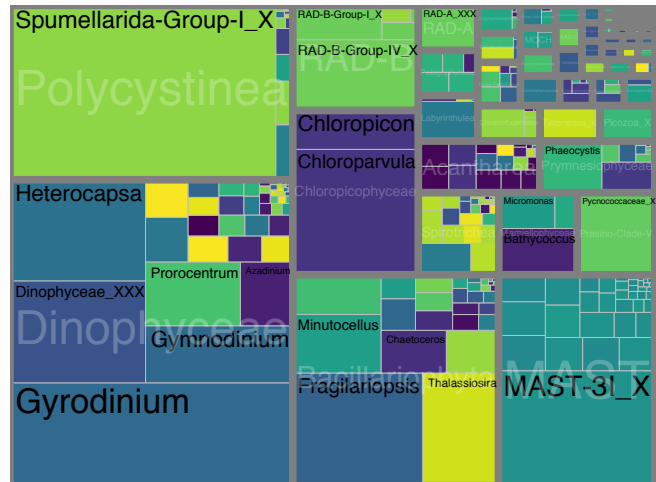
STW - aphotic - Genus



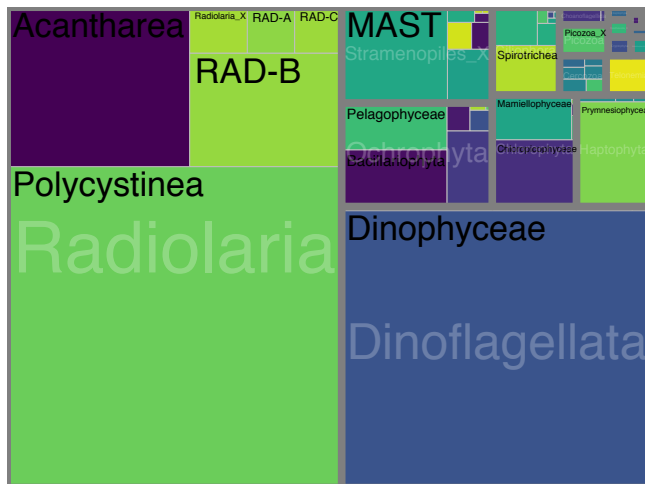
STF - aphotic - Class



STF - aphotic - Genus



SAW - aphotic - Class



SAW - aphotic - Genus

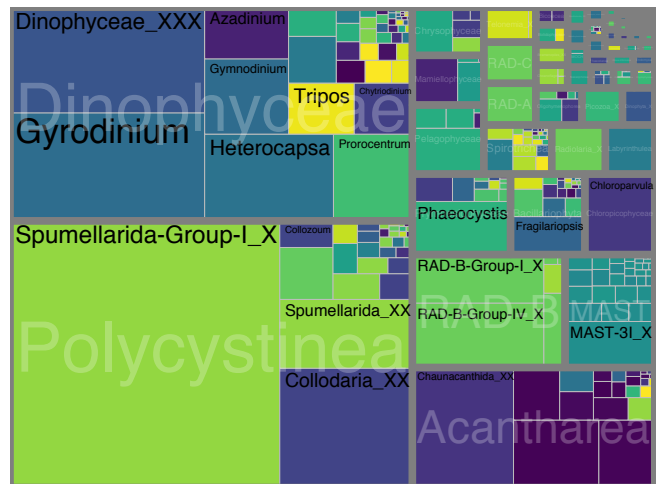


Figure S8. Supplementary Figure S8. Treemaps showing the mean relative abundance of main protistan groups divisions, class and genus in the aphotic zone of the STW, STF, and SAW sites. The area of each taxonomic group in the treemap represents the read abundances affiliated to each group standardized to the median sequencing depth across samples [$\text{median sum otus} * (\text{otu reads} / \text{sum (otu reads)})$]



Figure S9. Supplementary Figure S9. Treemaps showing the mean relative abundance of main protistan classes and species within major taxonomic divisions in the euphotic zone of STW, STF, and SAW water masses.



Figure S10. Supplementary Figure S10. Mean standardized read abundance of most abundant ASVs and assigned species in the euphotic zone of different regions and voyages, color coded for their class affiliation.

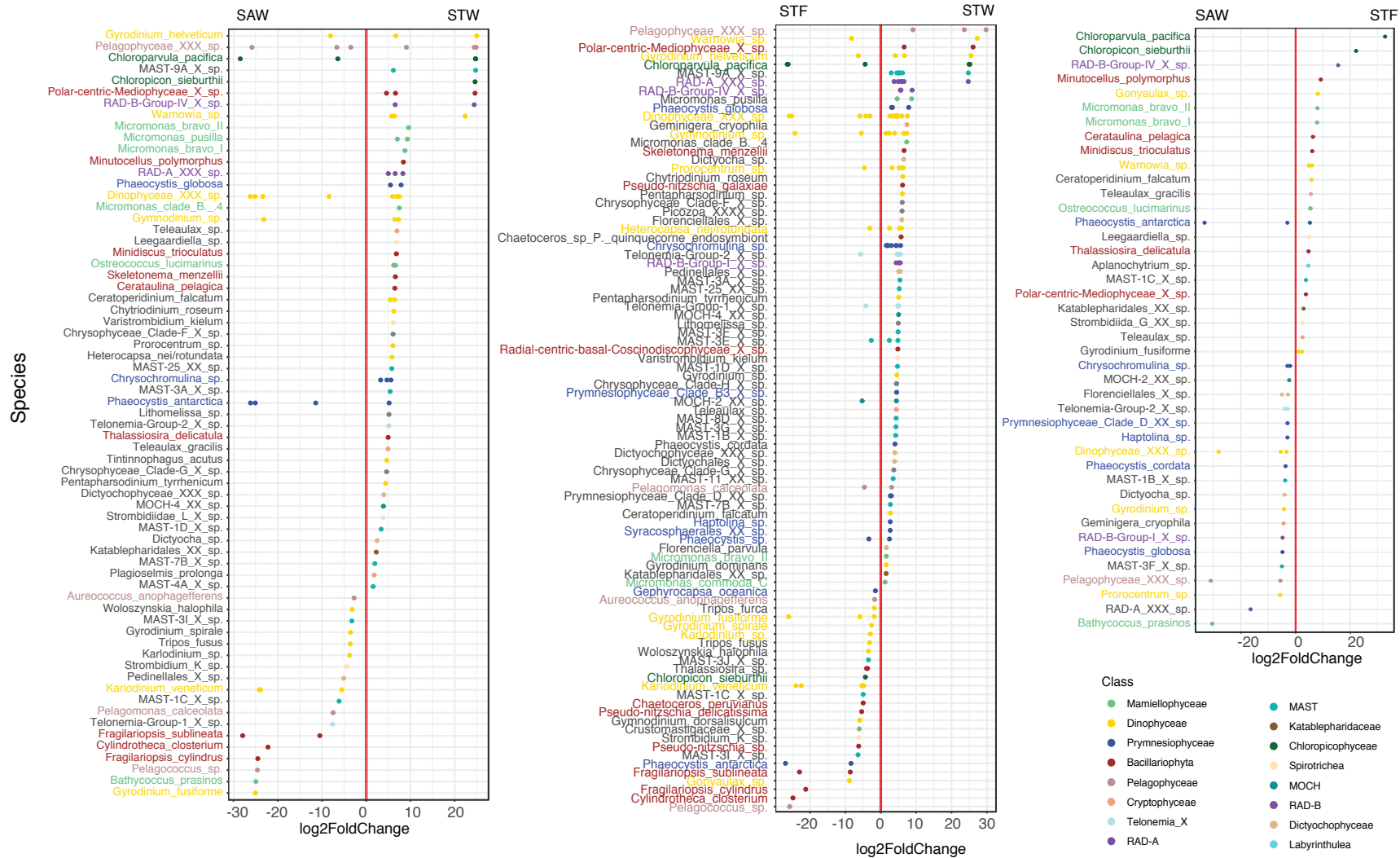


Figure S11. Supplementary Figure S11. Results from DESeq2 analysis depicting the species (Y-axis) with significantly different distribution between the euphotic zone of STW, SAW and STF waters. Difference in the distribution is expressed as the log2 fold change of the difference (X-axis). Each dot correspond to a different ASV color coded by their class affiliation.

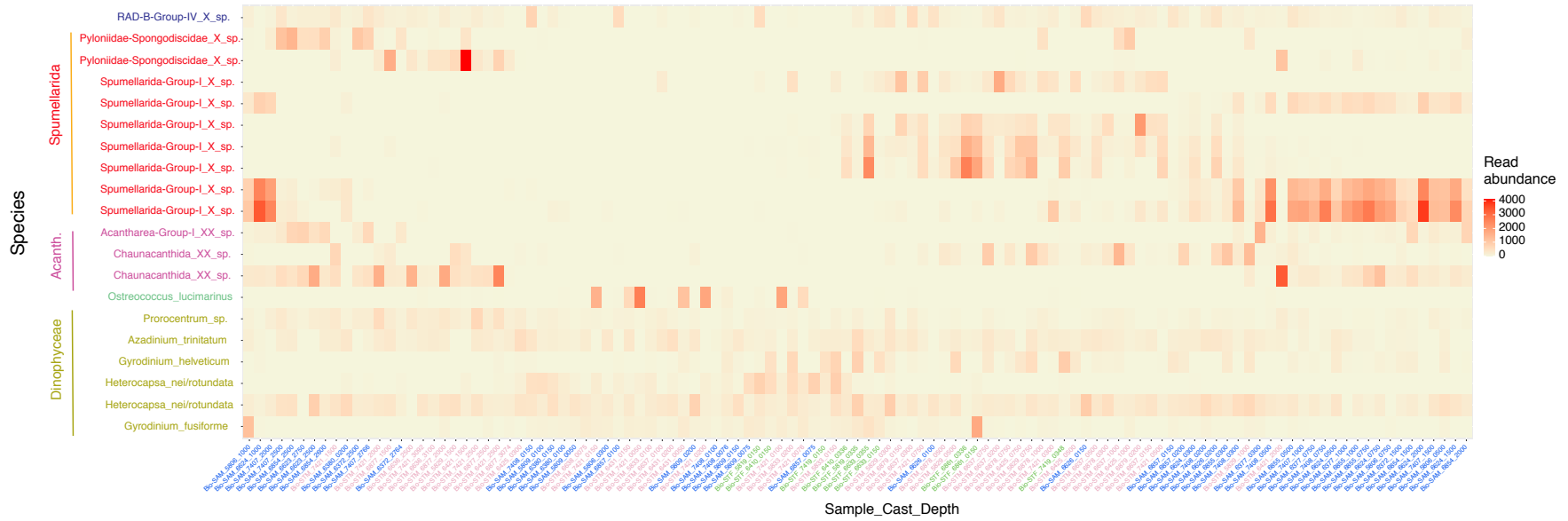


Figure S12. Supplementary Figure S12. Heatmap showing the abundance patterns of the top20 most abundant species in the aphotic zone of the Biophysical Mooring samples. Read abundance were normalized to mean sequencing depth. Samples are clustered using nMDS and Jaccard distance, and color coded according to the sampling site they were collected from (Bio-STM, purple; Bio-STF, green; Bio-SAM, blue). Species were organized and color coded by class affiliation.

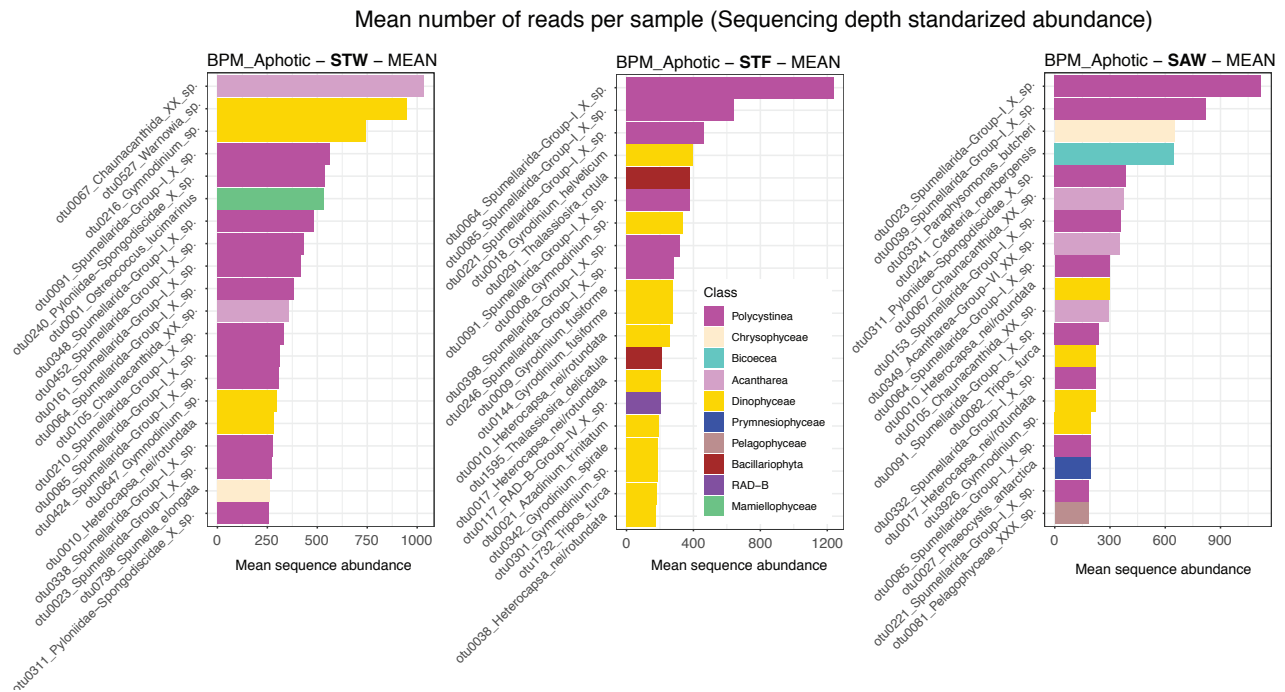


Figure S13. Supplementary Figure S13. Mean standardized read abundance of most abundant ASVs and assigned species in the aphotic zone of different water masses. Bars corresponding to each species color coded for their class affiliation.

BPM-Aphotic – Protist-synd

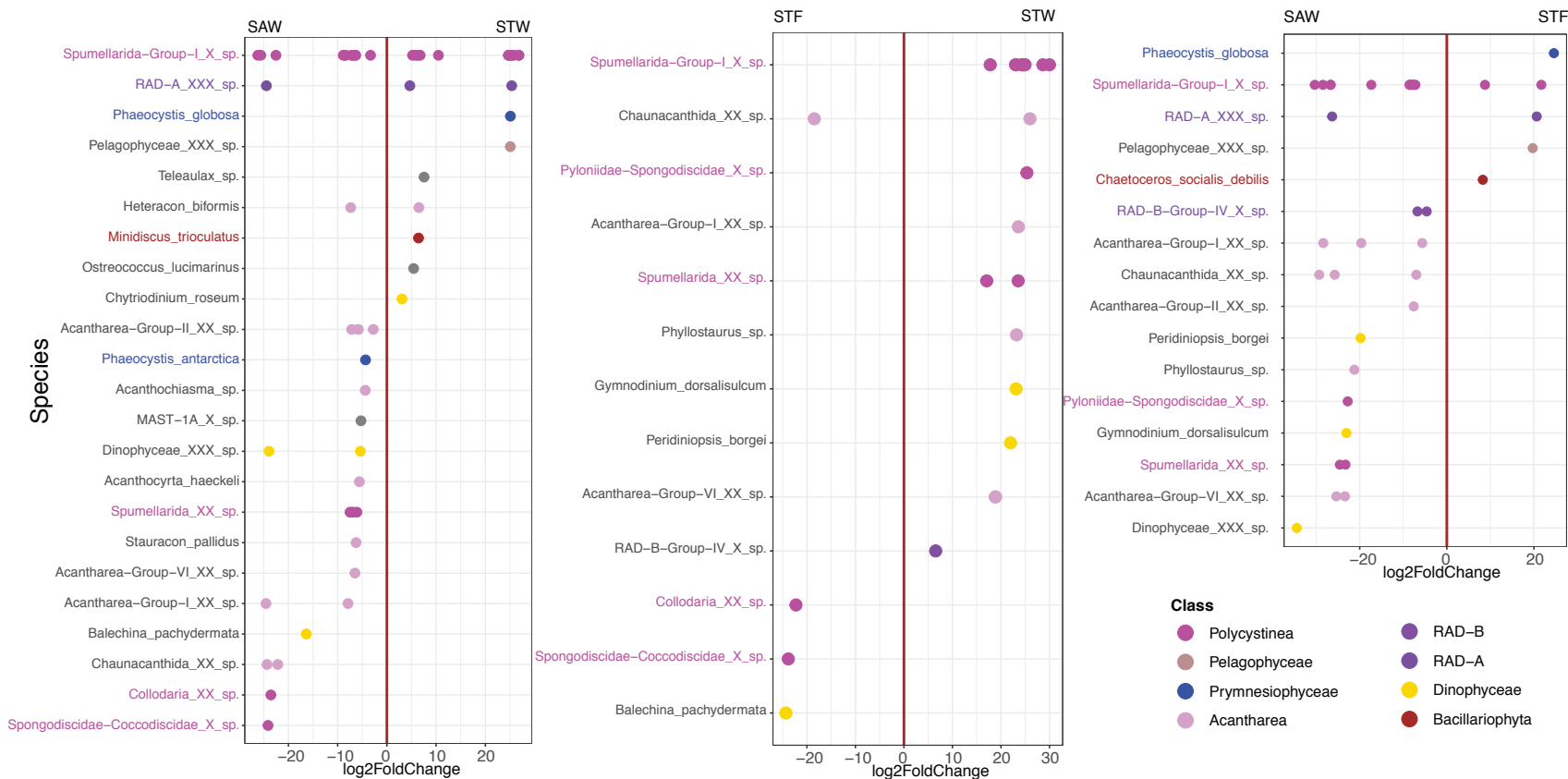


Figure S14. Supplementary Figure S14. Results from DESeq2 analysis depicting the species (Y-axis) with significantly different distribution between the aphotic zone of STW, SAW and STF waters. Difference in the distribution is expressed as the log₂ fold change of the difference (X-axis). Each dot correspond to a different ASV color coded by their class affiliation.

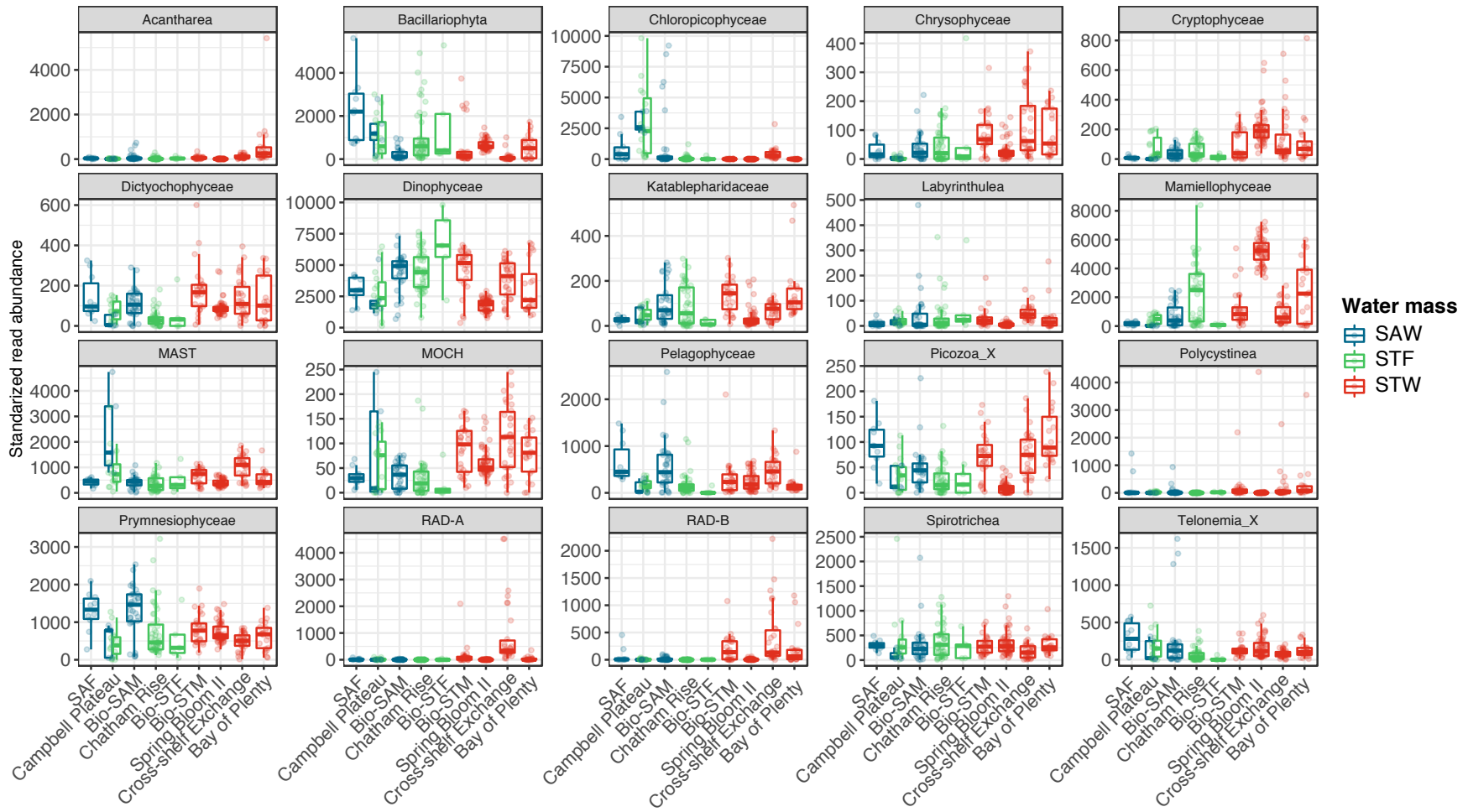


Figure S15. Supplementary Figure S15. Box-plots showing standardized read abundance of twentieth most abundant protist classes in the euphotic zone of different regions and color coded by different water masses. Box-plots show the median, the first and third quartiles (lower and upper hinges) and the values within the $\pm 1.5 * IQR$ (IQR, interquartile range) (line). Points represent values of single samples

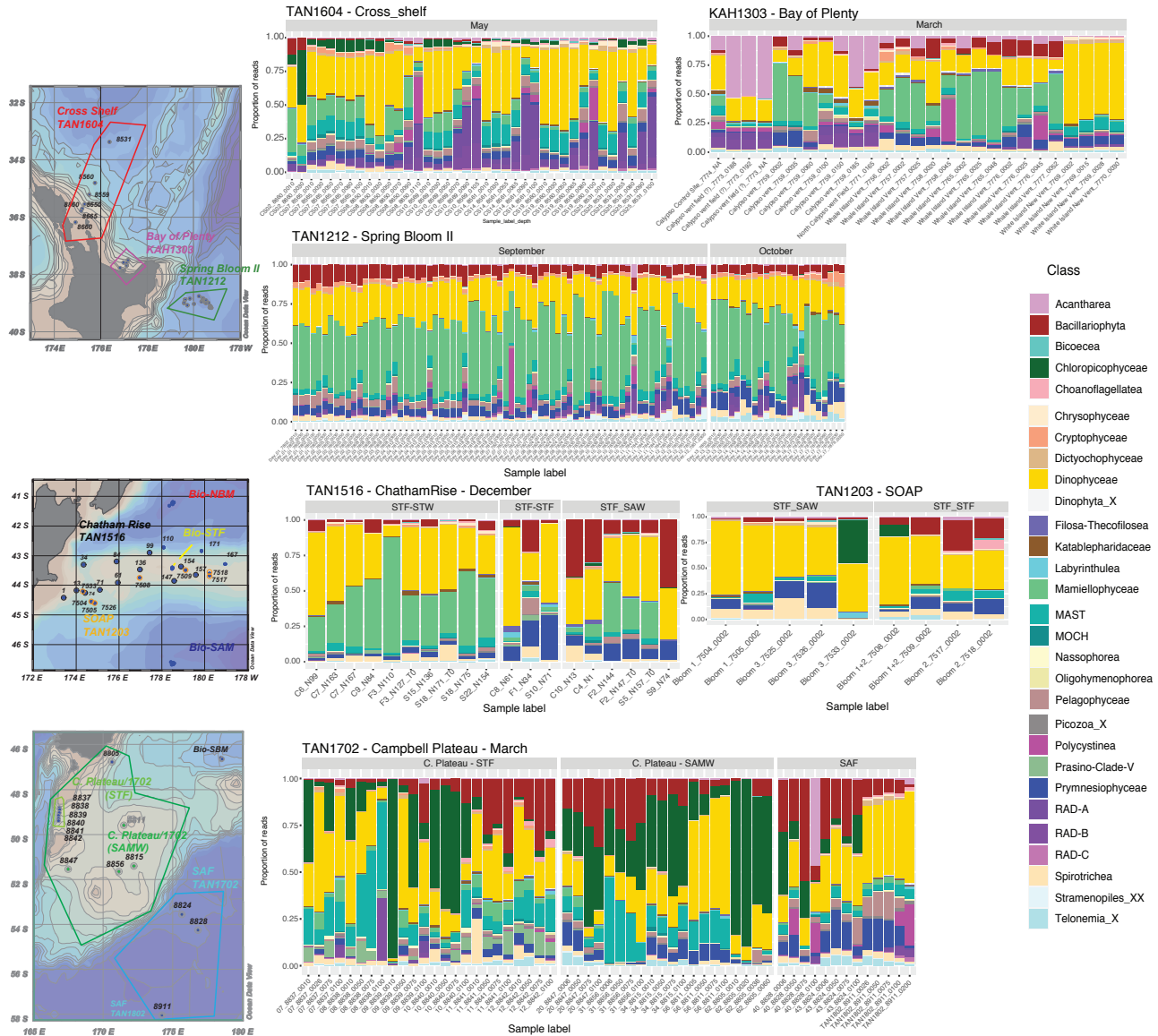
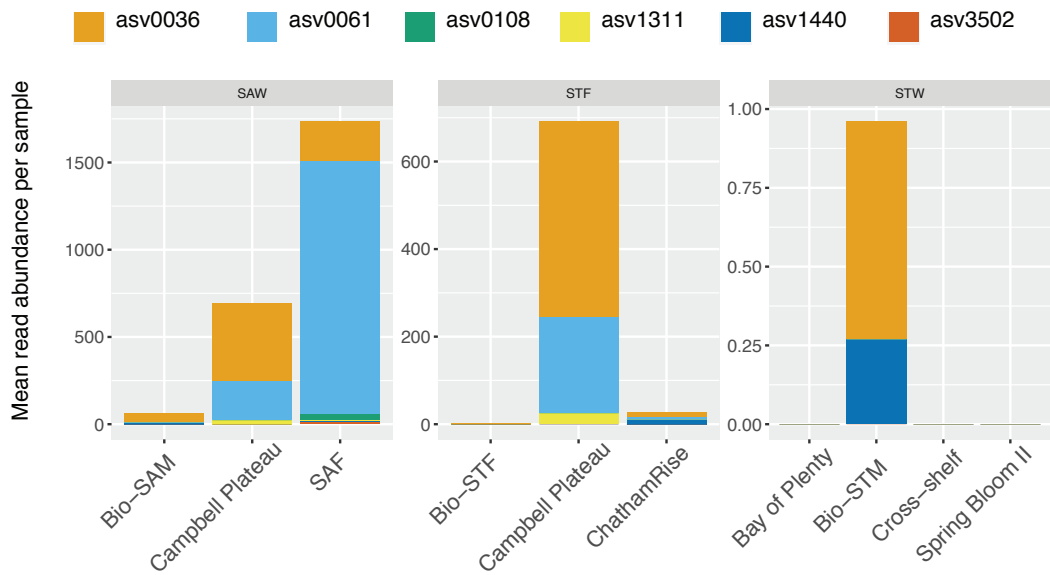


Figure S16. Supplementary Figure S16. Relative read abundance of main protistan classes in samples collected during different voyages and regions within the STW, STF and SAW water masses.

A

	asv_0061 F...	KJ866919.1... AF525665....	AF290085....	asv_0108 F...	asv_1440 F...	asv_3502 F...	asv_1311 F...	asv_3242 F...	asv_0036 F...
asv_0061 F_sublineata		5	5	5	5	10	7	7	6
KJ866919.1 F_kerguele...	5		0	0	0	5	4	6	5
AF525665.1 F_sublineata	5	0		0	0	5	4	6	5
AF290085.1.1672_U F...	5	0	0		0	5	4	6	5
asv_0108 F_sublineata	5	0	0			5	4	6	5
asv_1440 F_sublineata	10	5	5	5			3	6	4
asv_3502 F_sublineata	7	4	4	4	4		3	4	3
asv_1311 F_sublineata	7	6	6	6	6	6		4	2
asv_3242 F_sublineata	6	5	5	5	5	4	3		1
asv_0036 F_sublineata	6	5	5	5	5	5	3	1	

B



C

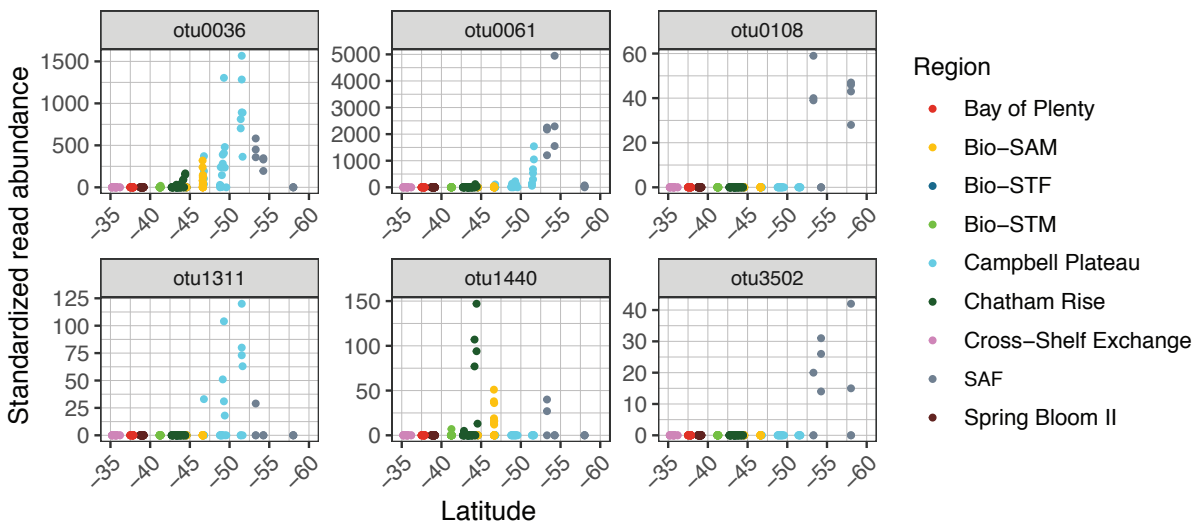


Figure S17. Supplementary Figure S17. (A) Matrix showing the number of nucleotide mismatches in the v4 region of the 18S rRNA between the most abundant ASVs assigned to *F. sublineata* and the annotated sequences for this and *F. kerguelensis* species included in PR2 reference database. (B) mean relative abundance of these ASVs in the euphotic zone of different water masses and regions. (C) distribution of these ASVs abundance (standardized read abundance) in each euphotic sample in relation to latitude and color coded by region.